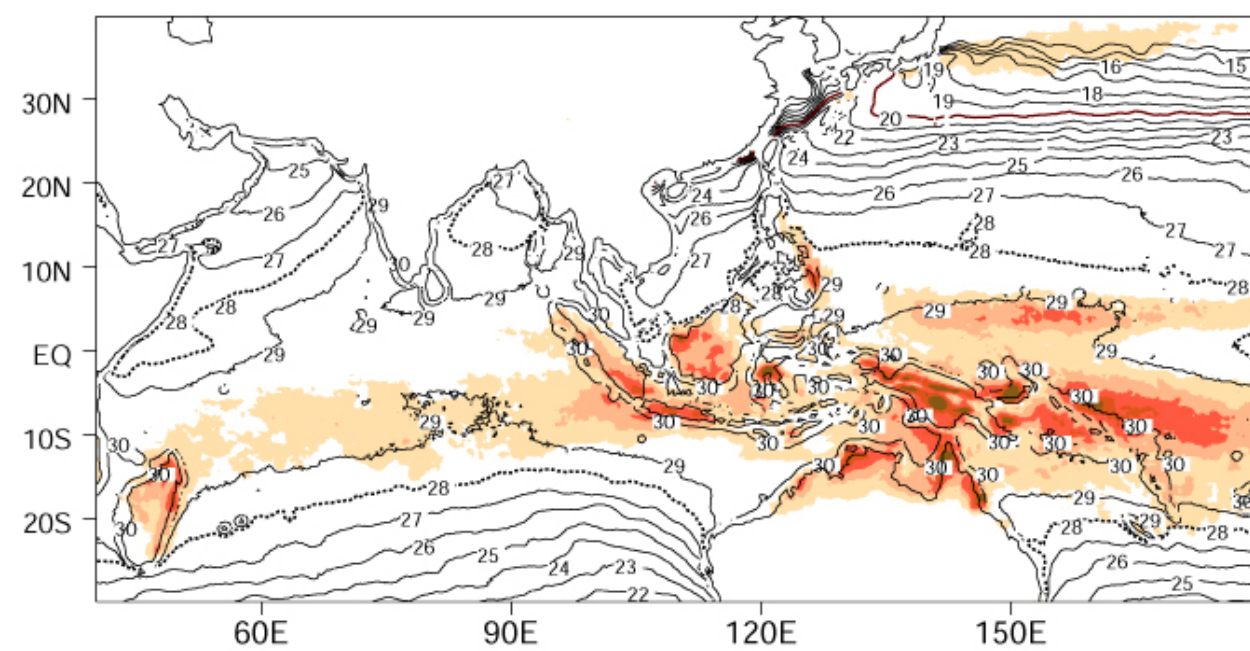


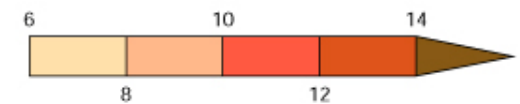
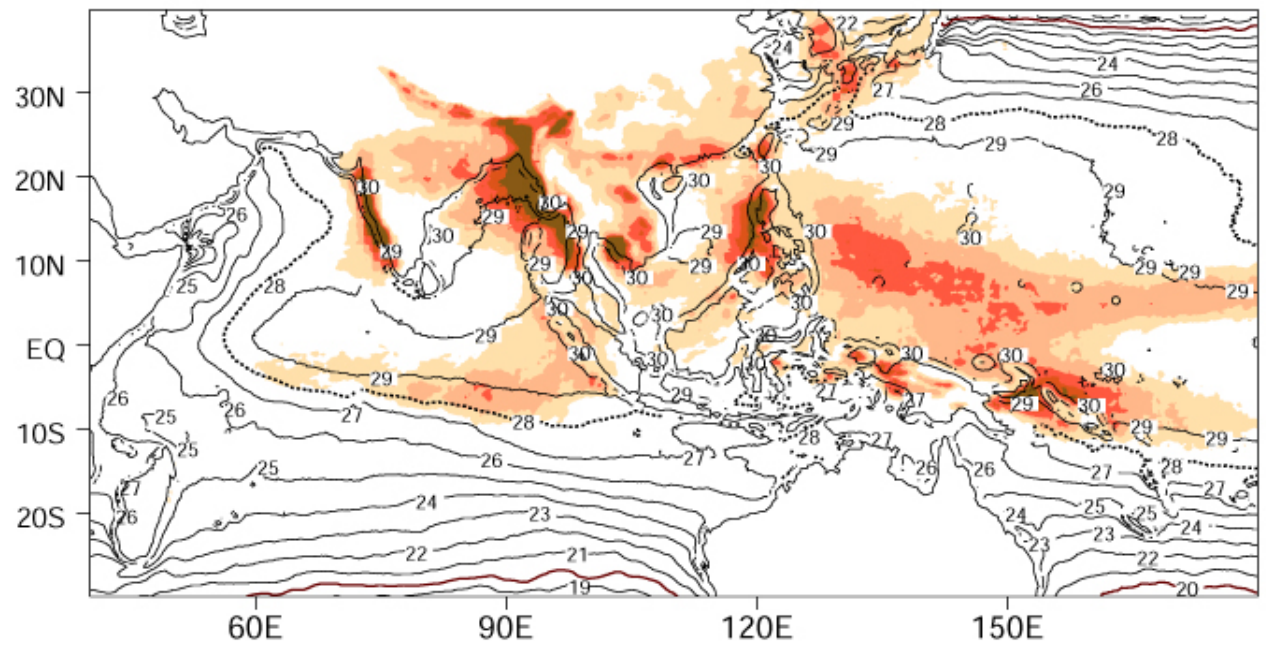
# Dynamics and processes of Asian-Australian Monsoon

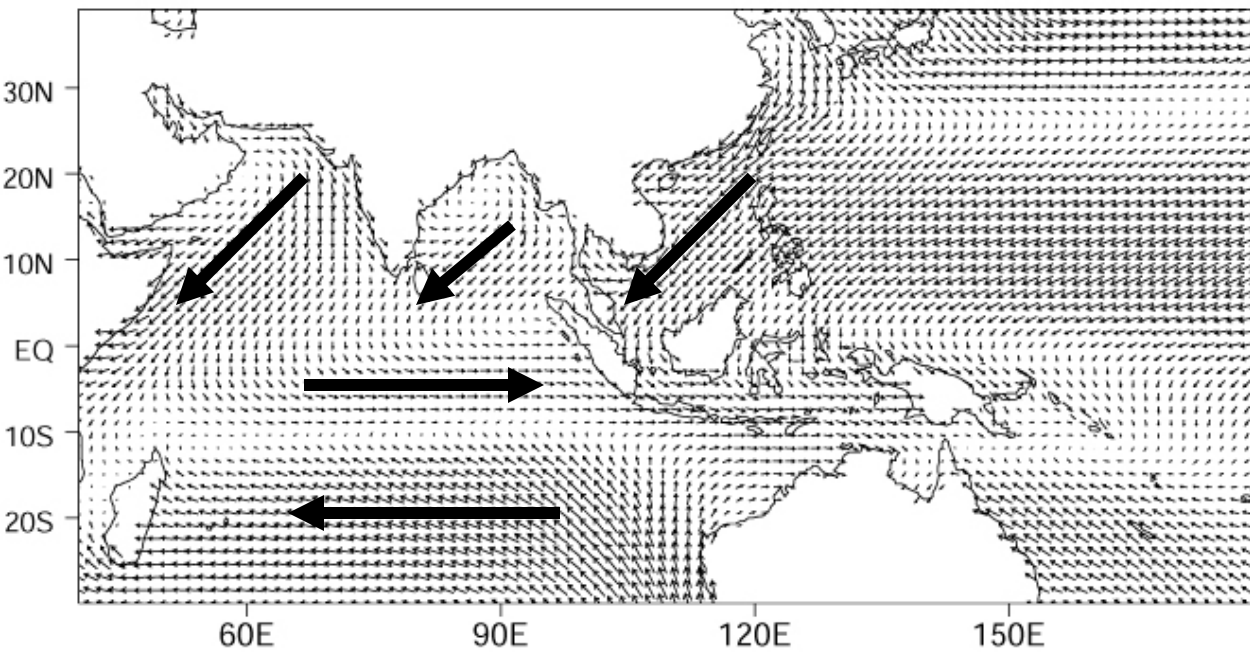
**H. Annamalai**



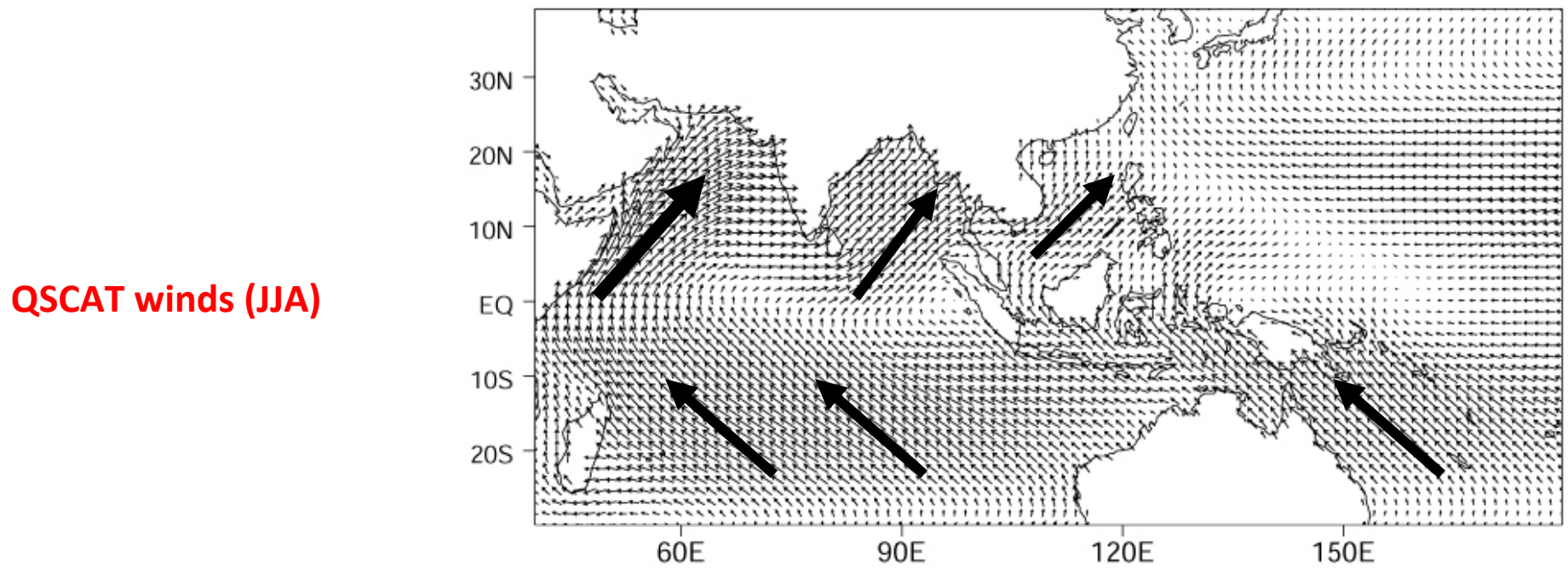
TRMM SST, Pr (DJF)

TRMM SST, Pr (JJA)

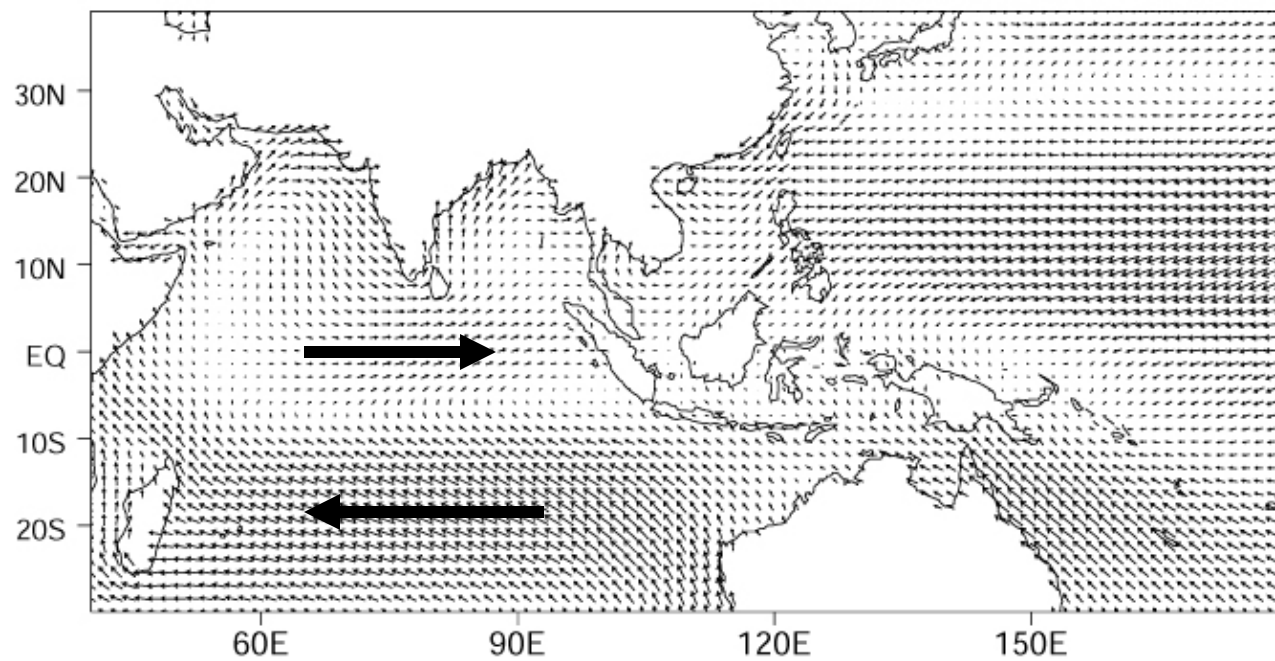




**QSCAT winds (DJF)**

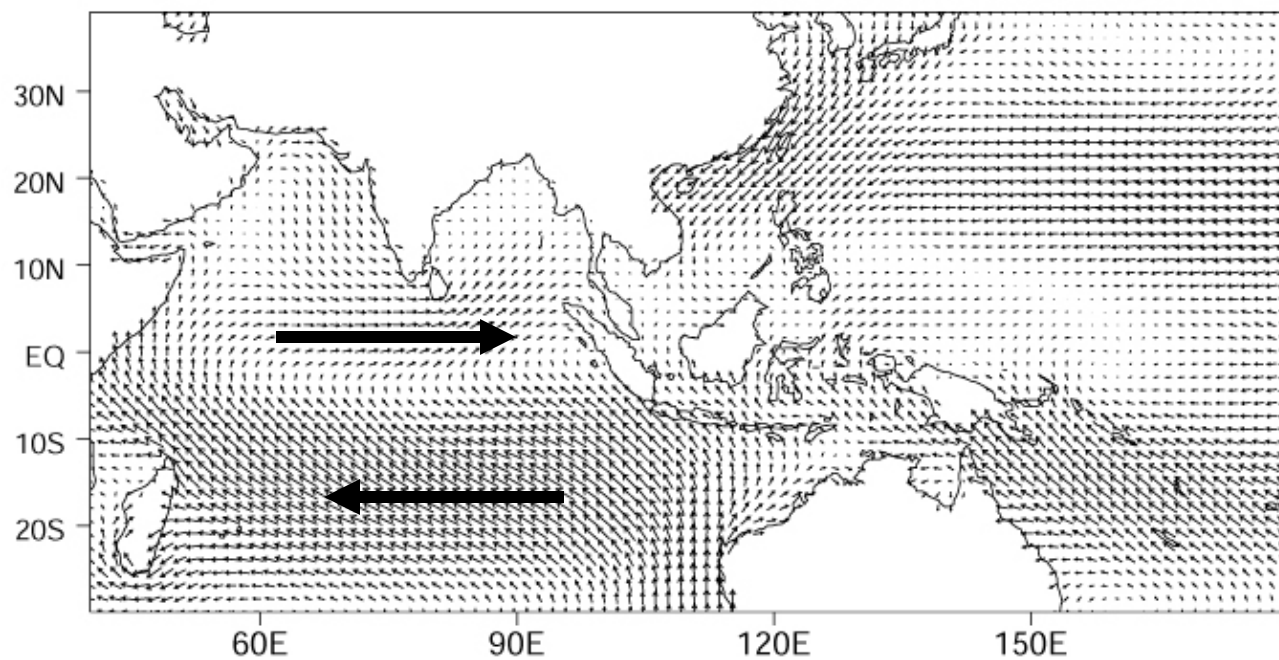


**QSCAT winds (JJA)**



QSCAT winds (MAM)

QSCAT winds (SON)





# Why are there monsoons?

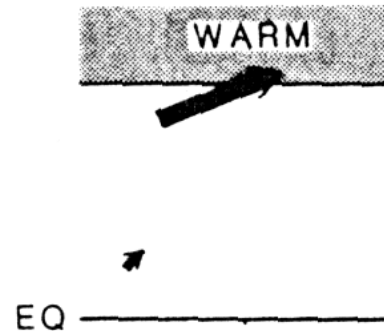
## *Monsoon Annual Cycle – Robust ?*

- land-sea thermal contrast ( $T_s/P_s$  gradient)
- Orography (Tibet – elevated heat source)  
(East Africa – frictional force  
cross-equatorial flow)
- Earth's rotation (Coriolis force)
- Moisture from the tropical Indian Ocean

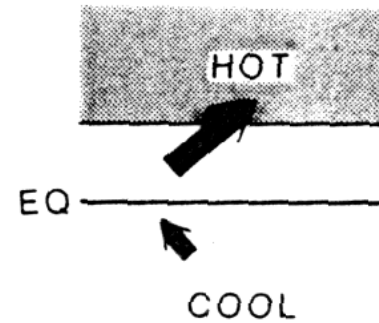
(a)  
EARTH'S  
ROTATION  
UNIMPORTANT



(b)  
MID-LATITUDE  
'MONSOON'

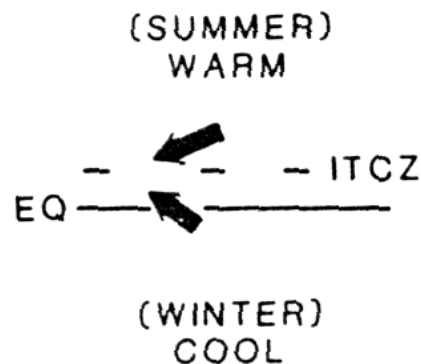


(c)  
NEAR-EQUATORIAL  
MONSOON

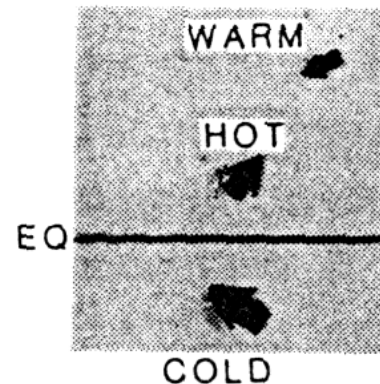


**“land – shaded”**

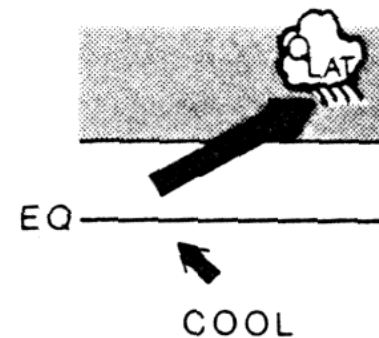
(d)  
NO LAND



(e)  
NO OCEAN



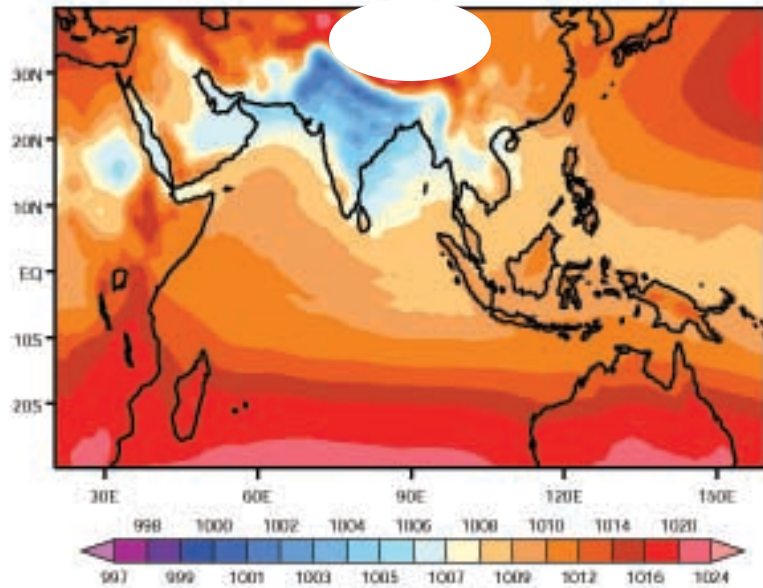
(f)  
LATENT  
HEATING



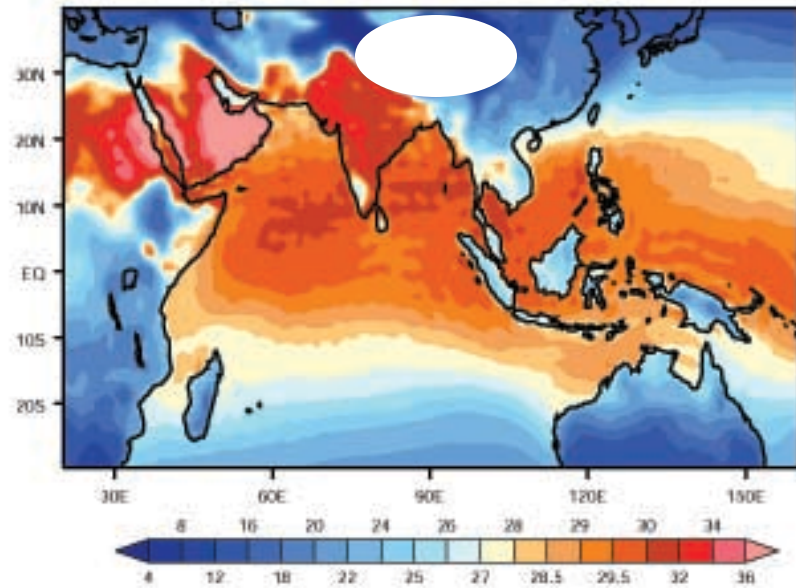
**Young, 1987 – a chapter in Fein and Stephens (1987)**  
**Hoskins and Wang (2004) – a chapter in Bin Wang (2004) –**  
**Vorticity and thermodynamics perspective**

## Patterns for May

### Mean Sea Level Pressure



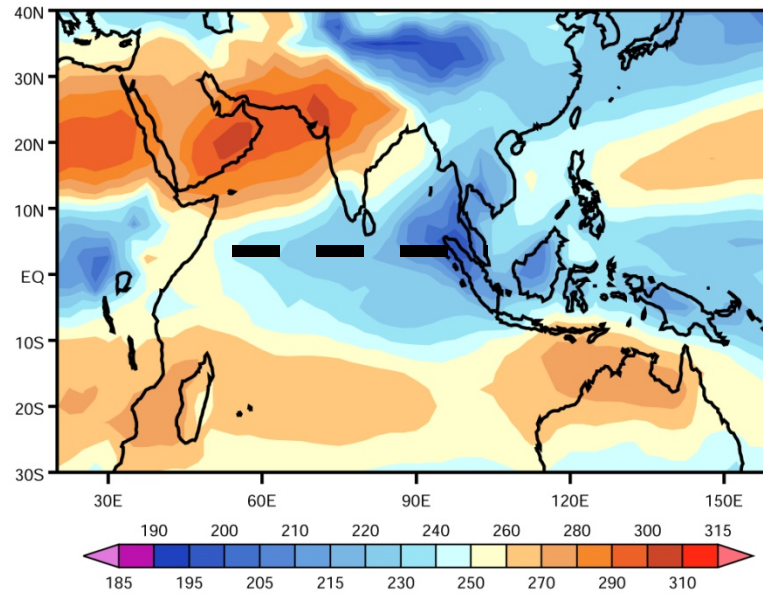
### Surface Temperature



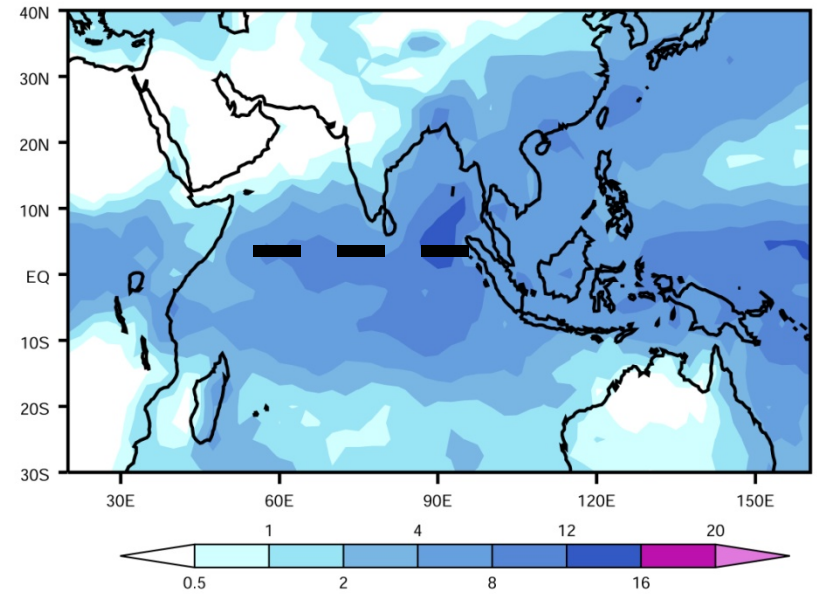
- (i) Slowly, due to surface heating and attendant sensible heat flux, shallow low-pressure forms over land
- (ii) This low, gradually generates low-level wind inflow that draws moisture from the ocean to the south
- (iii) The moisture is carried upward by the convection, condenses, and warms the upper troposphere, where an anticyclone forms

## Patterns for May

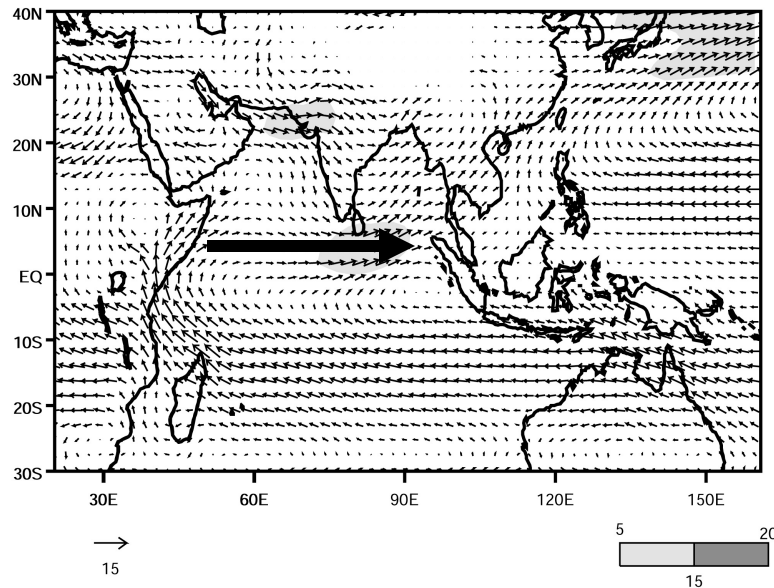
### OLR



### Precipitation (Xie-Arkin)



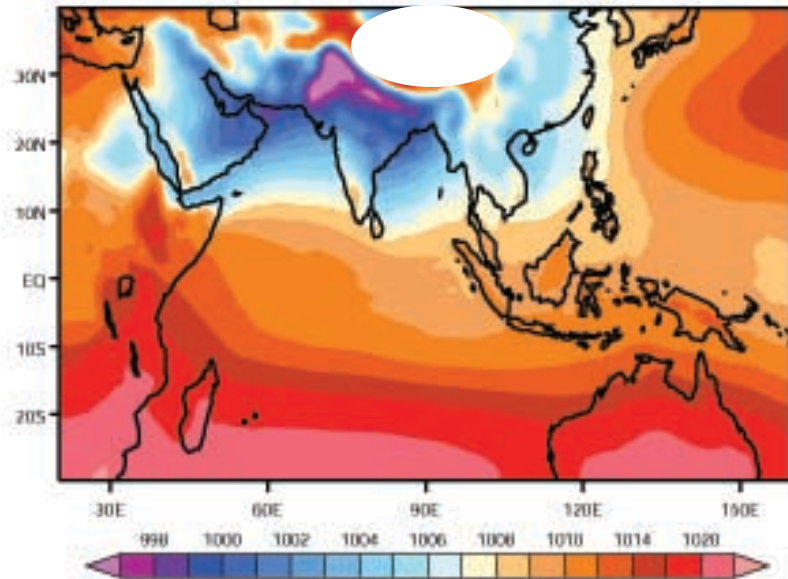
### 850hPa



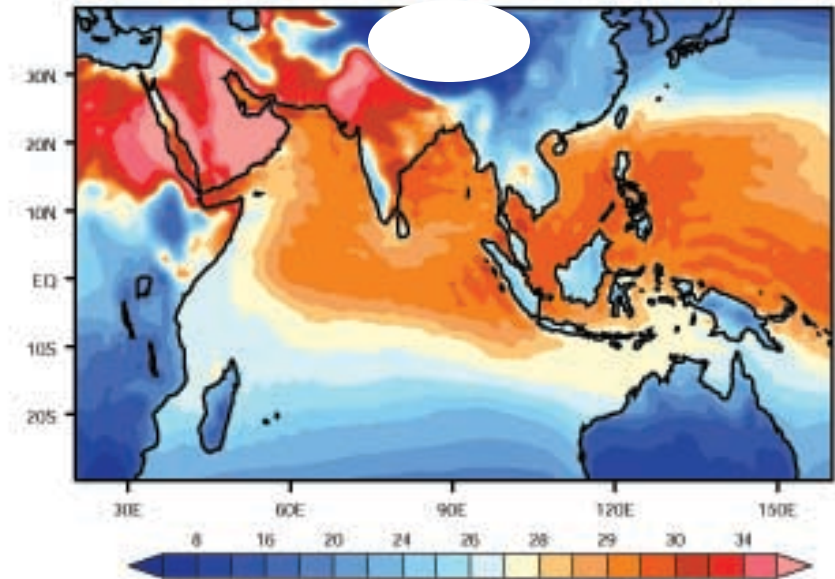


## Patterns for June

### Mean Sea Level Pressure

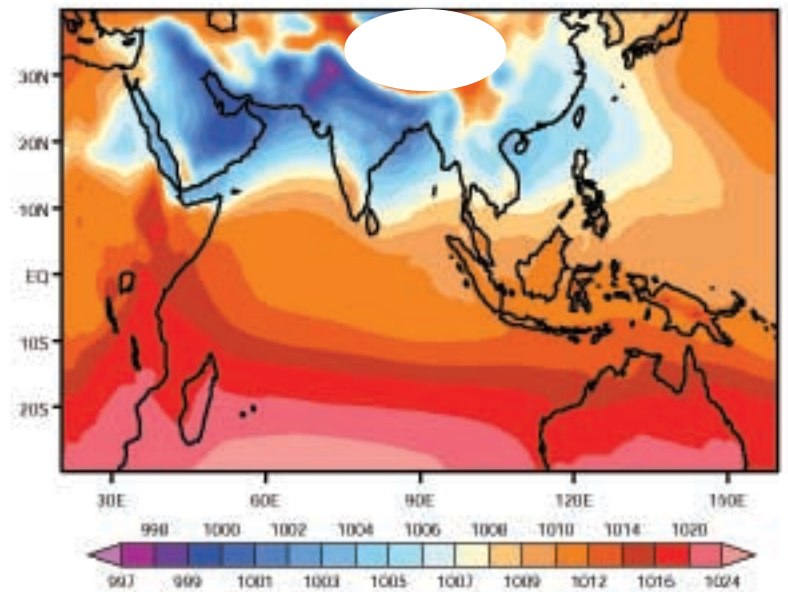


### Surface Temperature

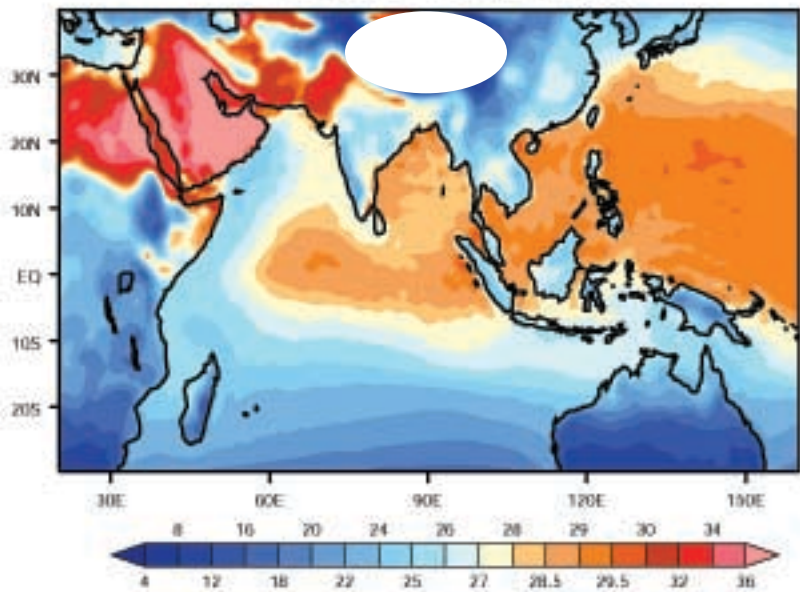


## Patterns for August

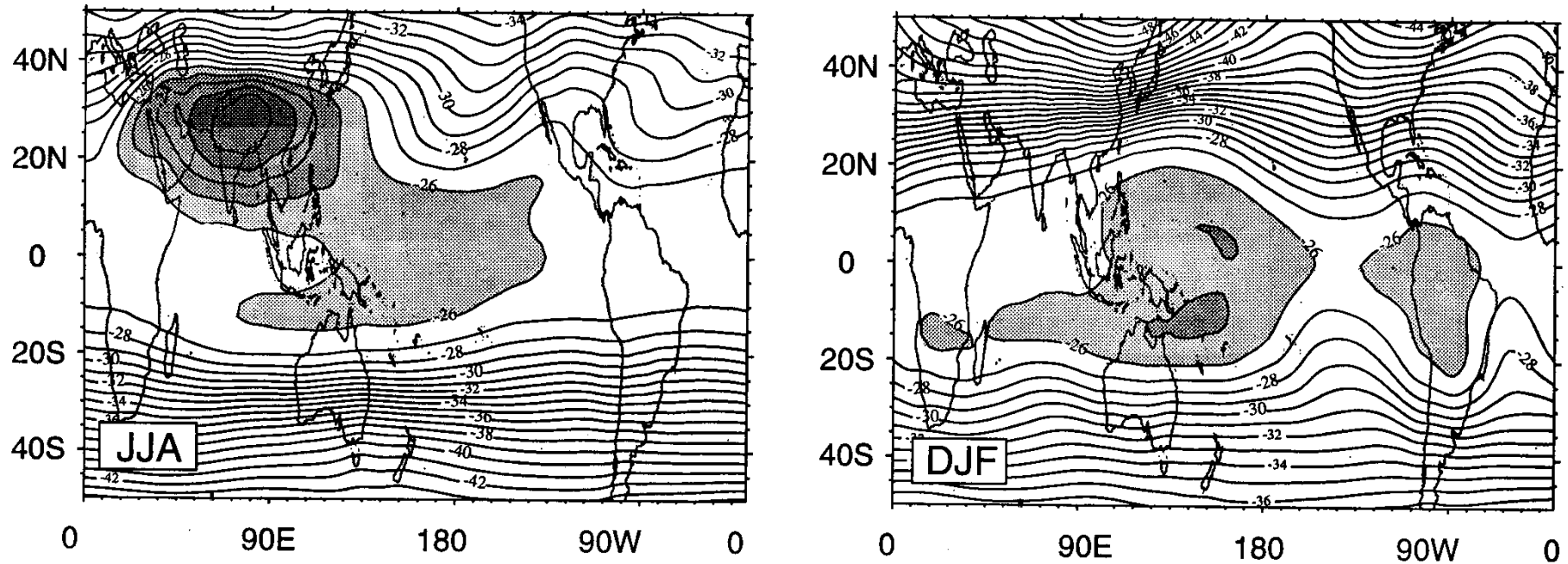
### Mean Sea Level Pressure



### Surface Temperature



## Mean Upper-Tropospheric Temperature: 200-500 mbar

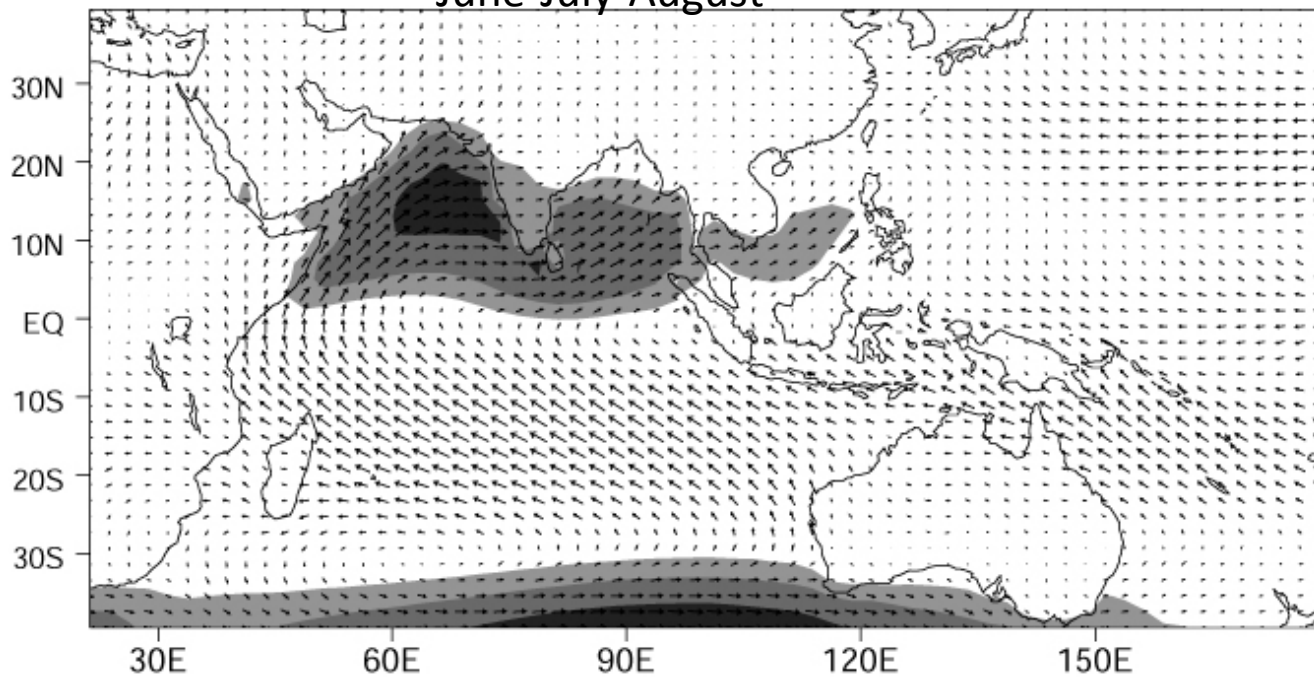


**Figure 6a.** Mean upper tropospheric (200–500 mbar) temperature (degrees Celsius) for the boreal summer (JJA), and boreal winter (DJF), averaged between 1979 and 1992. The boreal summer plot is based on calculations first made by *Li and Yanai* [1996]. Mean columnar temperatures warmer than  $-25^{\circ}\text{C}$  are shaded.

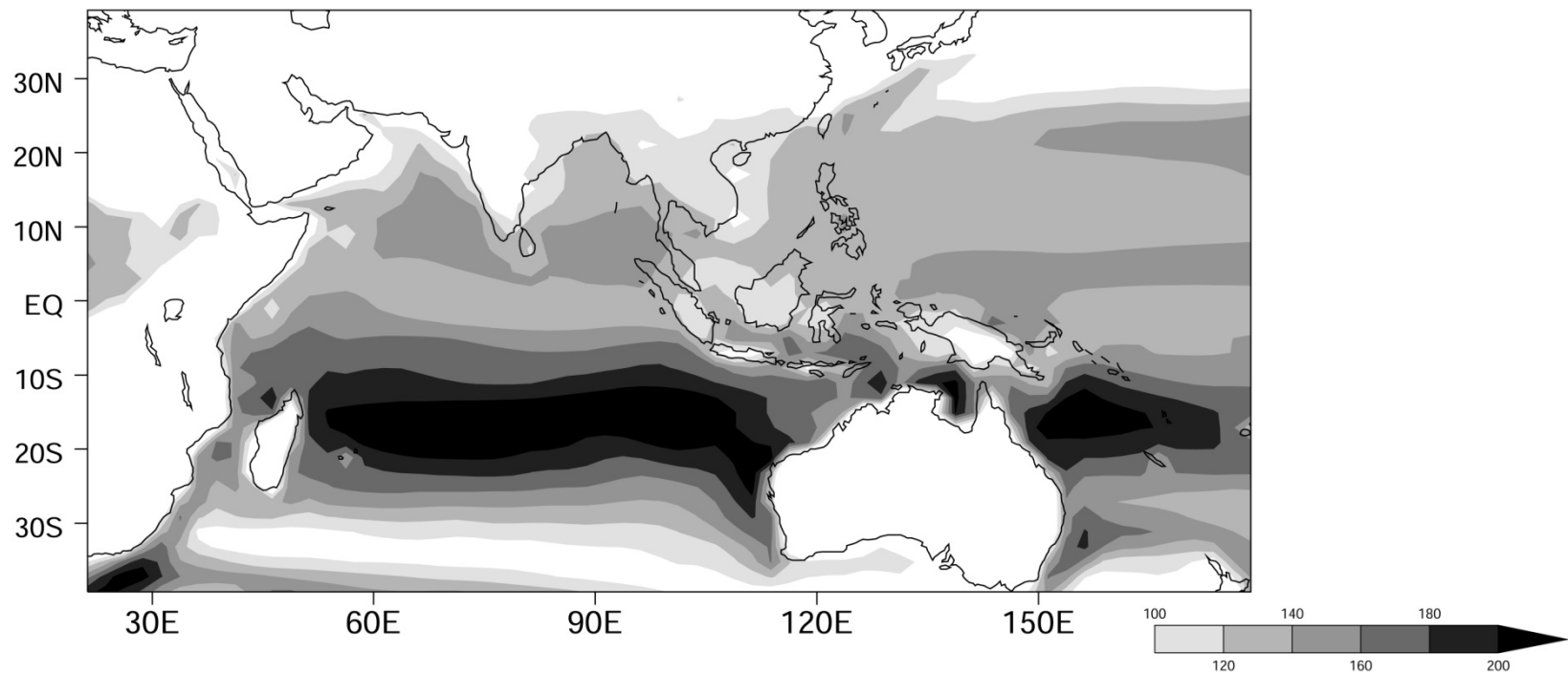
Webster et al. (1998, J. Geophys. Res)

June-July-August

10 m wind

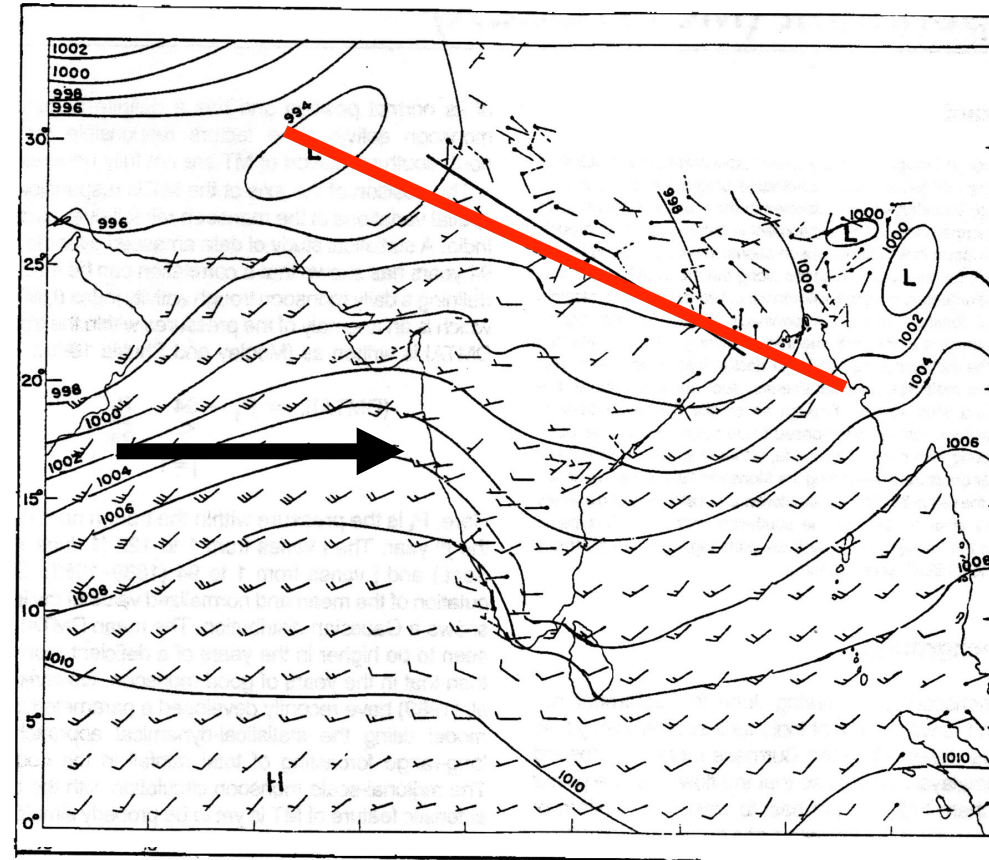
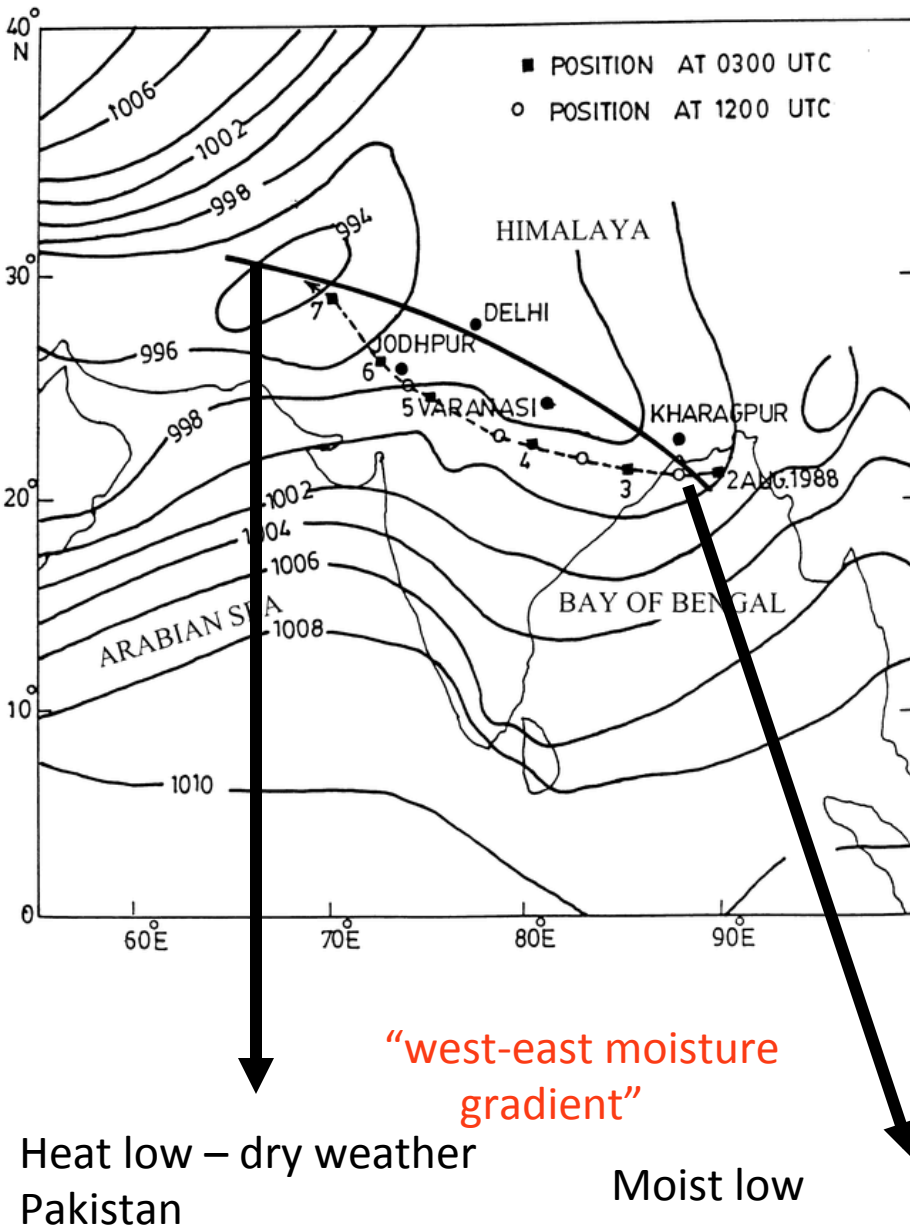


Evaporation  
(mm)





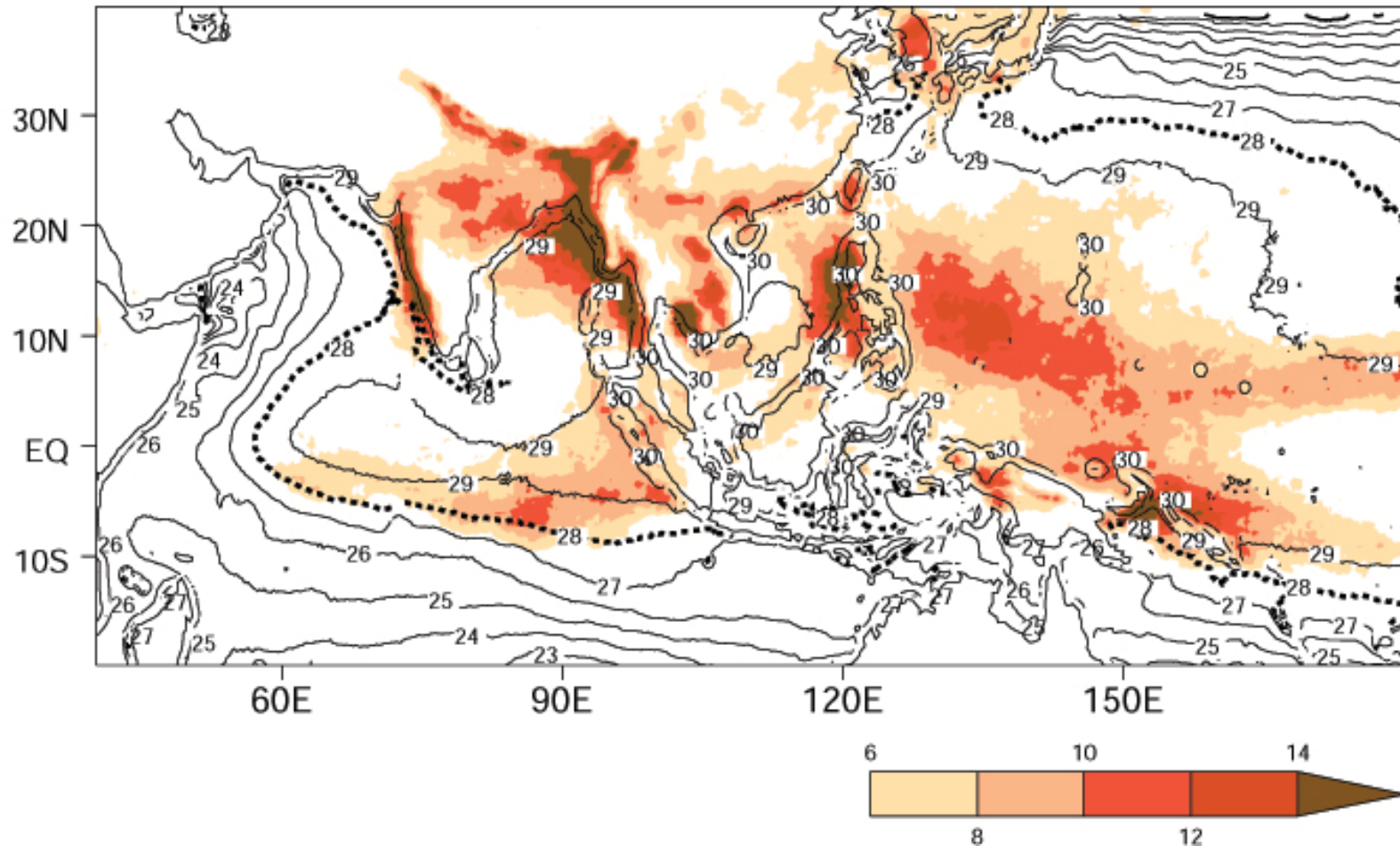
# Normal position of the monsoon trough



"analogous to the equatorial trough"  
 "seasonal displacement of the TCZ"  
 "shear vorticity – Ekman pumping"

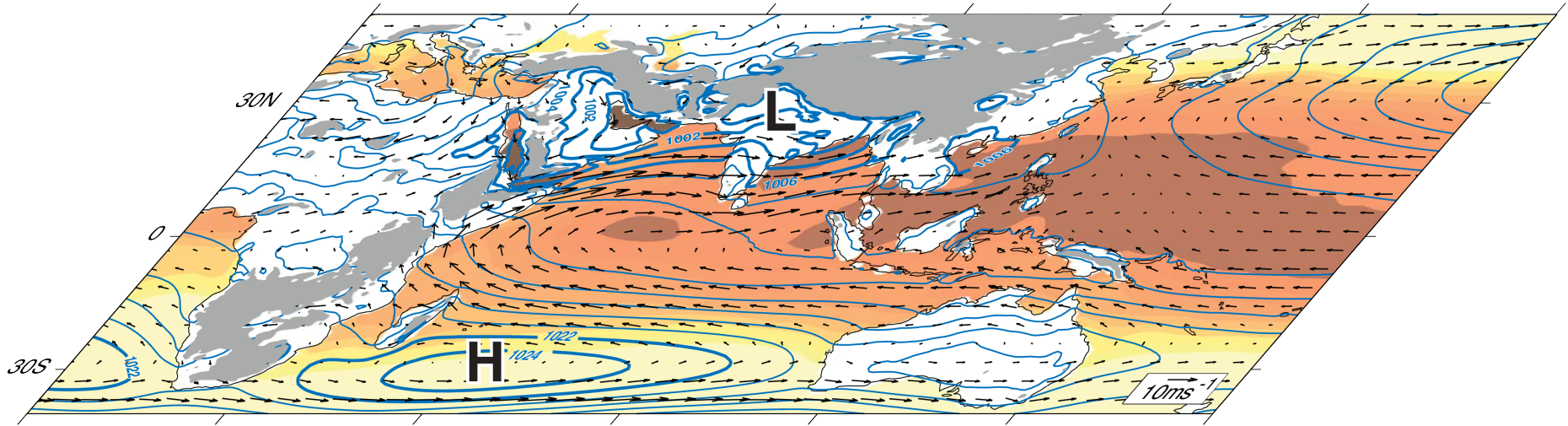
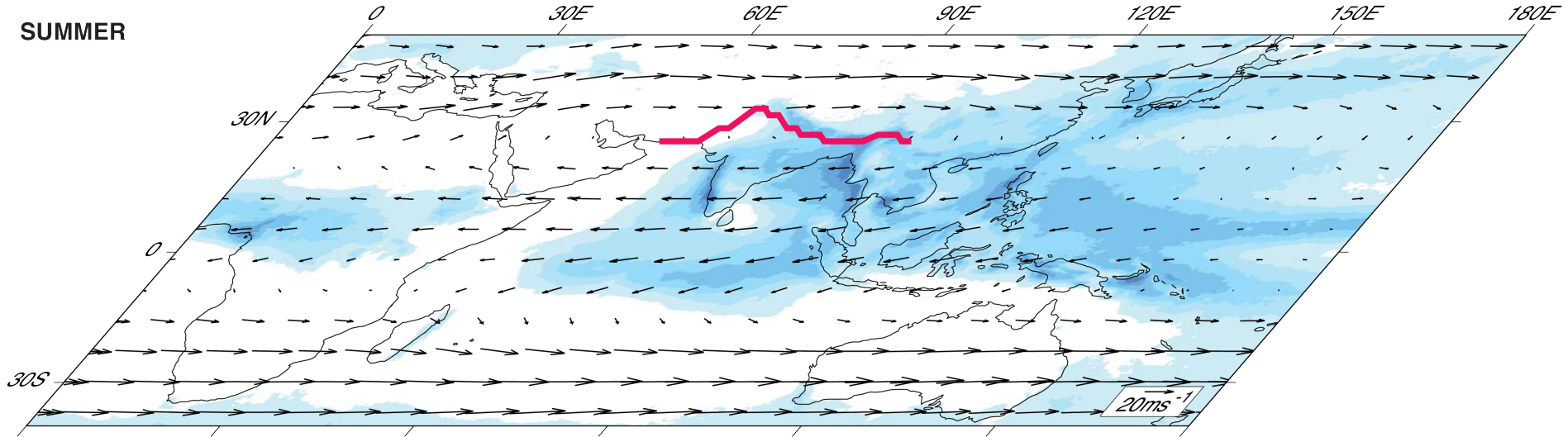


# JJAS – Precipitation and SST Climatology

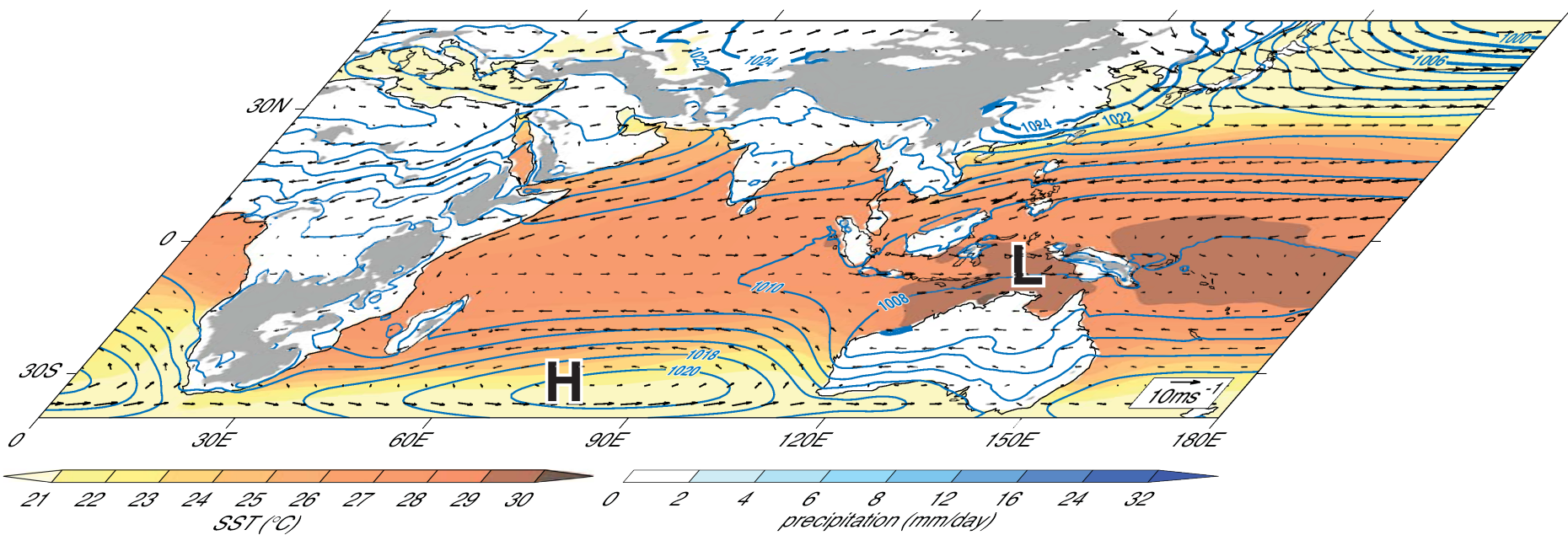
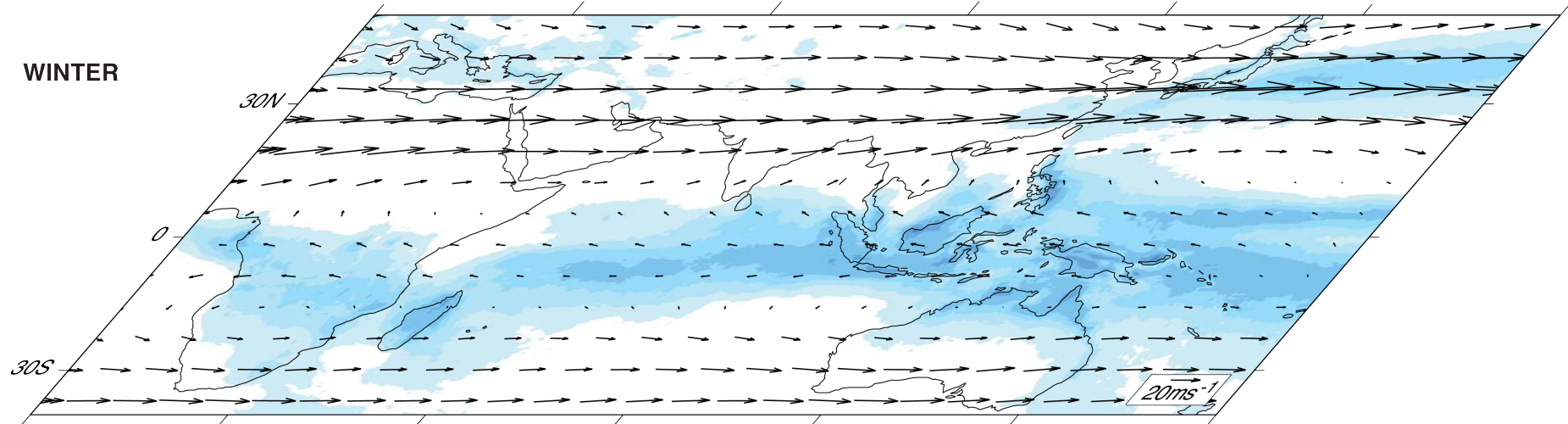


**“Three regional heat sources”**

SUMMER

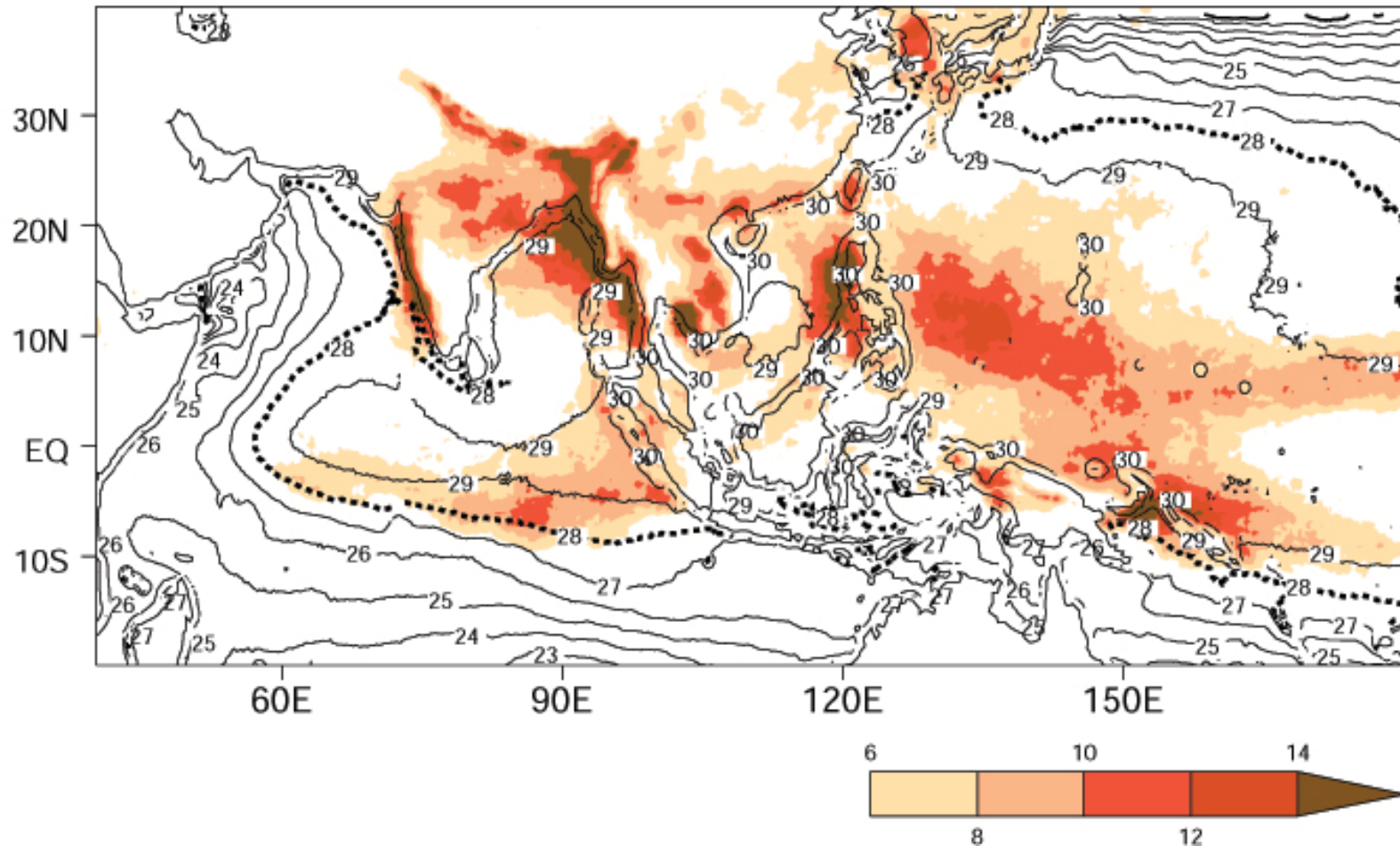


WINTER





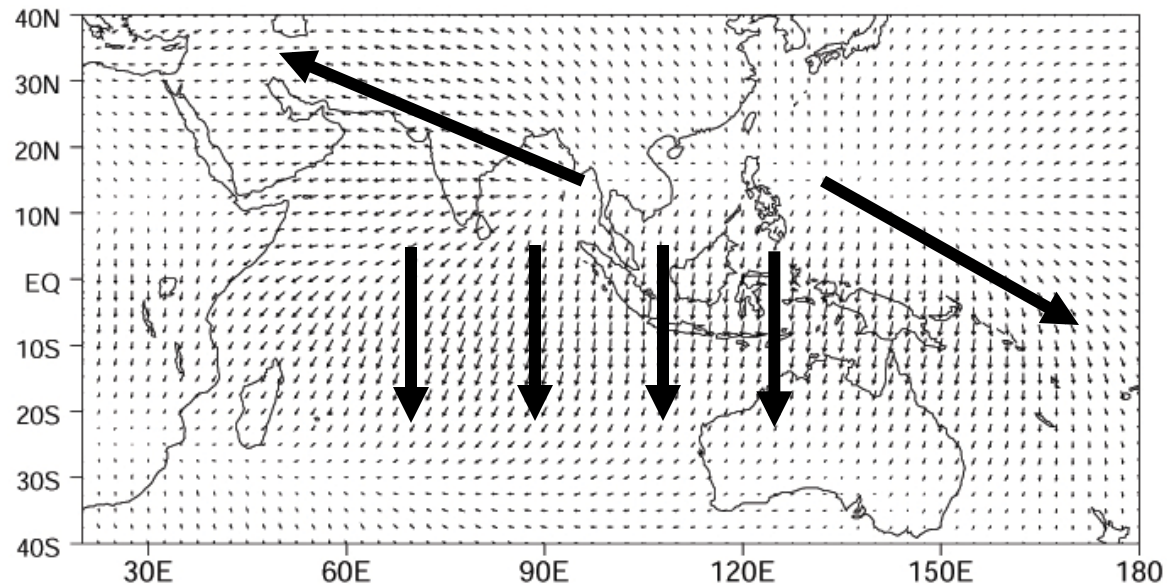
# JJAS – Precipitation and SST Climatology



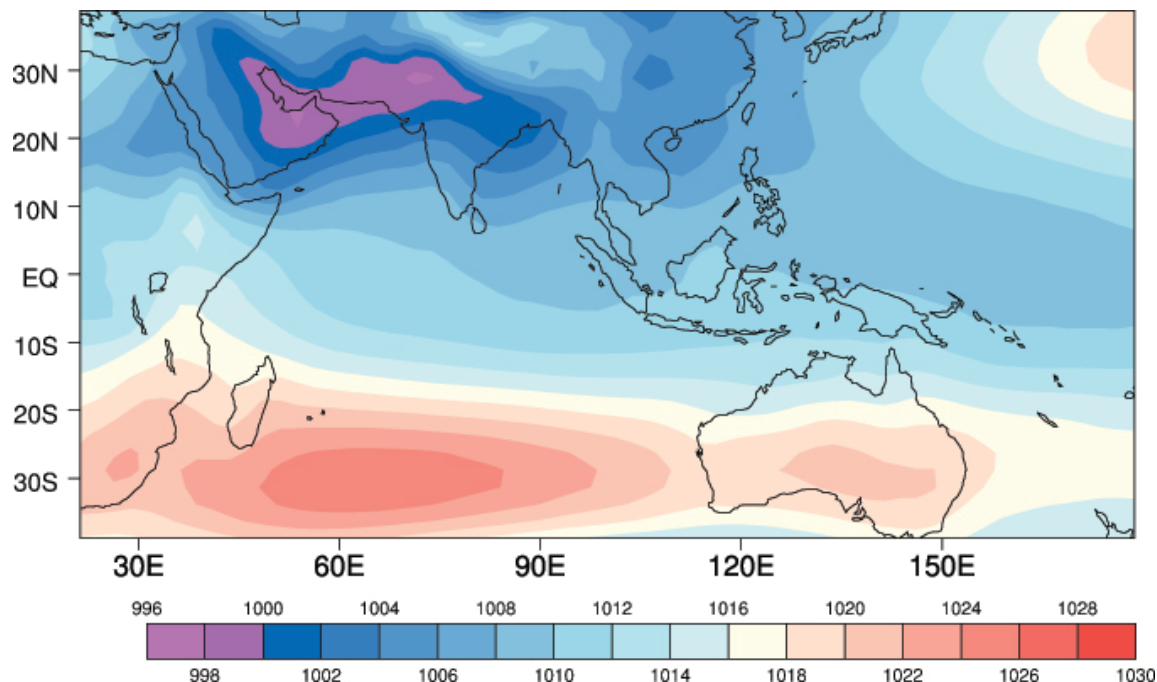
**“Three regional heat sources”**

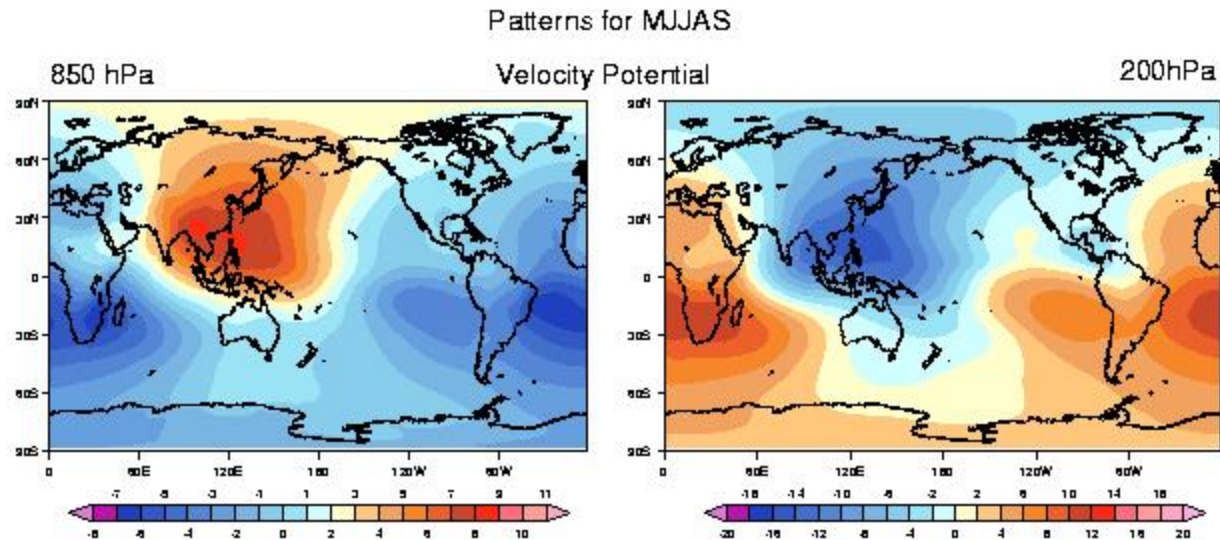


JJAS Climatology – divergent wind at 200hPa



JJAS Climatology – Sea Level Pressure

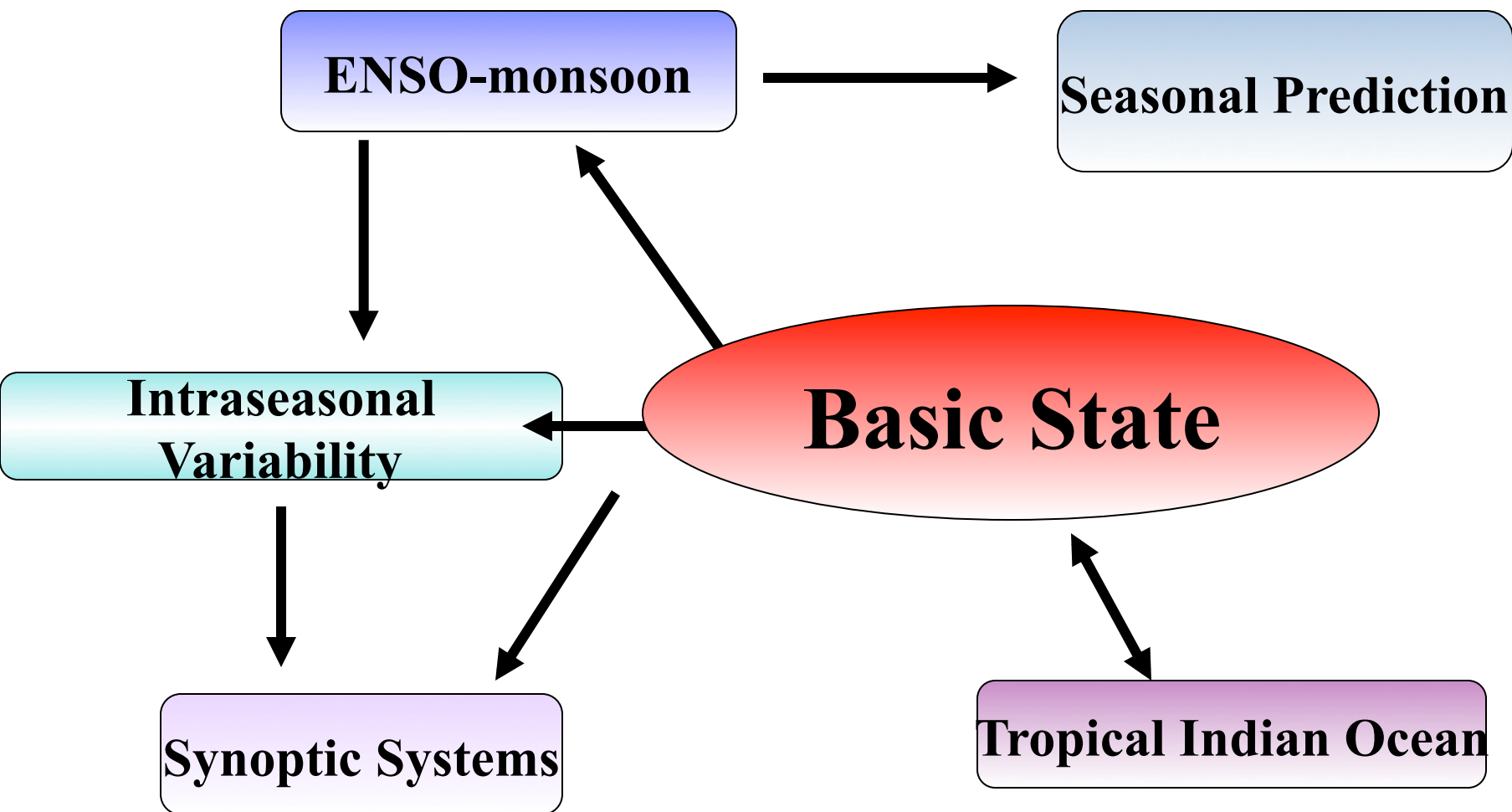




**(i) Lateral monsoon (Hadley) Circulation**

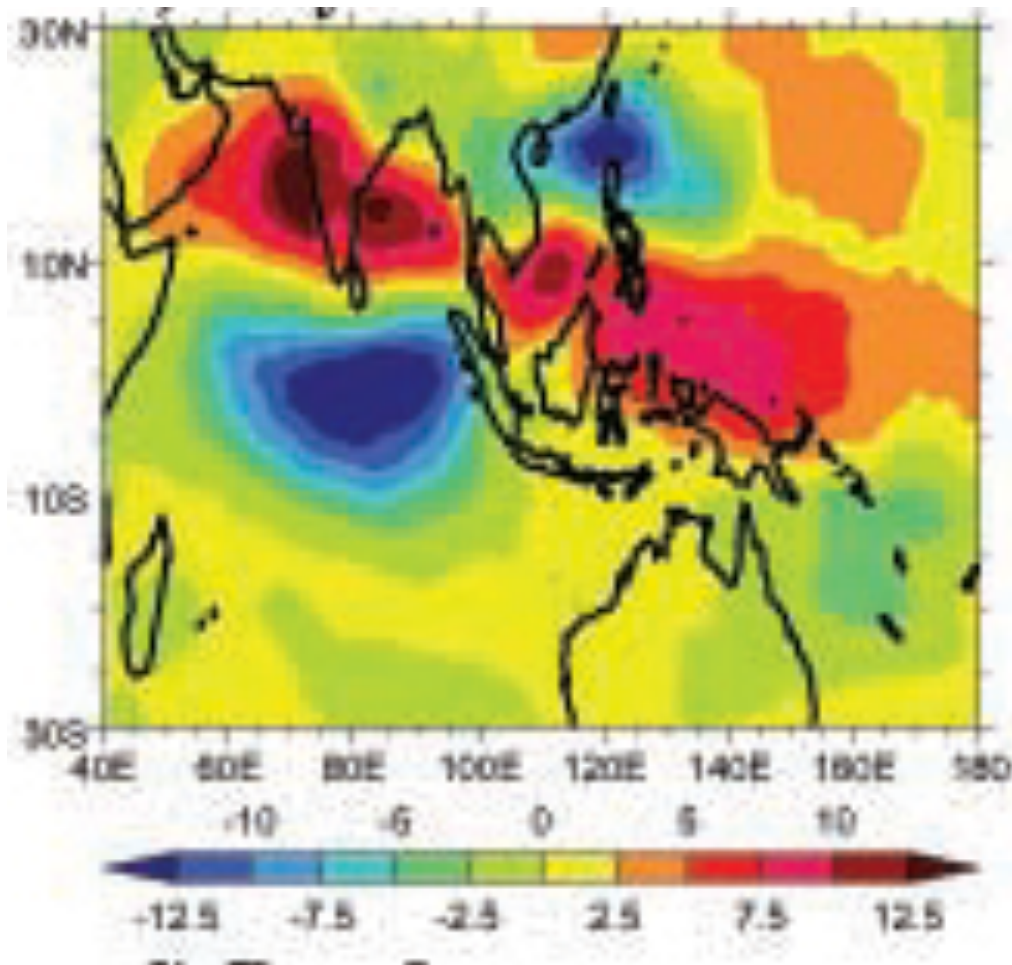
**(ii) Planetary Walker Circulation – ENSO**

**(iii) Transverse monsoon circulation**



**“Importance of understanding and simulating the mean monsoon – basic state”**

# Observed Boreal Summer ISV



“internal dynamics”

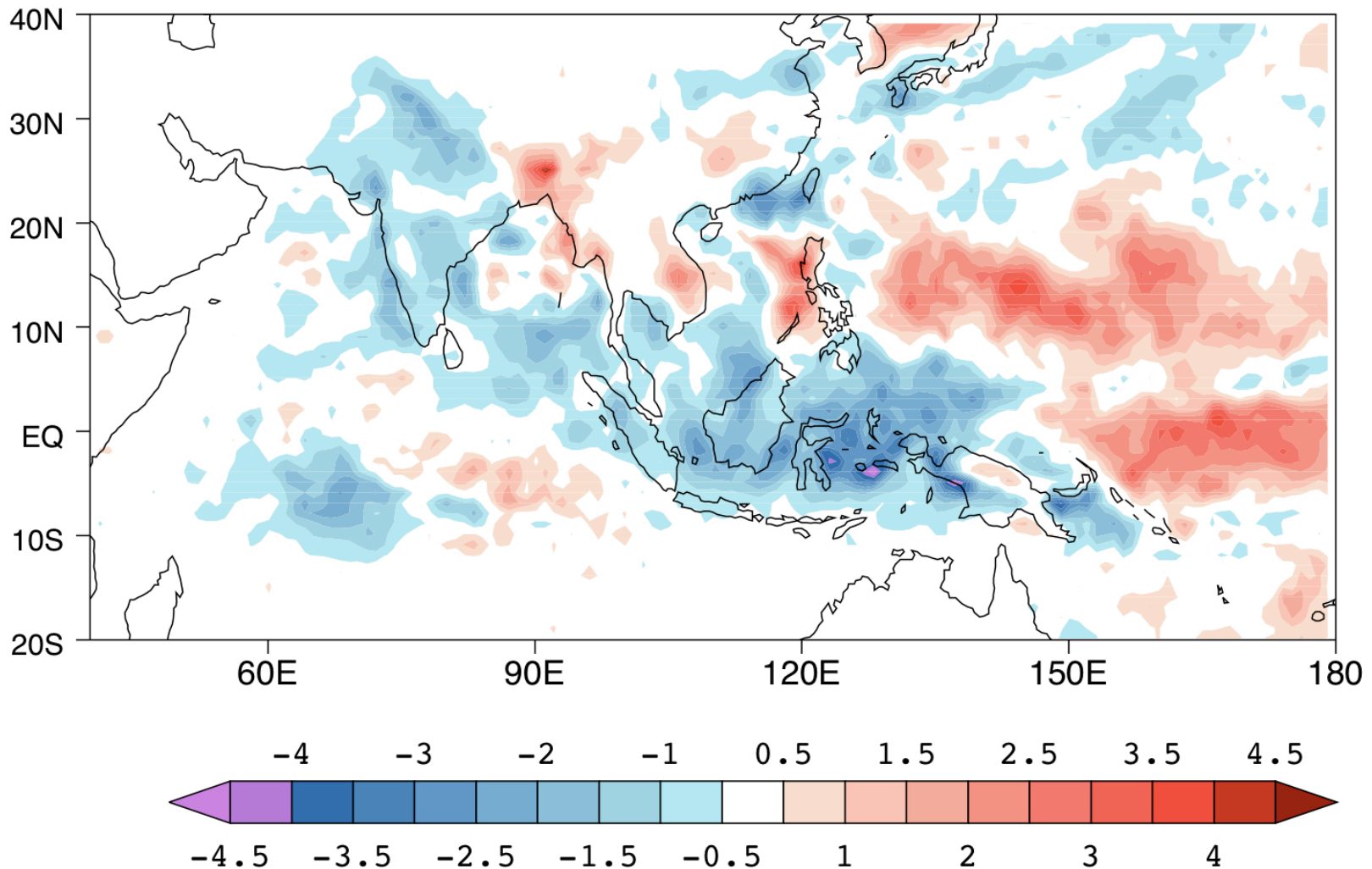
OLR anomalies  $\text{W/m}^2$

Annamalai and Sperber (2005, JAS)  
Lau and Chan (1986, JAS)

“one-phase of the ISV”



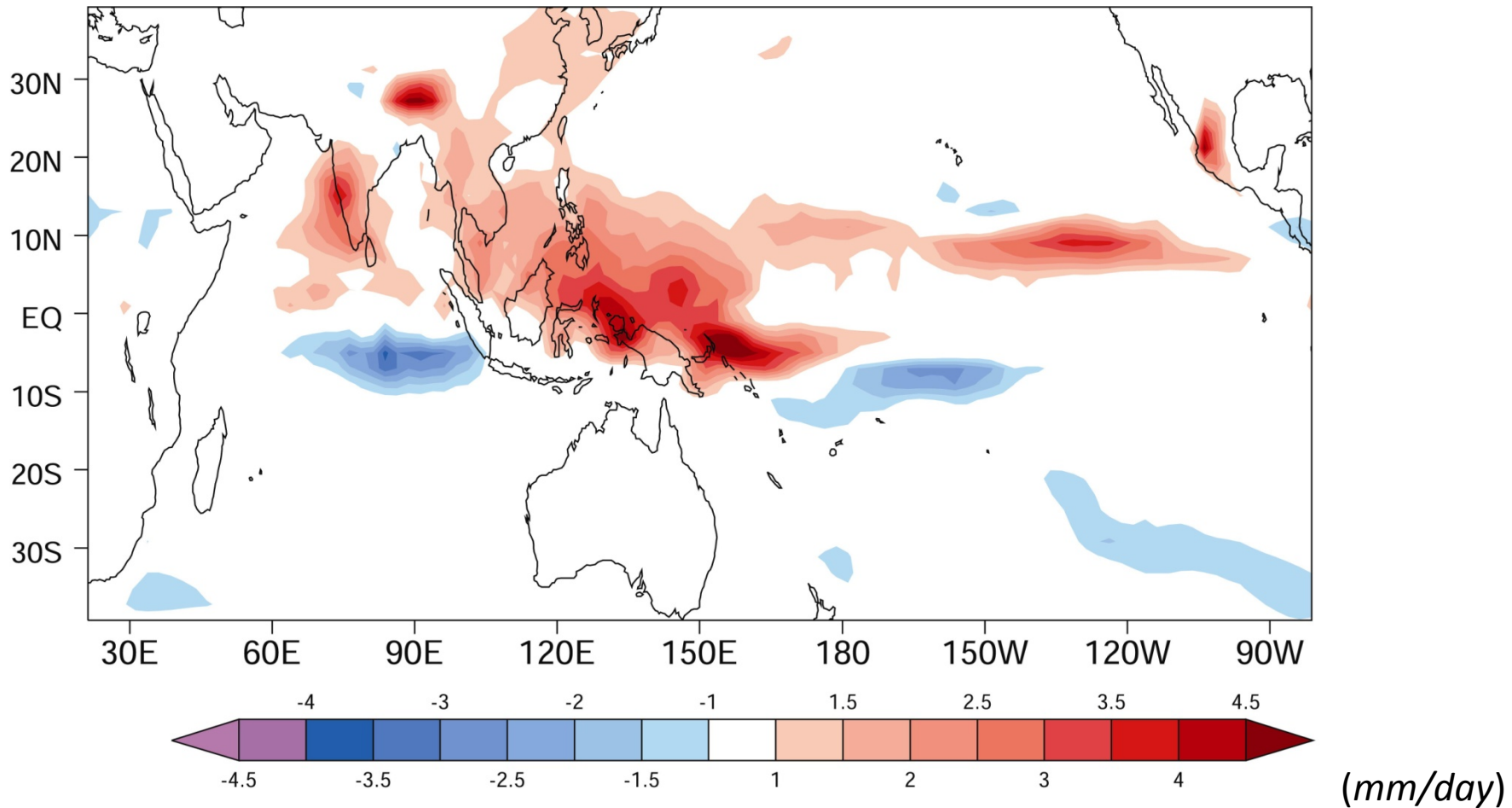
## JJAS rainfall anomalies (2002/04/09) - TRMM



“Boundary forcing”

# JJAS Precipitation response in CM\_2.1

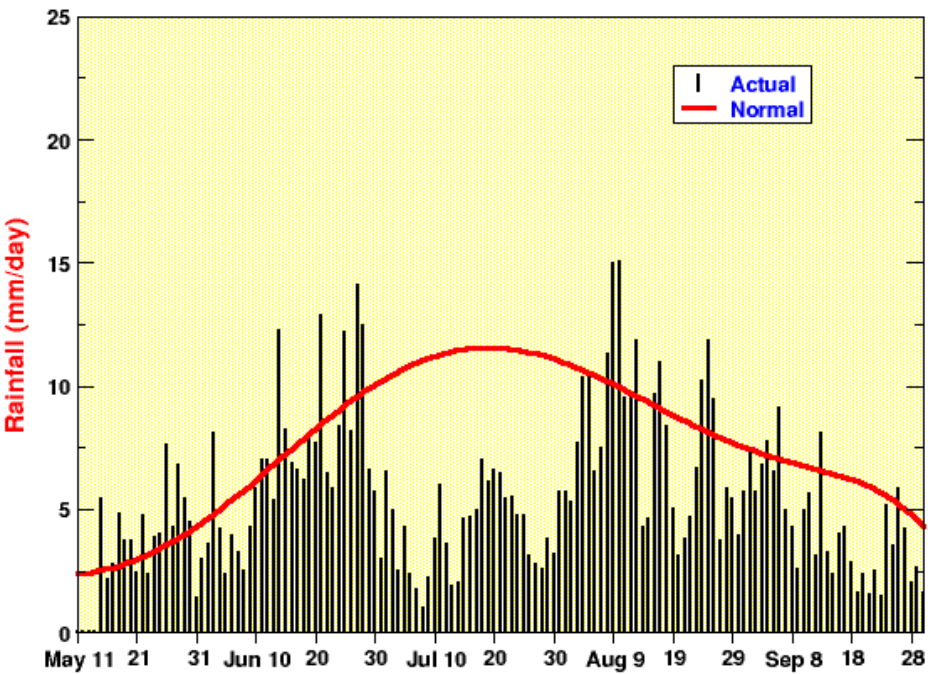
*4xCO<sub>2</sub> minus 20c3m*



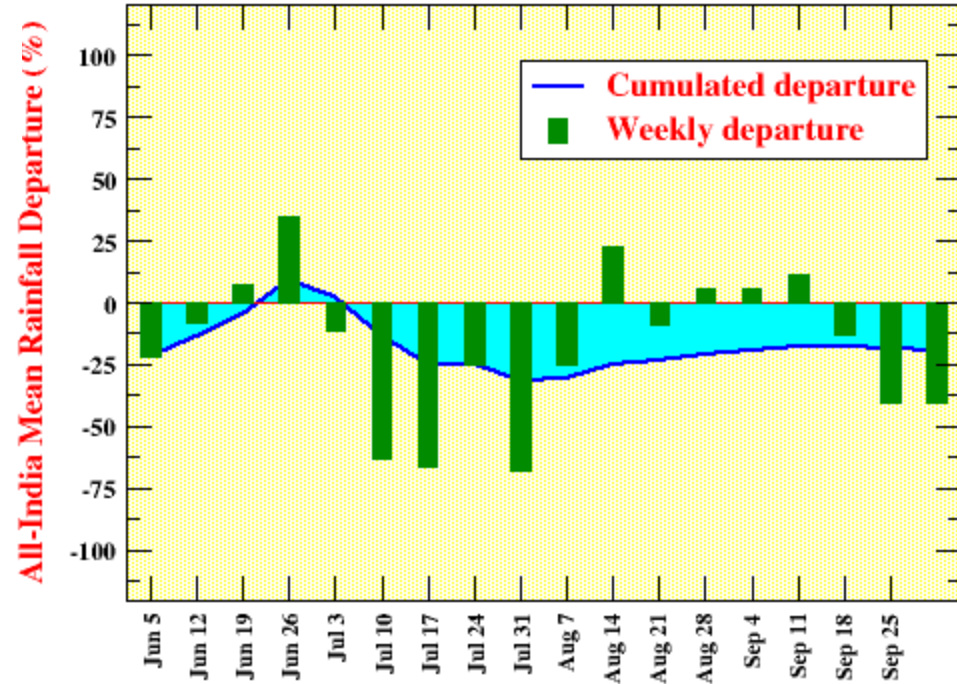
## Overarching Hypothesis

*Interaction between equatorial waves and moist physics  
needs to be understood for attributing the causes for  
precipitation anomalies over “mean ascent” regions*

## Daily Observed rainfall over India - 2002

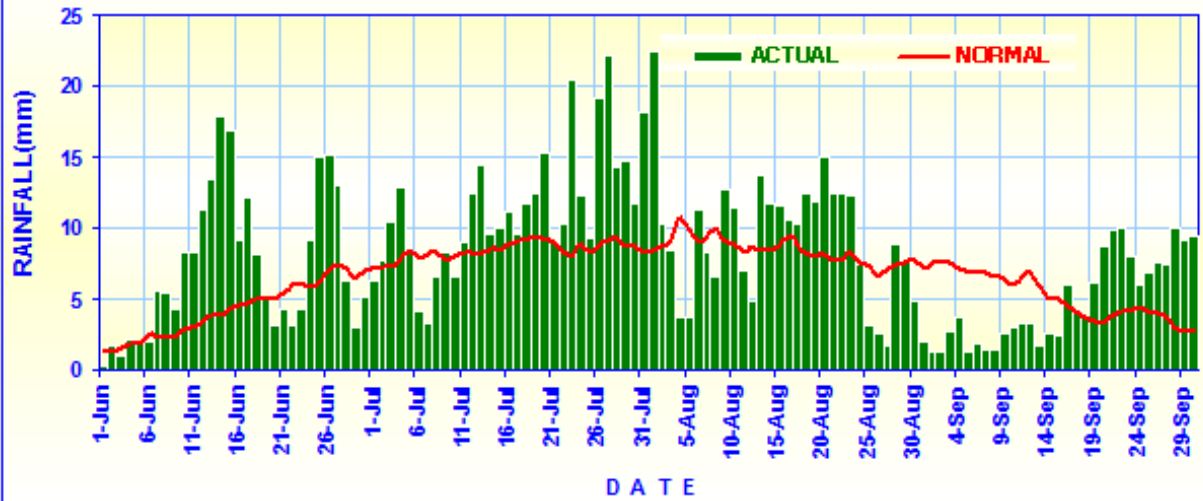


## Weekly rainfall departure

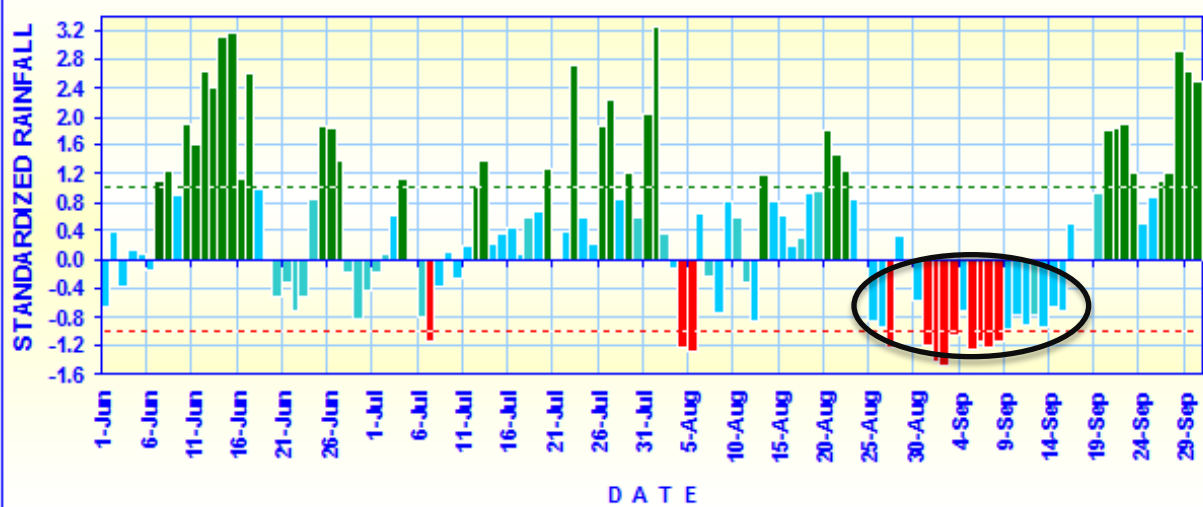




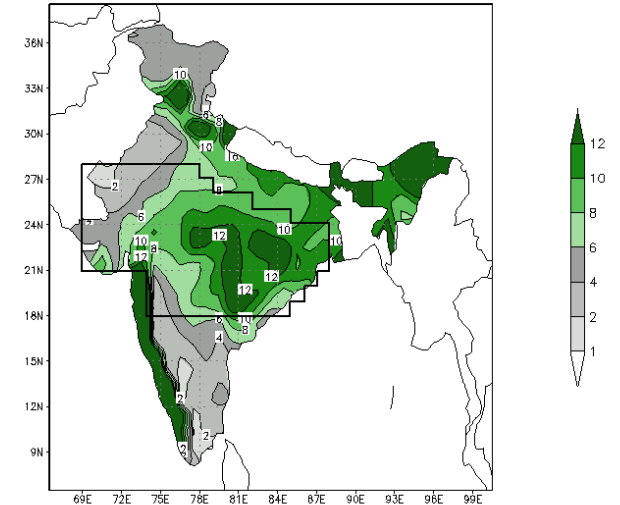
### AVERAGE RAINFALL (mm) OVER THE MONSOON ZONE (2013)

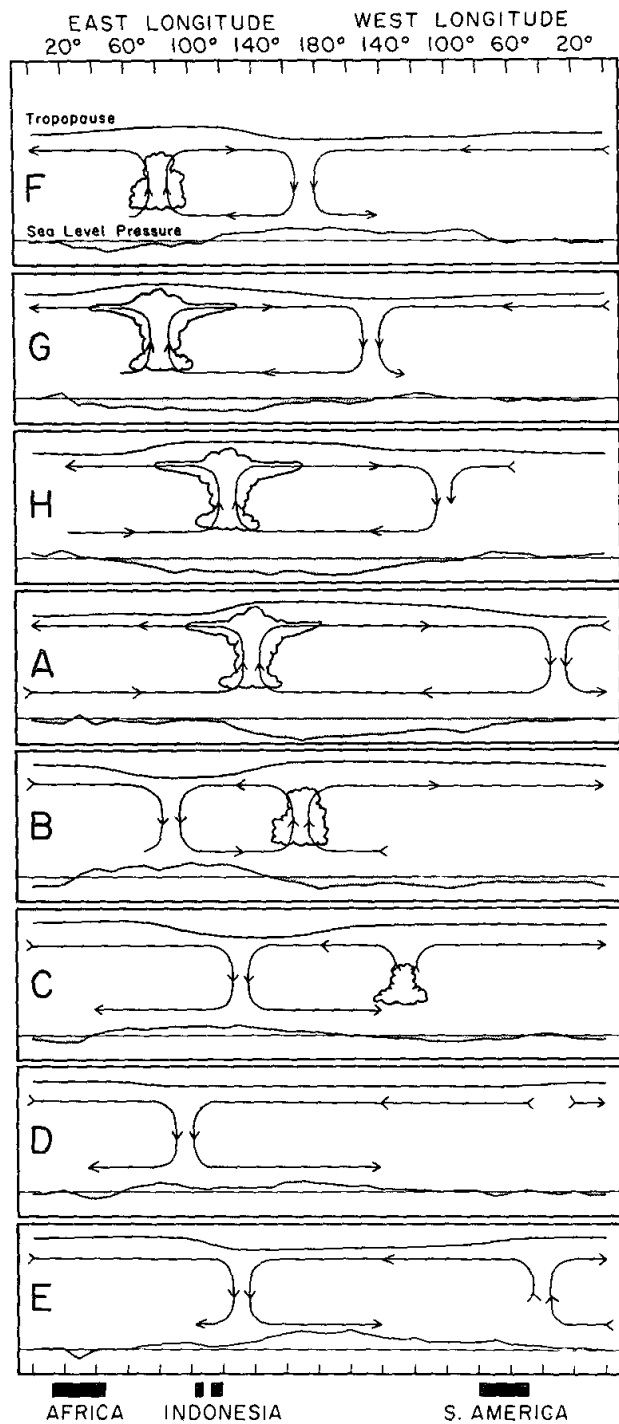


### STANDARDISED RAINFALL FOR MONSOON ZONE (2013)



### MEAN SEASONAL RAINFALL FOR JUL+Aug (mm/day)





**Madden and Julian (1971, 1972, 1994)**

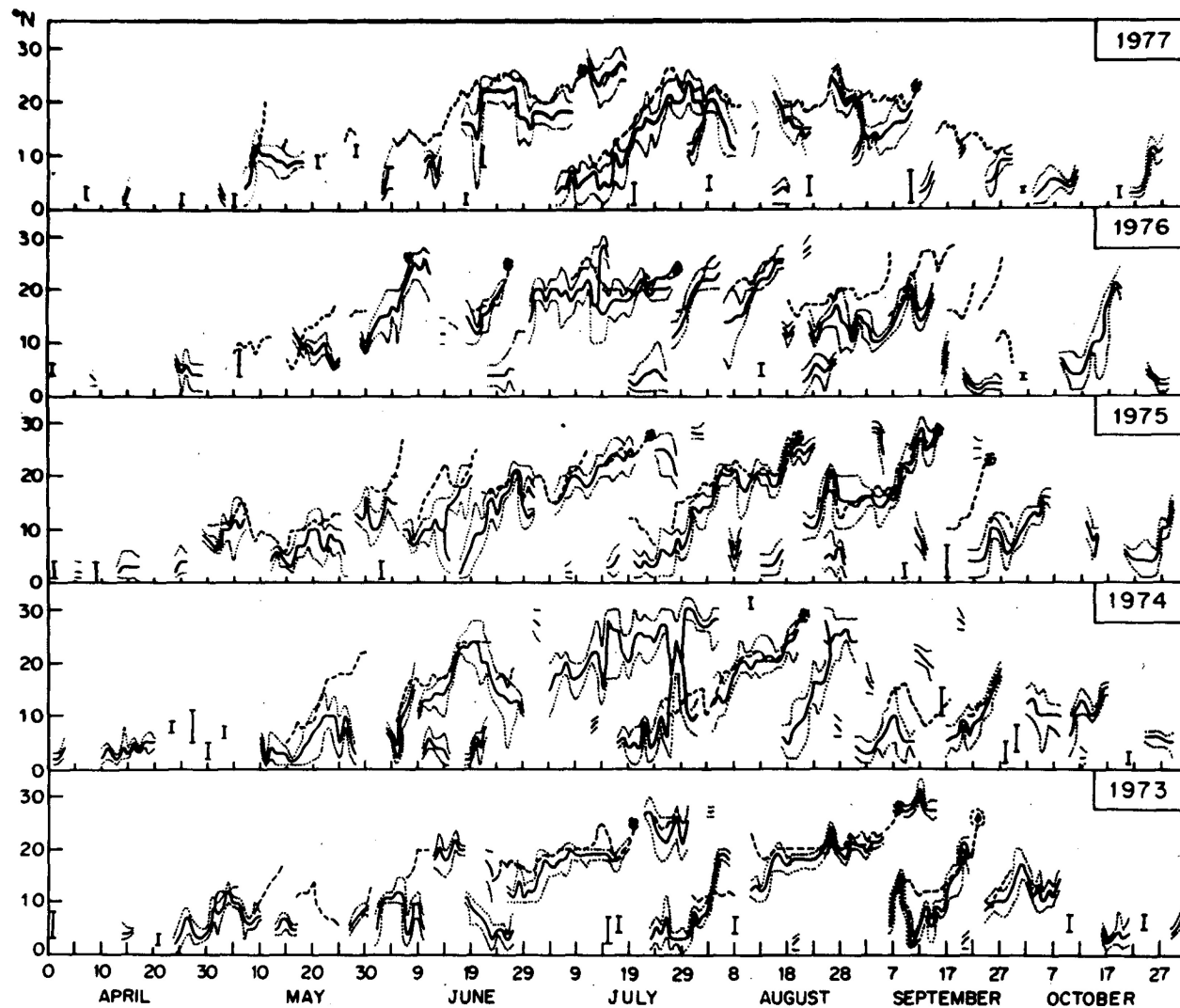
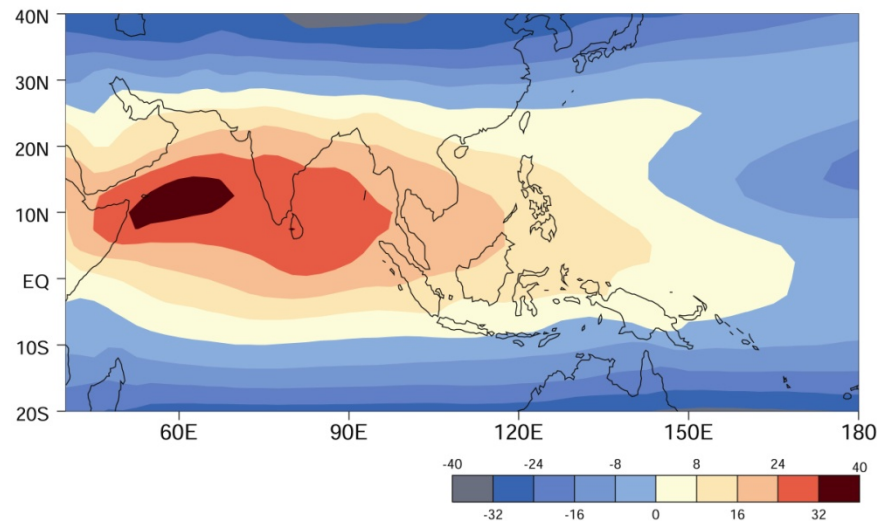


FIG. 4. Daily variation of the latitude of the axis of the MCZ (solid line); northern and southern limits (dotted line) of the MCZ; and the location of the 700 mb trough (dashed line) at 90°E during 1973-77.

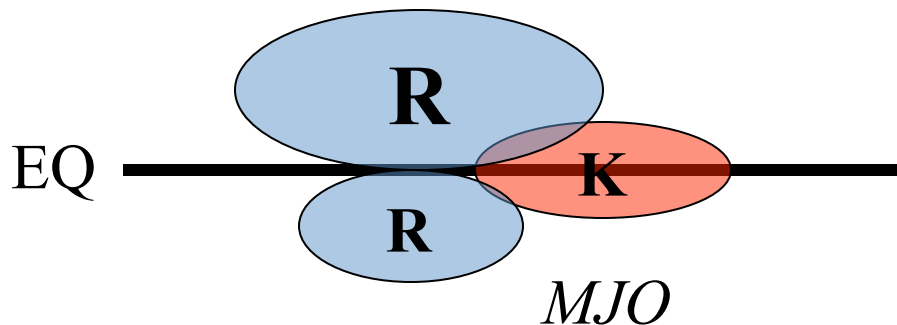
**Sikka and Gadgil (1980)**

# Mean Monsoon and Intraseasonal Variability

## Zonal Vertical Shear

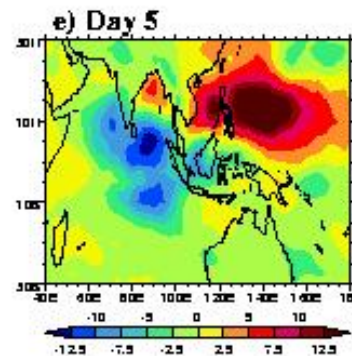
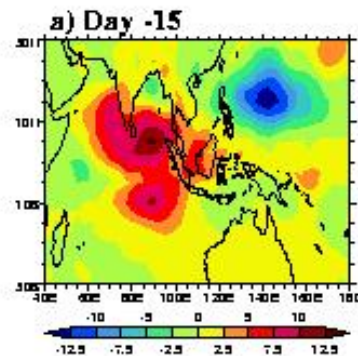


Lau and Peng (1990)  
Wang and Xie (1997)  
Annamalai and Sperber (2005)

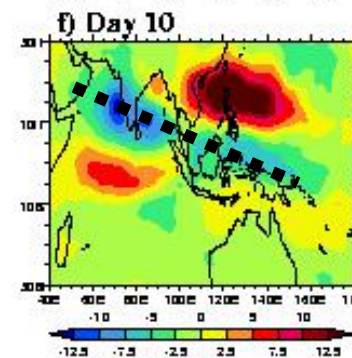
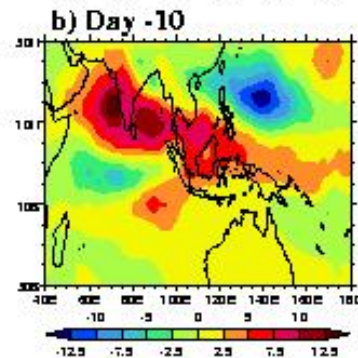




Initiation



Rossby-Kelvin packet

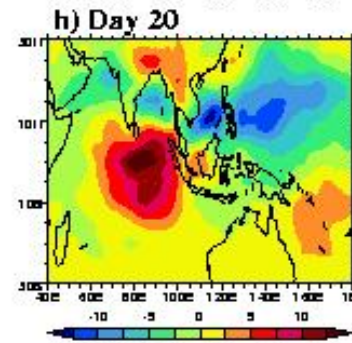
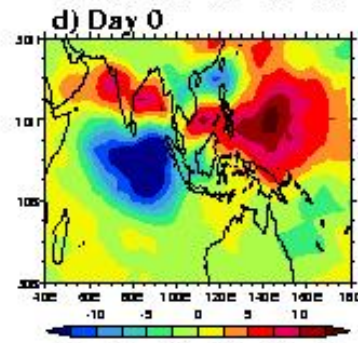
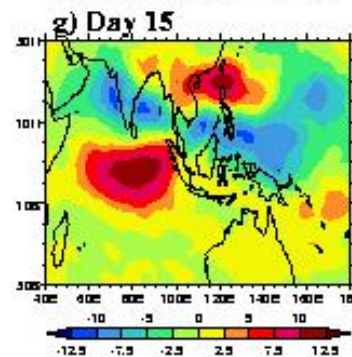
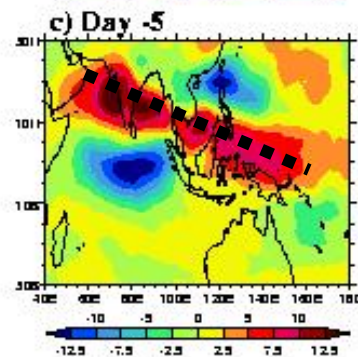


Poleward - India

Eastward - W. Pacific

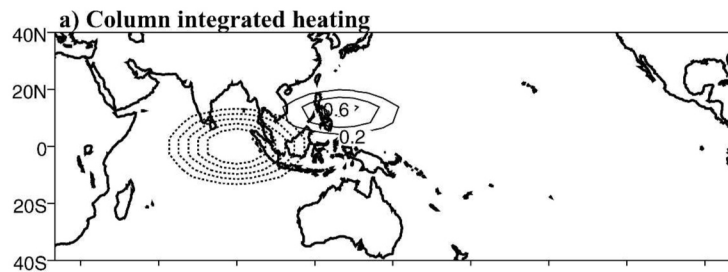
Annamalai and  
Sperber (2005)

“tilted rain band”

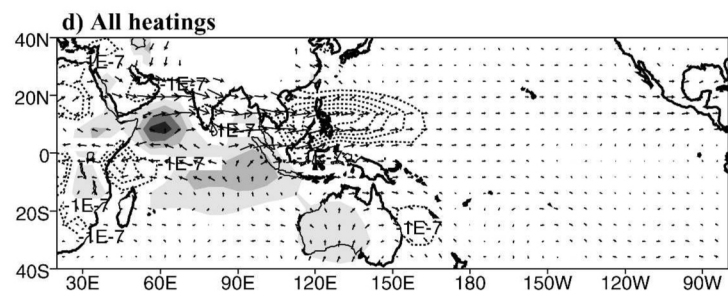
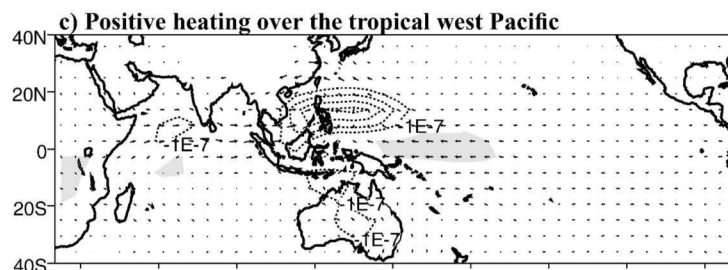
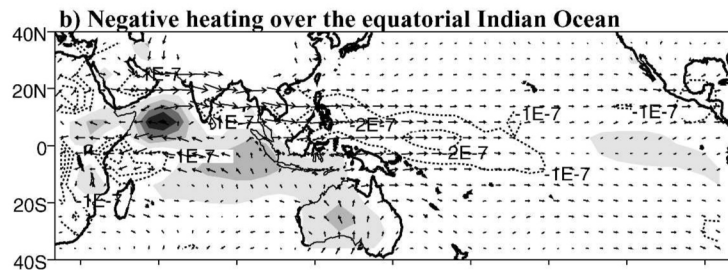


Amplification

Poleward - W. Pacific



**Prescribed heating**



**850 hPa wind / divergence**



**Linear model (Watanabe and Jin 2002)**

FIG. 6. Day -15 (a) column integrated heating anomalies ( $\text{K day}^{-1}$ ), steady-state response of 850-hPa wind ( $\text{m s}^{-1}$ ), and divergence ( $\text{s}^{-1}$ ) to day -15 heating: (b) negative heating over the equatorial Indian Ocean, (c) positive heating over the tropical west Pacific, and (d) total response [sum of (b) and (c)].

**Annamalai and Sperber 2005**

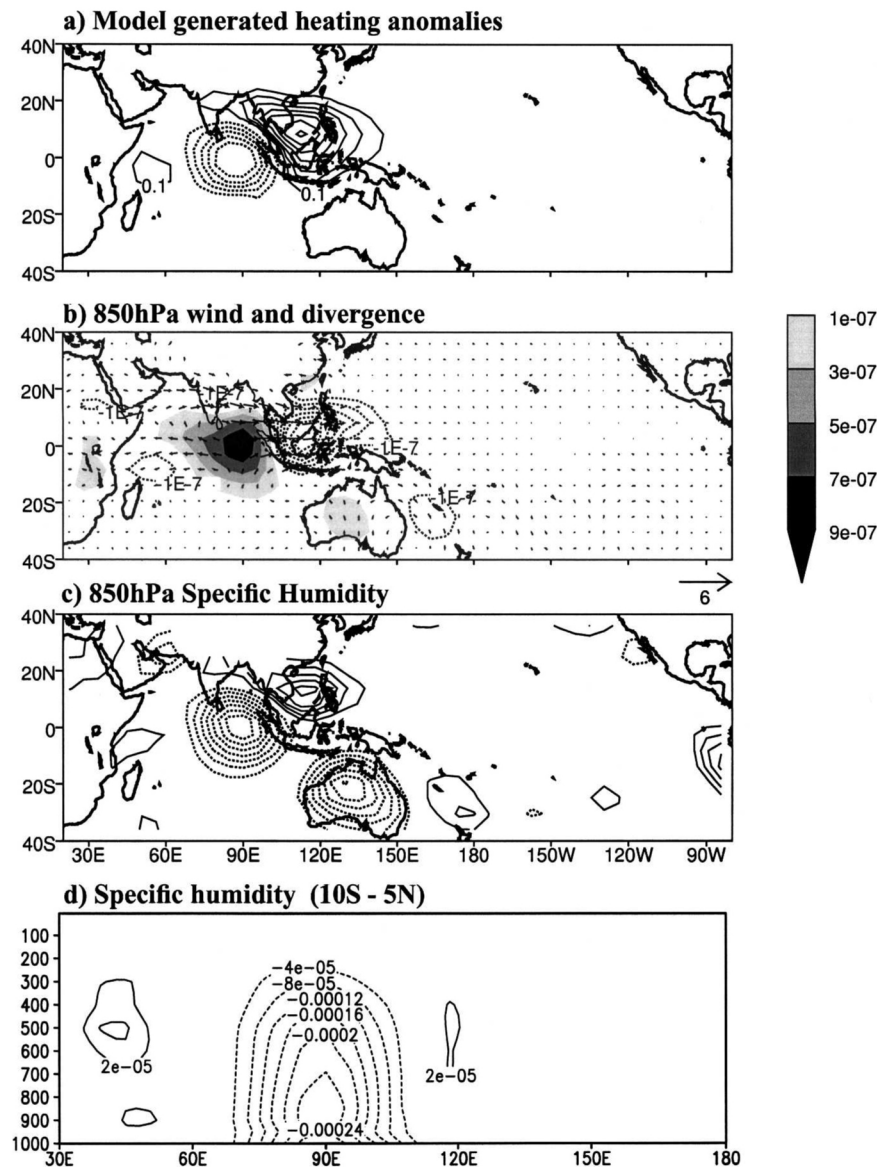
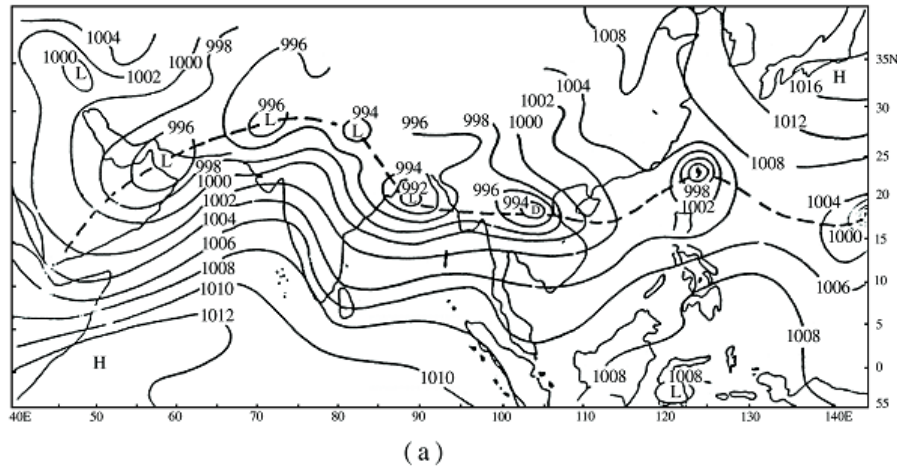


FIG. 7. (a) Model-generated heating anomalies ( $\text{K day}^{-1}$ ) for prescribed cold SST anomalies over the equatorial Indian Ocean, (b) steady-state response of 850-hPa wind ( $\text{m s}^{-1}$ ) and divergence ( $\text{s}^{-1}$ ), (c) 850-hPa specific humidity ( $\text{kg kg}^{-1}$ ), and (d) specific humidity averaged over  $10^{\circ}\text{S}$ – $5^{\circ}\text{N}$  from 1000 to 100 hPa. The negative (positive) contour interval is 0.2 (0.1)  $\text{K day}^{-1}$  in (a) and is dashed (solid). The negative contour interval in (b) is  $1.0 \times 10^{-7}$  and the positive values are shaded progressively. The contour interval for positive (negative) values is  $2 \times 10^{-5}$  ( $4 \times 10^{-5}$ ) in (c)–(d).

Model-generated heating

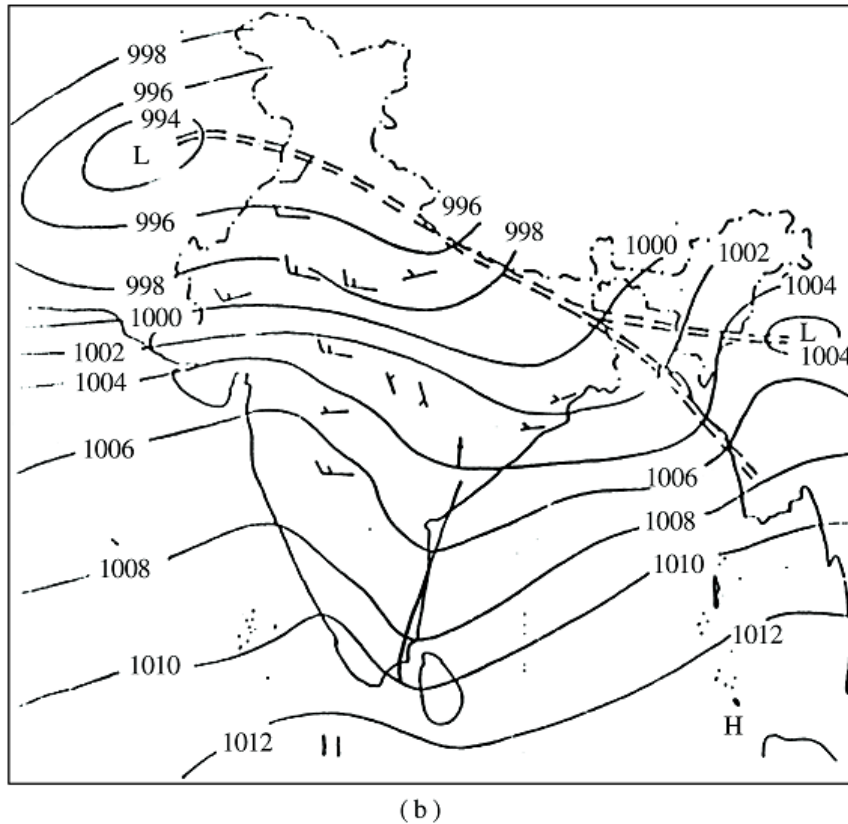
Annamalai and Sperber 2005

Active Phase



“favorable for successive  
genesis of lows/depressions”

Break Phase

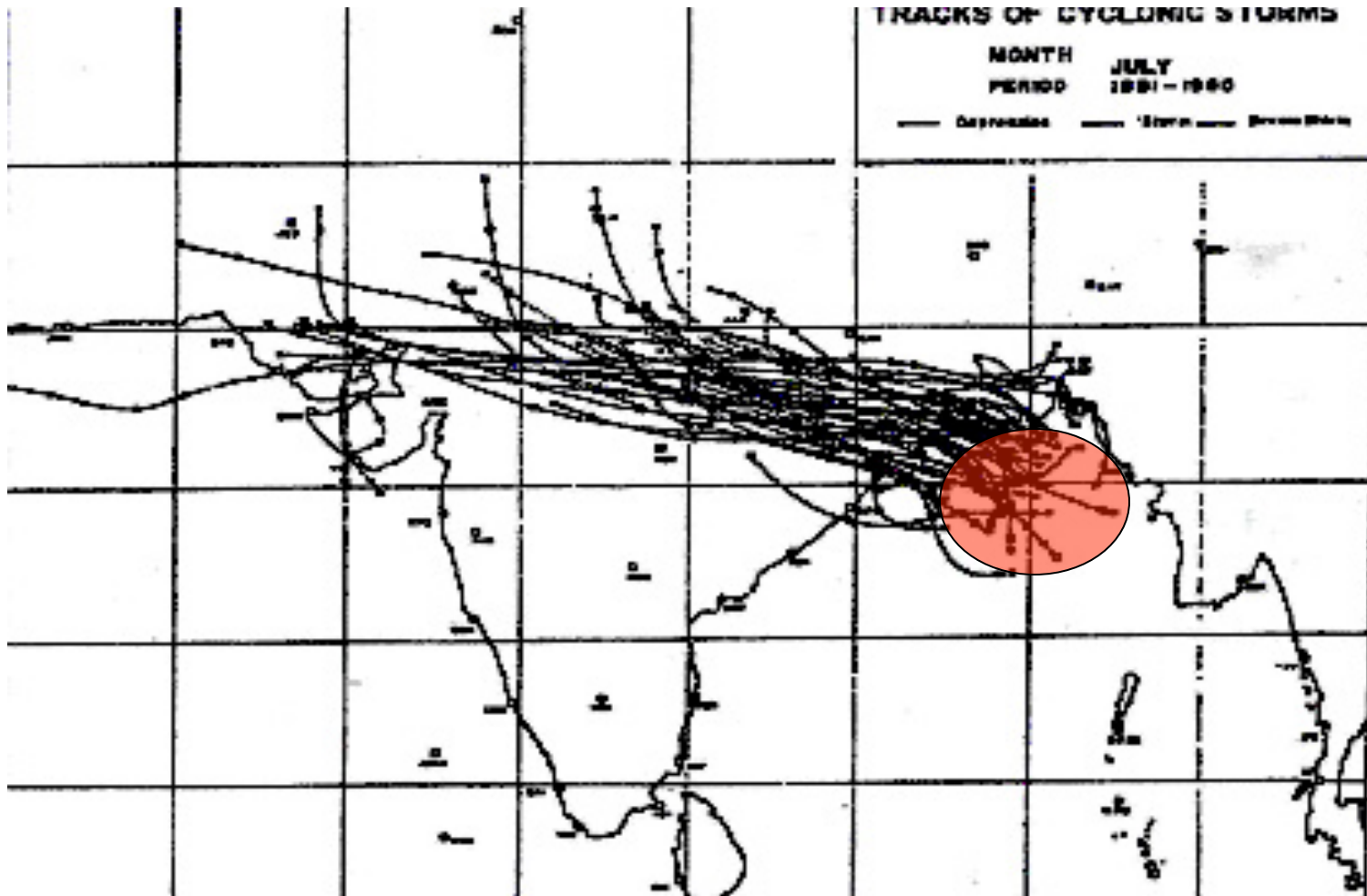


“foot hills of the Himalayas  
lead to floods in the rivers”

“in both cases, some regions over India face flood situations”

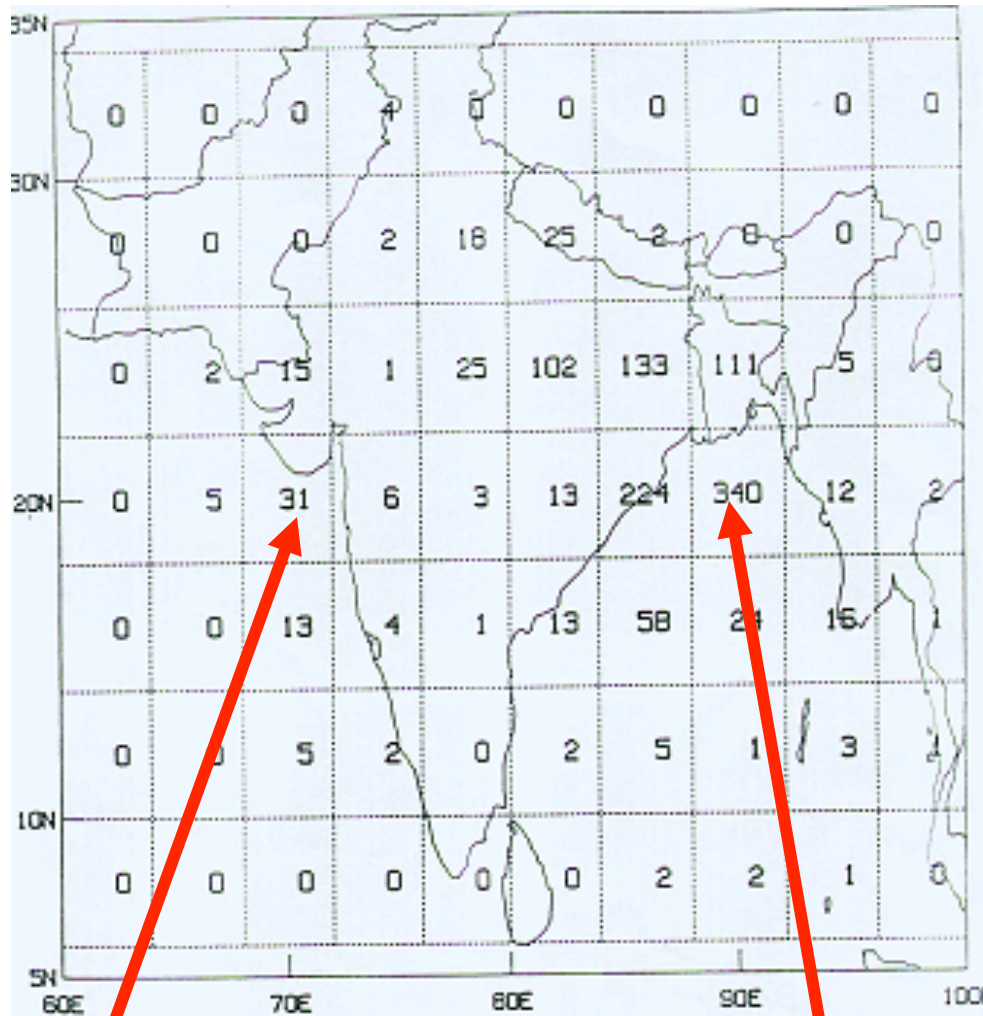


# Mean Monsoon and Synoptic Systems



**Role of the basic flow**

# Number of lows and depressions



## Monsoon lows and depressions

Number of low-pressure systems (lows, depressions, and cyclonic storms) which formed over 4° lat/long blocks over the south Asian region in the summer monsoon season during 1888-1983

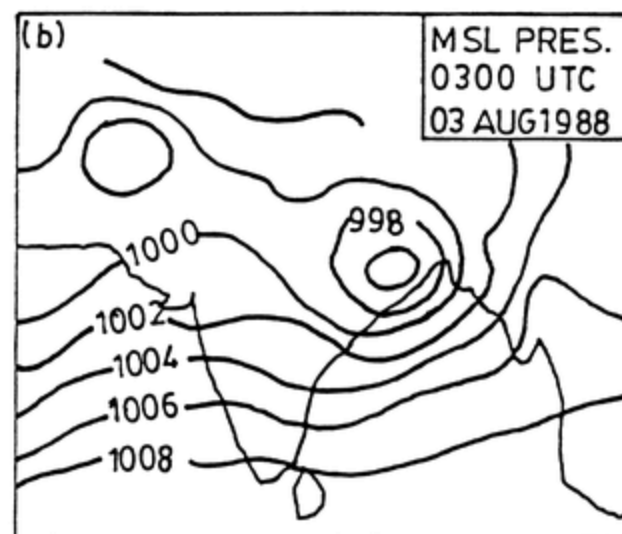
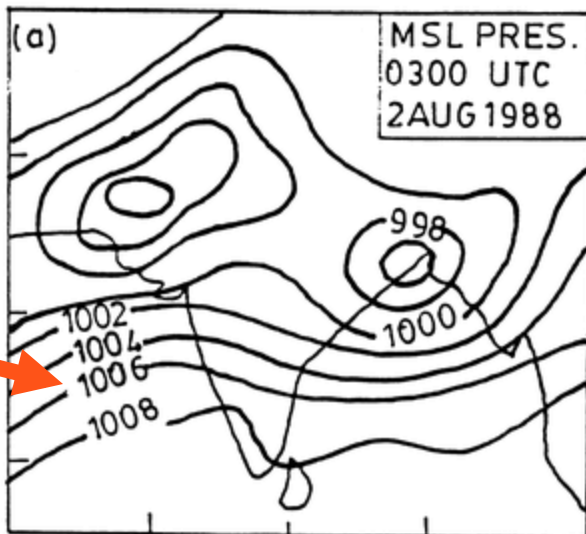
### Main features:

Maximum numbers of low-pressure systems form over the northern Bay of Bengal.

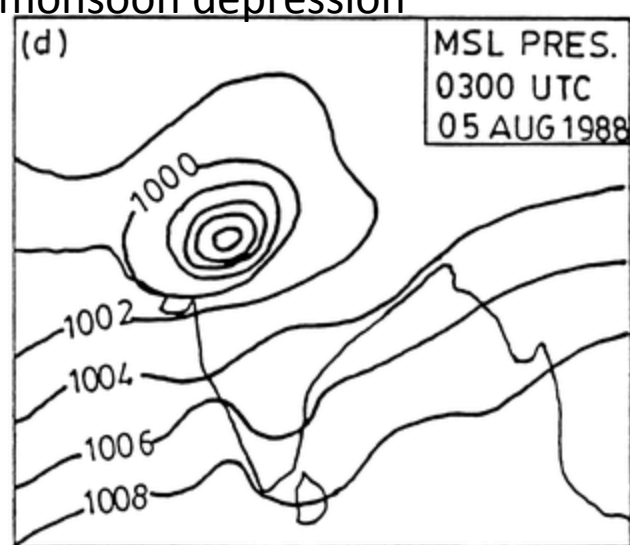
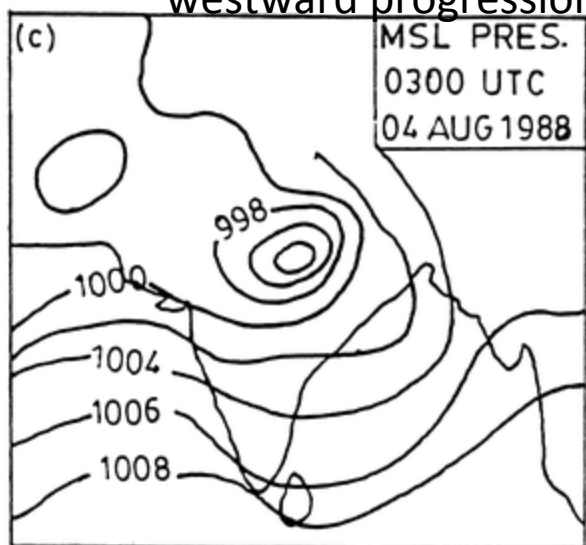
Climatologically, 7 depressions and 1.5 cyclonic storms form over the Bangladesh-India-Pakistan region.

Mean: 13      Standard deviation: 2.2

"closey packed  
Isobars"



"westward progression of monsoon depression"

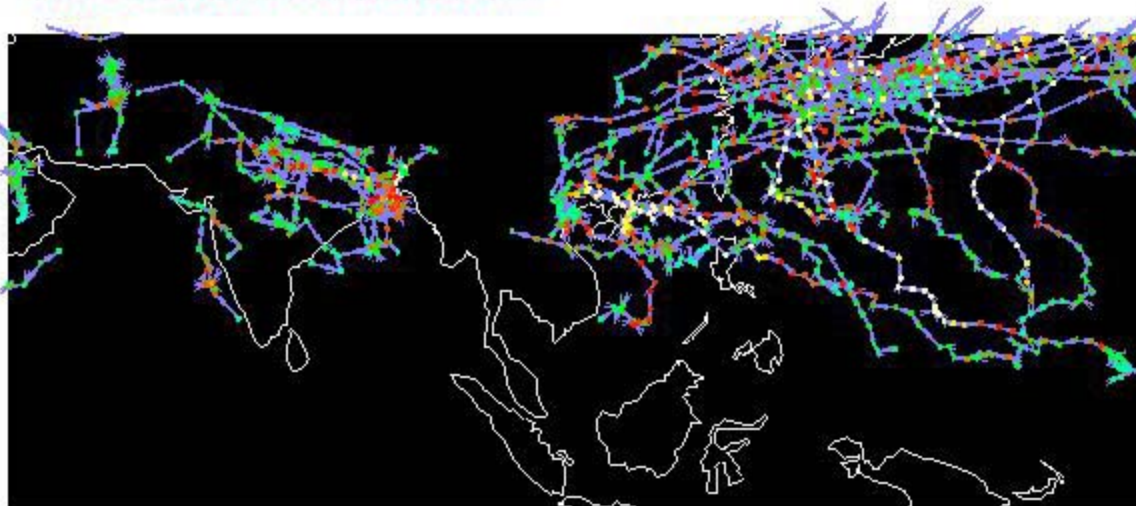


"westward progression of a monsoon depression – vorticity budget – divergence term  
- Maximum rainfall SW quadrant - Rossby waves – east-west divergent circulation"



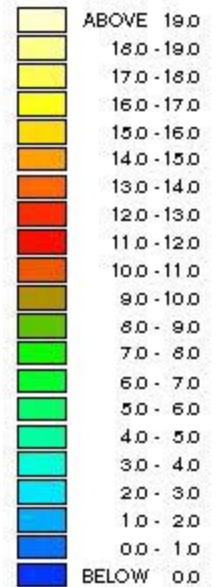
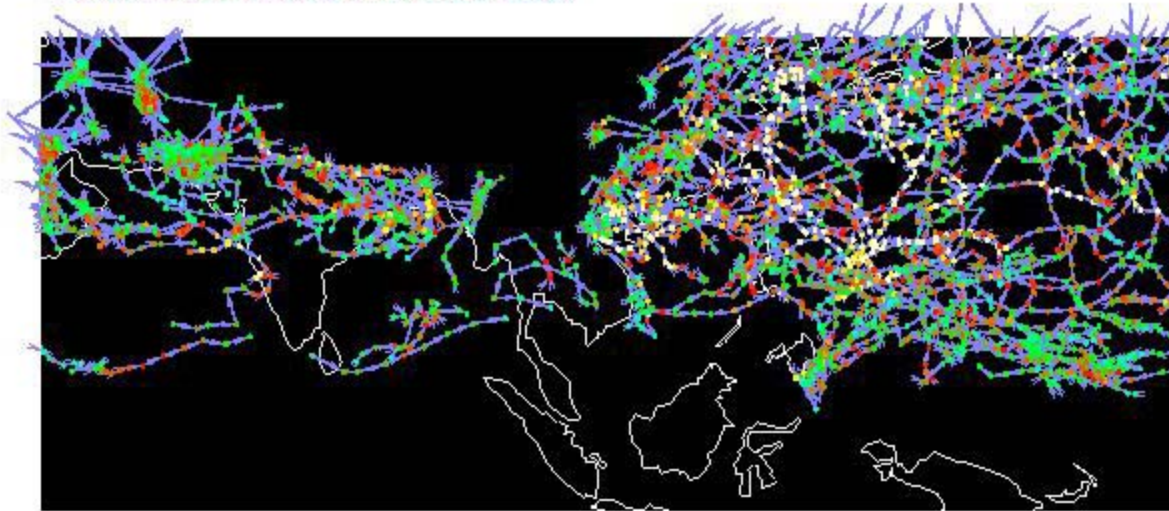
## Tracks of Synoptic Systems (1993 – Boreal summer)

ERA, VOR850, ave, MJJAS 1993, T>=2 days, D<=5 deg., Str>=10V U



850hPa  
Vorticity

ERA, VOR850, ave, MJJAS 1994, T>=2 days, D<=5 deg., Str>=10V U



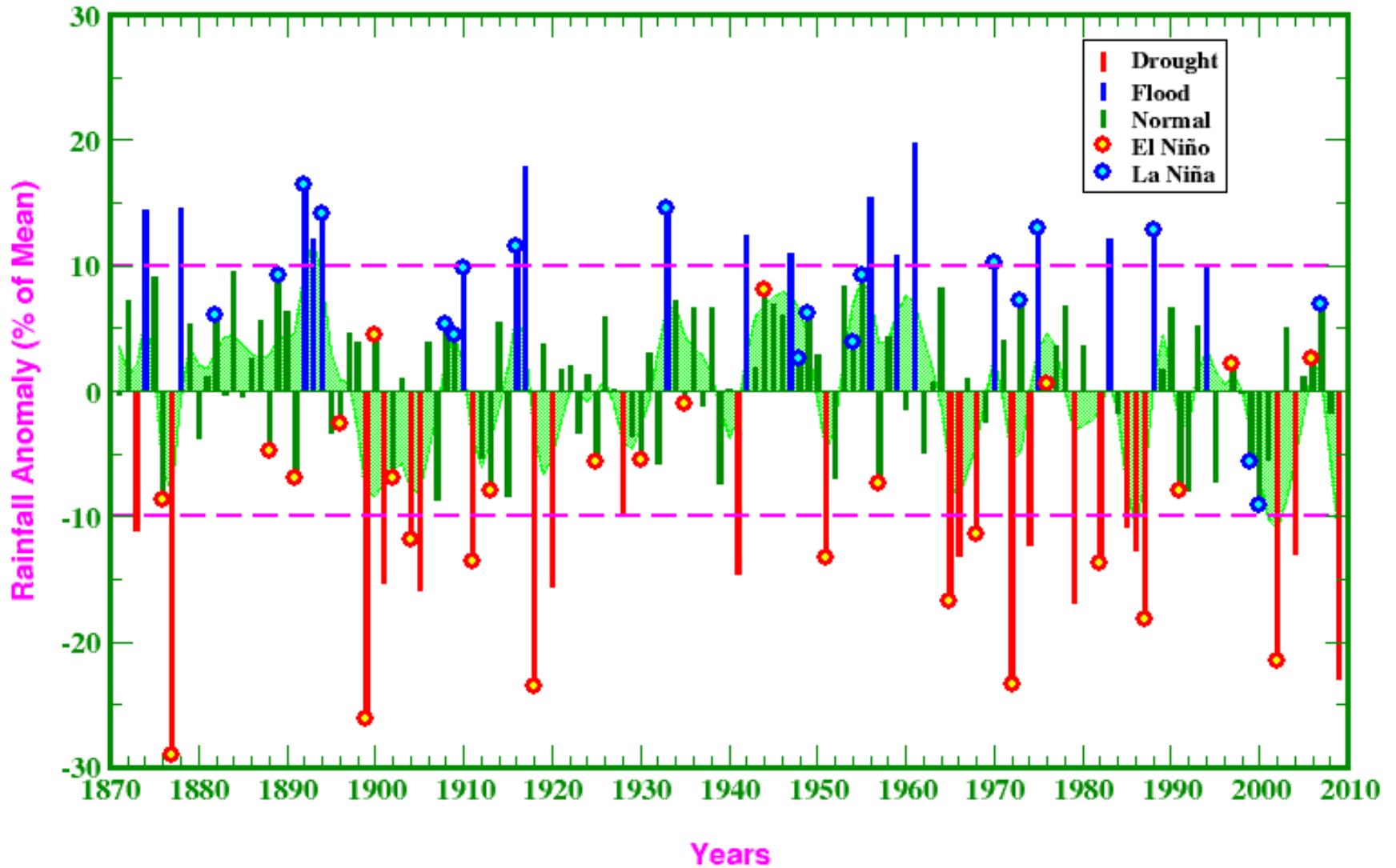
## Tracks of Synoptic Systems (1994 - Boreal summer)

Strong – Indian  
and West Pacific  
monsoon



# All-India Summer Monsoon Rainfall, 1871-2009

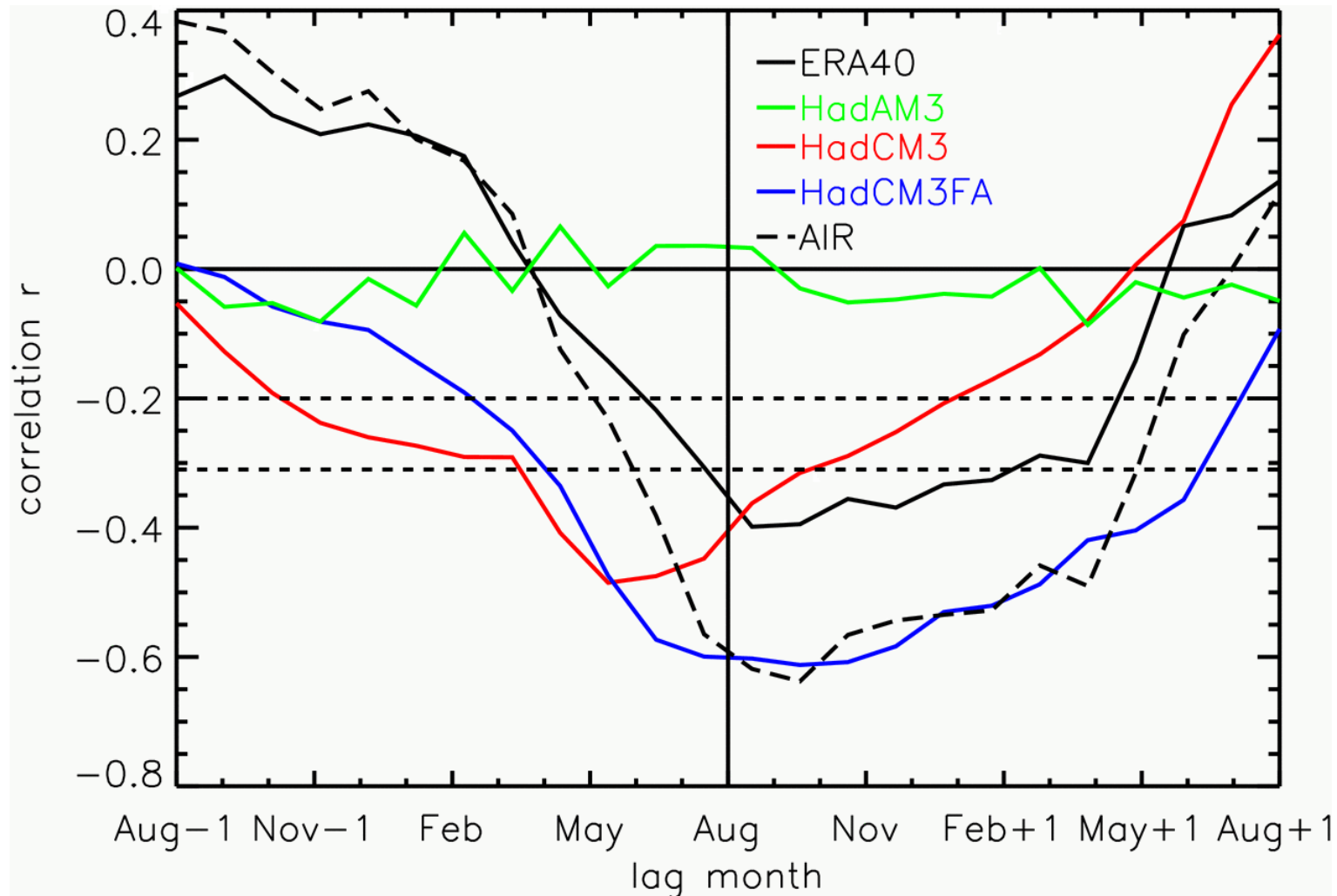
*(Based on IITM Homogeneous Indian Monthly Rainfall Data Set)*



Severe weak monsoon years – associated with El Nino

# ENSO-Monsoon Teleconnections

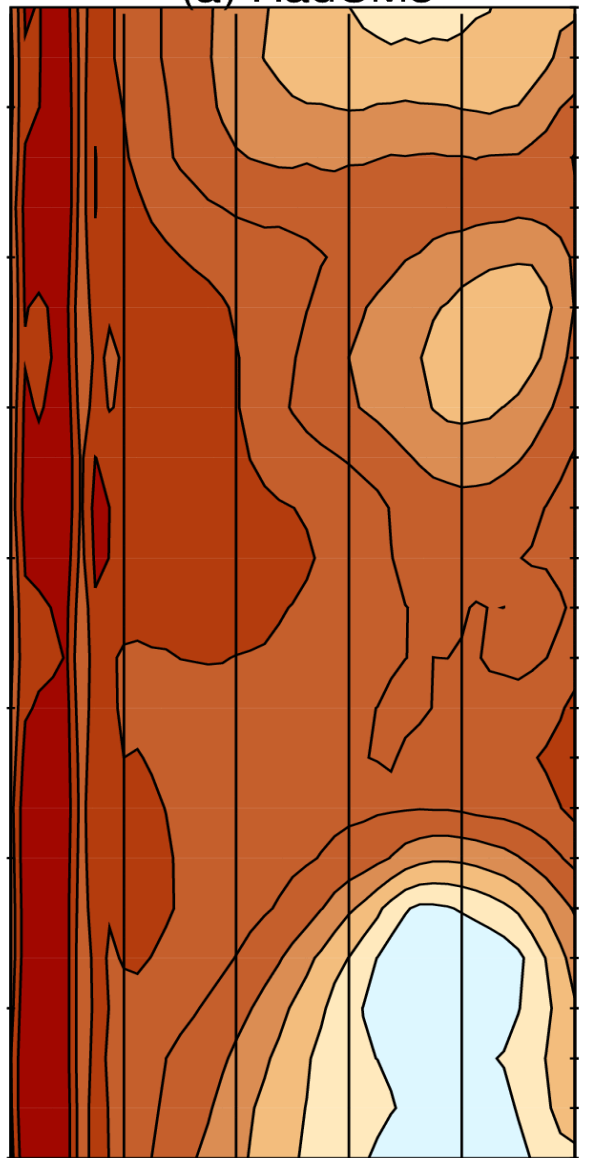
Summer All India Rainfall (AIR) lag-correlated with Nino-3 SSTs



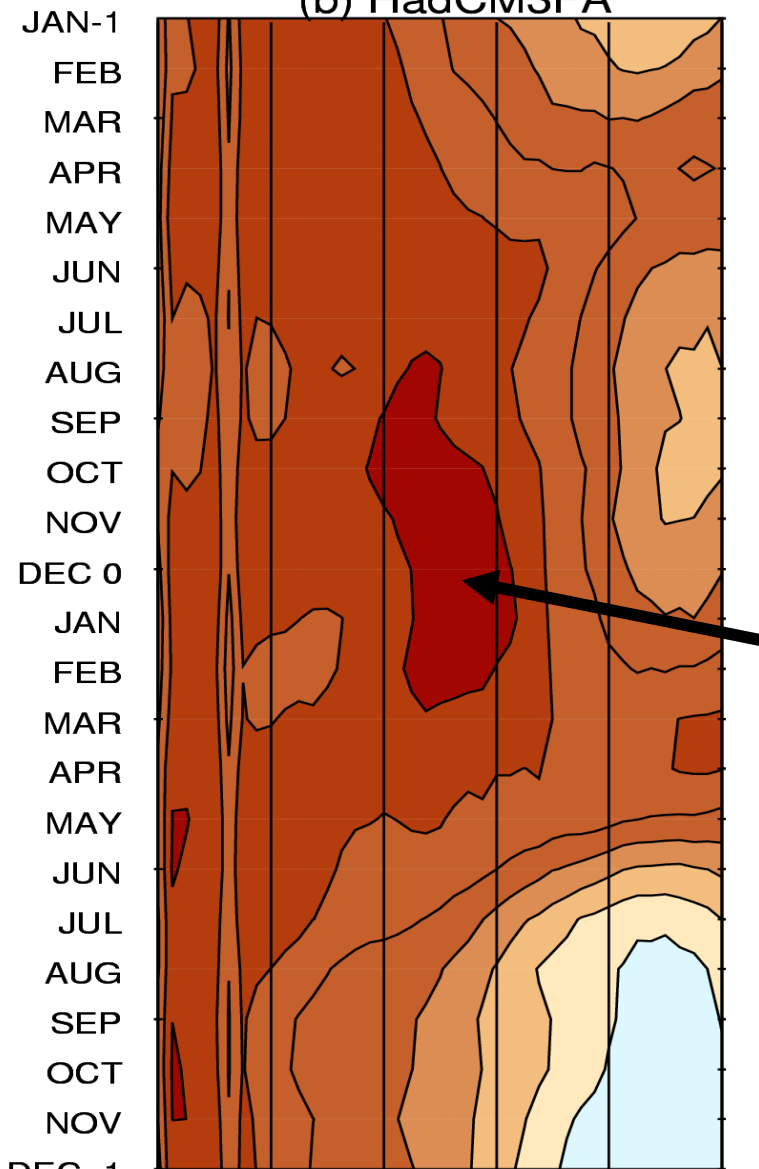
“highlights the role of the basic state”

*Turner et al. 2005, QJRMS*

(a) HadCM3



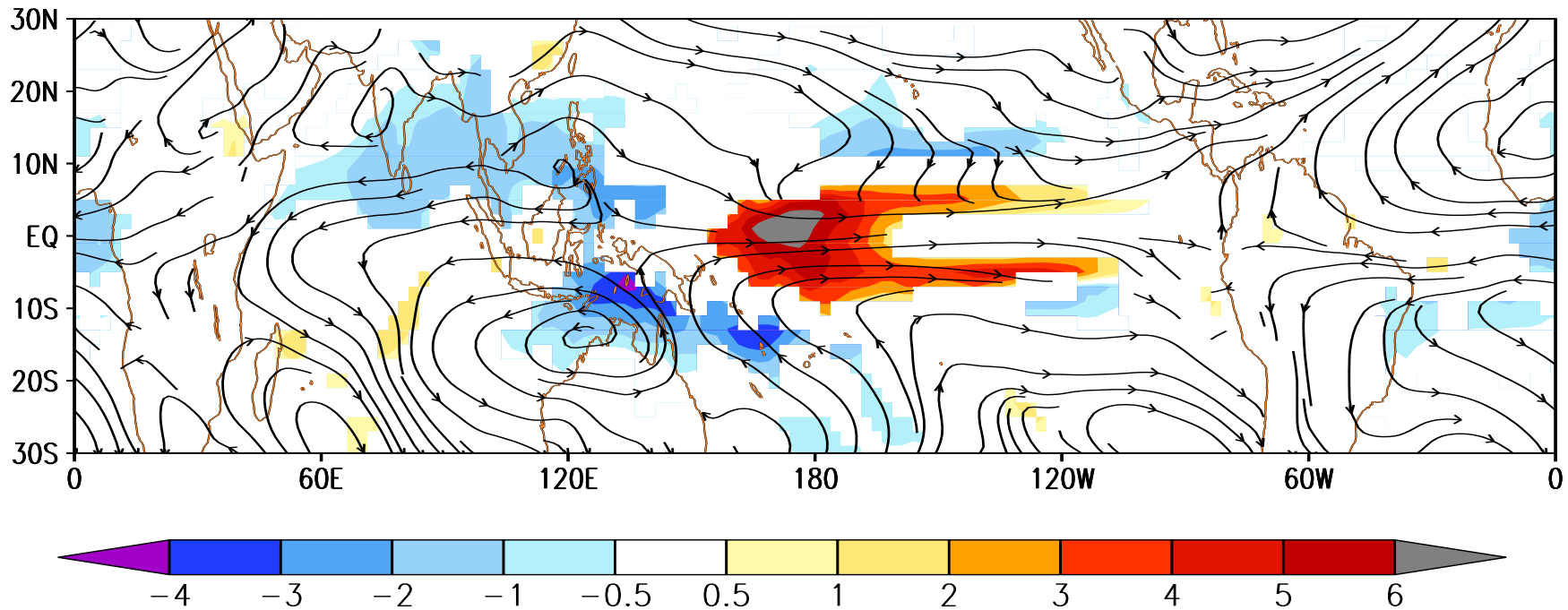
(b) HadCM3FA



Evolution of  
SST  
through  
composite  
El Nino

Note that  
warmest  
water  
moves into  
the central  
Pacific with  
improved  
basic state

**May** averaged CM2.1 composite of anomalous 850 hPa stream line and rainfall



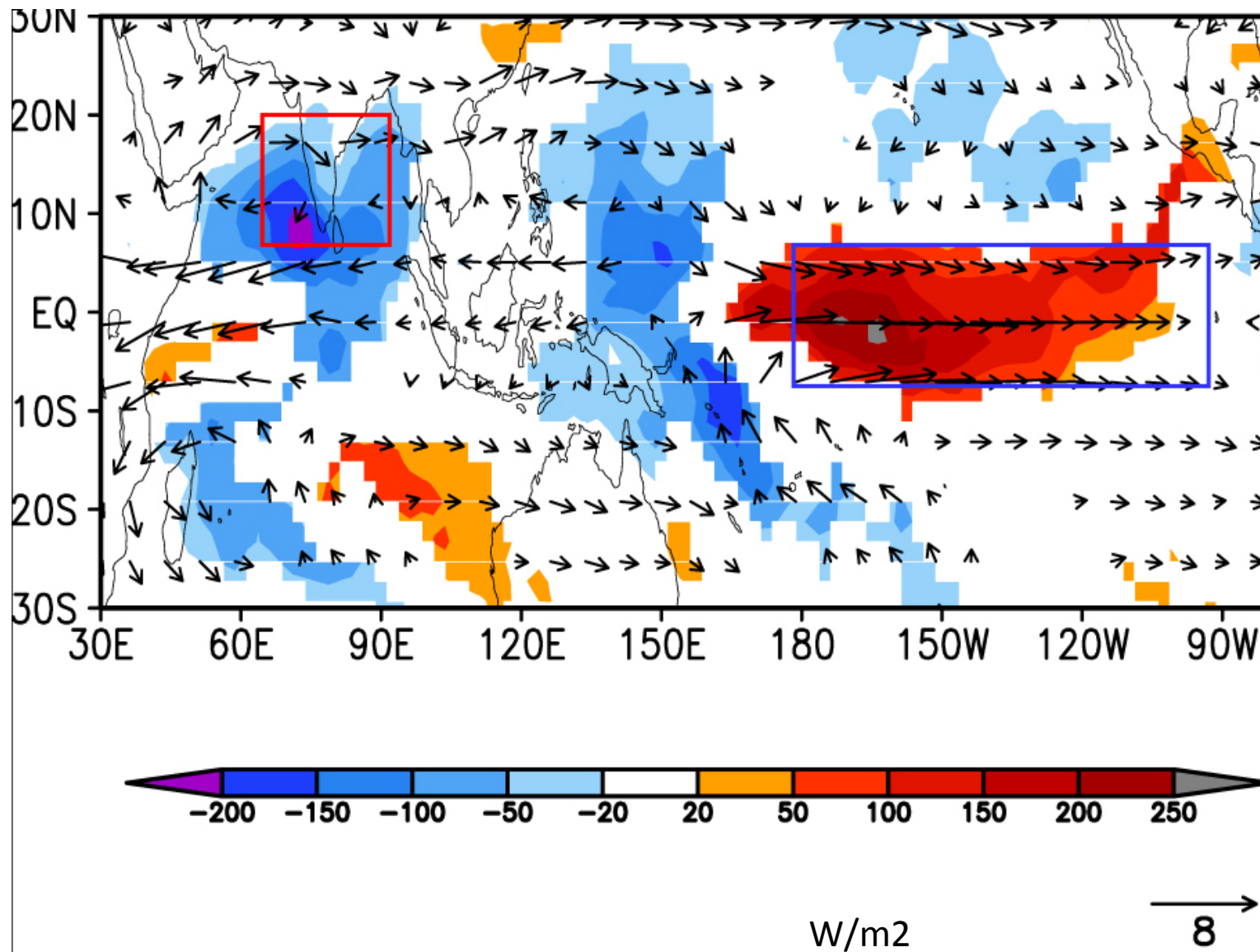
Severe weak monsoons over south Asia co-occurred with developing phase of El Nino

**NIO** – anticyclonic vorticity – within 2-3 days of SST forcing – rainfall after about 20 days

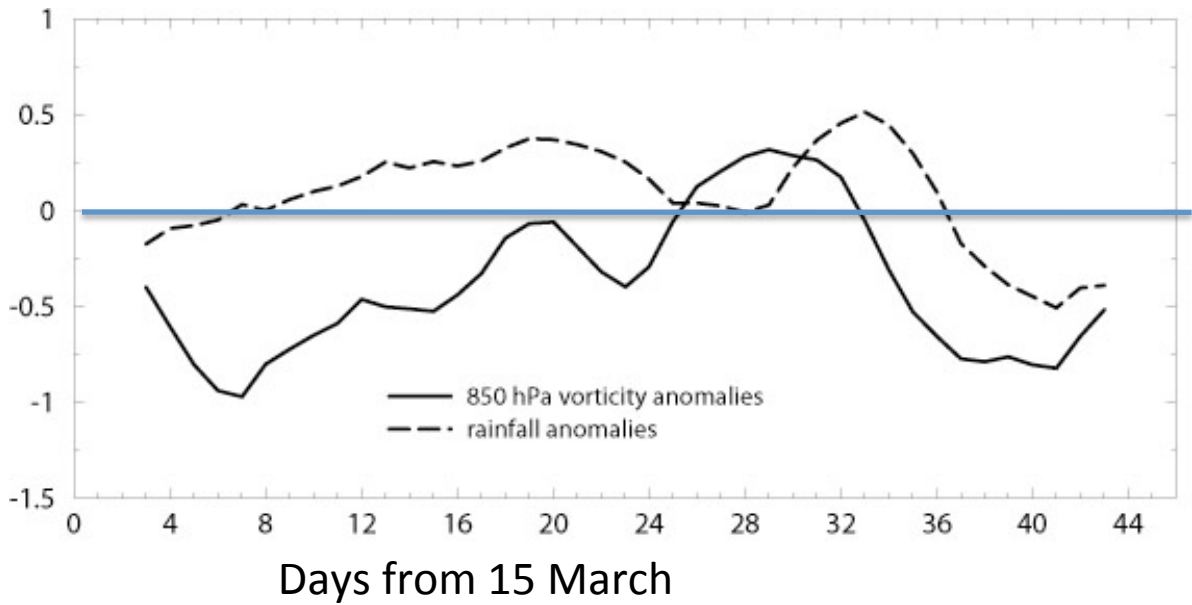
**Dry air advection** from north is instrumental in initiating the dryness



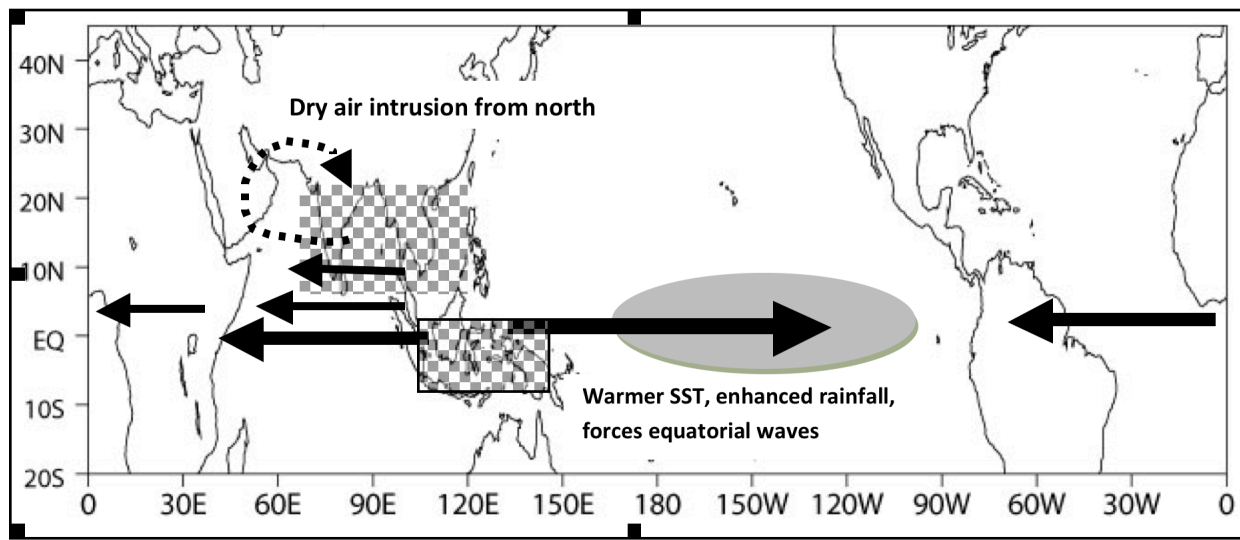
## May rainfall and 850 hPa wind response to El Nino SST forcing



AM2.1 solutions – Forced with CM2.1 composite SST anomalies (El Nino)



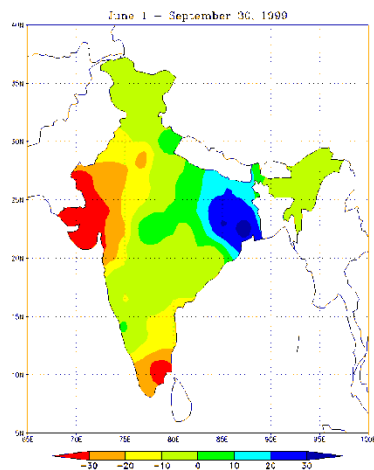
Rainfall over S. Asia  
850 hPa Vorticity  
(west of rainfall maximum)



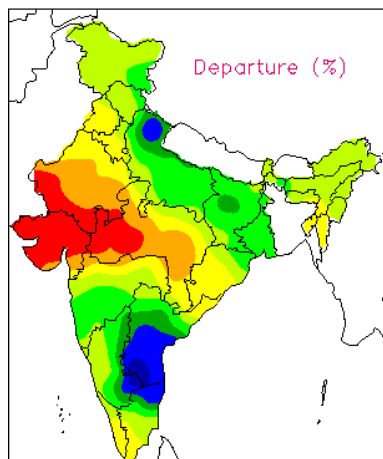
Summary – Case II

“dry advection leads rainfall anomalies – long lead time – useful for prediction”

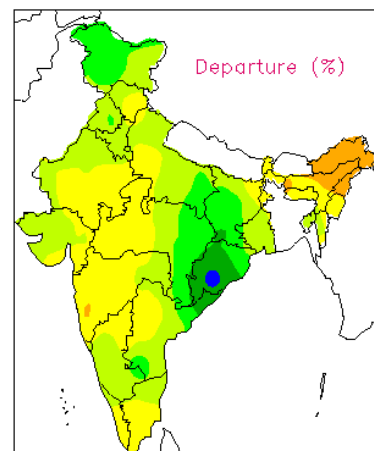
1999



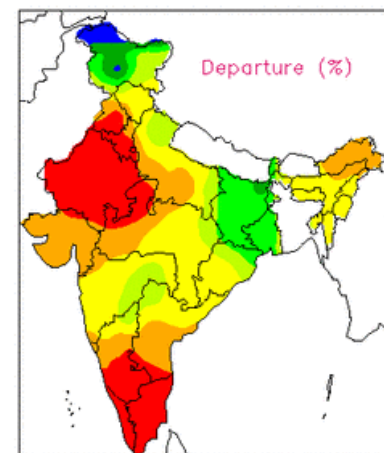
2000



2001

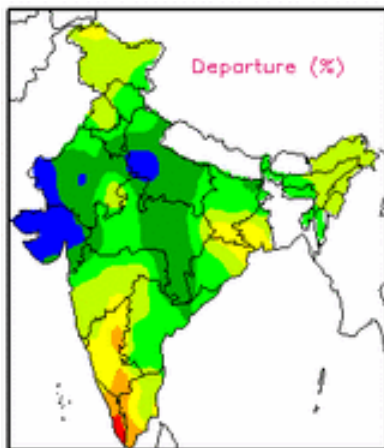


2002

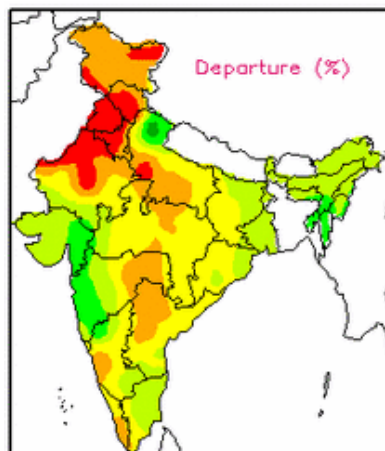


“spatial variation of rainfall anomalies in selected years”

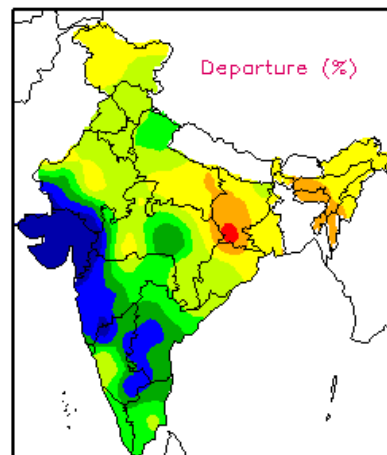
2003



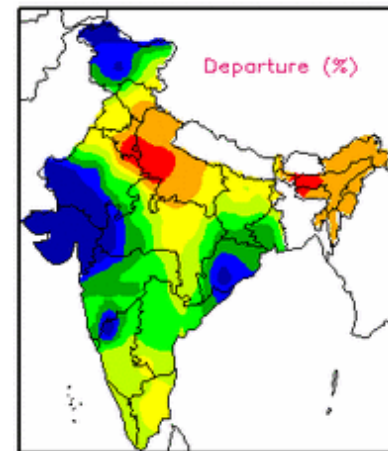
2004



2005



2006



## Arabian Sea

Strong winds  
(Findlater Jet)

Strong vertical transport

Weak near-surface  
stratification

Cool SST

$P-E < 0$

Weak convective  
activity

## Bay of Bengal

Weak winds

Weak vertical transport

Strong near-surface  
stratification

Warm SST

$P-E > 0$

Strong convective  
activity