

# Ocean tipping points - an overview

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BCCR Earth system modelling support  
BCCR Strategic project LOES



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*Thank you for the invitation!*

**Tipping point:**

(IPCC SROCC Glossary, 2019)

A ***level of change in system properties***

beyond which a system reorganises, often in a non-linear manner,

and does not return to the initial state even if the drivers of the change are abated.

For the climate system, the term refers to a critical threshold at which global or regional climate changes from one stable state to another stable state.

Tipping points are also used when referring to impact:

the term can imply that an impact tipping point is (about to be) reached in a natural or human system.

**Regime shift:**

(Biggs et al., *Ecology and Society*, 2018,  
<https://doi.org/10.5751/ES-10264-2303092018>)

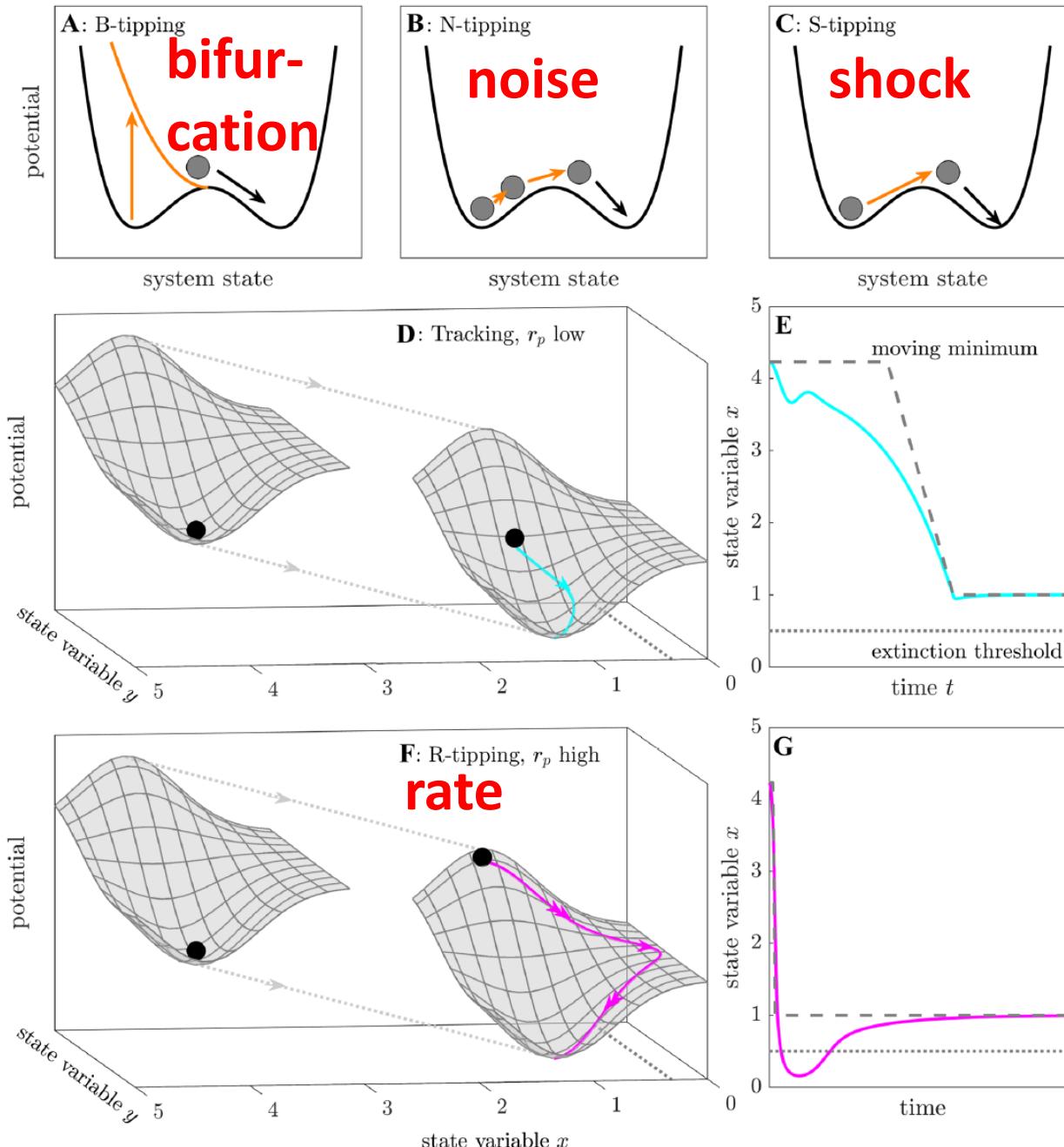
Changes in ecosystems and social-ecological systems (SESSs) are often experienced as relatively slow and incremental, but sometimes ***dramatically large, persistent, and often unexpected changes*** take place = regime shifts.

# Tipping types

Vanselov et al.,  
2021

Theoretical Ecology  
<https://doi.org/10.1007/s12080-021-00522-w>

cup = system properties  
ball = state of system



## R-tipping - rate induced

Claire Kiers, PhD, 2020,

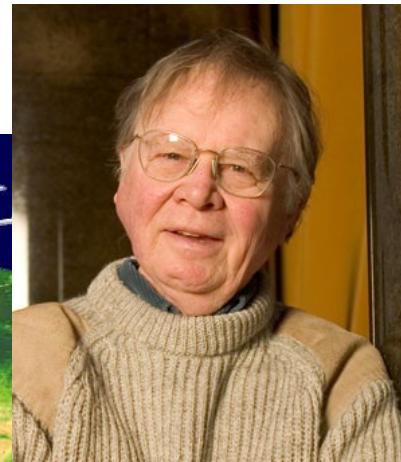
<https://doi.org/10.17615/2d03-0651>

*“... Here we get two distinct outcomes: the dishes come with the tablecloth or they get left behind.*

*The deciding factor between these outcomes is not how far the tablecloth moves, but how fast. ...”*

... easier to find cases for  
regime shifts than to  
quantify the thresholds ...

# The great ocean conveyor....



“Wally” Broecker

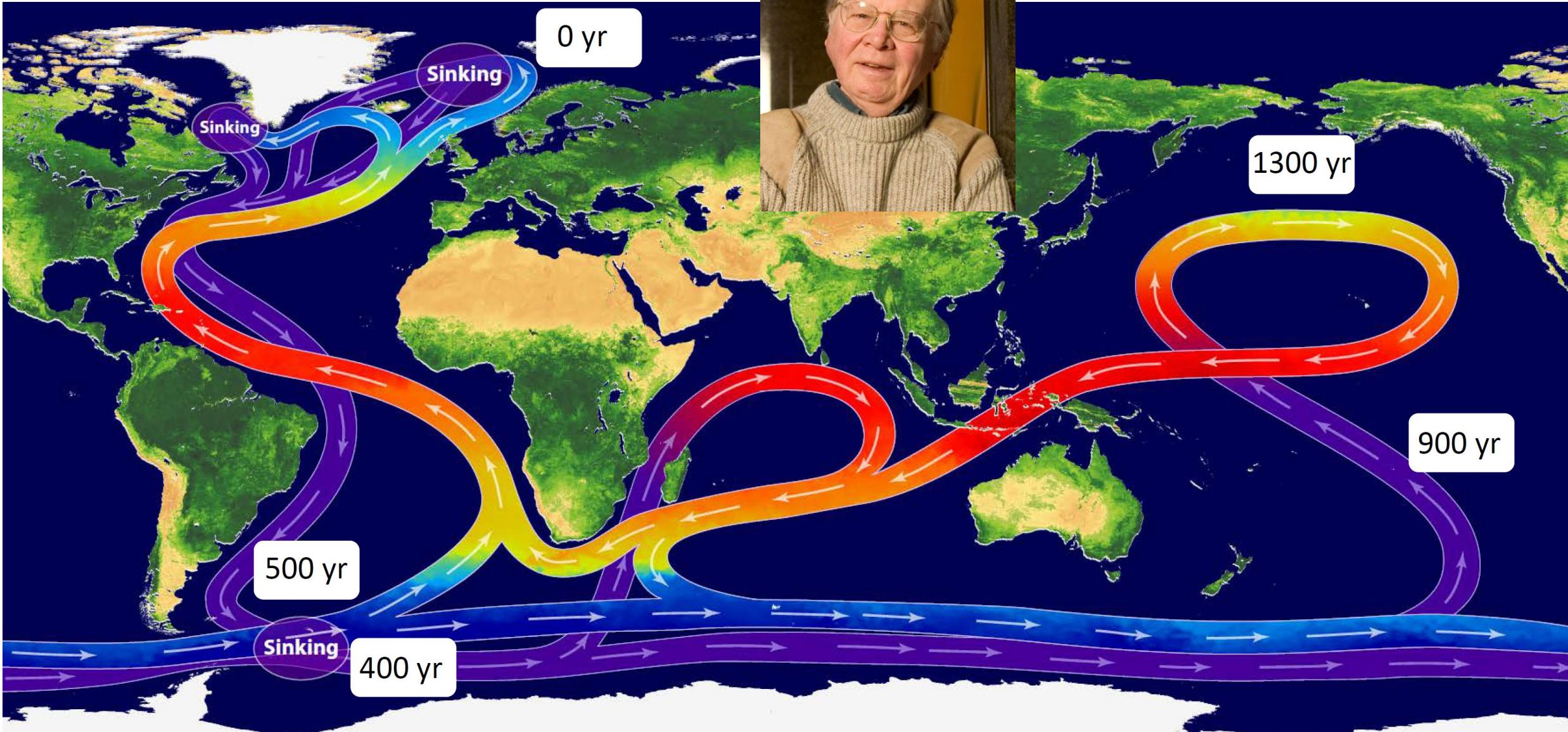
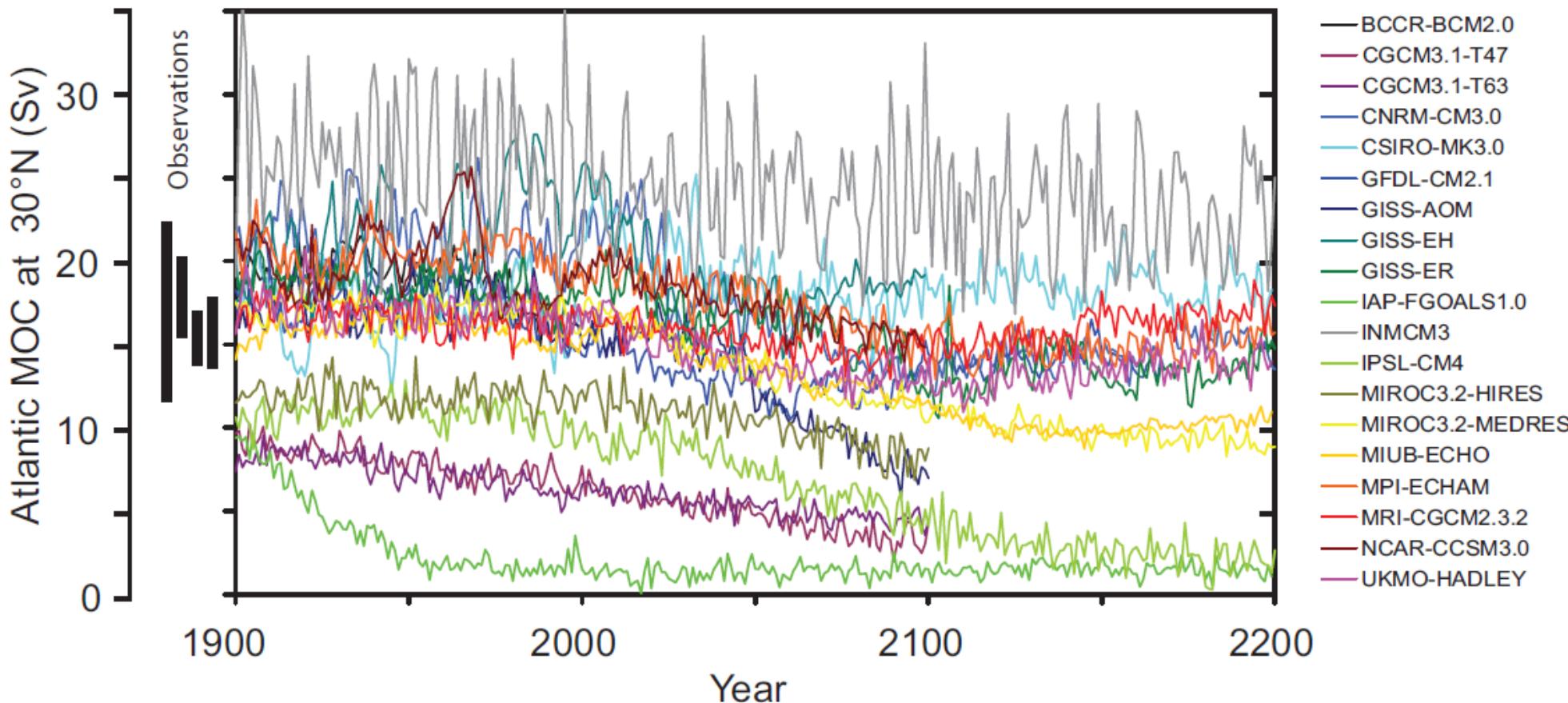


Image credit: V. Byfield, National Oceanography Centre (NOC), licensed under CC BY 3.0. The original figure was modified with mean age indications of water masses since their last contact with atmosphere. (Published in this form in: Heinze et al., 2021, <https://doi.org/10.1073/pnas.2008478118>)

# The concern for an AMOC collapse:

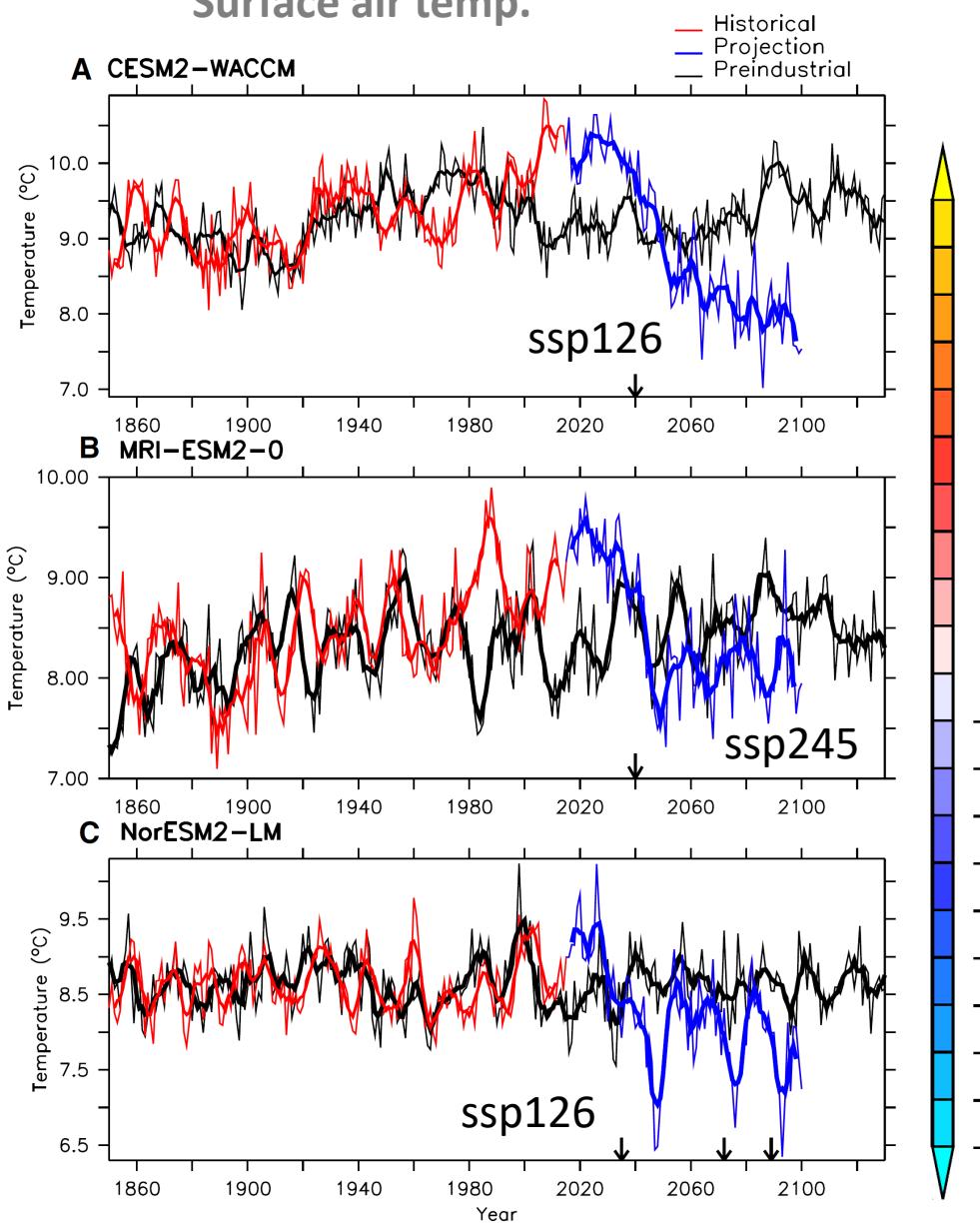


Meehl, G.A., T.F. Stocker, W.D. Collins, P. Friedlingstein, A.T. Gaye, J.M. Gregory, A. Kitoh, R. Knutti, J.M. Murphy, A. Noda, S.C.B. Raper, I.G. Watterson, A.J. Weaver and Z.-C. Zhao, 2007: Global Climate Projections. In: *Climate Change 2007: The Physical Science Basis*. Contribution of Working Group I to the Fourth Assessment Report of the Intergovernmental Panel on Climate Change [Solomon, S., D. Qin, M. Manning, Z. Chen, M. Marquis, K.B. Averyt, M. Tignor and H.L. Miller (eds.)]. Cambridge University Press, Cambridge, United Kingdom and New York, NY, USA.

E.g. Fabbri al., *Environ. Sci. Technol.*, 2021, SI, <https://dx.doi.org/10.1021/acs.est.0c02928>, SI1:  
**Addition of freshwater in the North Atlantic (due to sea ice and Greenland ice sheet melting, river inputs and ocean precipitation) may reduce the density-driven sinking of North Atlantic waters until the Atlantic thermohaline circulation is significantly slowed down or even stopped**

# Abrupt cooling of the N.Atl. sub-polar gyre:

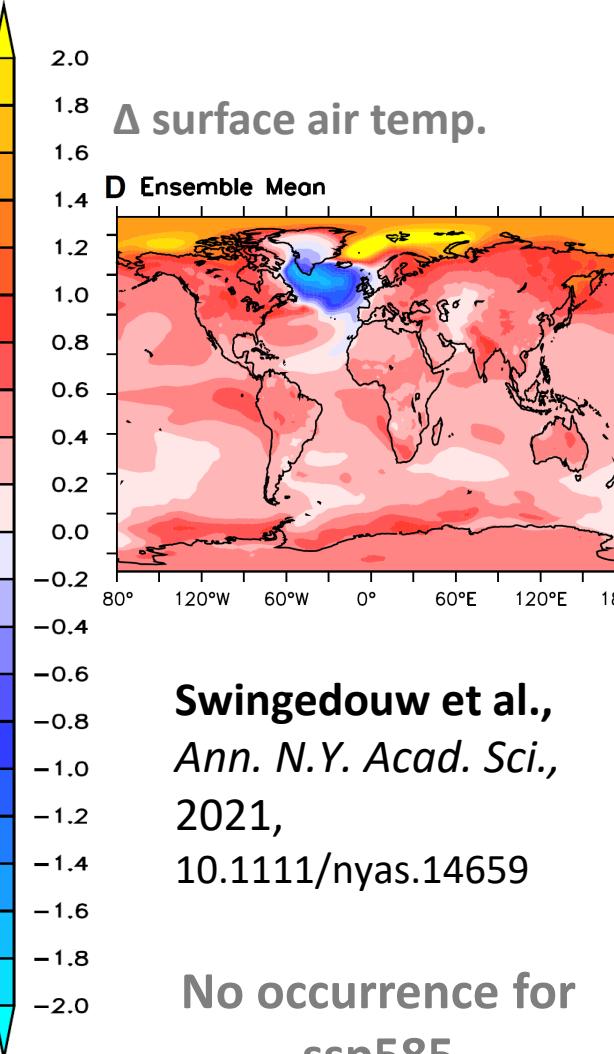
Surface air temp.



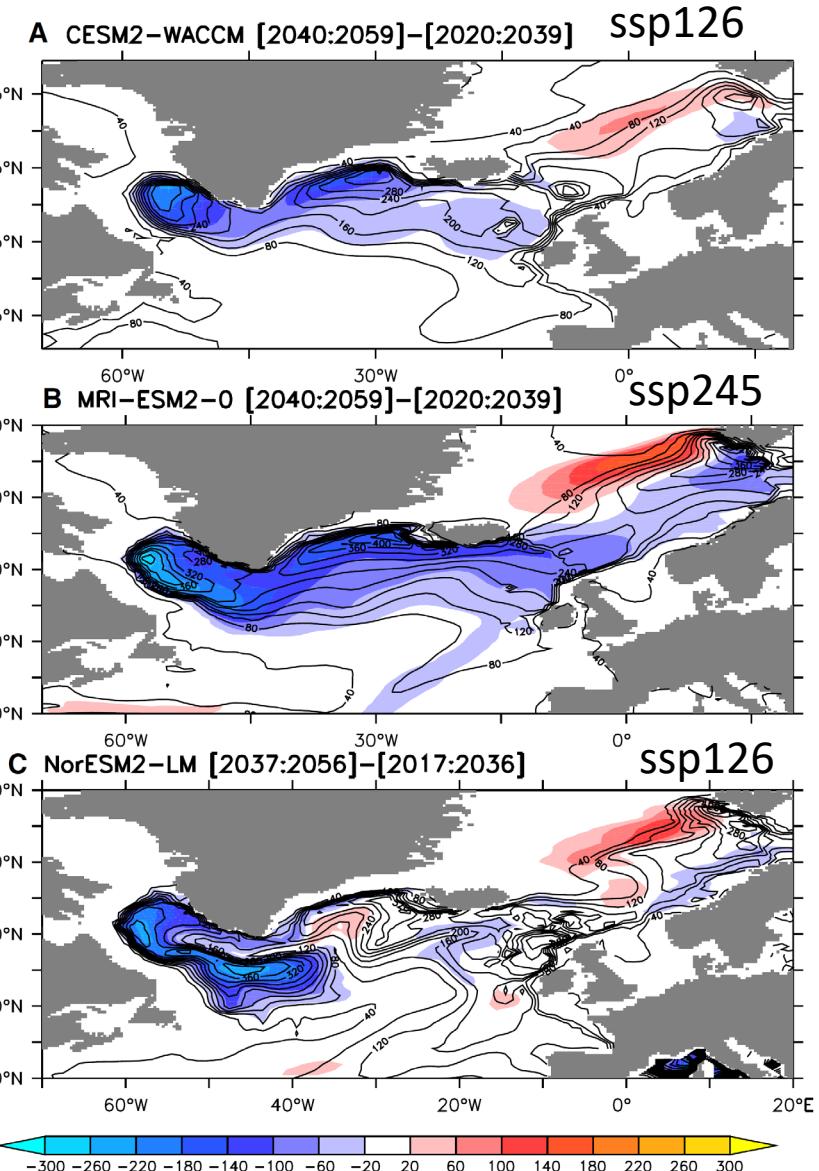
CMIP6

Δ surface air temp.

D Ensemble Mean



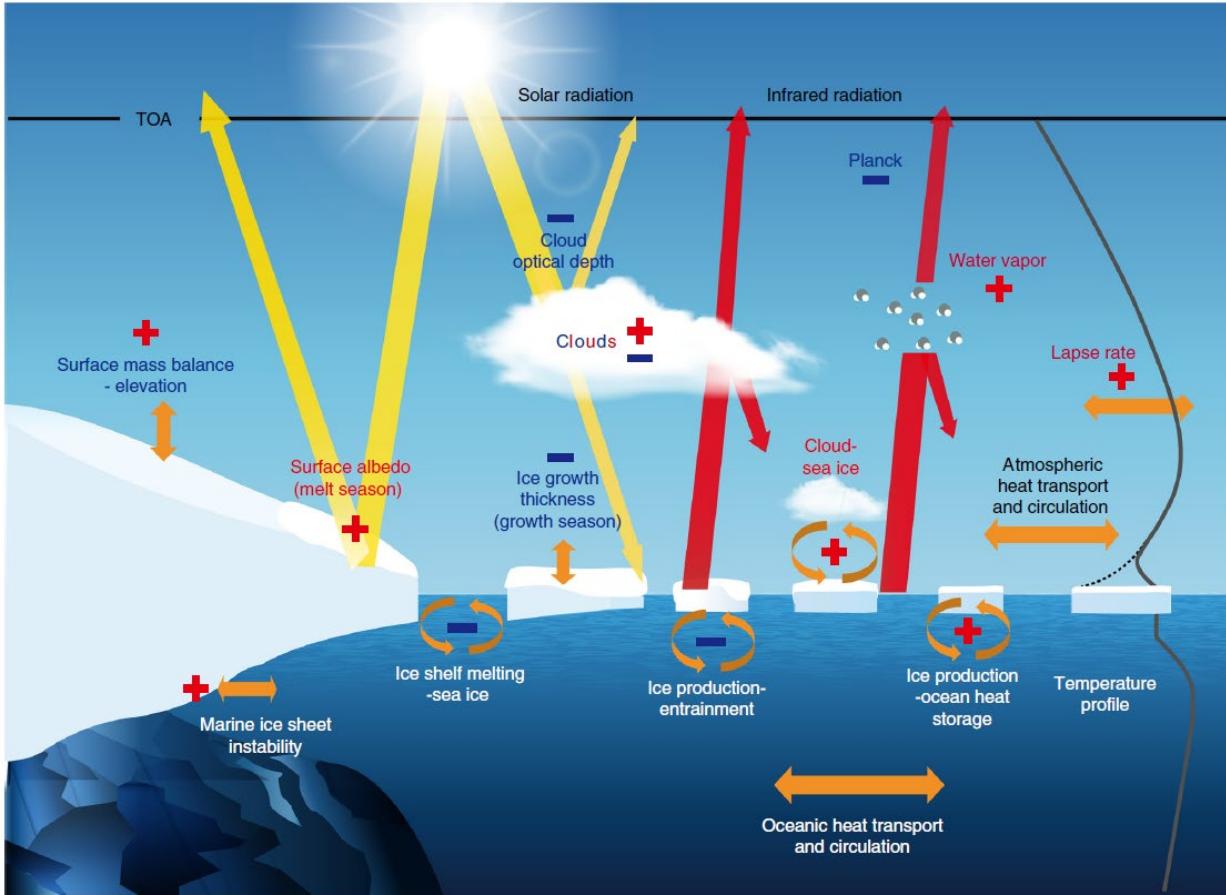
Δ ocean mixed layer depth



Swingedouw et al.,  
Ann. N.Y. Acad. Sci.,  
2021,  
10.1111/nyas.14659

No occurrence for  
ssp585

# Arctic and Antarctic sea ice changes:



Goosse et al, 2018, Nat. Comm., DOI: 10.1038/s41467-018-04173-0

## Offshore permafrost and gas hydrate stability:

Depends on combination of T and p.

- >800 Gt of  $CH_4$  accumulated in the ESAS seabed? Potential for abrupt  $CH_4$  release.
- Source: destabilizing hydrates or from free gas accumulations beneath permafrost.

See Shakhova et al., Geosciences, 2019, 9, 251; doi:10.3390/geosciences9060251

See, e.g., Fabbri et al., Environ. Sci. Technol., 2021, SI,  
<https://dx.doi.org/10.1021/acs.est.0c02928>

## Arctic summer sea ice loss:

Ice-albedo feedback.

(Comment: but heat loss at open waters.)

## Abrupt sea ice increase in Southern Ocean:

Warming reduces deep ocean convection (Indian sector).

Formation of a fresh surface layer and hence of sea ice.

See, e.g., Bathiany et al., J. Clim., 2016, doi: 10.1175/JCLI-D-15-0466.1

## Arctic winter sea ice loss:

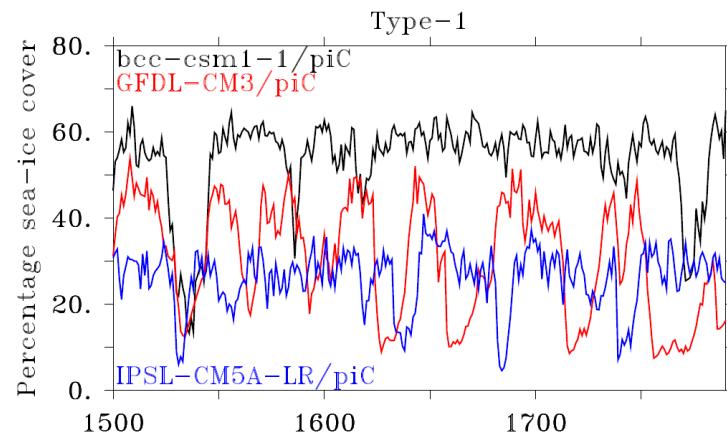
“... Sea ice areal coverage remains high as long as sea ice still forms, and then drops to zero wherever the ocean warms sufficiently to no longer form ice during winter. ...”

## “How reversible is sea ice loss?”

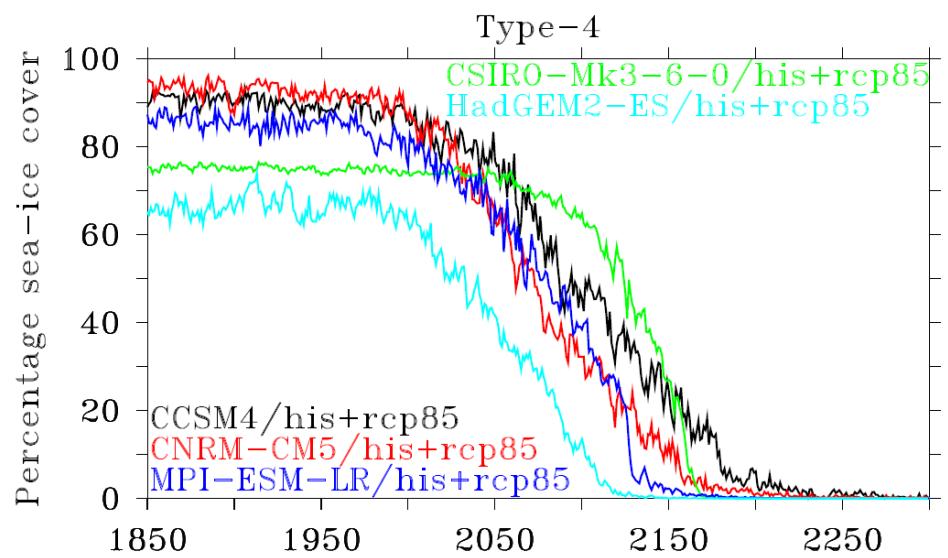
See: J. K. Ridley, J. A. Lowe, and H. T. Hewitt,  
*The Cryosphere*, 6, 193–198, 2012  
doi:10.5194/tc-6-193-2012

# Arctic and Antarctic sea ice, abrupt changes from models:

Type 1: Unforced  
Bimodal Switches  
in Sea Ice Cover.  
*(Southern Ocean)*.

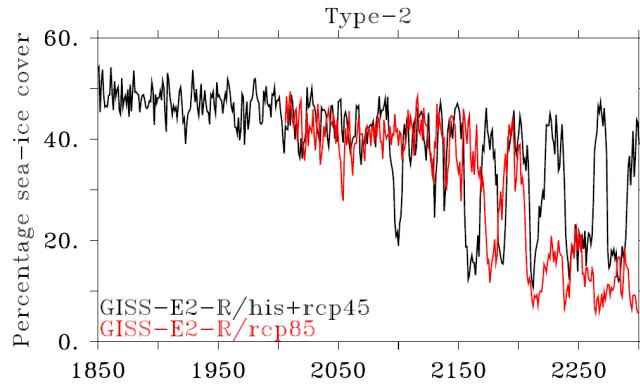


Type 4: Arctic  
Winter Sea Ice  
Collapse.  
*(Arctic Ocean)*.

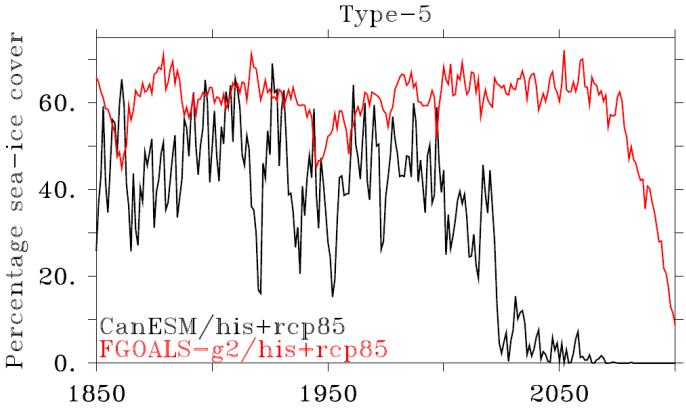


Drijfhout et al., PNAS, 2015,  
[www.pnas.org/cgi/doi/10.1073/pnas.1511451112](http://www.pnas.org/cgi/doi/10.1073/pnas.1511451112)

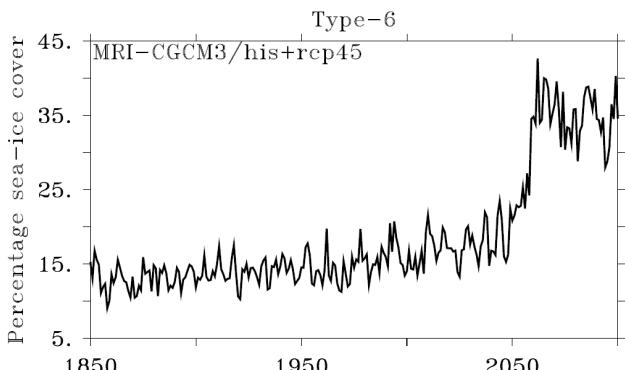
Type 2: Forced  
Bimodal  
Switches in  
Sea Ice Cover.  
*(Southern  
Ocean)*.



Type 5: Abrupt  
Regional Sea  
Ice Loss.  
*(Various  
regions)*.



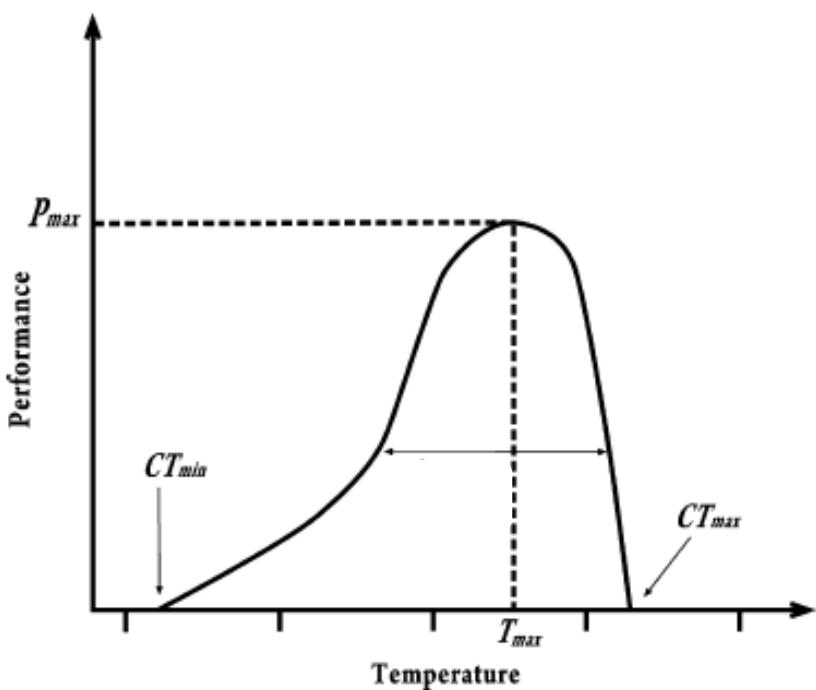
Type 6:  
Abrupt Sea  
Ice Increase.  
*(Southern  
Ocean)*.



# Environmental thresholds:

## Warming

Asymmetric thermal performance curve of organisms



Monaco & Helmuth, Advances in Marine Biology, Vol. 60, 2011  
DOI: 10.1016/B978-0-12-385529-9.00003-2 A

## Ocean acidification

$\text{CaCO}_3$  saturation levels for different polymorphs

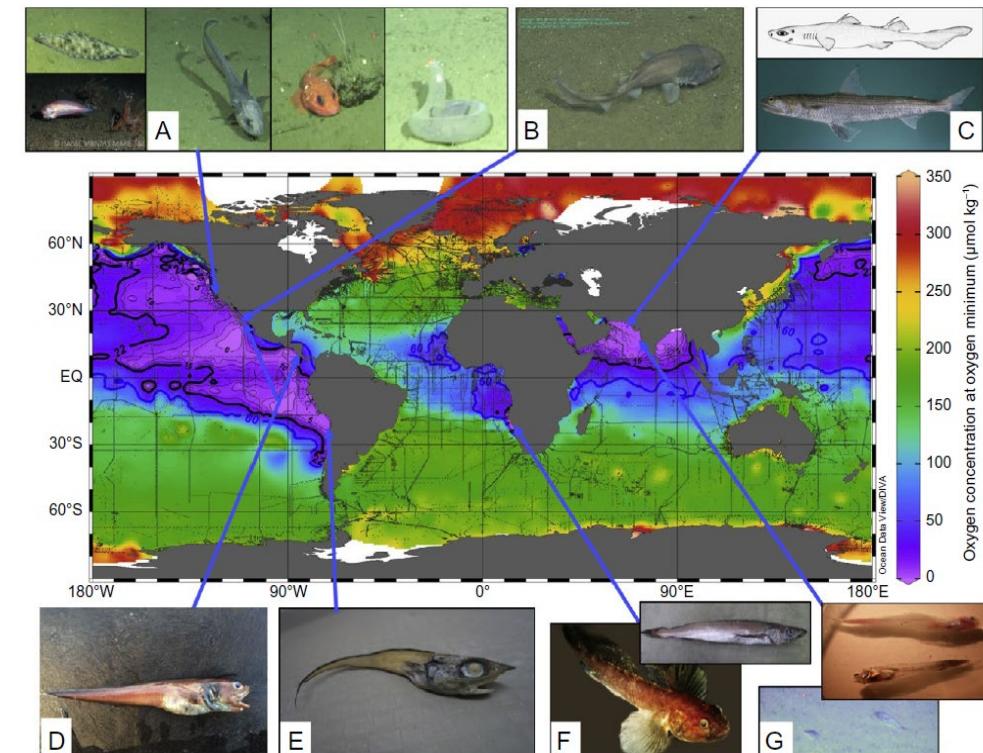
Taxa	Response	Mean Effect
Calciifying algae	Survival Calcification Growth Photosynthesis Abundance	-28% -80%
Corals	Survival Calcification Growth Photosynthesis Abundance	-32% -47%
Coccolithophores	Survival Calcification Growth Photosynthesis Abundance	-23% -25%
Mollusks	Survival Calcification Growth Development Abundance	-34% -40% -17% -25%
Echinoderms	Survival Calcification Growth Development Abundance	-10% -11%
Crustaceans	Survival Calcification Growth Development Abundance	
Fish	Survival Calcification Growth Development Abundance	
Fleshy algae	Survival Calcification Growth Photosynthesis Abundance	+22%
Seagrasses	Survival Calcification Growth Photosynthesis Abundance	
Diatoms	Survival Calcification Growth Photosynthesis Abundance	+17% +12%

Legend:  
 Not tested or too few studies (Grey)  
 Enhanced <25% (Green)  
 95% CI overlaps 0 (Yellow)  
 Reduced <25% (Pink)  
 Reduced >25% (Red)

For -0.5 pH change

## De-oxygenation

Lethal  $\text{O}_2$  boundary for fishes ca.  $<60 \mu\text{mol/L}$

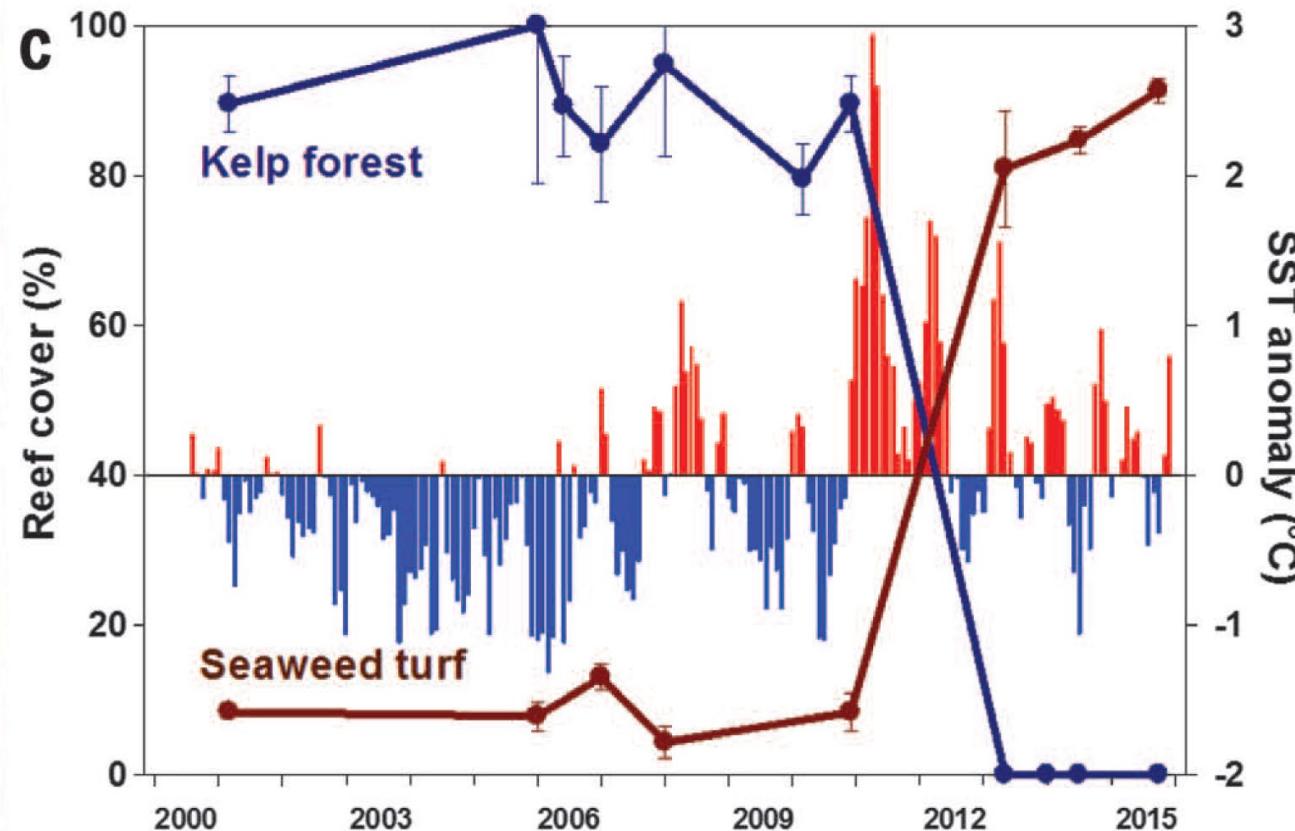
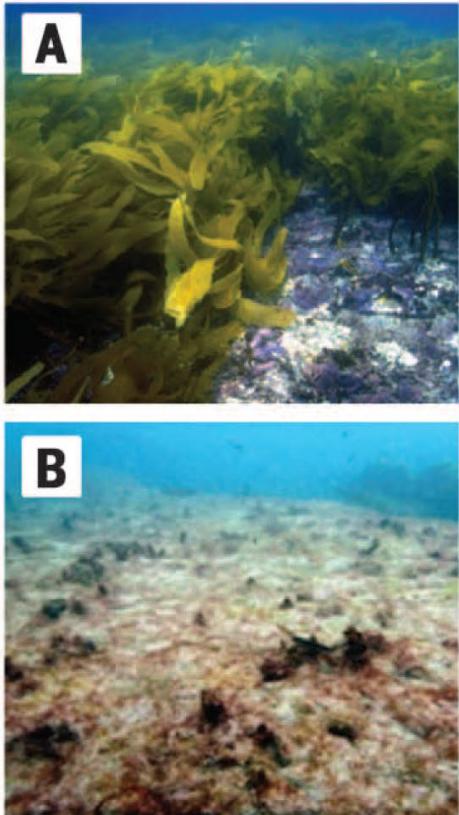


Kroeker et al., Global Change Biology (2013) 19, 1884–1896, doi: 10.1111/gcb.12179

Gallo & Levin, 2016, Advances in Marine Biology, Vol. 74, ch. 3, <http://dx.doi.org/10.1016/bs.amb.2016.04.001>

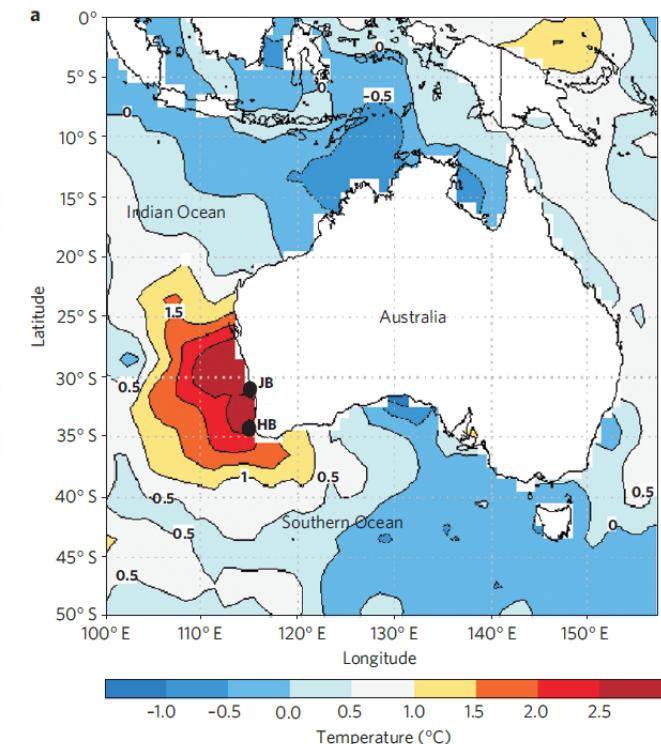
# Ocean heat waves trigger change:

## TROPICAL COASTAL ECOSYSTEM



Wernberg et al., Science, 2016,  
doi: 10.1126/science.aad8745

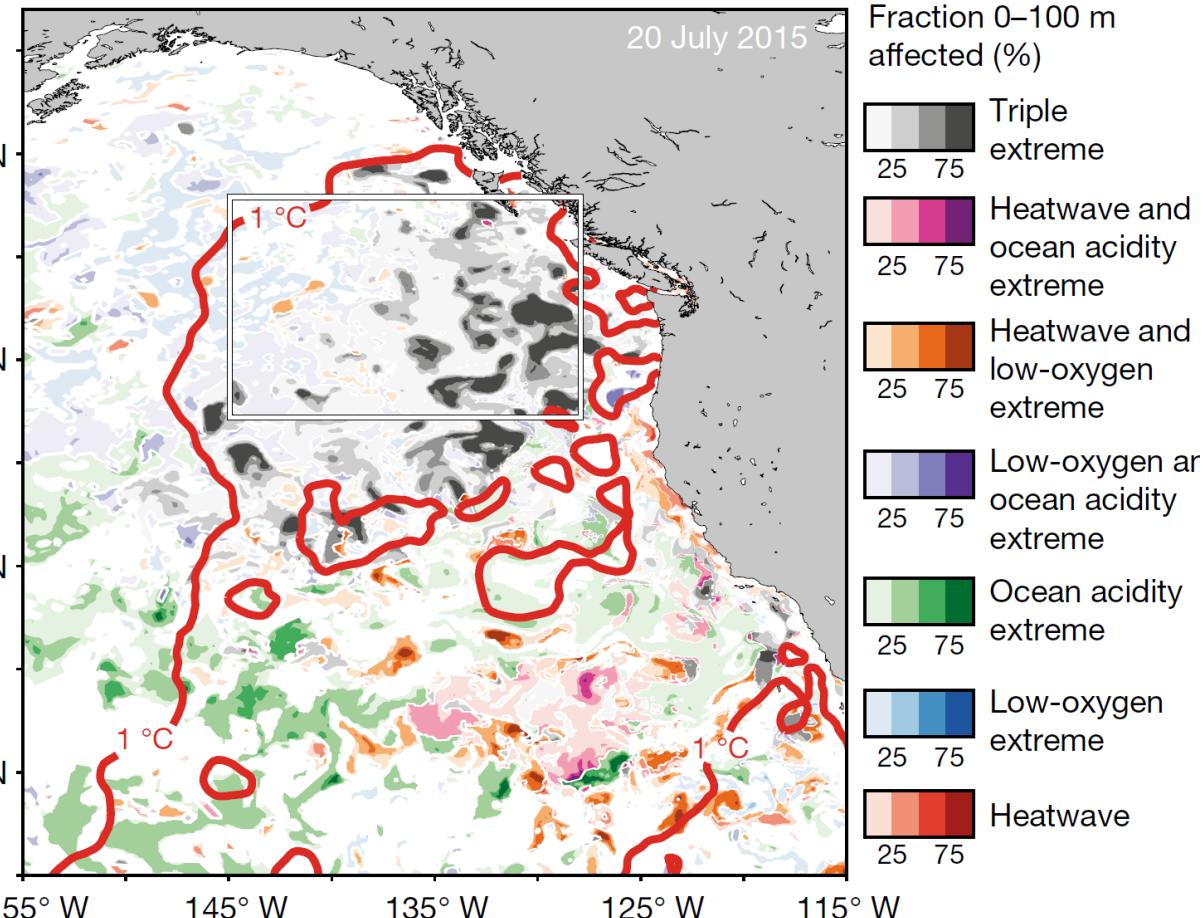
Wernberg et al., *Nature Clim. C.*,  
2012, doi:10.1038/NCLIMATE1627



2011 heatwave

# Marine extreme events:

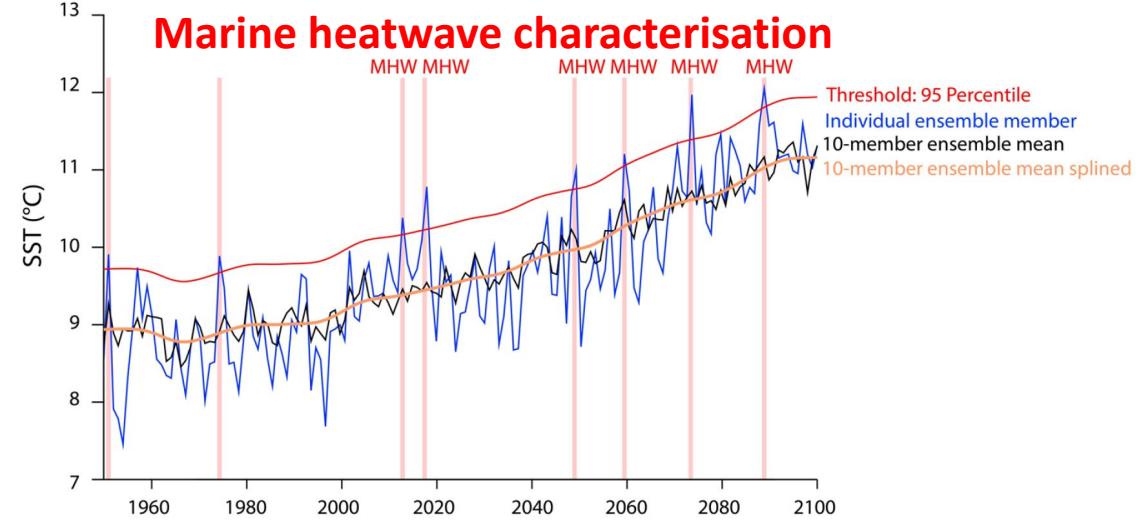
Compound extremes N.E. Pacific  
(model simulation)



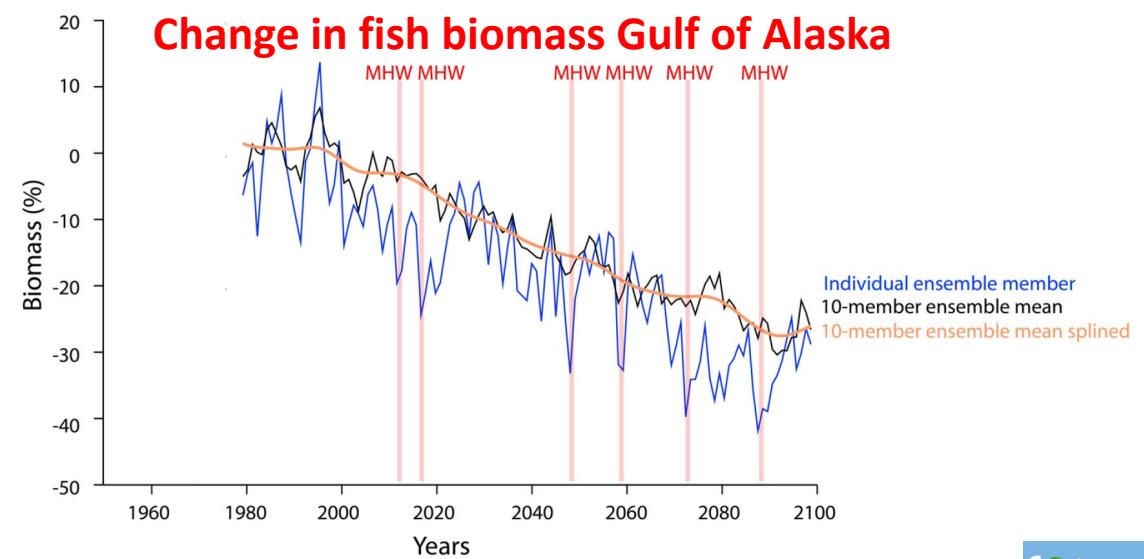
Gruber et al., *Nature*, 2021,  
<https://doi.org/10.1038/s41586-021-03981-7>



Earth system model with a fish impact model



Change in fish biomass Gulf of Alaska

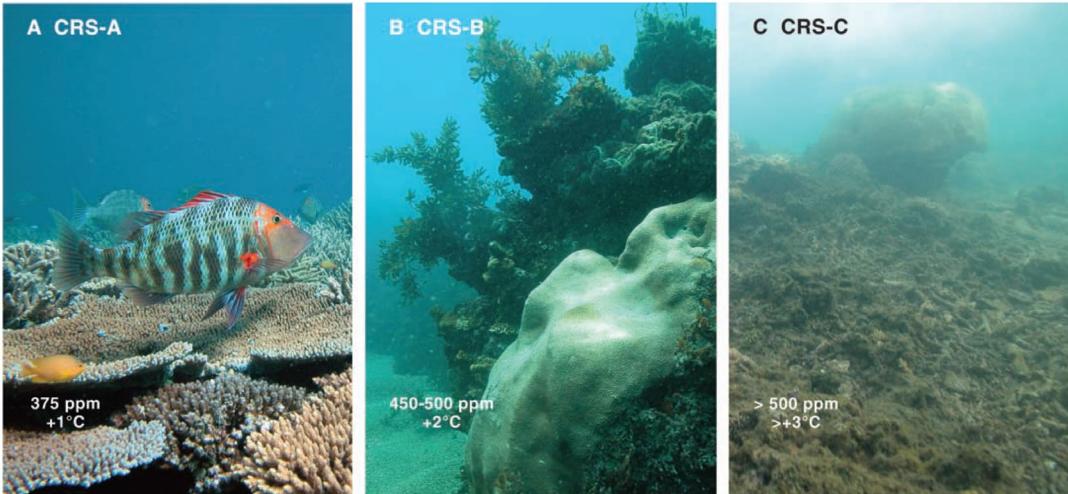


Cheung & Frölicher, *Sci. Reports*,  
2020, <https://doi.org/10.1038/s41598-020-63650-z>



# Loss of $\text{CaCO}_3$ , living organisms - examples:

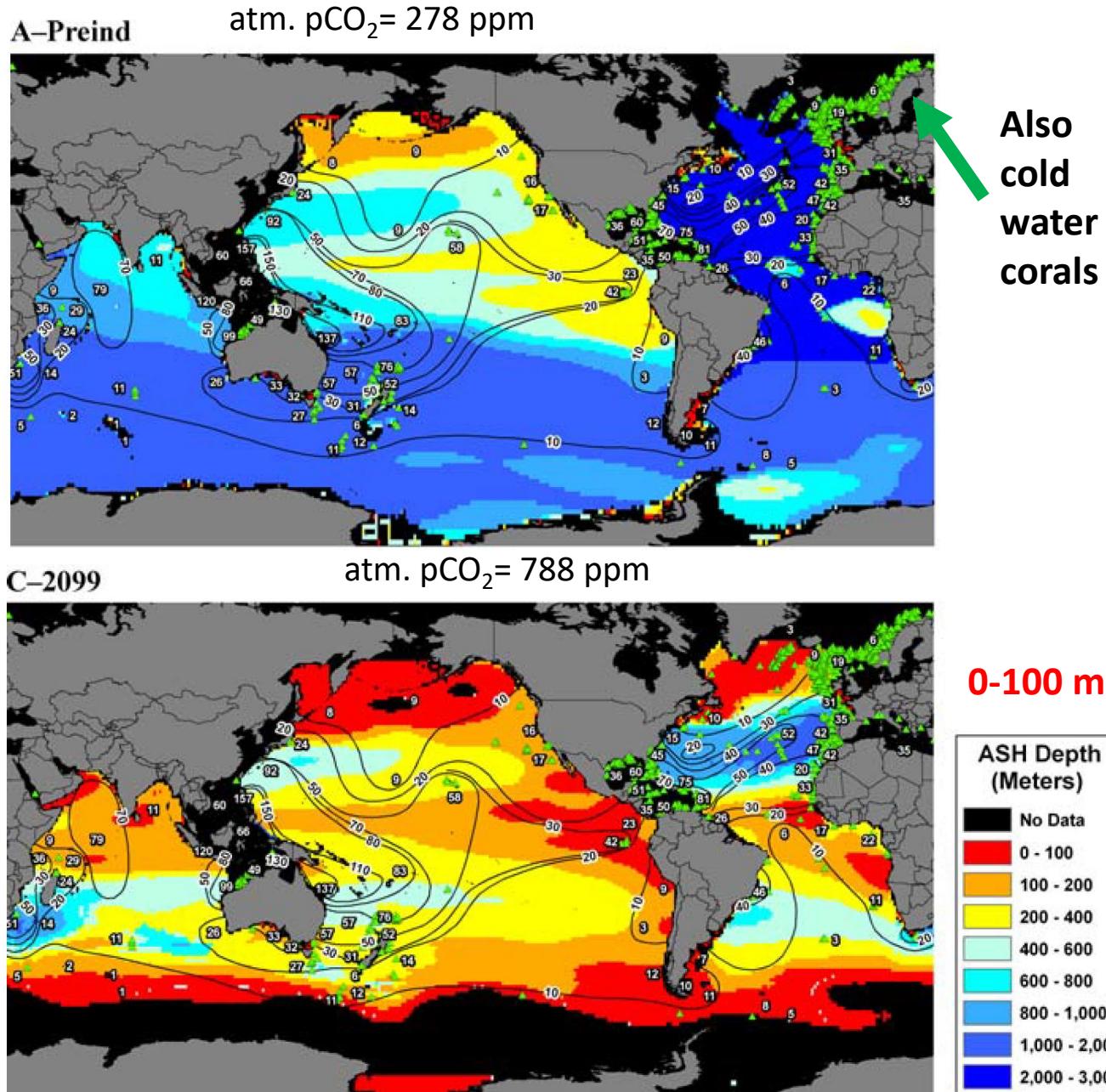
Warm water corals:  $\Delta T$ ,  $\Delta \text{pH}$



Hoegh-Guldberg et al., Science, 2007,  
doi: 10.1126/science.1152509

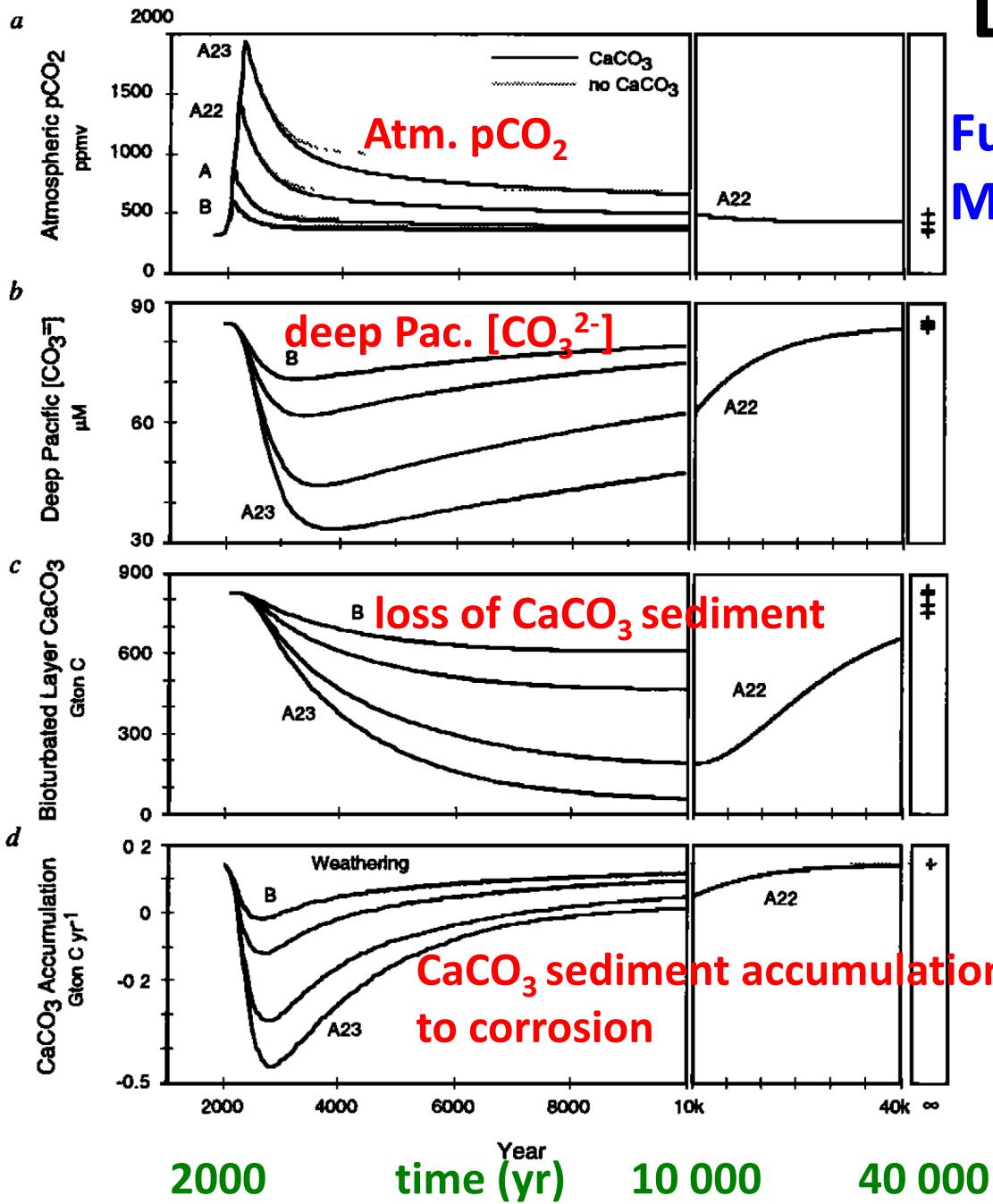
Showling of  
aragonite  
saturation  
horizon

Guinotte & Fabry,  
Ann N Y Acad Sci,  
2008, doi:  
10.1196/annals.1439.013



# Loss of sedimentary

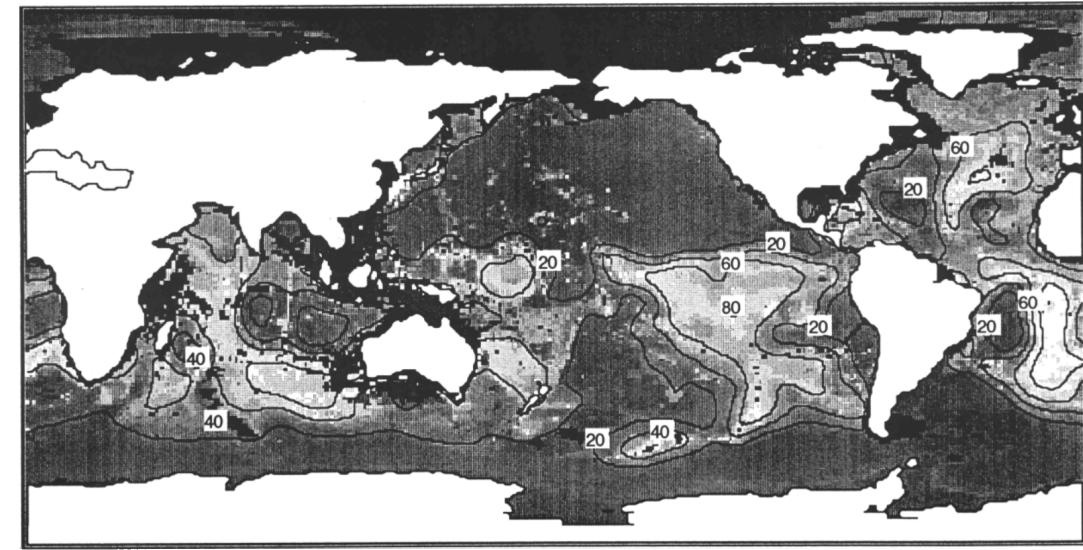
Future,  
MODEL



CaCO<sub>3</sub>:

Quasi-irreversibility on human time scales,  
huge legacy effect also for atm. pCO<sub>2</sub>

CaCO<sub>3</sub> surface sediment like “snow” on ocean mountains,  
this surface sediment is already doomed to vanish for long



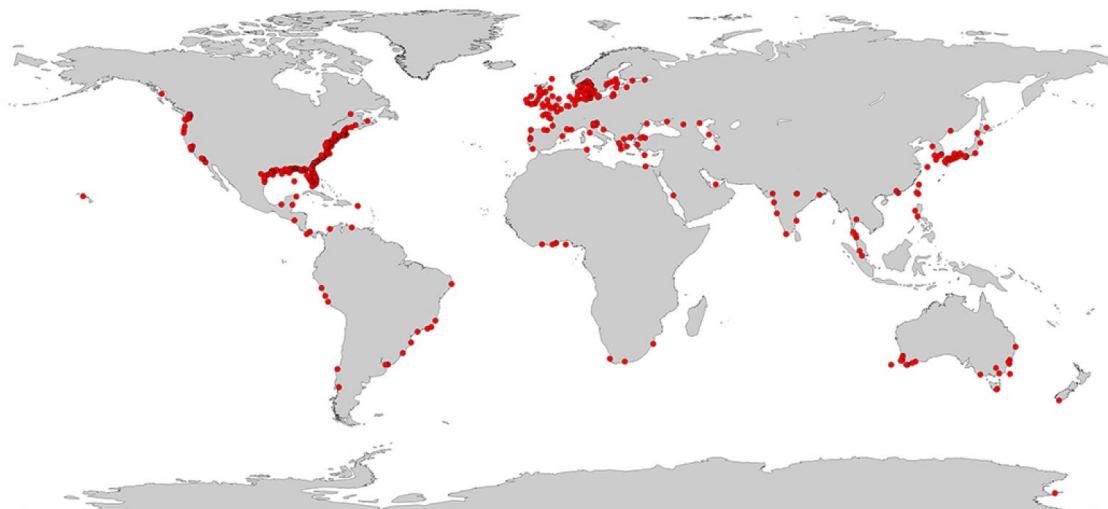
Archer, *Glob. Biogeochem. Cy.*, 1996,  
<https://doi.org/10.1029/95GB03016>

Present,  
OBSERVATION

Archer et al., *Geophys. Res. L.*, 1997,  
[doi https://doi.org/10.1029/97GL00168](https://doi.org/10.1029/97GL00168)

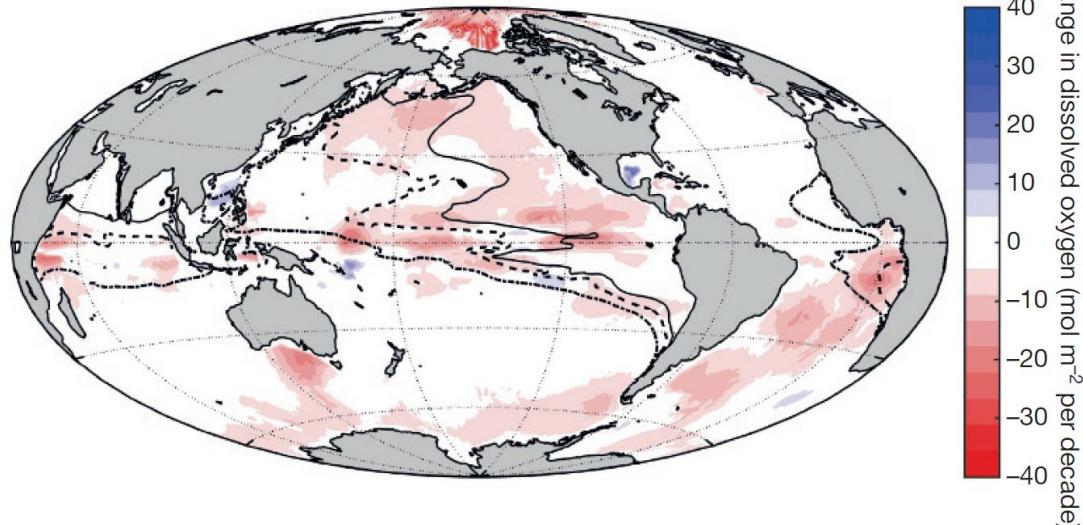
# O<sub>2</sub> loss:

Coastal waters where [O<sub>2</sub>] ≤ 61 mmol kg<sup>-1</sup> past 50 yrs

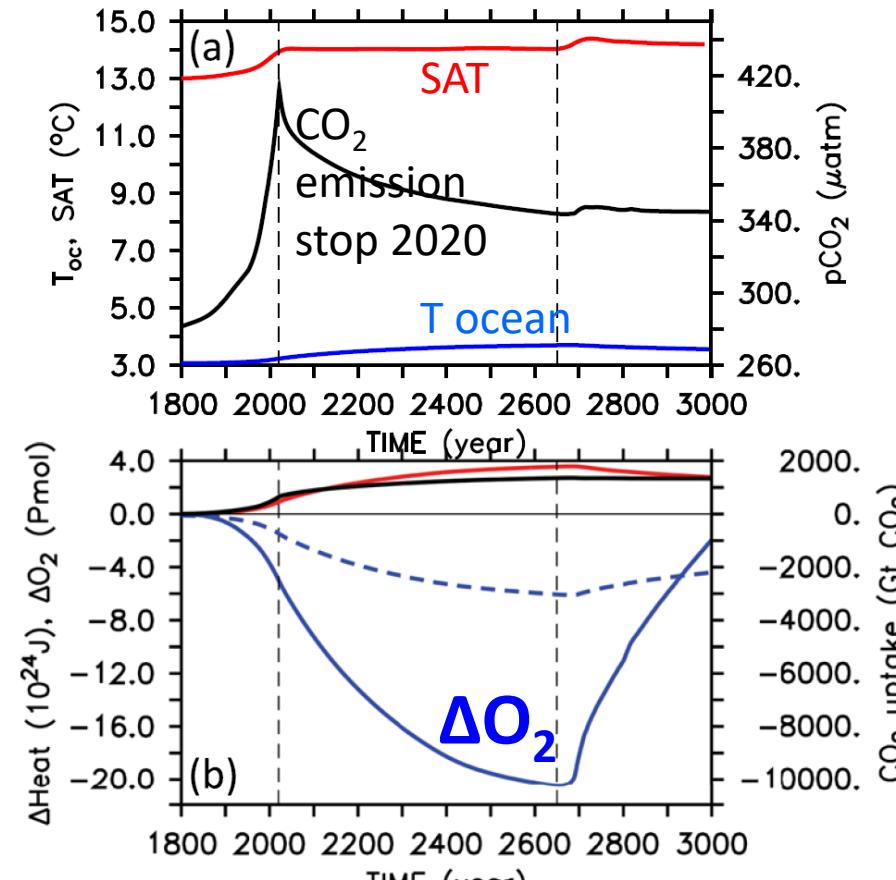


Breitburg et al., *Science*, 2018, doi: 10.1126/science.aam7240

Decadal ΔO<sub>2</sub> mol/m<sup>2</sup> 1960-2010



Schmidtko et al., *Nature*, 2017, doi:10.1038/nature21399



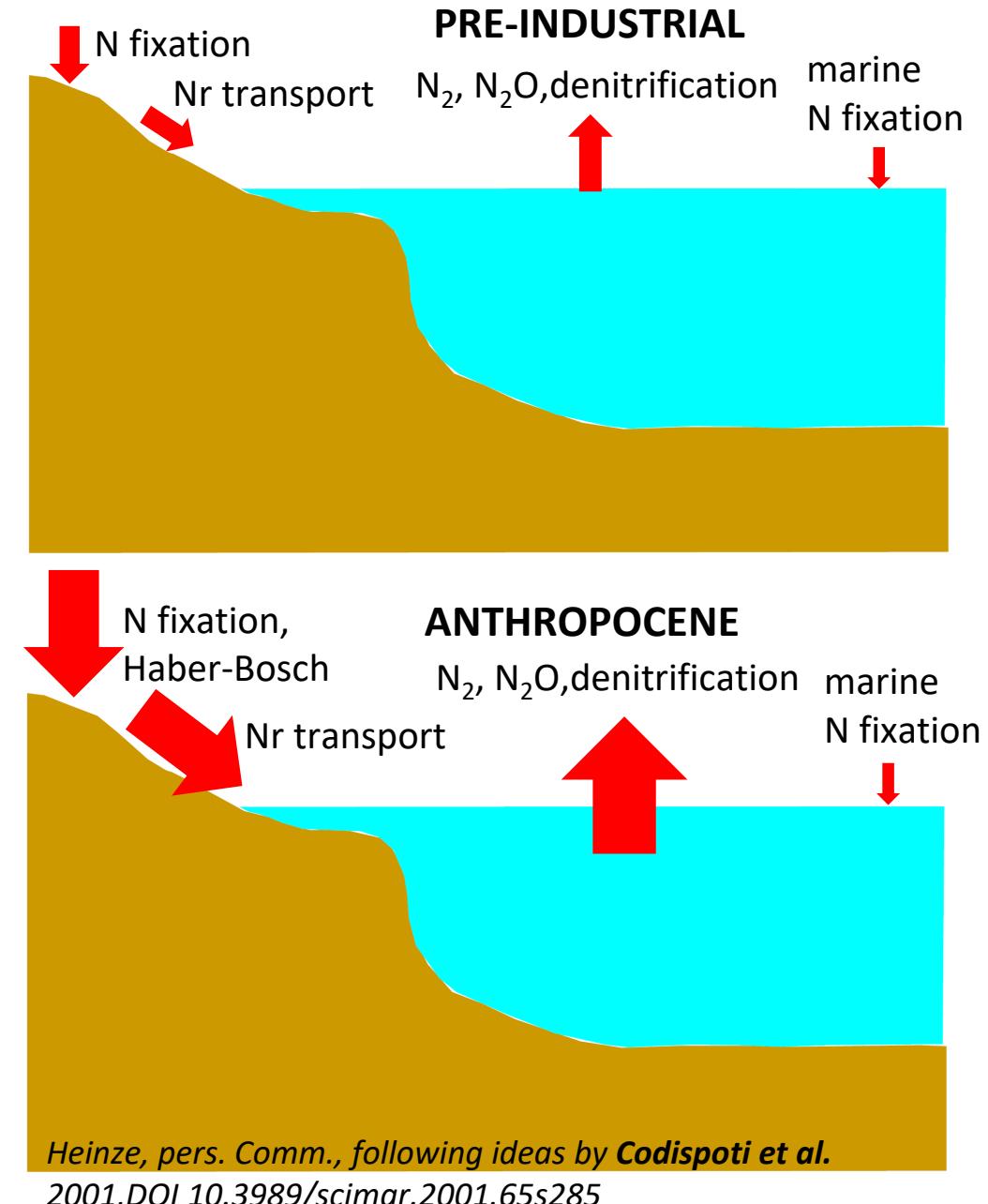
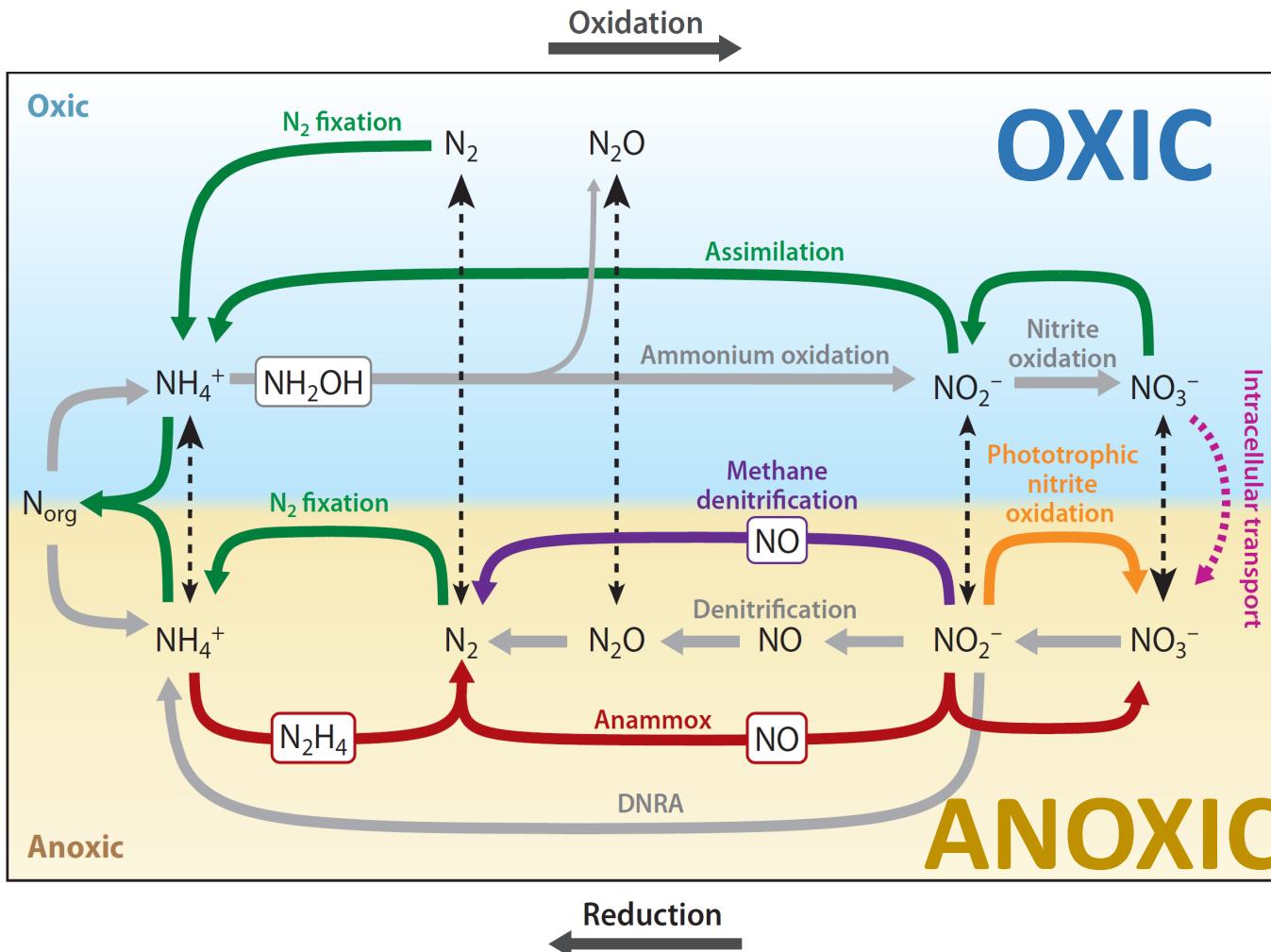
Quasi-irreversibility on human time scales



Oschlies et al., *Nature Comm.*, 2021, doi: <https://doi.org/10.1038/s41467-021-22584-4>

# $\text{N}_2\text{O}$ production under low $\text{O}_2$ :

Switch between biogeochemical and ecological regimes!



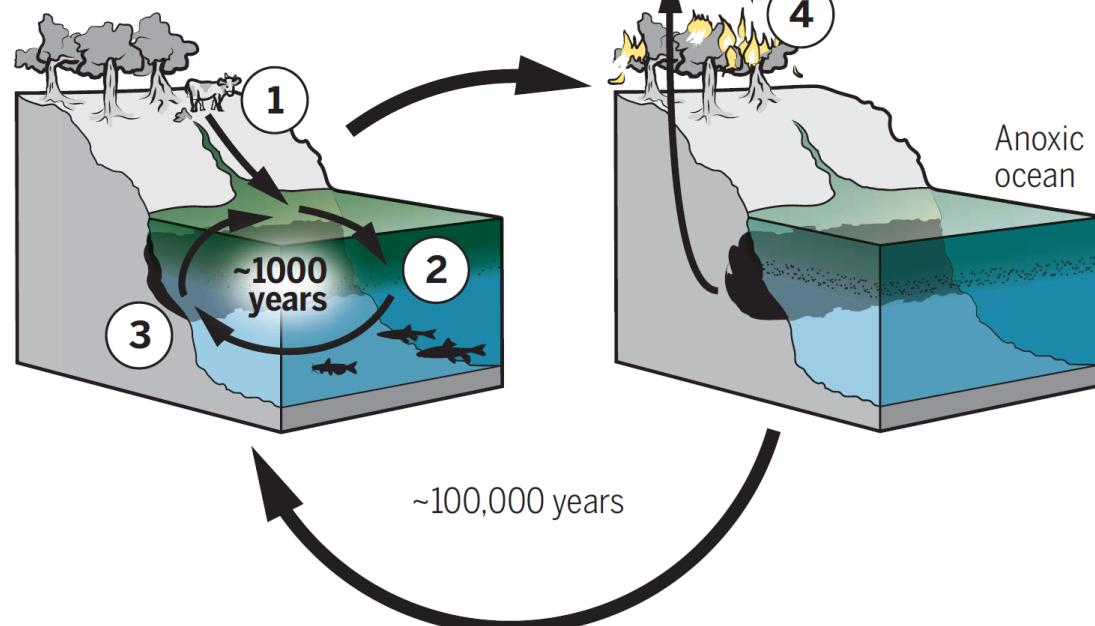
# $\text{PO}_4^{3-}$ & anoxia runaway following deoxygenation:

## Oxygen crises in the ocean

Major environmental crises can lead to oxygen depletion (anoxia) throughout the ocean's oxygen minimum zones, which become deadly for animal life.

### Positive feedbacks

High phosphate inputs (1) enhance marine production, reducing oxygen in the OMZ (2). The resulting phosphorus release from sediments (3) further increases marine production and carbon burial.

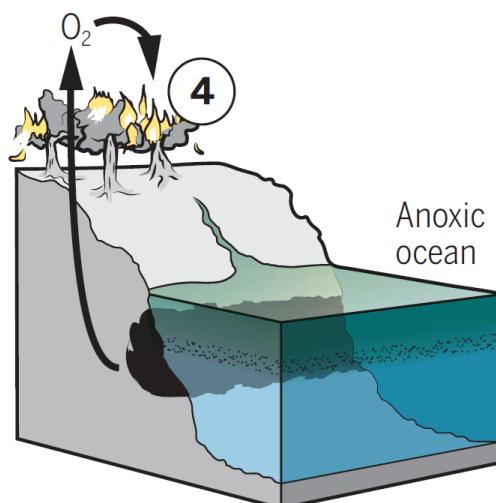


### Closing the cycle

Over time periods of 100,000 years or more, the oceans recover from their deoxygenated state.

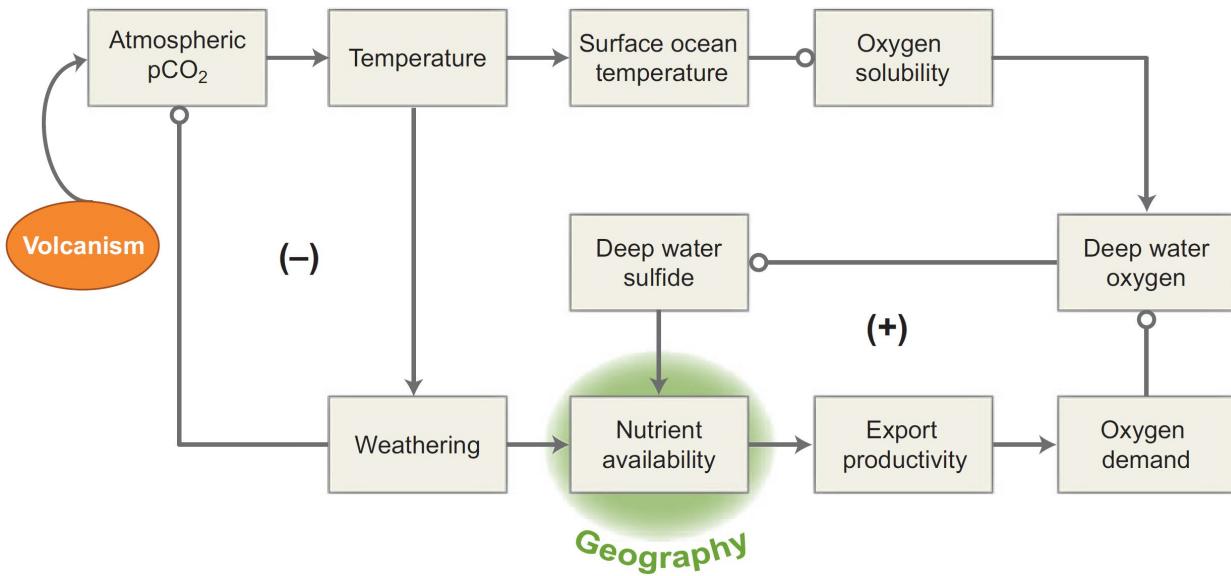
### Limiting processes

More burial of carbon leads to an increase in atmospheric oxygen, causing more wildfires that reduce forest vegetation (4). This lowers the phosphorus input to the oceans.



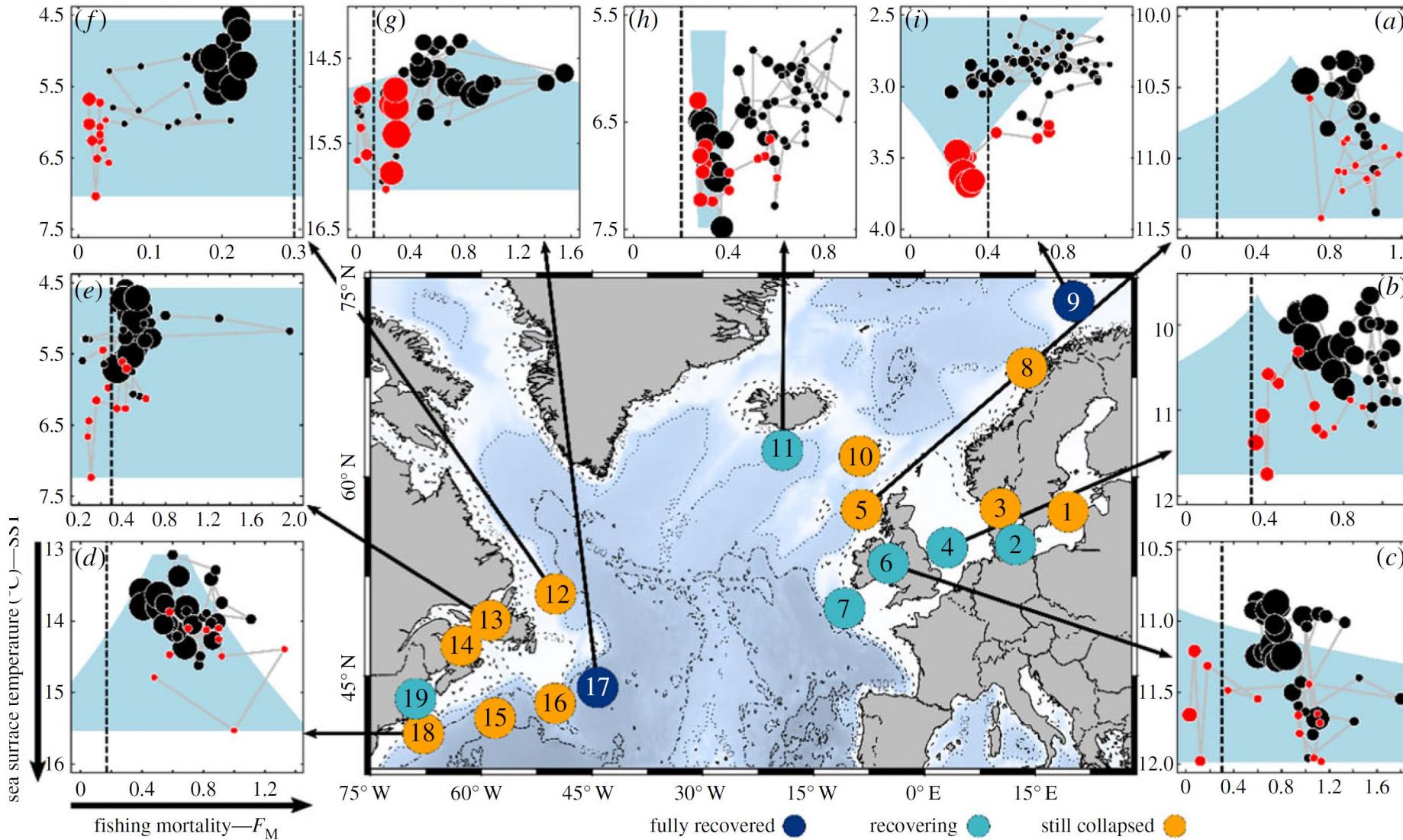
Watson, 2016, *Science*, doi:  
10.1126/science.aa j2321

### Climate



Meyer & Kump, 2008, *Annu. Rev. Earth Planet. Sci.*, doi:10.1146/annurev.earth.36.031207.124256

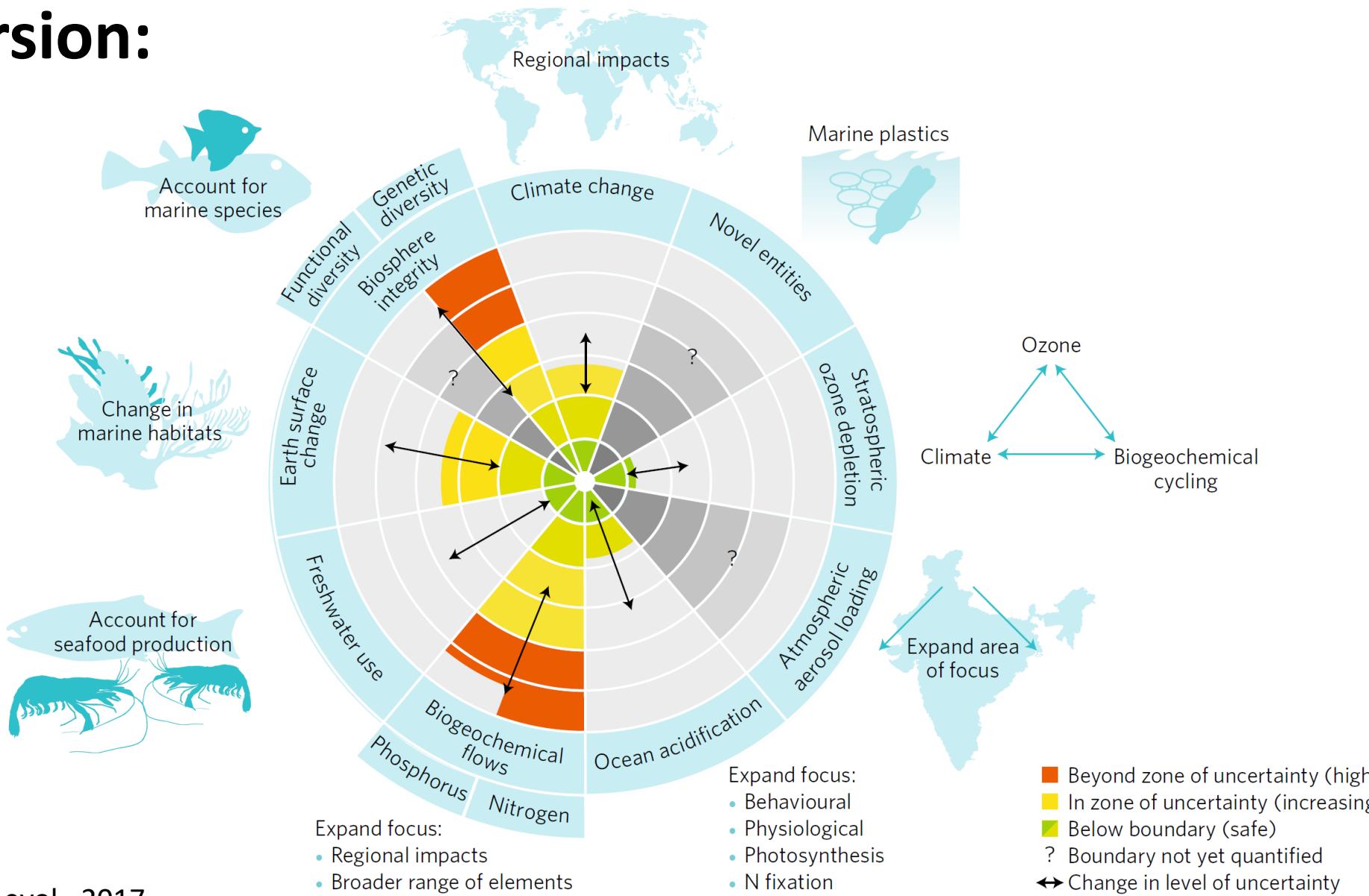
# Fish impact, tipping through climate change and overfishing:



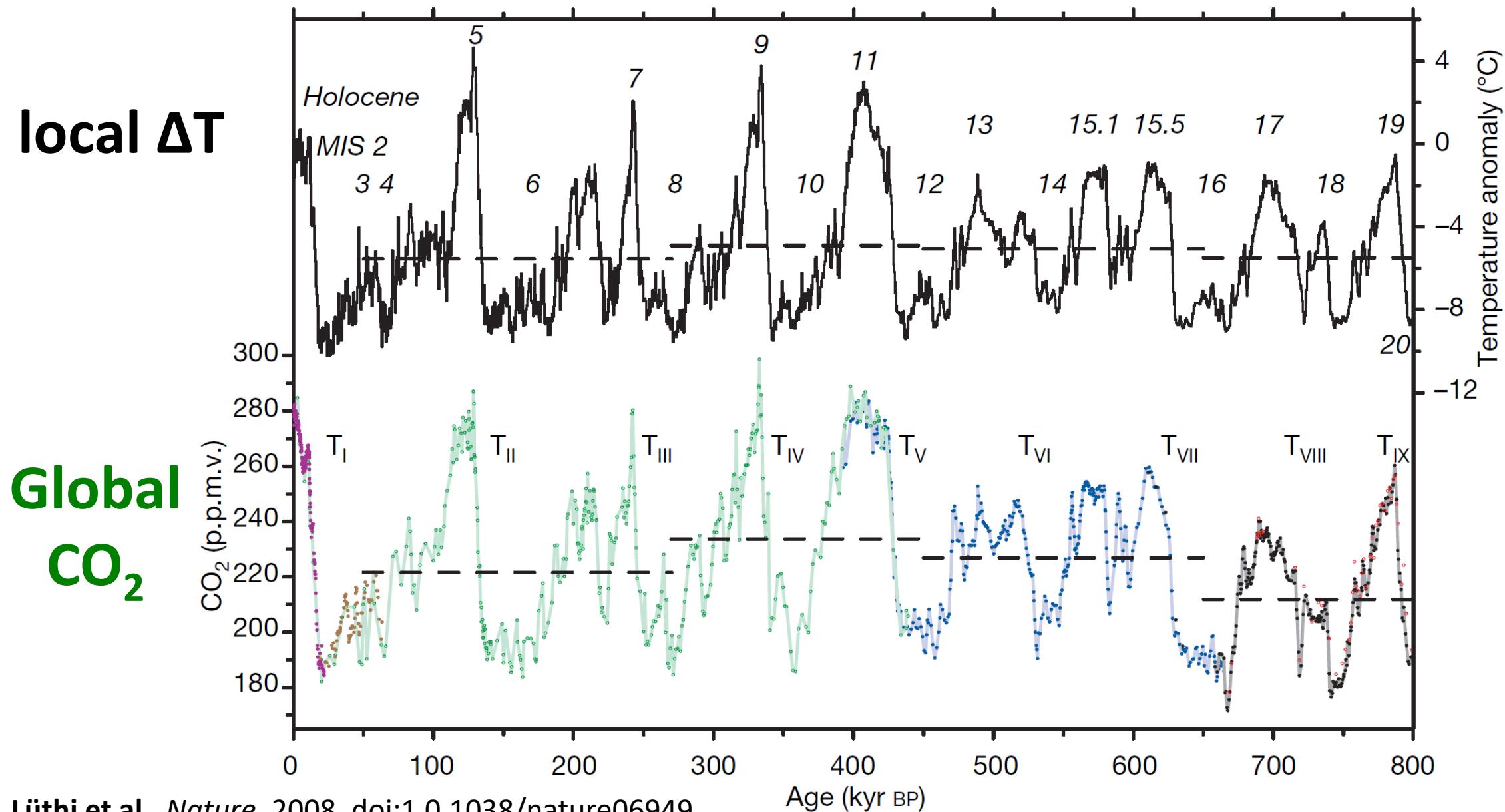
Statistical ("cusp") analysis of 19 Atlantic Cod stocks ca. 1960-2010

(cusp involving turning point as alternative to change point analysis)

# Planetary boundaries and safe operating space – marine version:



# Glacial-Interglacial changes: ...causes (?) ... stable (?) states ...



Change in system component	Potentially abrupt	Irreversibility if forcing reversed (time scales indicated)	Impacts on natural and human systems; global vs. regional vs. local	Projected likelihood and/or confidence level in 21st century under scenarios considered
<b>Ocean</b>				
Atlantic Meridional Overturning Circulation (AMOC) collapse (Section 6.7)	Yes	Unknown	Widespread; increased winter storms in Europe, reduced Sahelian rainfall and agricultural capacity, variations in tropical storms, increased sea levels on Atlantic coasts	<i>Very unlikely</i> , but physically plausible
Subpolar gyre (SPG) cooling (Section 6.7)	Yes	Irreversible within decades	Similar to AMOC impacts but considerably smaller	<i>Medium confidence</i>
Marine heatwave (MHW) increase (Section 6.4)	Yes	Reversible within decades to centuries	Coral bleaching, loss of biodiversity and ecosystem services, harmful algal blooms, species redistribution	<i>Very likely (very high confidence)</i> for physical change <i>High confidence</i> for impacts
Arctic sea ice retreat (Section 3.3)	Yes	Reversible within decades to centuries	Coastal erosion in Arctic (may take longer to reverse), impact on mid-latitude storms ( <i>low confidence</i> ); rise in Arctic surface temperatures ( <i>high confidence</i> )	<i>High confidence</i>
Ocean deoxygenation and hypoxic events (Section 5.2)	Yes	Reversible at surface, but irreversible for centuries to millennia at depth	Major changes in ocean productivity, biodiversity and biogeochemical cycles	<i>Medium confidence</i>
Ocean acidification (Section 5.2)	Yes	Reversible at surface, but irreversible for centuries to millennia at depth	Changes in growth, development, calcification, survival and abundance of species, for example, from algae to fish	<i>Virtually certain (very high confidence)</i>

**IPCC SROCC** Collins M., M. Sutherland, L. Bouwer, S.-M. Cheong, T. Frölicher, H. Jacot Des Combes, M. Koll Roxy, I. Losada, K. McInnes, B. Ratter, E. Rivera-Arriaga, R.D.

Susanto, D. Swingedouw, and L. Tibig, 2019: Extremes, Abrupt Changes and Managing Risk. In: IPCC Special Report on the Ocean and Cryosphere in a Changing Climate [H.-O. Pörtner, D.C. Roberts, V. Masson-Delmotte, P. Zhai, M. Tignor, E. Poloczanska, K. Mintenbeck, A. Alegría, M. Nicolai, A. Okem, J. Petzold, B. Rama, N.M. Weyer (eds.)]. In press.

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<b>Ocean deoxygenation and hypoxic events (Section 5.2)</b>	<b>Yes</b>		<b>Reversible at surface, but irreversible for centuries to millennia at depth</b>	
<b>Ocean acidification (Section 5.2)</b>	<b>Yes</b>		<b>Reversible at surface, but irreversible for centuries to millennia at depth</b>	
Ocean hypoxic events (Section 5.2)	Yes	for centuries to millennia at depth	biodiversity and biogeochemical cycles	<i>Medium confidence</i>
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Susanto, D. Swingedouw, and L. Tibig, 2019: Extremes, Abrupt Changes and Managing Risk. In: IPCC Special Report on the Ocean and Cryosphere in a Changing Climate [H.-O. Pörtner, D.C. Roberts, V. Masson-Delmotte, P. Zhai, M. Tignor, E. Poloczanska, K. Mintenbeck, A. Alegría, M. Nicolai, A. Okem, J. Petzold, B. Rama, N.M. Weyer (eds.)]. In press.