

האוניברסיטה העברית בירושלים  
THE HEBREW UNIVERSITY OF JERUSALEM

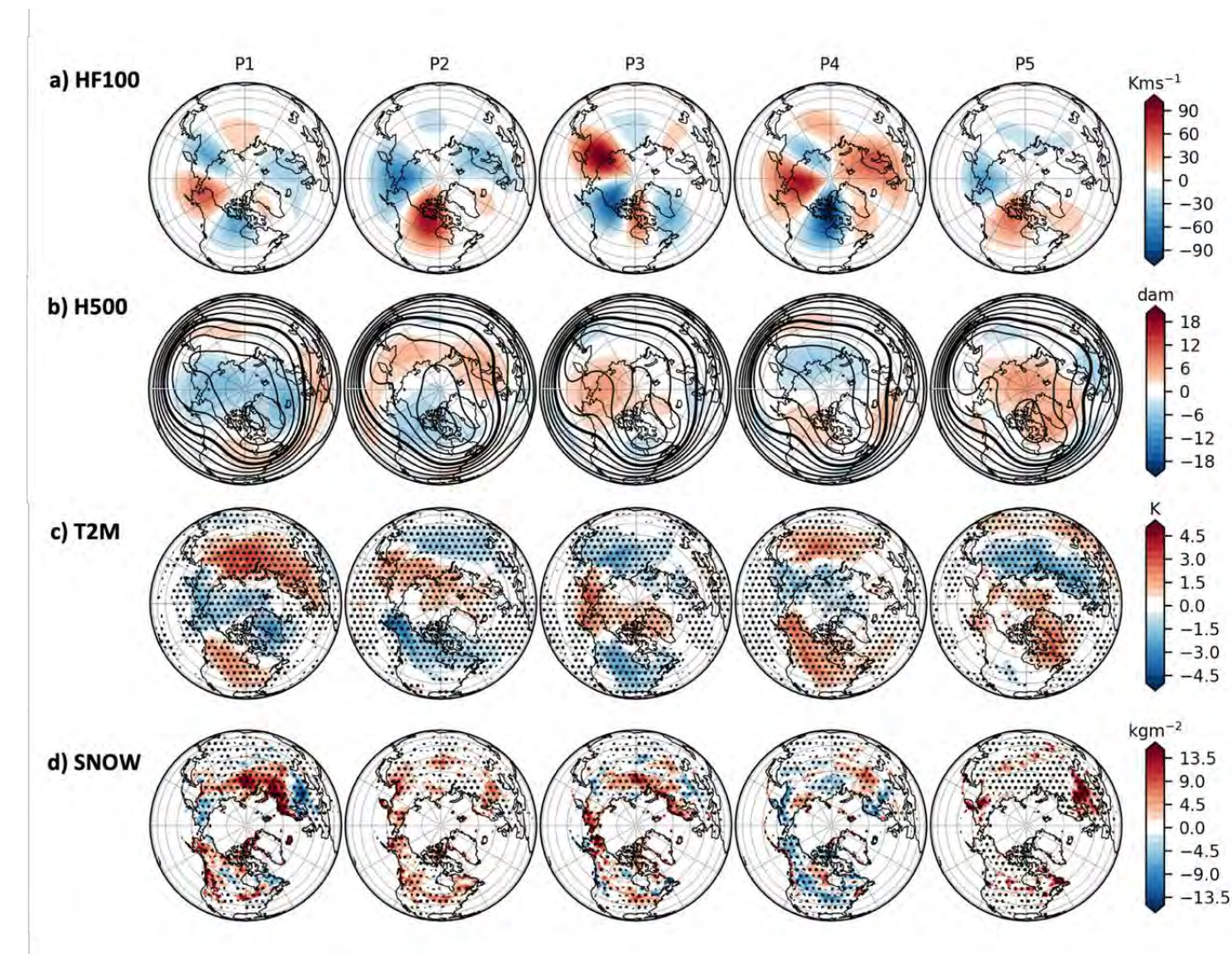
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## 1. Introduction

- Does the surface impact of polar vortex anomalies depend on the zonal structure of the vortex? Specifically, does the direction in which a polar vortex is displaced matter for its impact?
- We use the model of an idealized moist atmosphere (MiMA) to study the downward response to reflection and SSW events by applying asymmetric momentum torques to the zonal momentum budget.

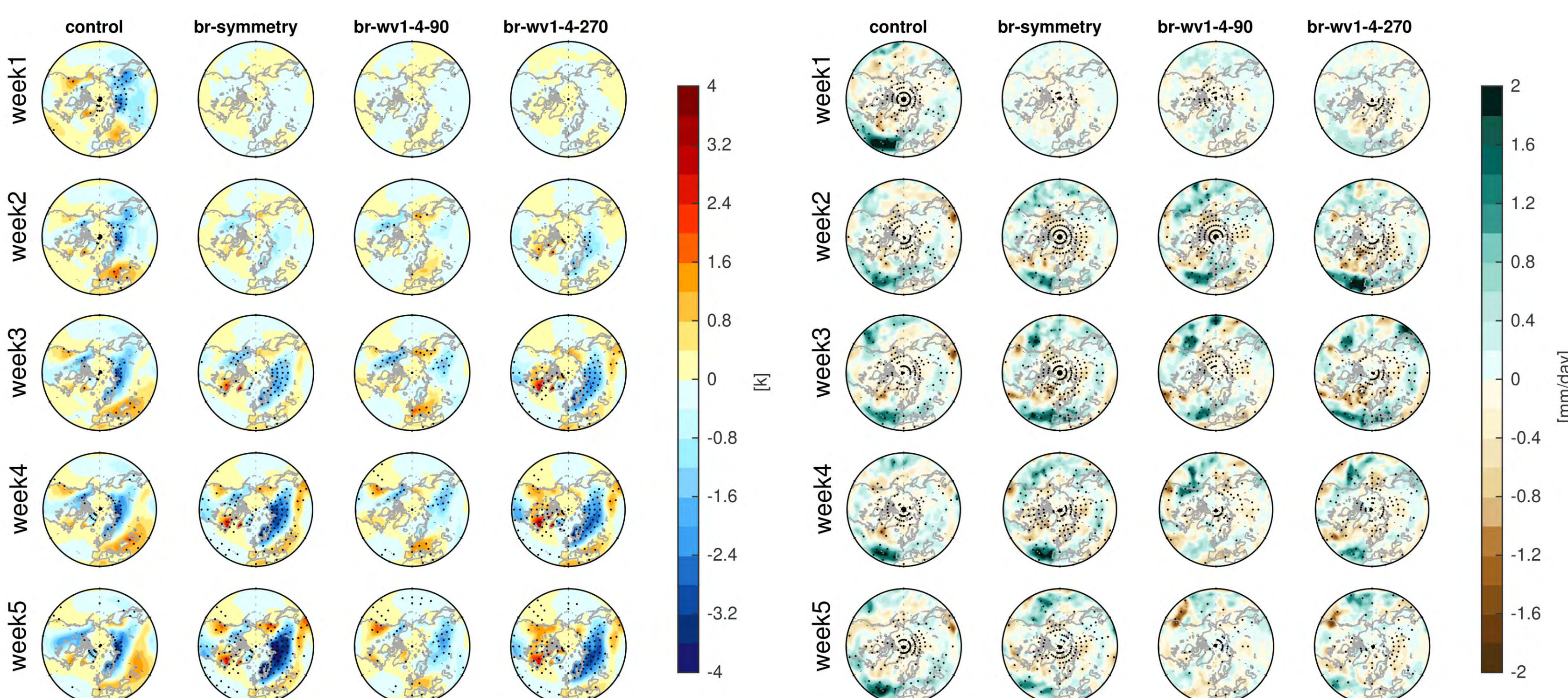
## 2. Research Background/Motivation



Agel, et.al. 2024 (submitted)

- P1 denotes strong vortex states while P2-P5 denotes moderate and weak vortex states.
- P2 and P3 can result in the cooling in North America and Europe.
- Our goal is to isolate the role of the vortex anomaly for the surface impact from any precursor wave pattern)

## 7. Surface Temperature & Precipitation



- Similar impacts in control and symmetric branch cases with cooling over north Eurasia and warmer Eastern Canada and subtropical Eurasia.
- The 90 case shows cooling over North America, while 270 case shows cooling over Eurasia.
- The rainfall dipole in Iberia/Scandinavia in control and symmetric is much weaker in 90 than in 270. In 90 case, the impact on Western North America much more pronounced.

## 3. MiMA experiments

- The applied momentum torque to model's zonal momentum budget:

$$F(\lambda, \phi, p, t) = \Lambda(\lambda) \Phi(\phi) H(p) \tau(t)$$

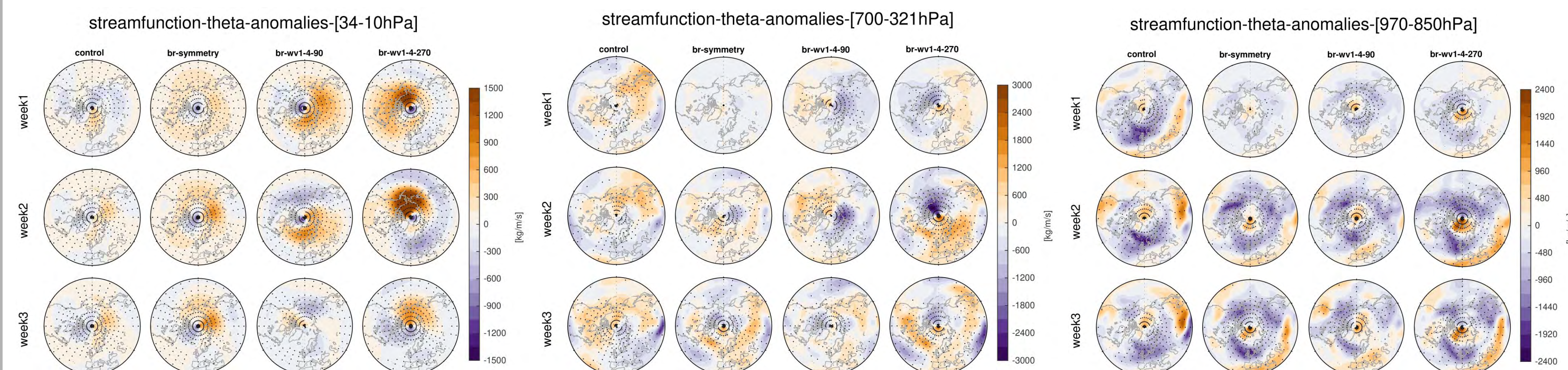
$$\Lambda(\lambda) = 1 + M_{FAS} * \sin(k * (\lambda - \lambda_0))$$

$$\Phi(\phi) = M_{FS} \sin(\pi \frac{\phi - \phi_L}{\phi_H - \phi_L})$$

$$H(p) = \begin{cases} \frac{P - P_b}{P_t - P_b}, & \text{if } P_b < P < P_t, \\ 1, & \text{if } P \leq P_t \\ 0, & \text{if } P \geq P_b \end{cases}$$

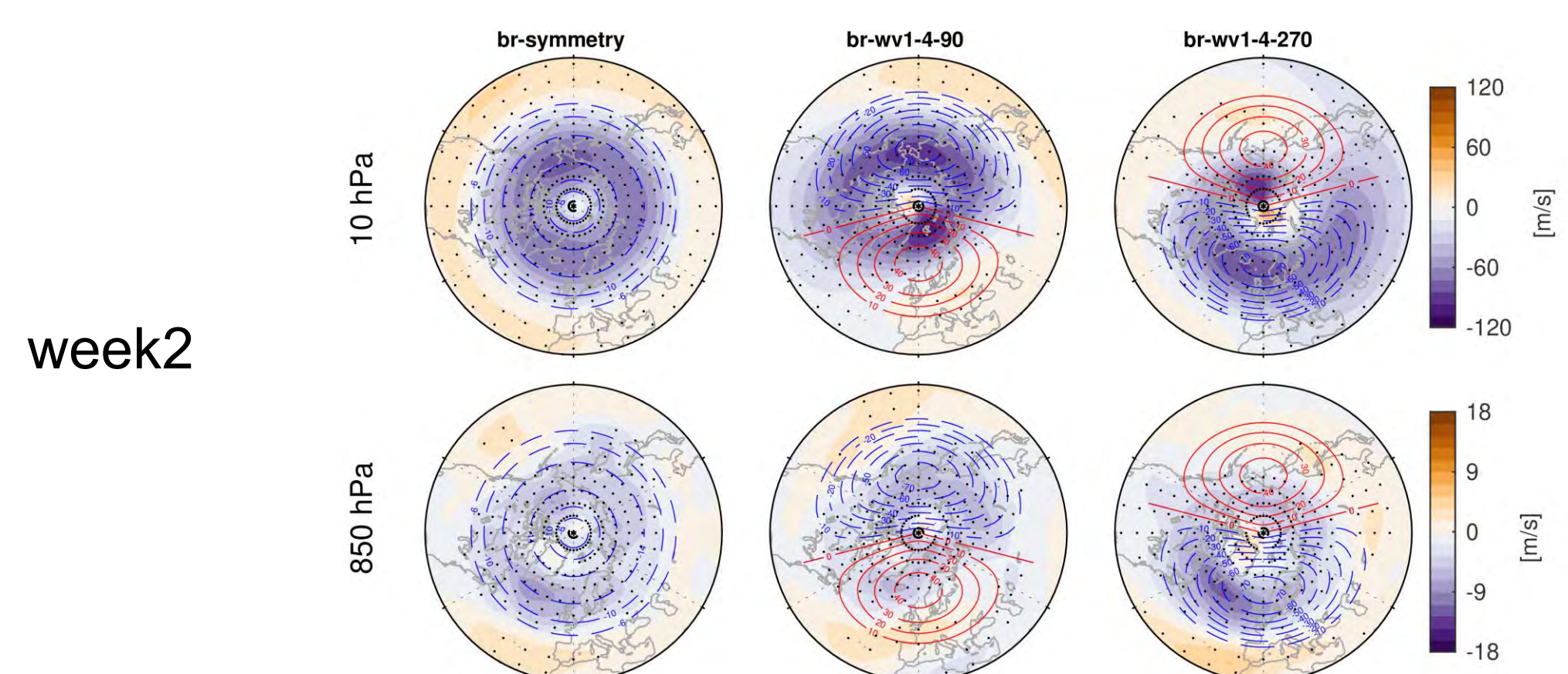
$$\tau(t) = \begin{cases} 1, & \text{if } 0 < t - t_0 \leq N_d \\ 0, & \text{otherwise} \end{cases}$$

## 8. Streamfunction



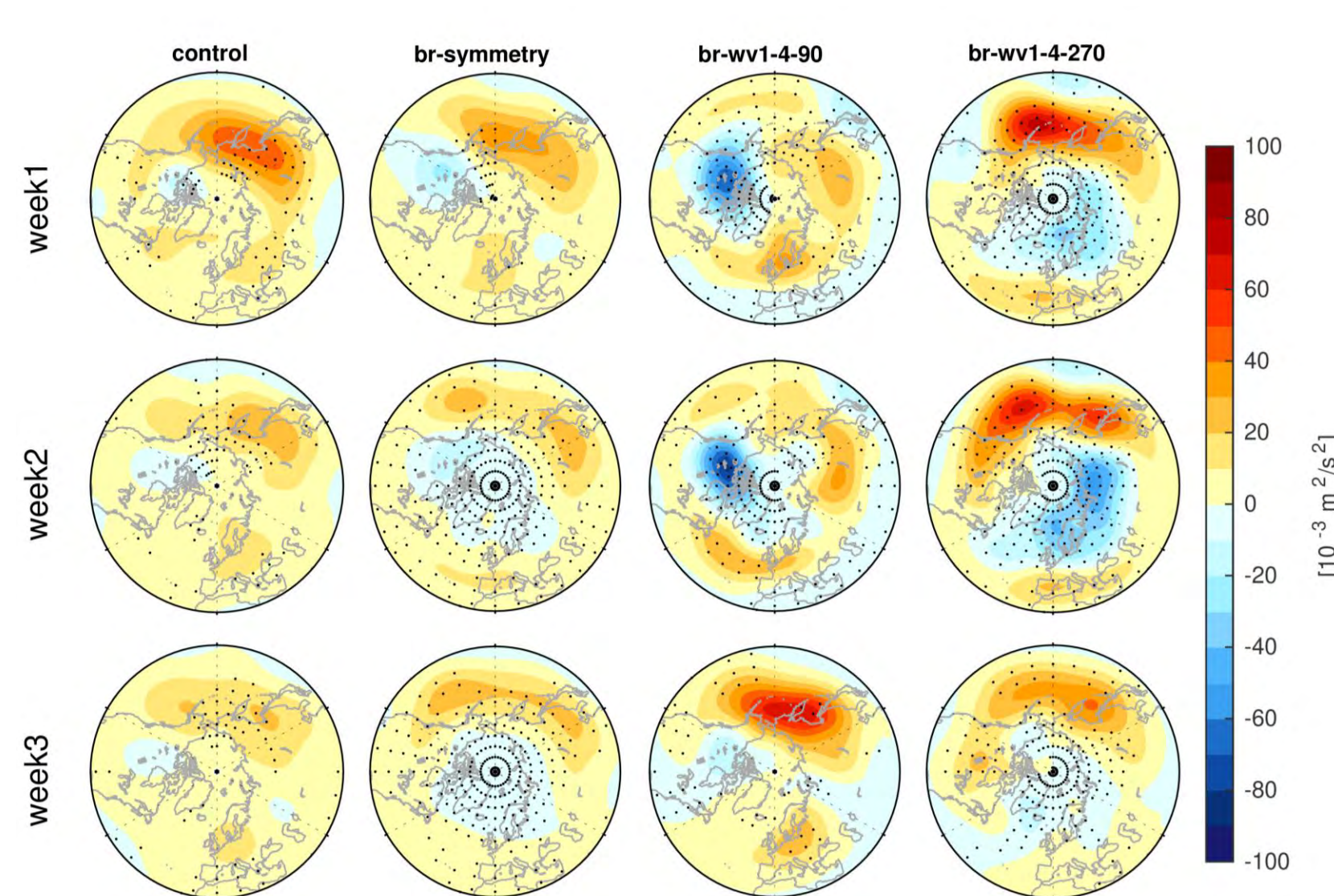
- The streamfunction zonal dipole in the stratosphere is balanced by an opposite signed dipole in the free-troposphere, though near the surface other factors play a role.

## 4. Wave forcing & Zonal wind



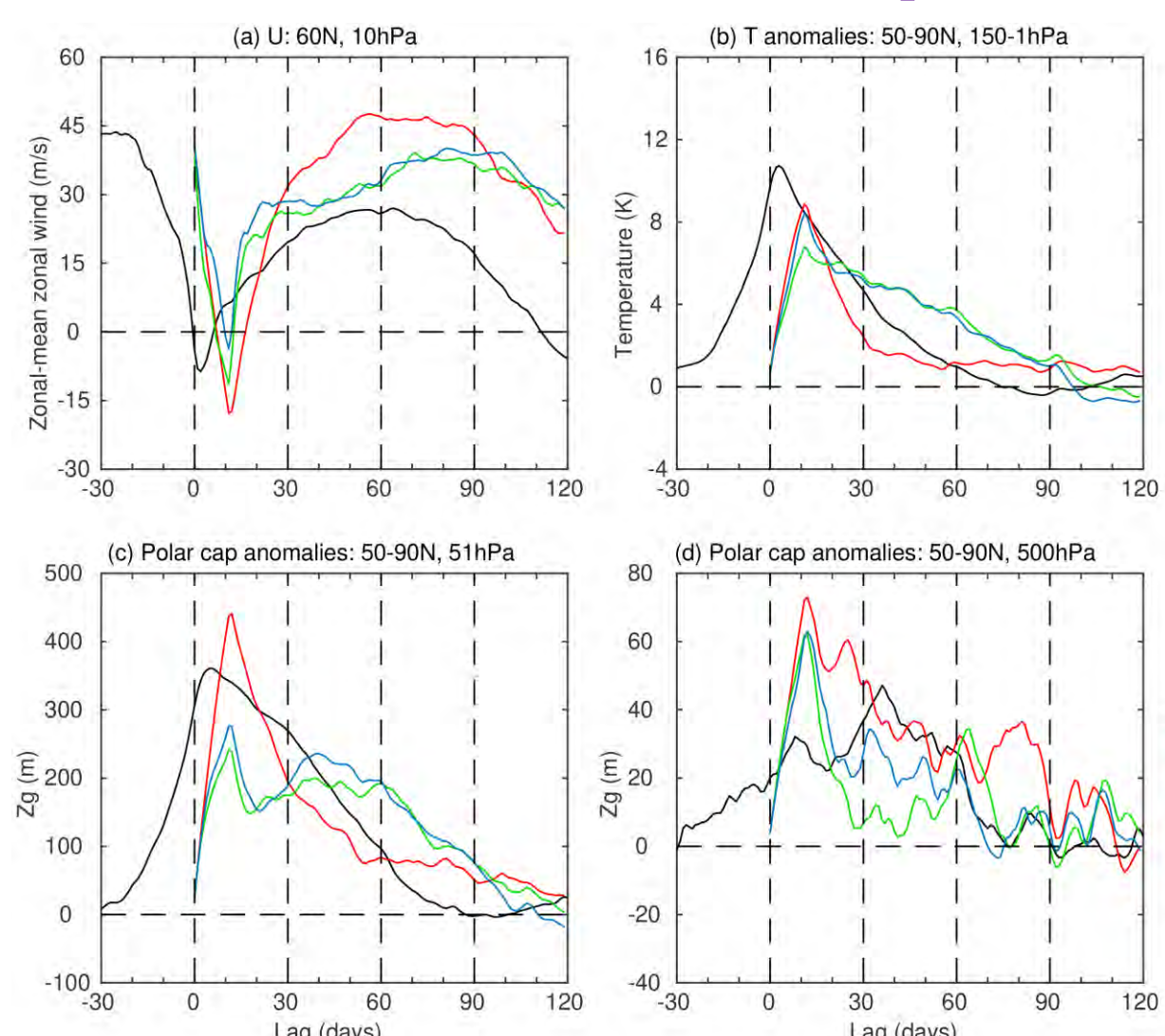
- The forcing allows us to alternately decelerates westerlies to easterlies in the zonal mean, or just in the western or eastern hemispheres. Surface impact is qualitatively different.

## 9. Fz, 93 hPa (raw field NOT anomalies)



- Negative Fz shows the downward propagation of quasi-stationary waves over north America for 90 and northern Eurasia for 270 case.

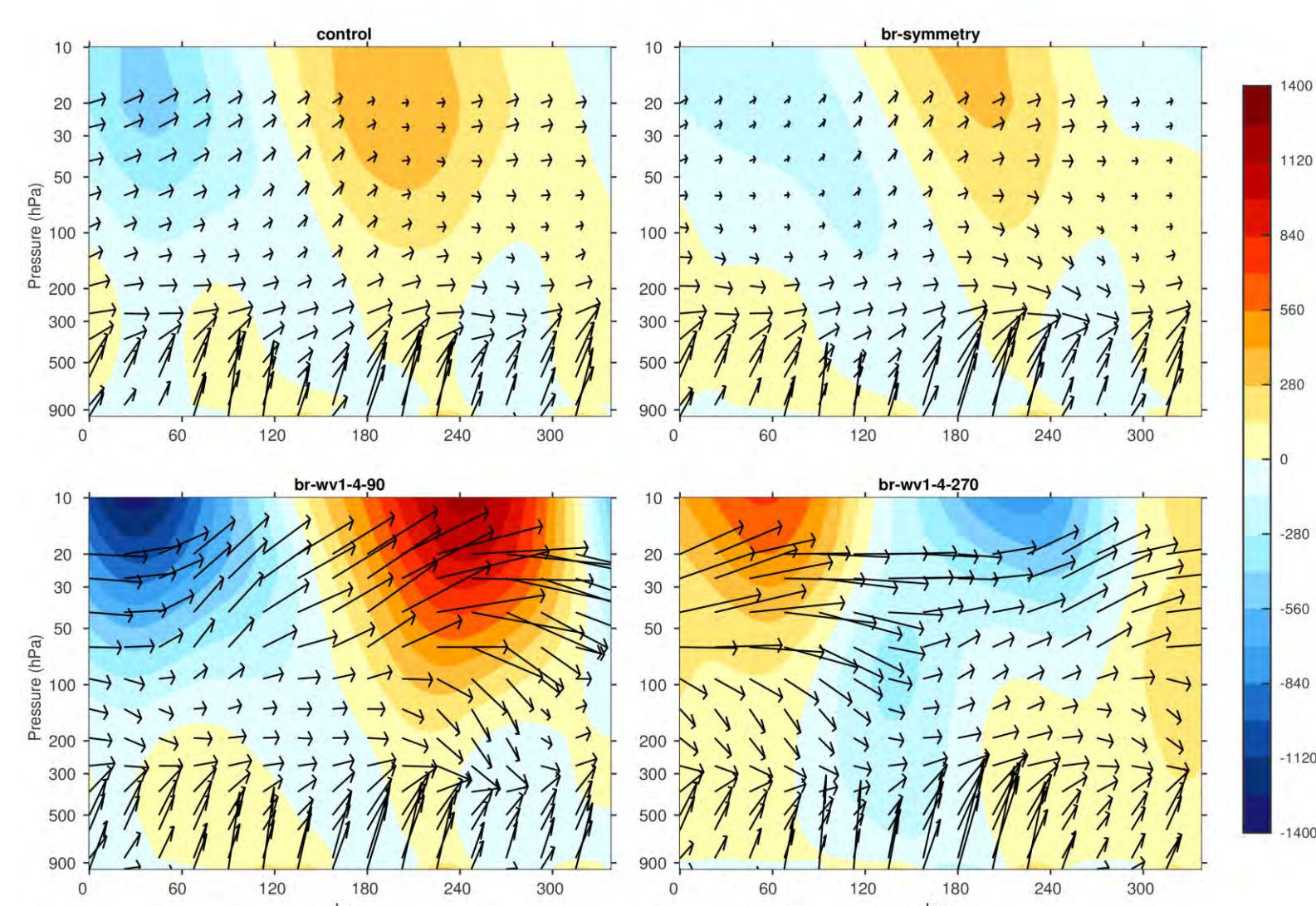
## 5. U & T & Polar cap



- In all cases the zonal mean structure of the SSW events are similar: U1060 easterlies, warmer polar cap, and higher geopotential height.

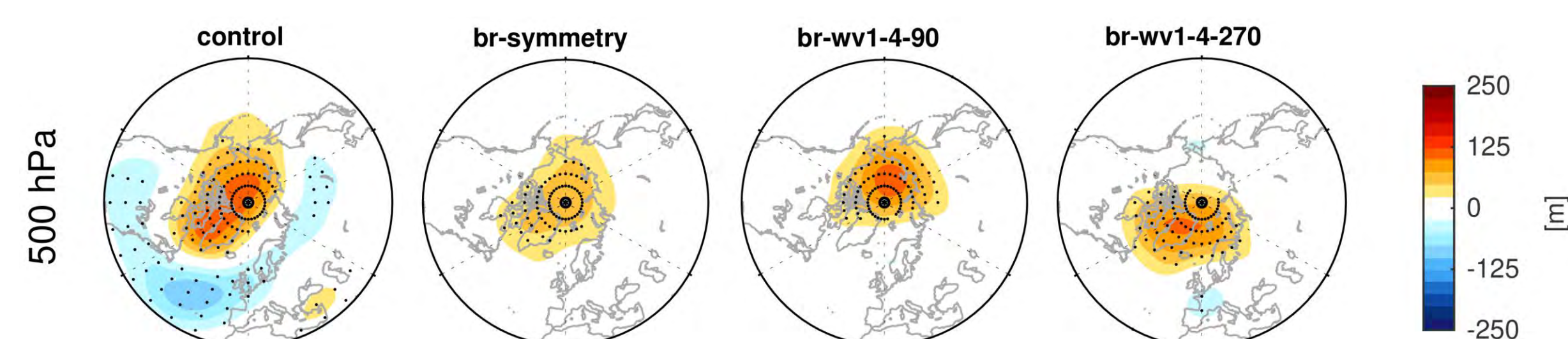
## 10. Fx-Fz

Plumb-flux-Fx-Fz-week2



- The eddy of geopotential height tilts eastward with height is helpful for upward propagation while downward tilting indicates downward propagation.
- For 270 case, the waves are reflected over north Europe while other runs, especially 90, reflected over North America.

## 6. Zg, 500 hPa, week1



- The control and symmetric have similar response while 90 and 270 cases have the opposite phase response in the troposphere.

## 11. Key Conclusions

- The asymmetries of wave forcing result in an asymmetric response in stratosphere and lower troposphere.
- The opposite phase of wave forcing have different impacts on the response of surface temperature and precipitation, which can be explained by the stream function of divergent part of meridional wind.
- The waves are reflected over North America and North Europe, which is strongly affected by the phase of applied wave forcing.