

Global and regional monsoons in a changing climate

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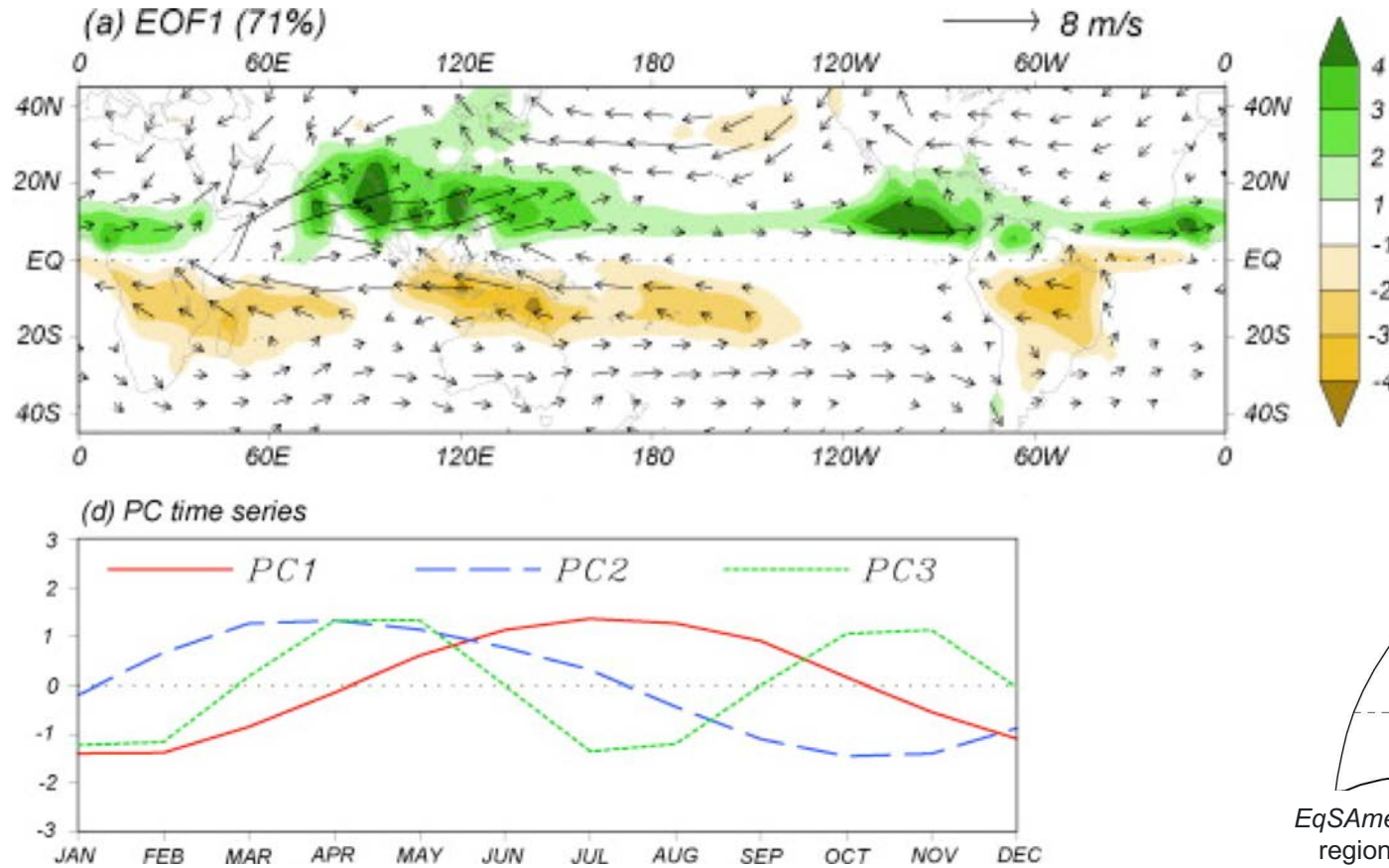
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Outline:

- Background/motivation: IPCC AR6 monsoon projections
 - Data & methods
- Hydrological sensitivity of global and regional monsoons
 - Summary & perspectives

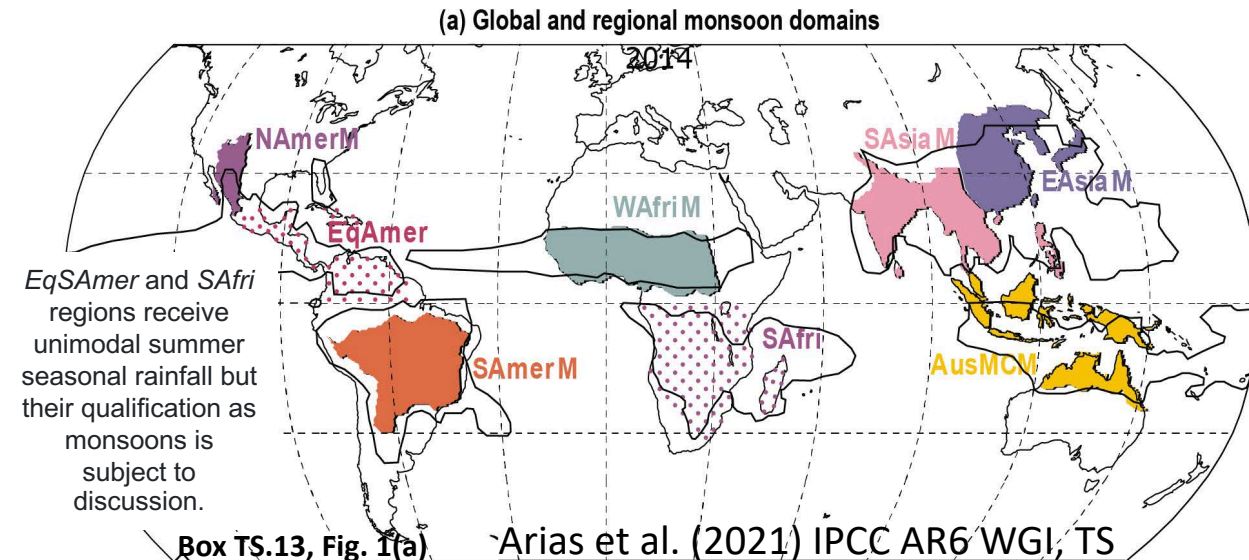
GLOBAL AND REGIONAL MONSOONS - background

Dominant mode of annual variation of precipitation (and wind) – hemispheric antisymmetric solstice mode (Wang and Ding, 2008)



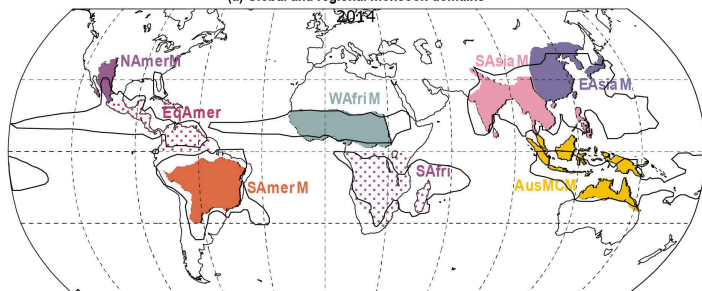
**IPCC AR6 - definition/convention:
Global (black contour) and regional
monsoons (colour shaded) domains**

The global monsoon (GM) is defined as the area with local summer-minus-winter precipitation rate exceeding 2.5 mm day^{-1} (see IPCC AR6 Annex V).



*Empirical Orthogonal Function (EOF) & normalized Principal Components (PC) of climatological monthly mean precipitation

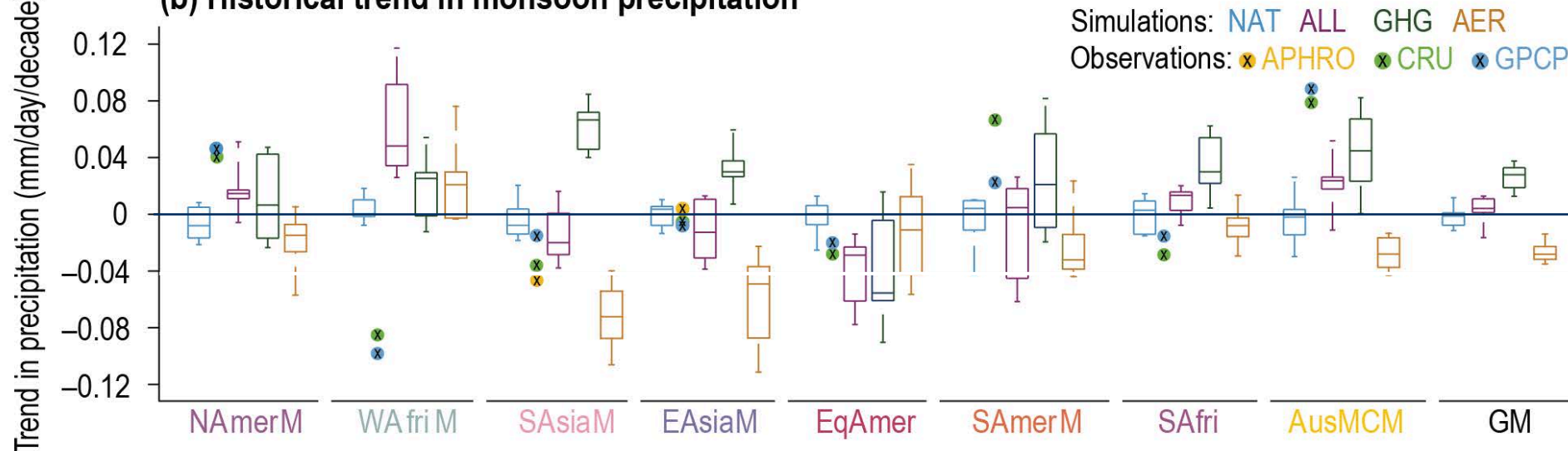
(a) Global and regional monsoon domains



Attribution of monsoon precipitation changes

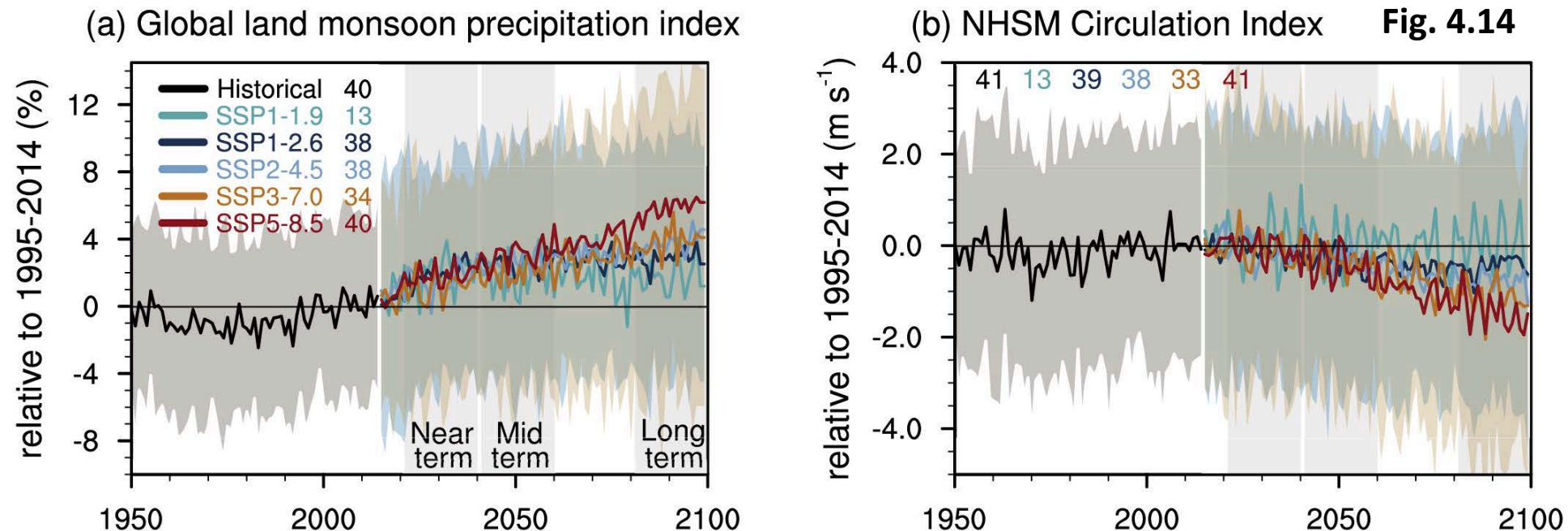
Box TS.13, Fig. 1(a,b)

(b) Historical trend in monsoon precipitation



Global and regional monsoons precipitation trends (2nd half of 20th century) based on DAMIP CMIP6 simulations with both natural and anthropogenic (ALL), greenhouse gas only (GHG), aerosols only (AER) and natural only (NAT) radiative forcing. Weighted ensemble means are based on nine Coupled model Intercomparison Project Phase 6 (CMIP6) models contributing to the MIP (with at least three members). Observed trends computed from CRU, GPCP and APHRO (only for *SAsiaM* and *EAsiaM*) datasets are shown as well.

Future changes of GM (precipitation & circulation indices)



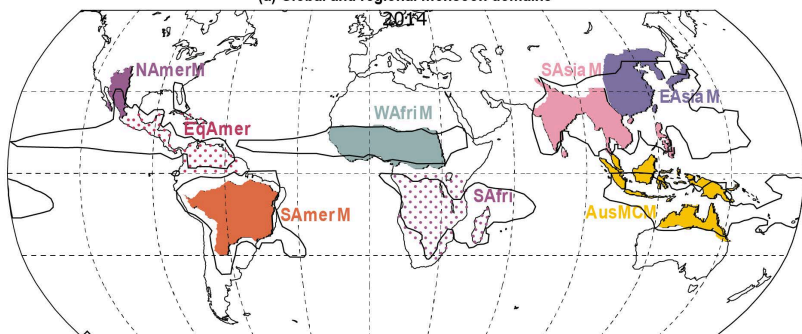
Time series of global land monsoon precipitation and Northern Hemisphere summer monsoon (NHSM) circulation index anomalies.

(a) Global land monsoon precipitation index anomalies (unit: %) defined as the area-weighted mean precipitation rate in the global land monsoon domain, as defined by [Wang et al. \(2013a\)](#) (b) Anomalies in NHSM circulation index (unit: m s^{-1}), defined as the vertical shear of zonal winds between 850 and 200 hPa averaged in a zone stretching from Mexico eastward to the Philippines (0° – 20°N , 120°W – 120°E ; [Wang et al., 2013a](#)). The curves show averages over the simulations, the shadings around the SSP1-2.6 and SSP3-7.0 curves show 5–95% ranges, and the numbers near the top show the number of model simulations used.

Near-term: changes in GM and related circulation will be affected by the combined effects of **model uncertainty** and **internal variability** (together larger than the forced signal) – *medium confidence*;

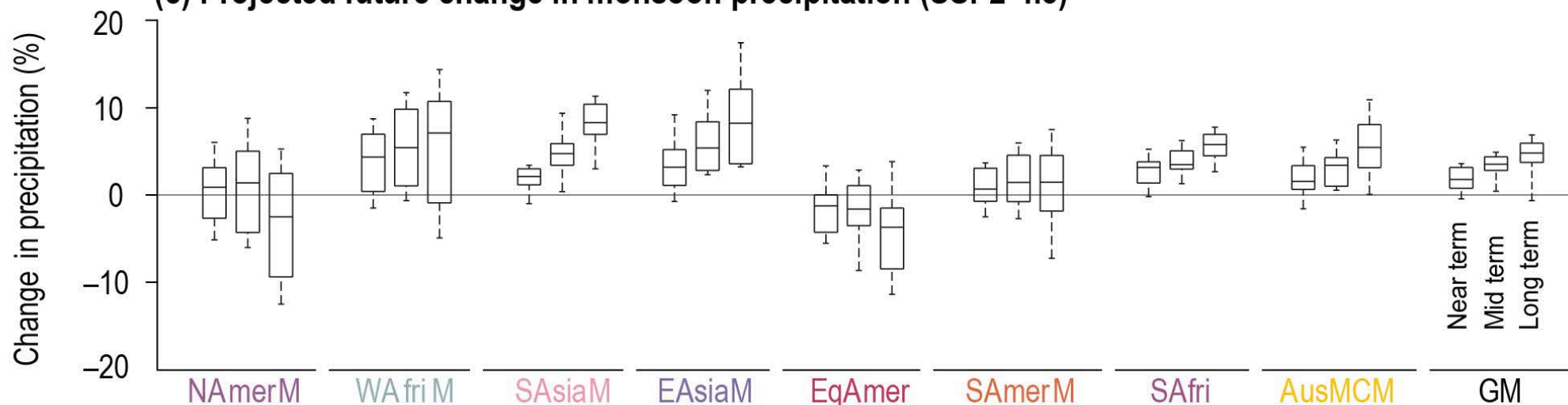
Long-term: it is likely that **GM land precipitation** will increase with GSAT rise (**1.3-2.4 % increase per $^{\circ}\text{C}$ of GSAT warming**), despite weakened monsoon circulation

(a) Global and regional monsoon domains



Future changes of monsoons precipitation

(c) Projected future change in monsoon precipitation (SSP2-4.5)



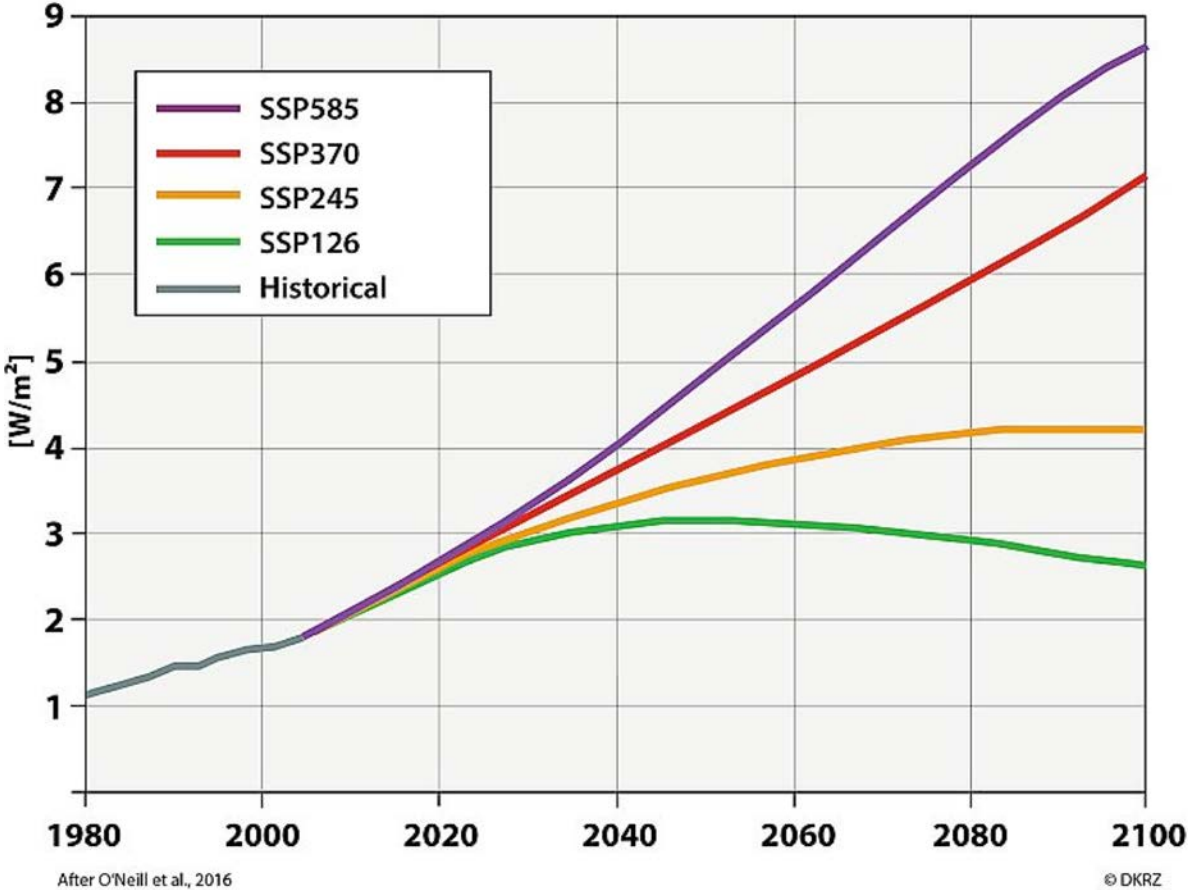
Box TS.13, Fig. 1(a,c)

Percentage change in projected seasonal mean precipitation over global and regional monsoons domain in the near term (2021–2040), mid-term (2041–2060), and long term (2081–2100) under SSP2-4.5 based on 24 CMIP6 models.

Data (CMIP6 models & experiments):

Model's name	Institute/Country
ACCESS-CM2; ACCESS-ESM1-5	CSIRO/Australia
AWI-CM-1-1-MR	AWI/Germany
BCC-CSM2-MR	BCC/China
CAMS-CSM1-0	CAM/China
CanESM5-1; CanESM5	CCCma/Canada
CAS-ESM2-0	CAS/China
CESM2-WACCM ; CESM2	NCAR/US
CMCC-CM2-SR5; CMCC-ESM2	CMCC/Italy
CNRM-ESM2-1	CNRM&CERFACS/France
CanESM5-1 ; CanESM5	CCCma/Canada
EC-Earth3 ; EC-Earth3-Veg; EC-Earth3-Veg-LR	EC-Earth Consortium/EU
FGOALS-f3-L; FGOALS-g3	LASG-IAP/China
GFDL-ESM4	GFDL/US
IITM-ESM	IITM/India
INM-CM4-8 ; INM-CM5-0	INM/Russia
IPSL-CM6A-LR	IPSL/France
KACE-1-0-G	NIMS-KMA/South Korea
MIROC6	MIROC/Japan
MPI-ESM1-2-LR	MPI-M/Germany
MRI-ESM2-0	MRI/Japan
NorESM2-MM ; NorESM2-LM	NCC/Norway
TaiESM1	Acad-SINICA/Taiwan

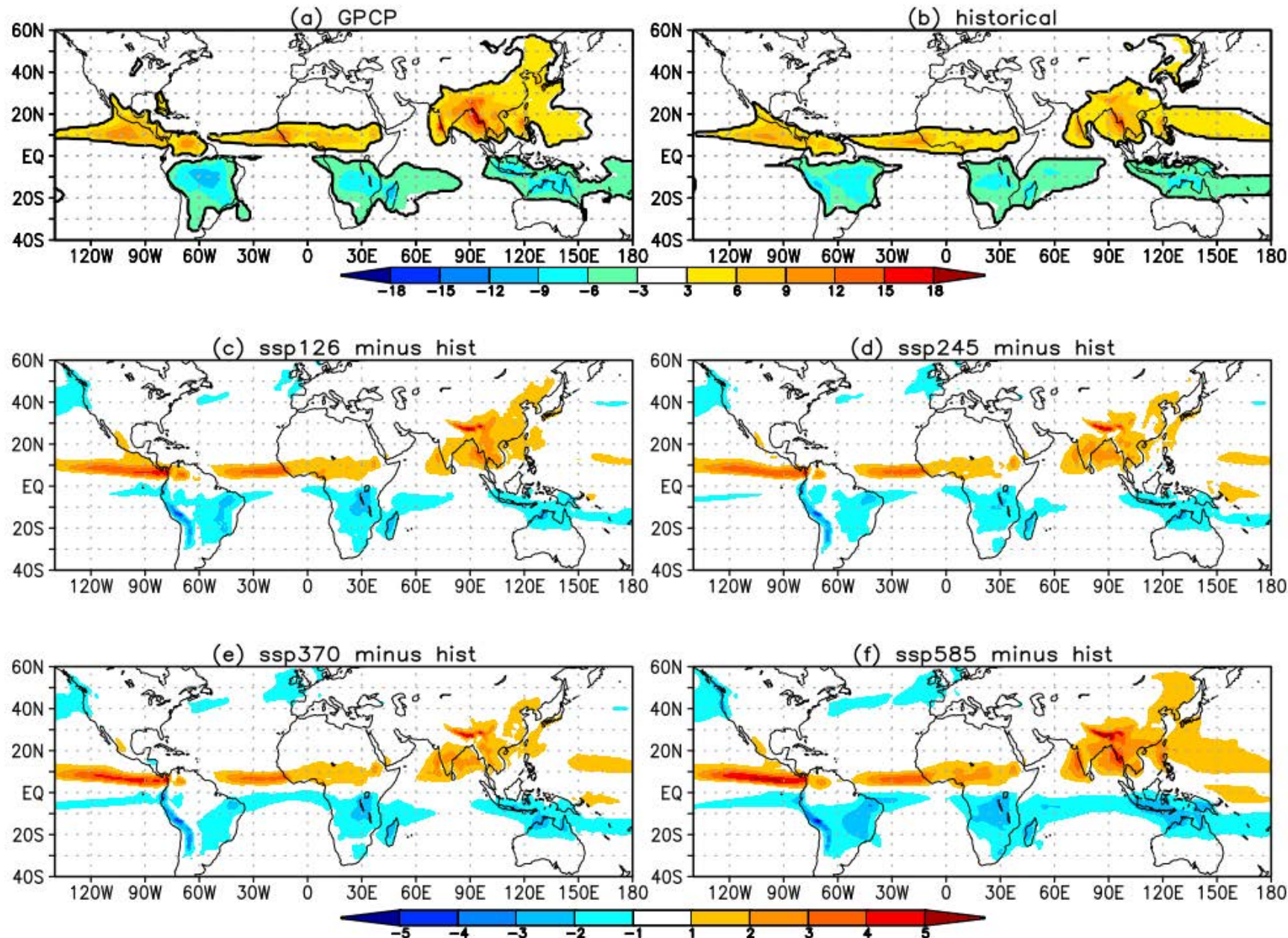
CMIP6 Scenarios - Anthropogenic Radiative Forcing [W/m²]



CMIP6 models considered (1 member for each model/simulation)
(in some of the figure only the models in bold are considered so far, to expand including all available)

JJA minus DJF precipitation (mm/d)

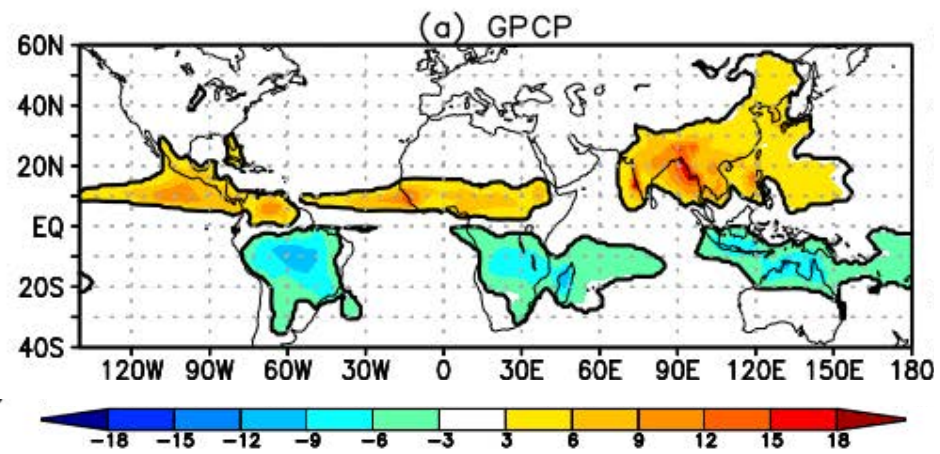
observation vs historical MMM (1985-2014 climatology) & change in projections with respect to hist climatology



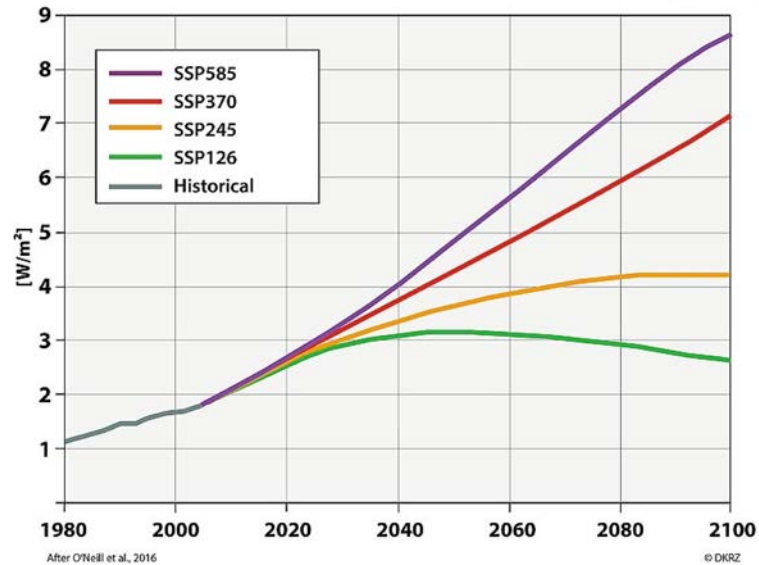
GM (black line) is identified as JJA minus DJF exceeding 2.5 mm/d

GM region is realistic in historical MMM but precipitation is underestimated in some places

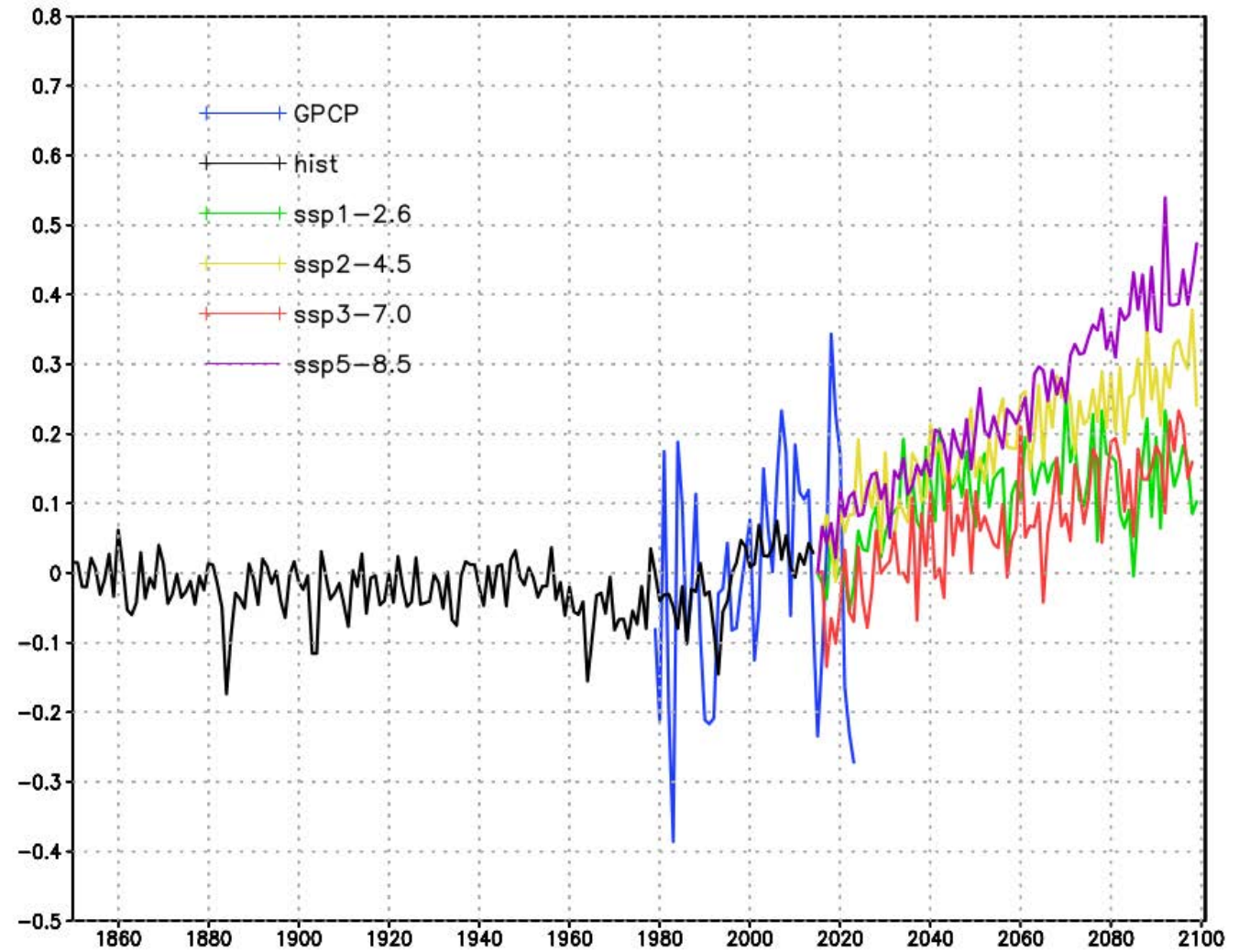
In future scenarios, precip is overall projected to increase but diversities among regional monsoons and mostly non-linearities between different scenarios seem to appear



CMIP6 Scenarios - Anthropogenic Radiative Forcing [W/m^2]



Global monsoon precipitation anomalies (mm/d)



Past and future evolution of GM precipitation anomalies (pre-ind) – clear non-linearities in future change comparing scenarios and associated end-of-century forcing

Methodology: hydrological sensitivity of global and regional monsoons

As in Katzenberger et al (2021):

Change in summer monsoon mean rainfall (mm/d) depending on change in GMT (K) –

Stepwise difference between future periods to the reference period (1985-2014) - 20-years future periods with 5-years time steps (i.e. 2000-2019, 2005-2024, ... 2080-2099)

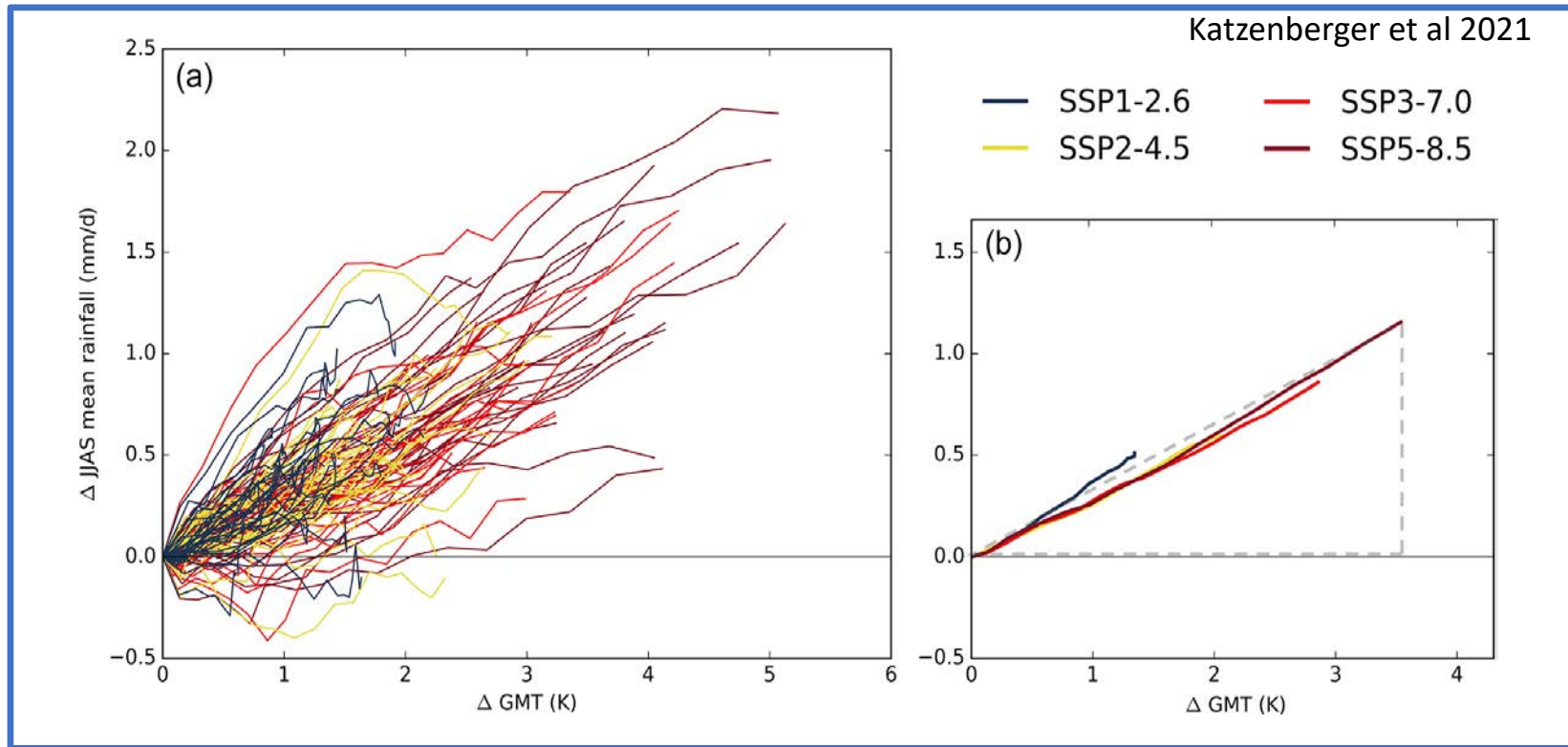
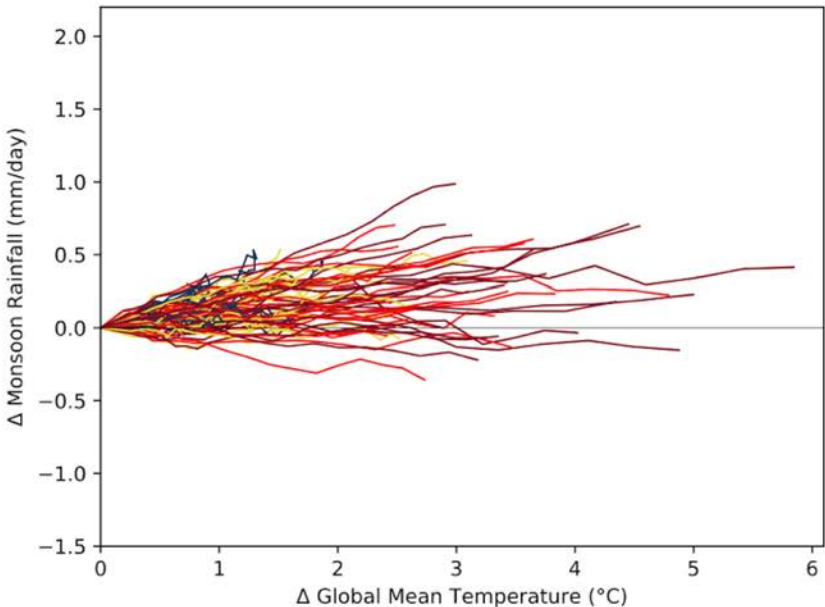


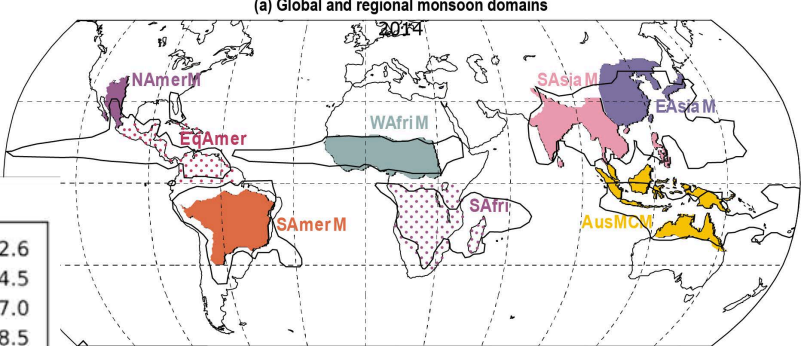
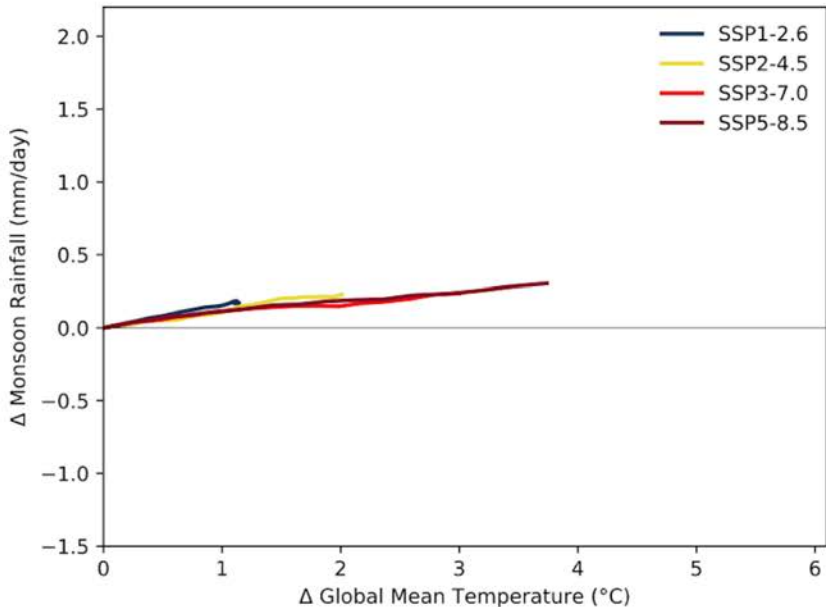
Figure here for Indian summer monsoon but we applied to GM and regional monsoons as defined in IPCC AR6

Hydrological sensitivity of global monsoon

(a) GM all members



(b) GM multi-models' mean



Linear regression fitting (LFT)
monsoon rainfall per GMT increase
(mm/d per K)

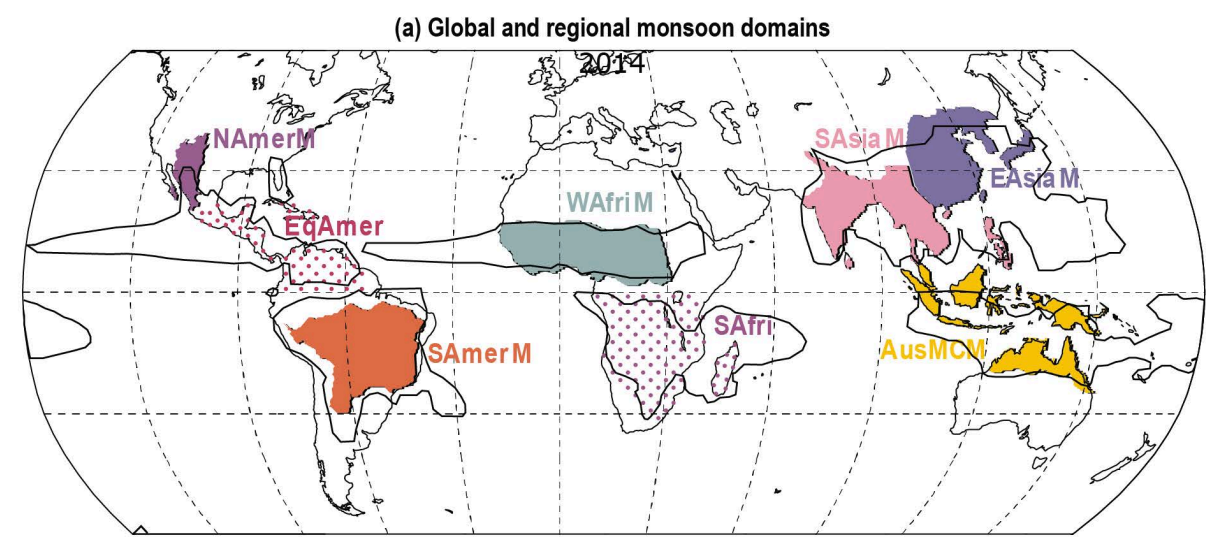
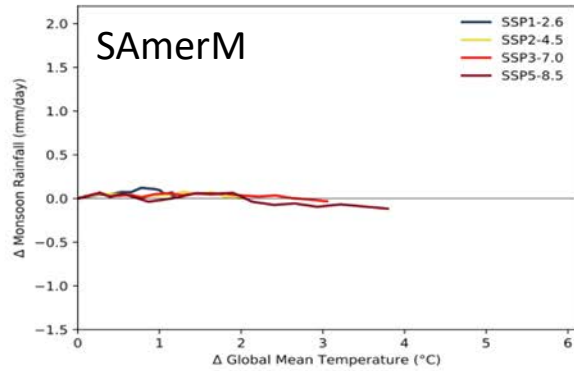
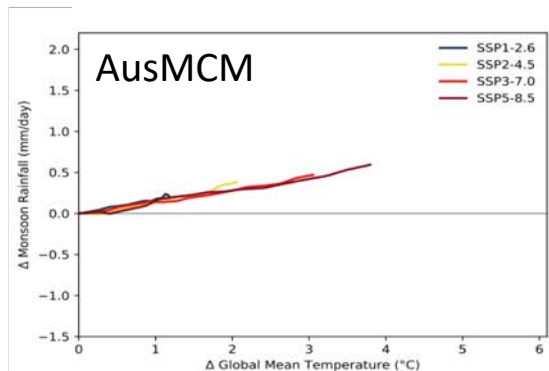
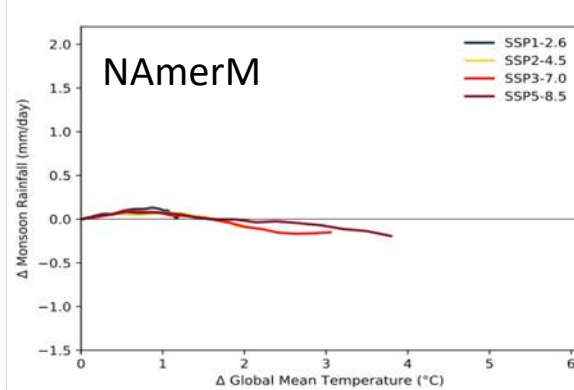
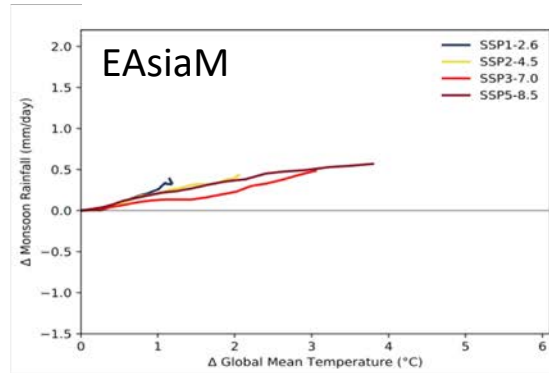
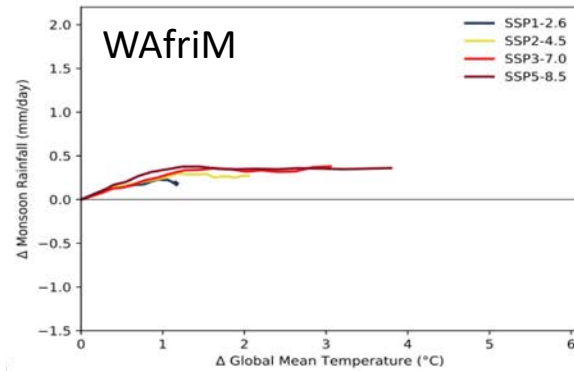
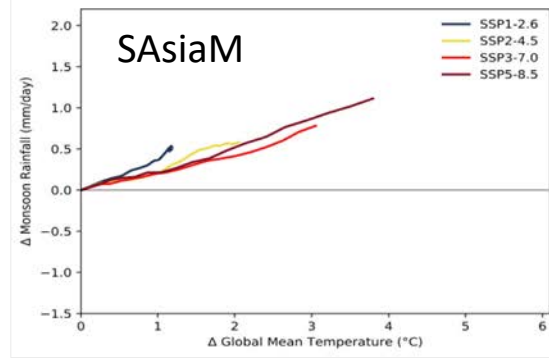
Value of the slopes (LFT) are the lowest for SSP3-7.0 but the global radiative forcing expected at the end of the century for that scenario is not – in GM, GM-NH, GM-EH

	SSP1-2.6	SSP2-4.5	SSP3-7.0	SSP5-8.5
GM	0.15	0.13	0.06	0.07
GM-NH	0.20	0.13	0.05	0.06
GM-SH	0.11	0.12	0.08	0.08
GM-EH	0.31	0.25	0.18	0.20
GM-WH*	0.04	0.04	-0.02	-0.03

*LFT is of limited use as the change of monsoon precipitation is non-linear with respect to the GMT increase

Hydrological sensitivity of regional monsoons

(summer precip increase as function of GWL)

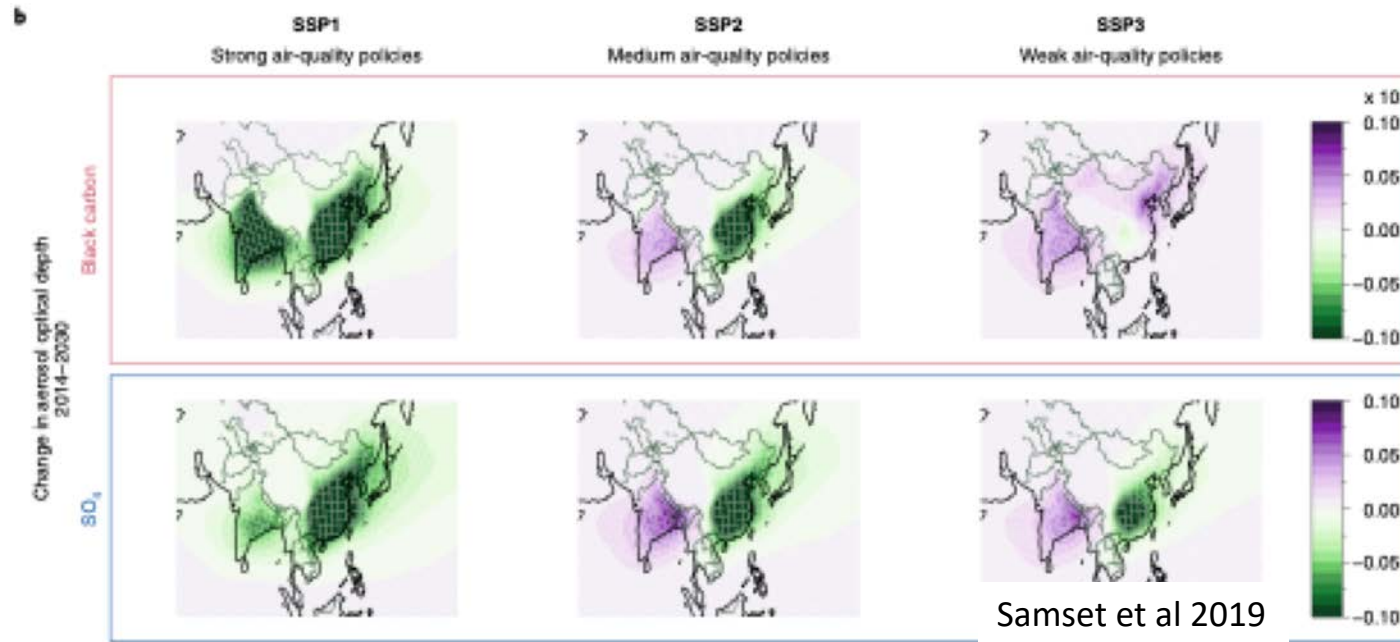


Linear regression fitting (LFT) –
monsoon rainfall per GMT increase (mm/d per K)

	SSP1-2.6	SSP2-4.5	SSP3-7.0	SSP5-8.5
SAsiaM	0.43	0.30	0.24	0.28
EAsiaM	0.33	0.21	0.13	0.15
AusMCM	0.17	0.19	0.13	0.13
NAmerM*	-0.08	-0.08	-0.11	-0.09
SAmerM*	0.04	0.01	-0.02	-0.03
WAfriM*	0.11	0.09	0.08	0.04

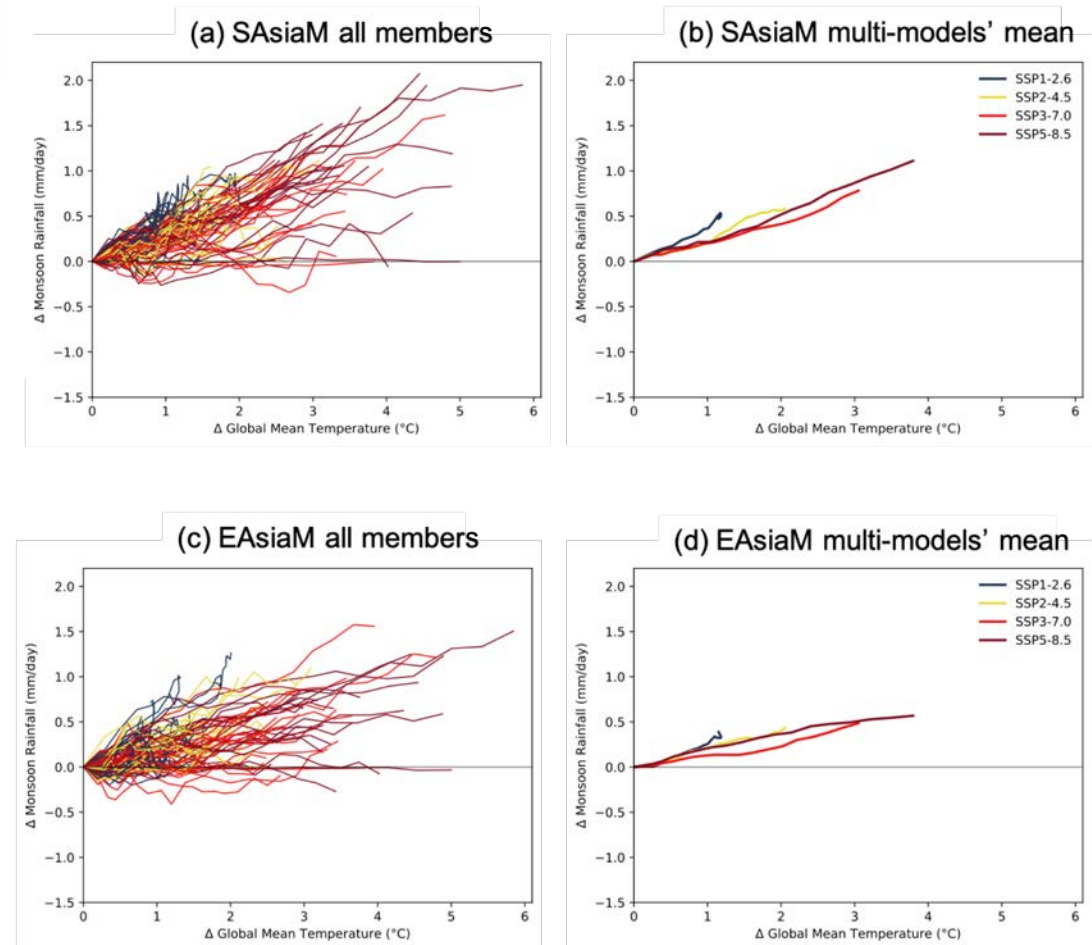
*LFT is of limited use as the change of monsoon precipitation is non-linear with respect to the GMT increase

Hydrological sensitivity: the case of Asian monsoons & role of aerosols



despite weaker air quality policies and increased black carbon and sulphates AOD in SSP3, precipitation increases less with GW compared to SSP2

Summer precipitation increase as function of GWL



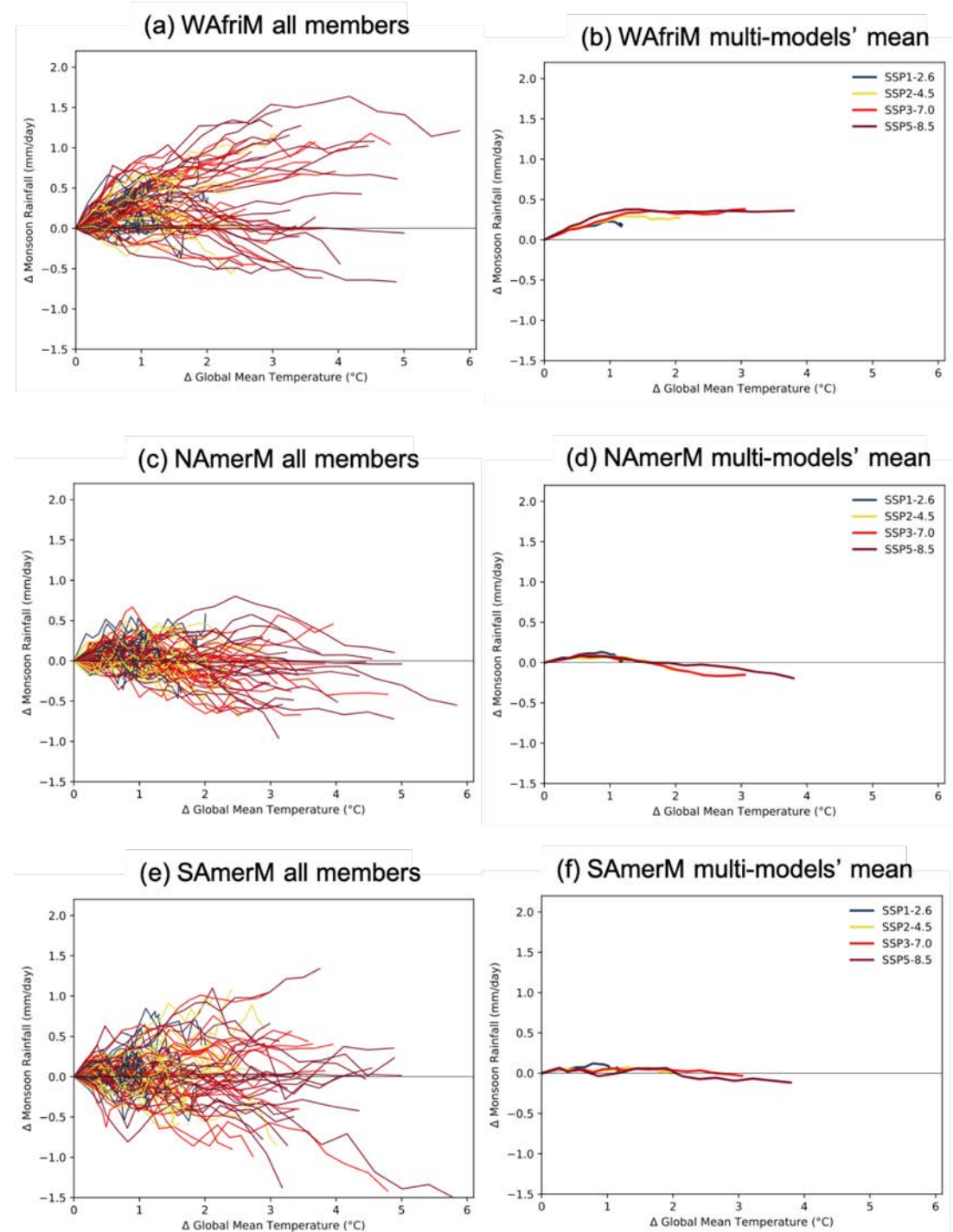
Complexity of aerosols-climate interactions, role of regional feedbacks & problems in representing all this in GCMs

Hydrological sensitivity of regional monsoons

(summer precip increase as function of GWL)

WafriM: anthropogenic emissions & change in the temperature of the sub-tropical North Atlantic Ocean relative to the global tropical oceans (Giannini and Kaplan, 2019)

NAmerM & SAmerM: summer precipitation decrease with GW increase (as MMM) but the models' agreement in the sign of the change of summer mean rainfall to GWL is the lowest



Summary & perspectives

- Monsoon precipitation increase with GW mostly for Global monsoon & Asian monsoons
- Non-linearity with radiative forcing because of diversities in response to type of anthropogenic aerosols (mostly EAsiaM)- complexity and role of local feedbacks
 - WafriM precipitation increase with GWL but with a threshold/plateau
 - Models' disagreement in American monsoons

What's next:

- disentangle role of anthropogenic aerosols & SST gradient for WafriM
- increase of the ensemble dimension to have more robust conclusions and mostly to see if disagreement in American monsoons is reduced
 - role of climate sensitivity?