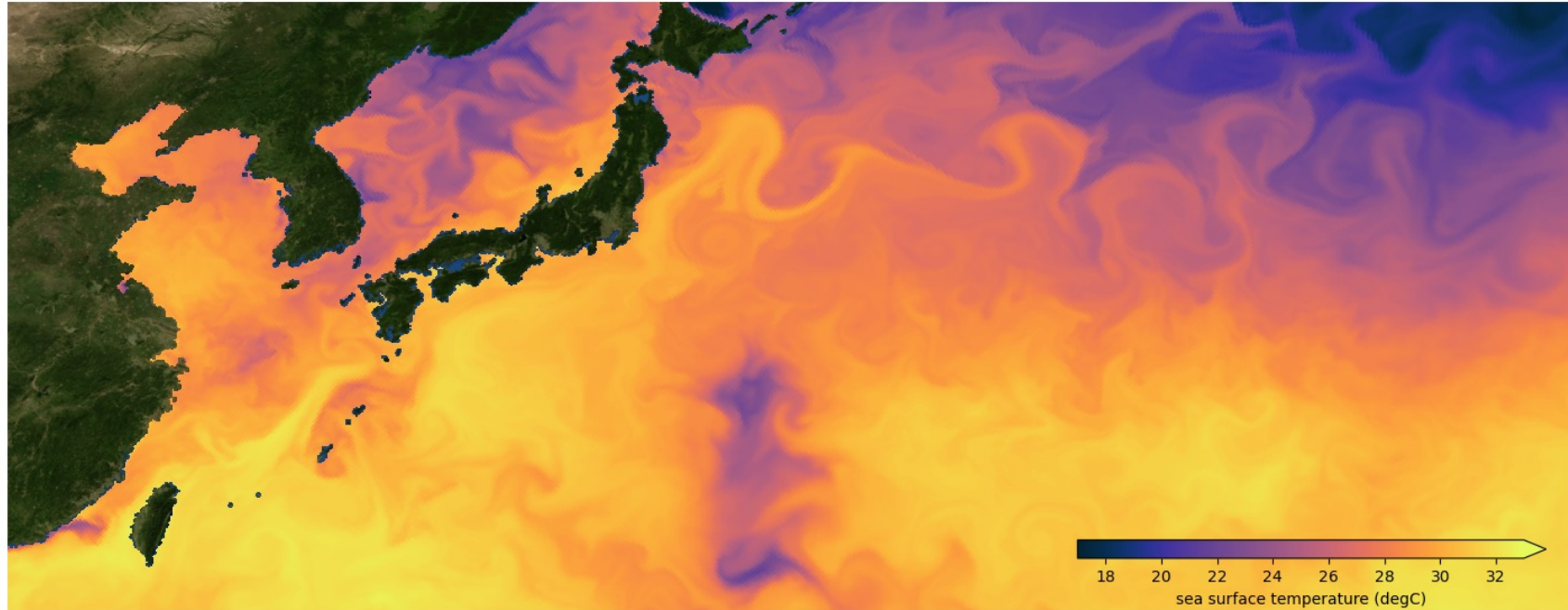


1. What is EERIE?

EERIE (European Eddy-Rich Earth-System Models) is a Horizon Europe project aiming to transform climate modeling by **explicitly resolving ocean eddies**—small-scale currents that play a crucial role in heat, carbon, and freshwater transport. Unlike conventional climate models that approximate these features, EERIE uses **high-resolution simulations** (~10 km) in both the ocean and the atmosphere that enable **more realistic processes and interactions**.



Courtesy of M. Aengenheyster

2. ENSO teleconnections

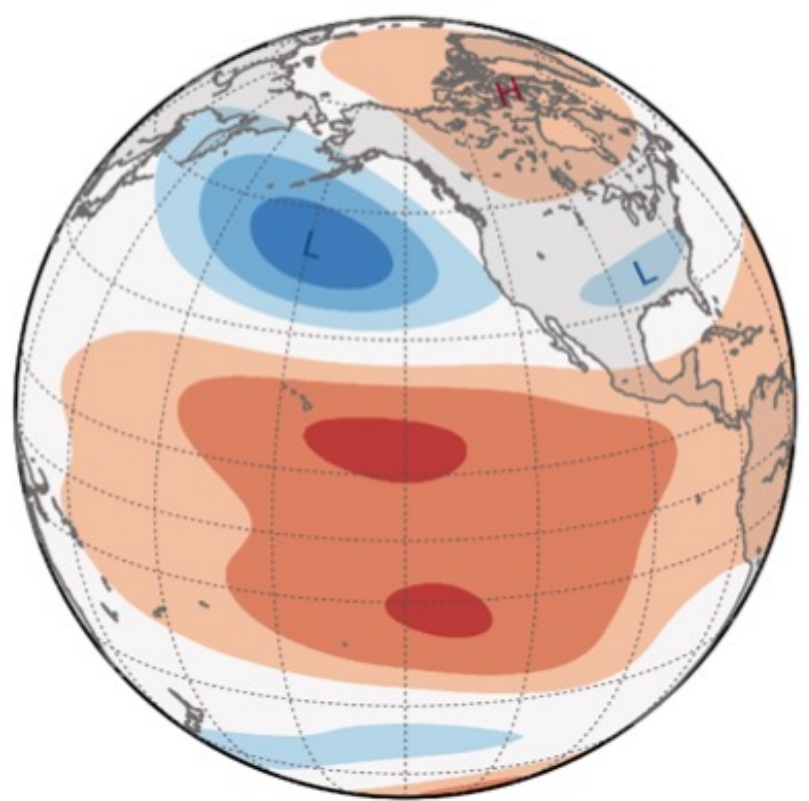
El Niño-Southern Oscillation (ENSO) is a prominent mode of climate variability that can impact remote regions through **atmospheric teleconnections**. The main mechanism for these teleconnections entails a large-scale Rossby Wave train triggered by upper-level divergence anomalies in the tropical Pacific, which then propagates towards the mid and high latitudes.

Our key questions:

How are ENSO teleconnections **represented** in the eddy-rich EERIE models?

Is there an **improvement** compared to existing state-of-the-art models?

... And **how** to assess this?



From B. Mezzina's PhD Thesis

3. Data & Methods

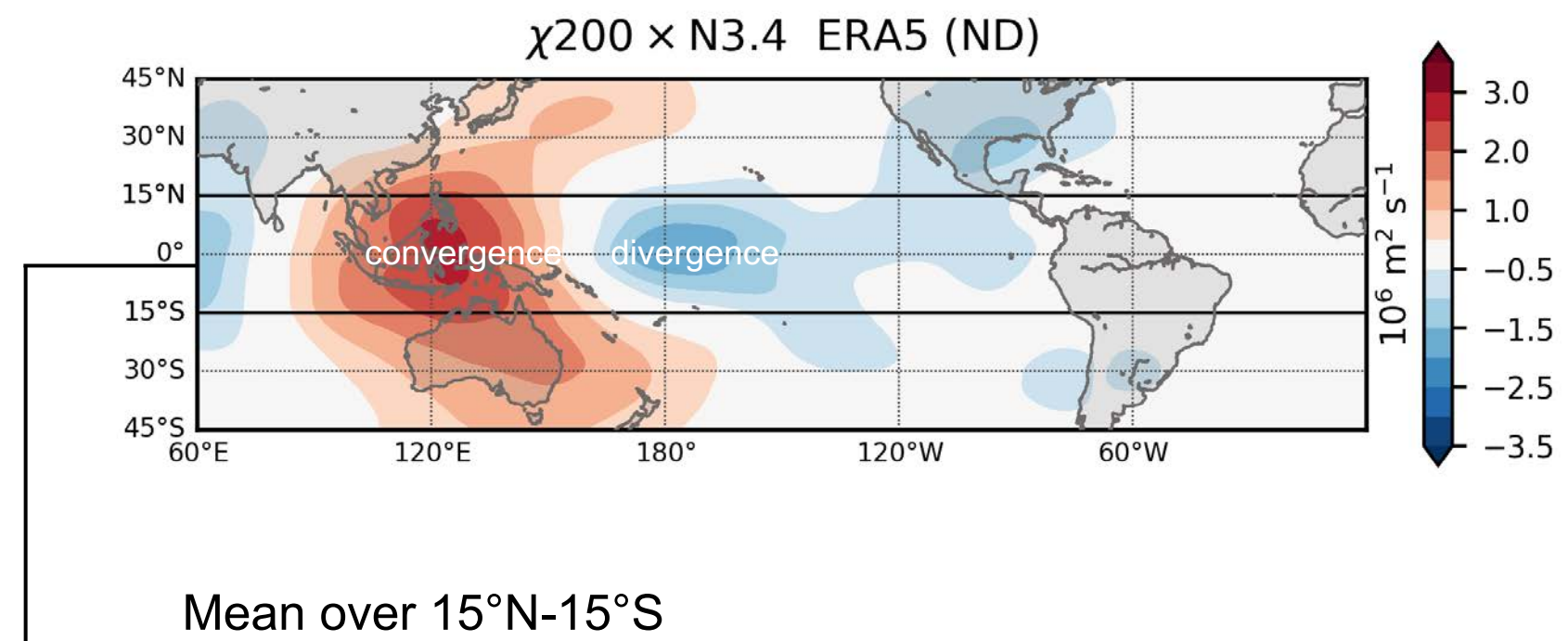
Simulations performed within EERIE:

Type	Institution	Model	Resolution (atm., ocean)	Members	Period
Coupled	AWI	IFS-FESOM	Tco1279 (~9 km), 4.5-13 km	1	1950-2014
	BSC	IFS-NEMO	Tco1279 (~9 km), eORCA12 (~8 km)	1	1950-1996
AMIP	ECMWF	IFS	Tco1279 (~9 km)	1	1980-2023
	ECMWF	IFS	Tco399 (~28 km)	10	1980-2023

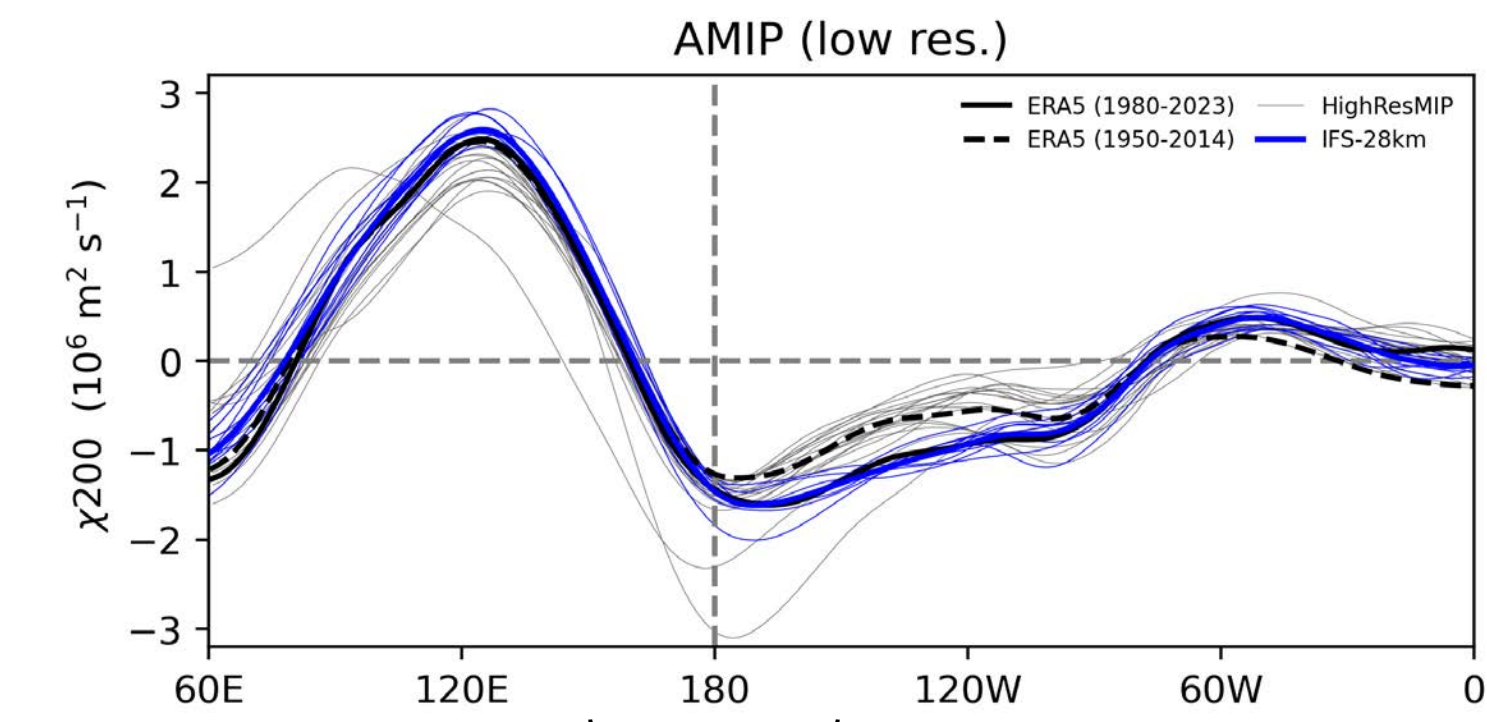
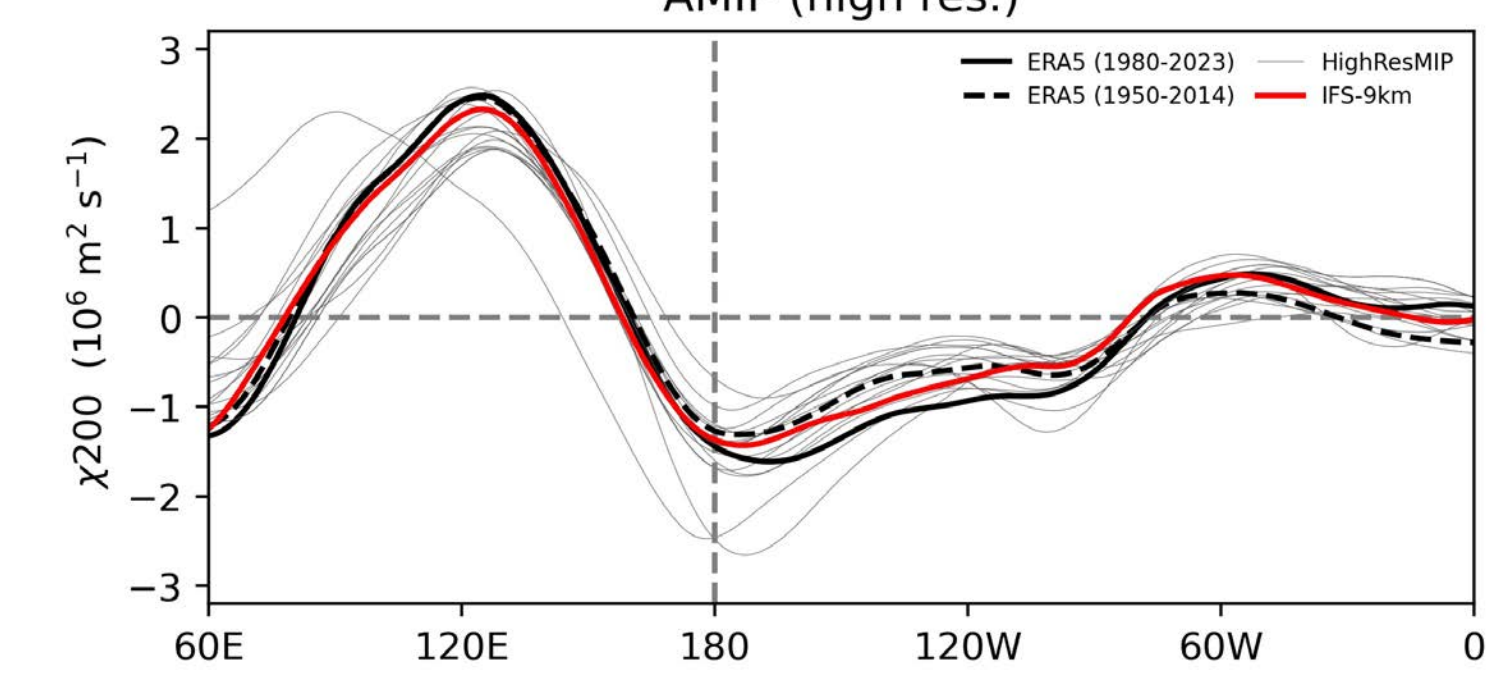
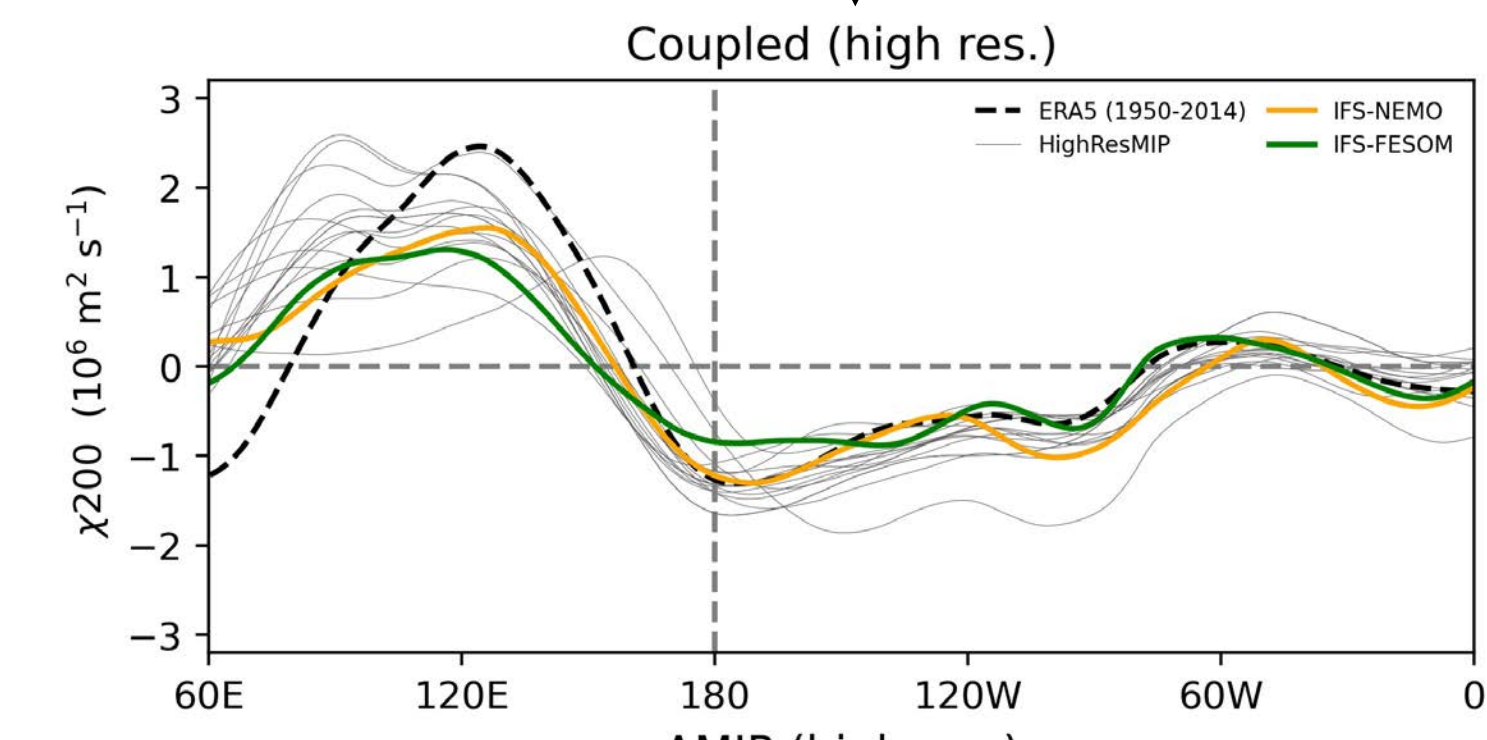
- *Comparison with HighResMIP*: simulations from 6 models, each with a “low resolution” (50-250 km) and “high resolution” (25-100 km) version, over 1950-2014
- *Comparison with reanalysis*: ERA5 (1980-2023 for the EERIE-AMIP runs, 1950-2014 for the rest)
- **Linear regressions** on DJF Niño3.4-index

4. Tropical Response

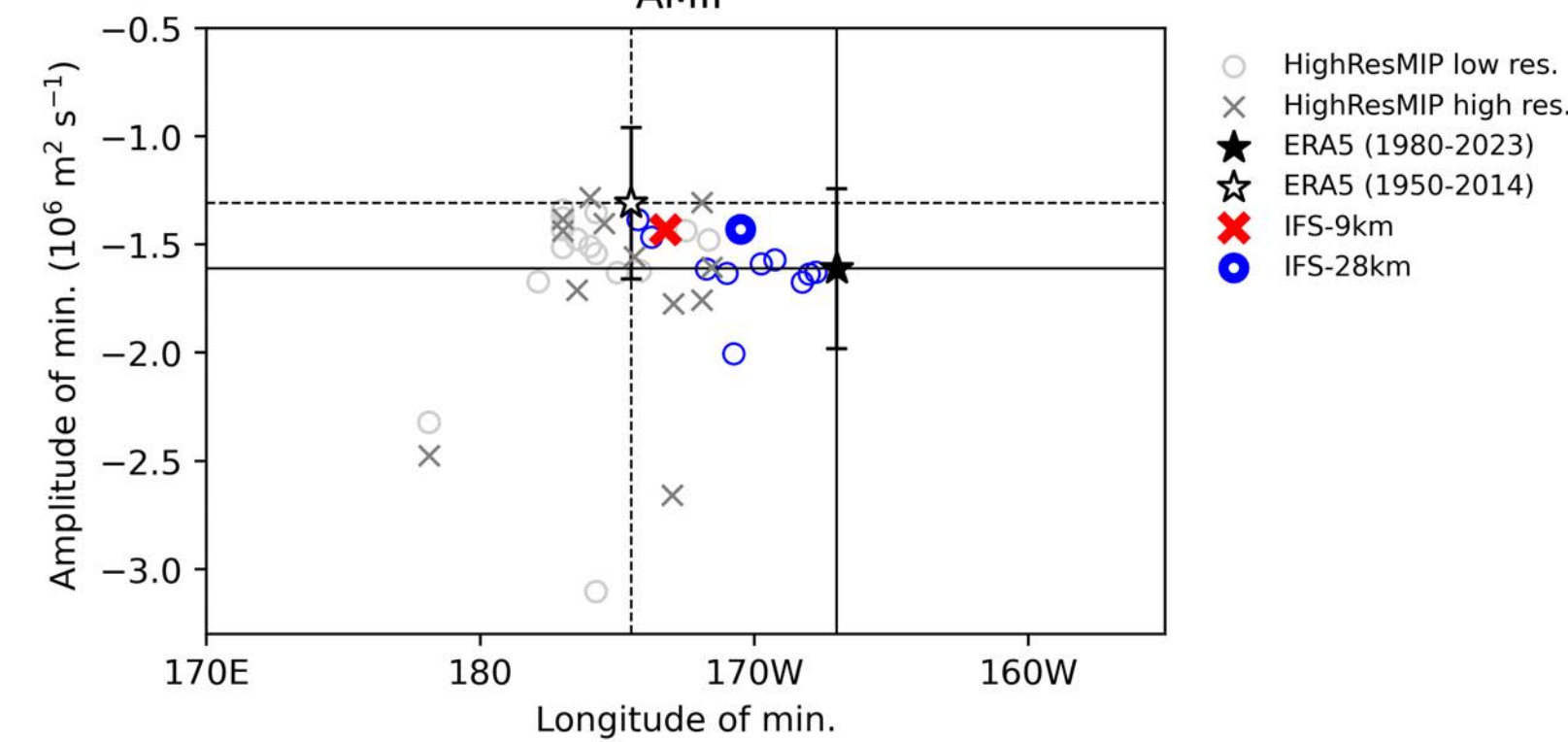
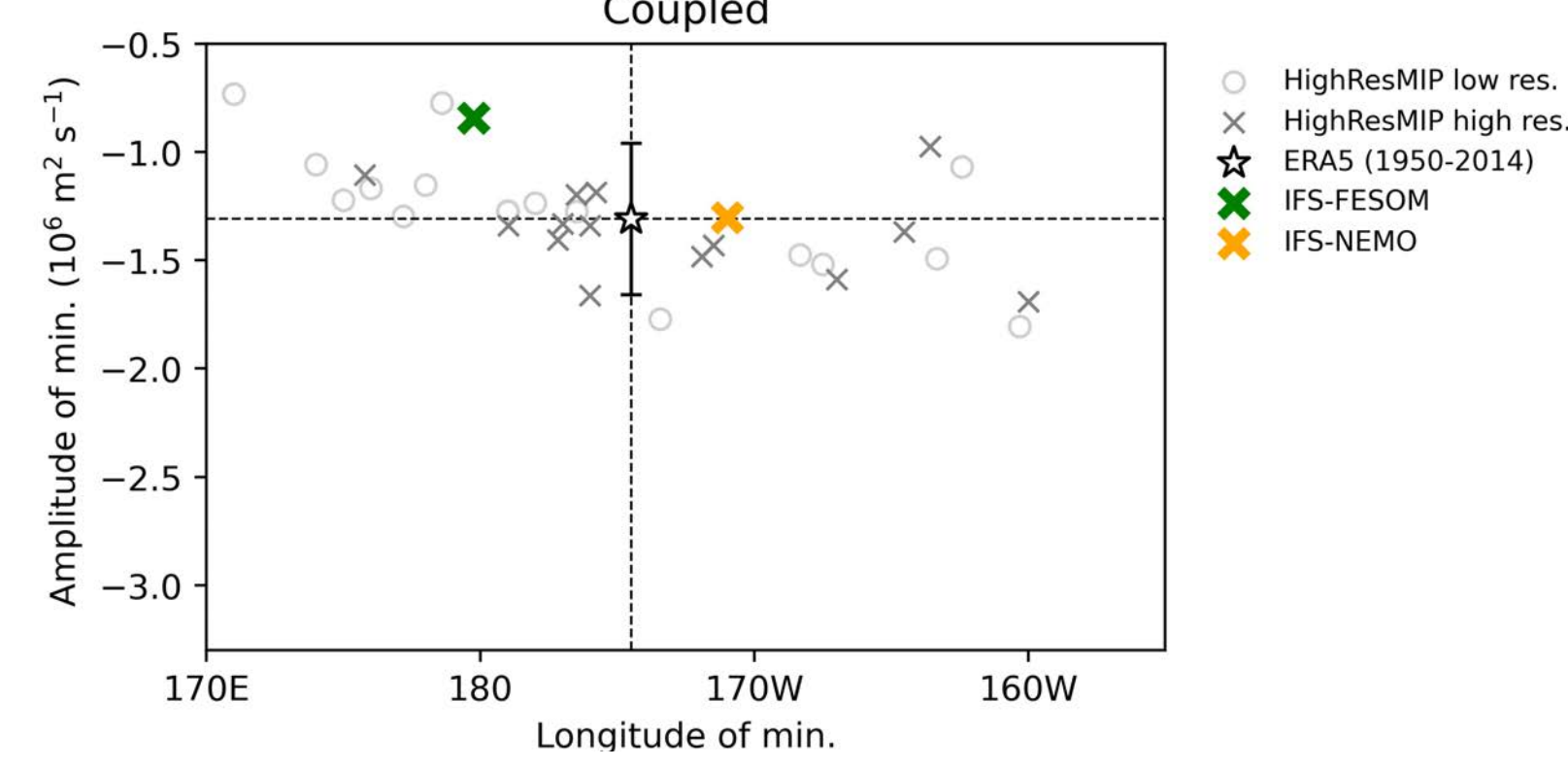
We can use the 200-hPa velocity potential to diagnose the upper-tropospheric flow anomalies related to ENSO:



Mean over 15°N-15°S



Find location & amplitude of minimum

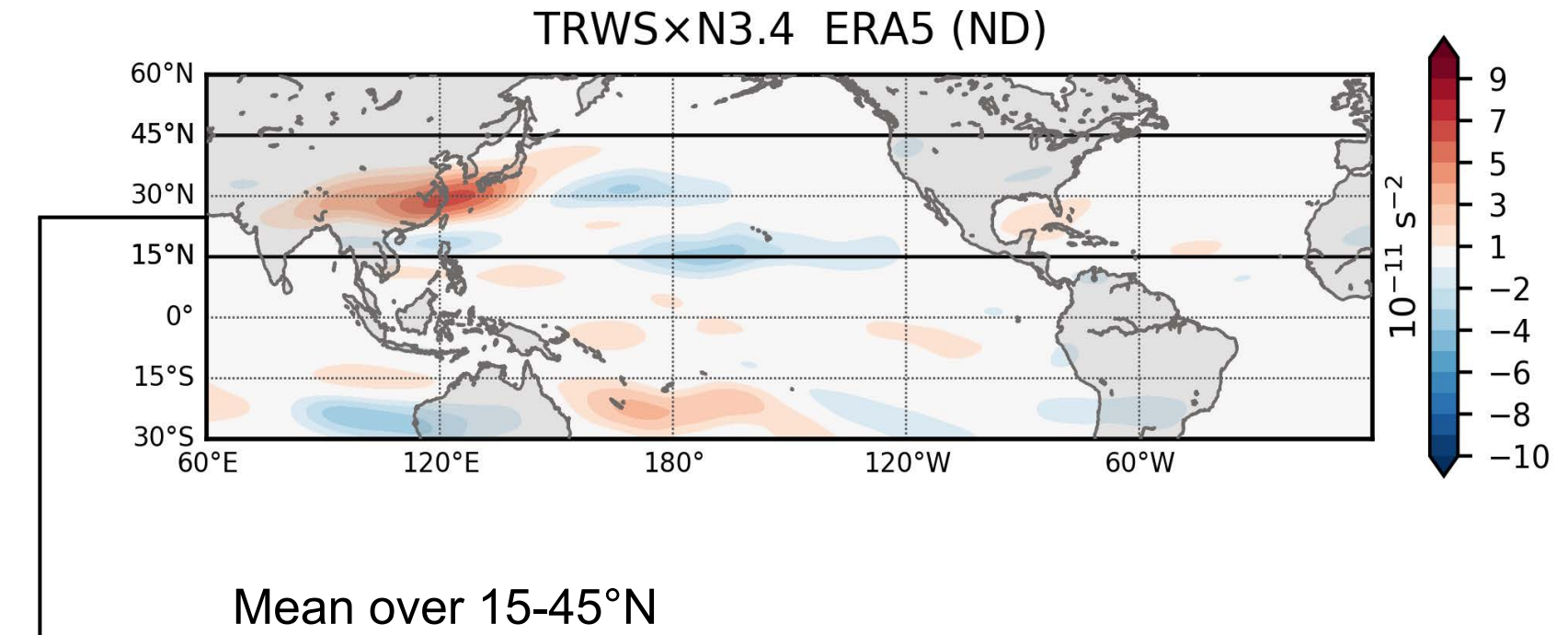


- **IFS-NEMO** accurately captures the minimum; **IFS-FESOM** underestimates the amplitude & is more shifted.
- Minor differences between **IFS-9km** and the ensemble mean from **IFS-28km**. Some **IFS-28km** members closer to ERA5.

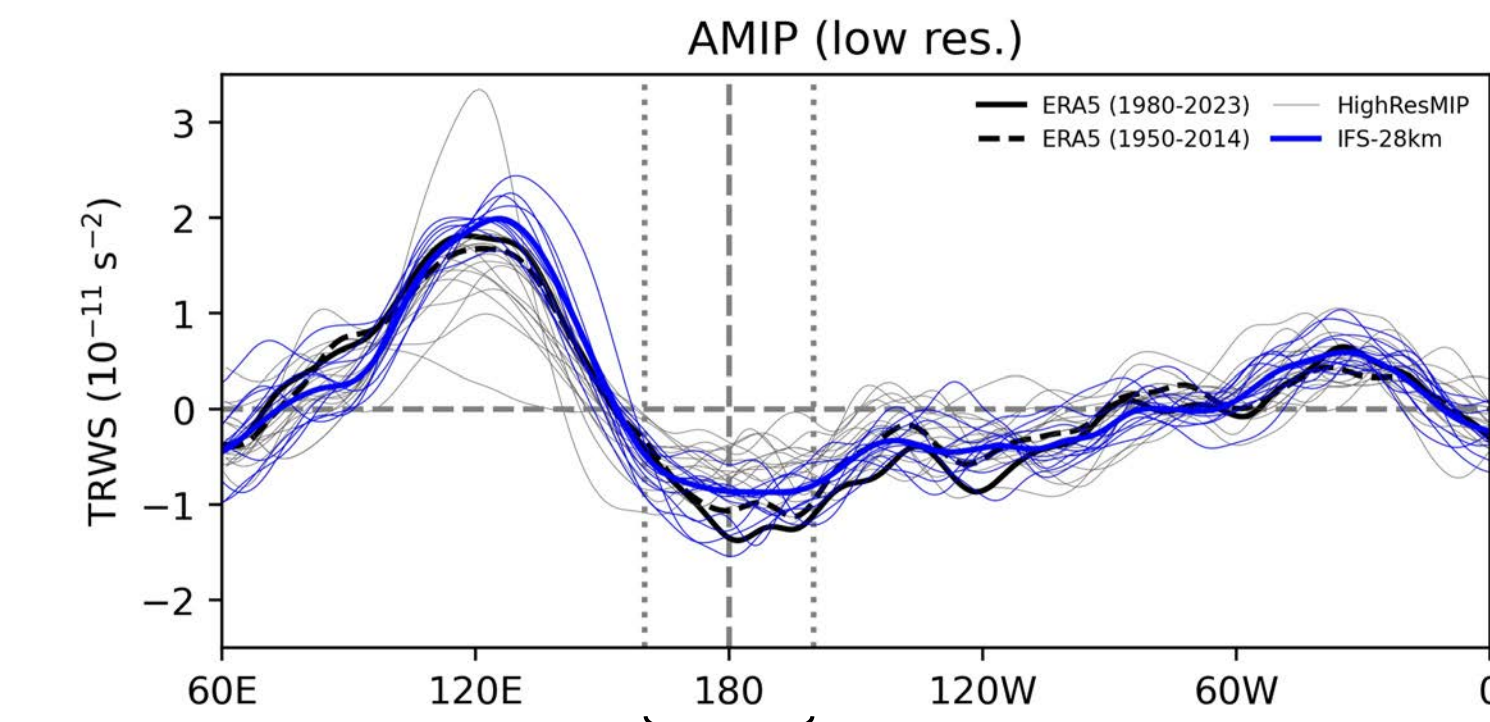
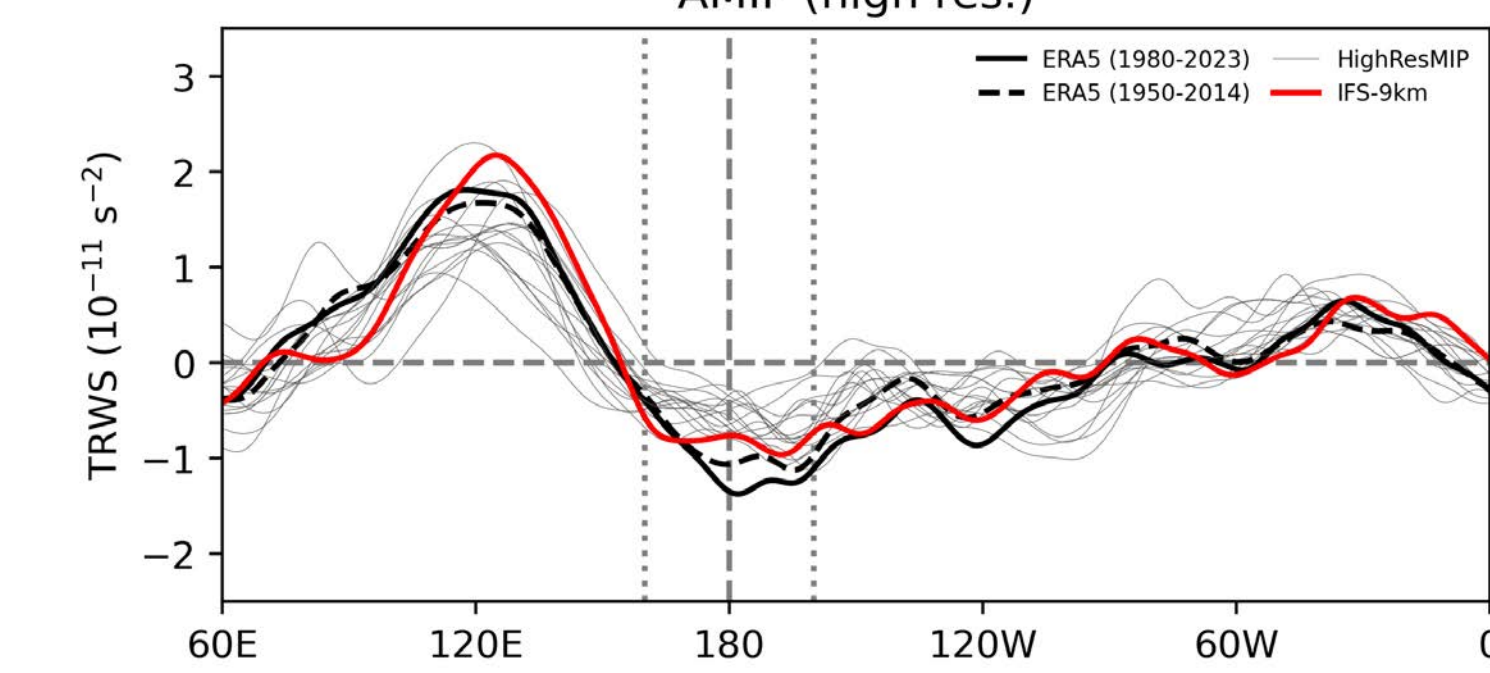
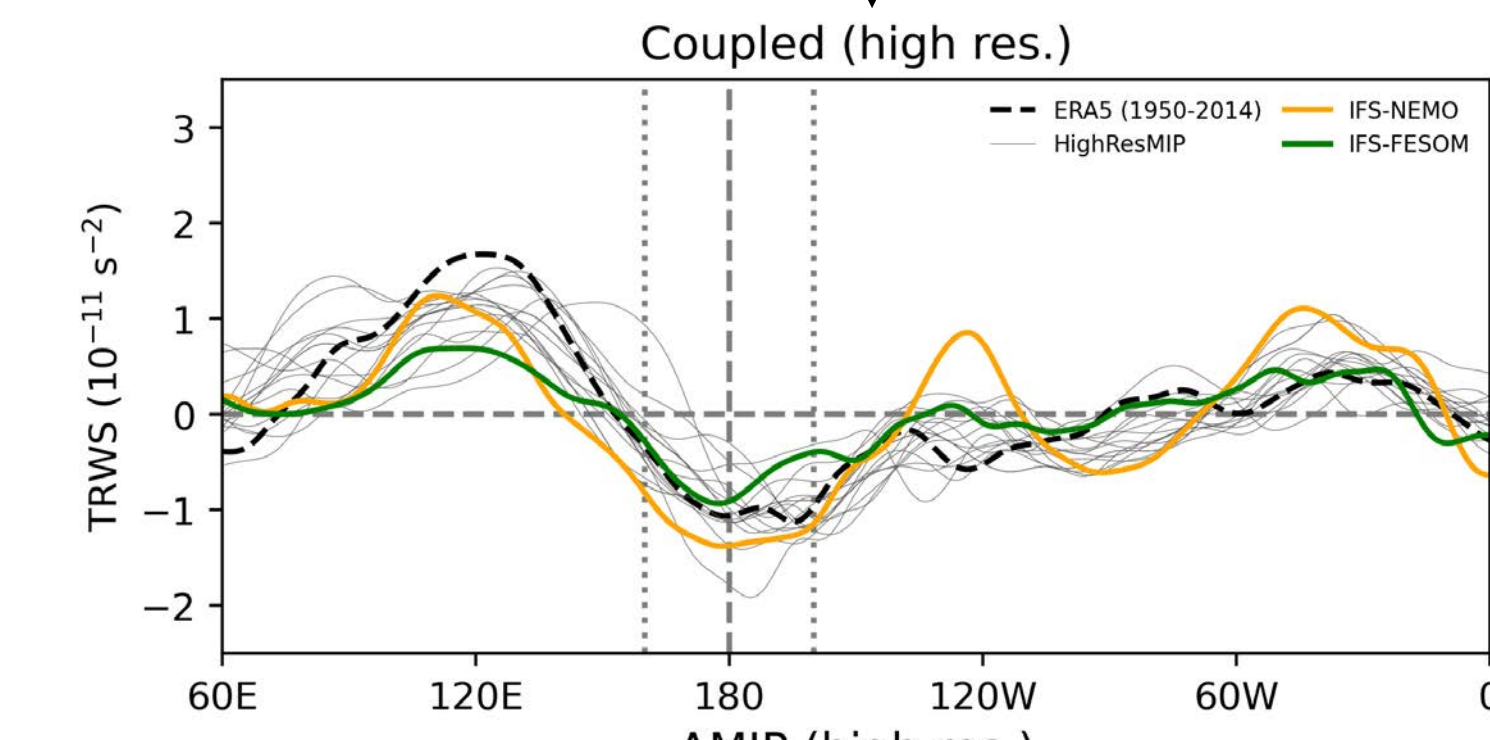
5. Rossby Wave Source (RWS)

We consider the tropical component of the RWS at 200hPa:

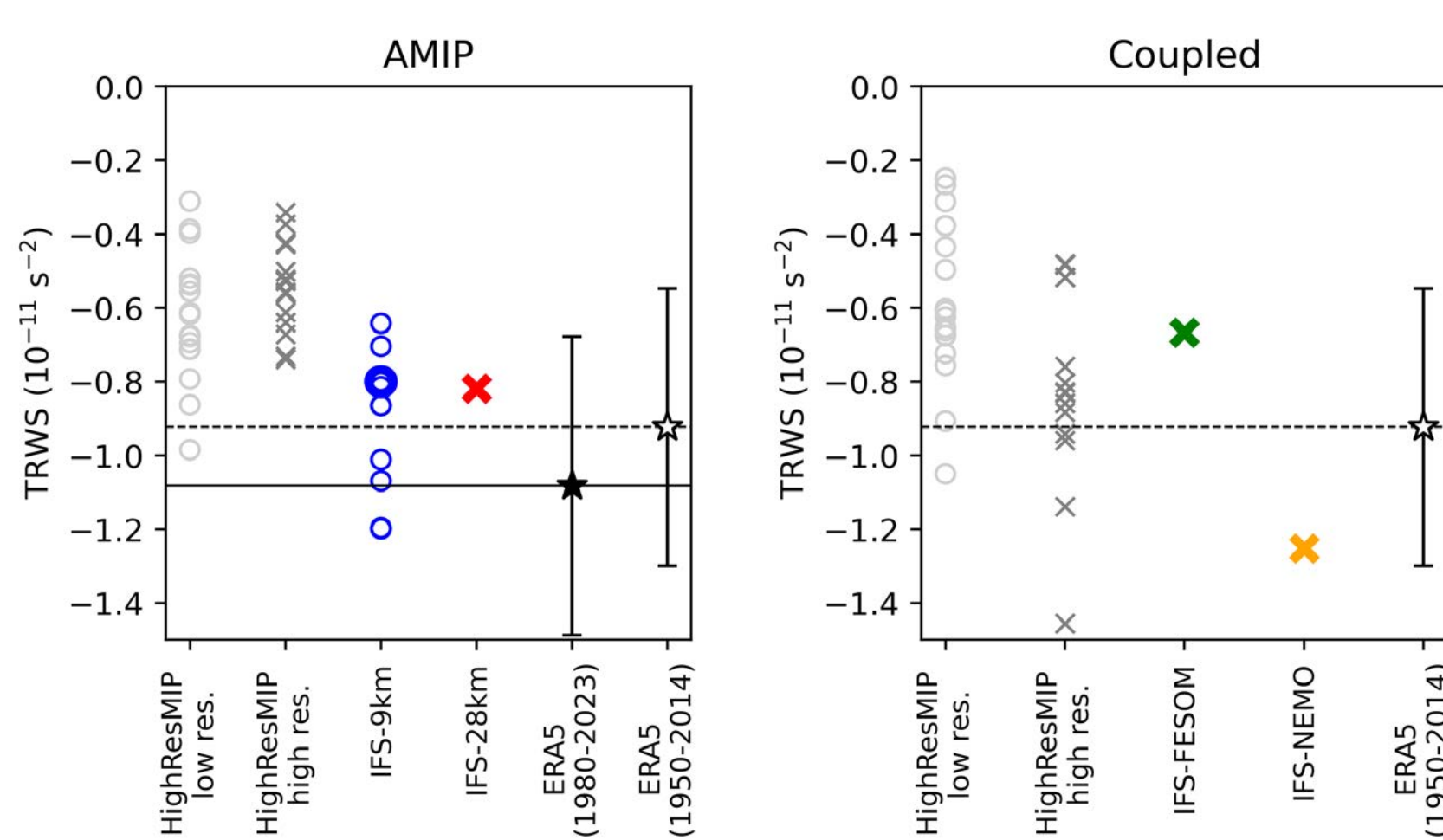
$$TRWS = -V'_\chi \cdot \nabla(\zeta + f)$$



Mean over 15-45°N



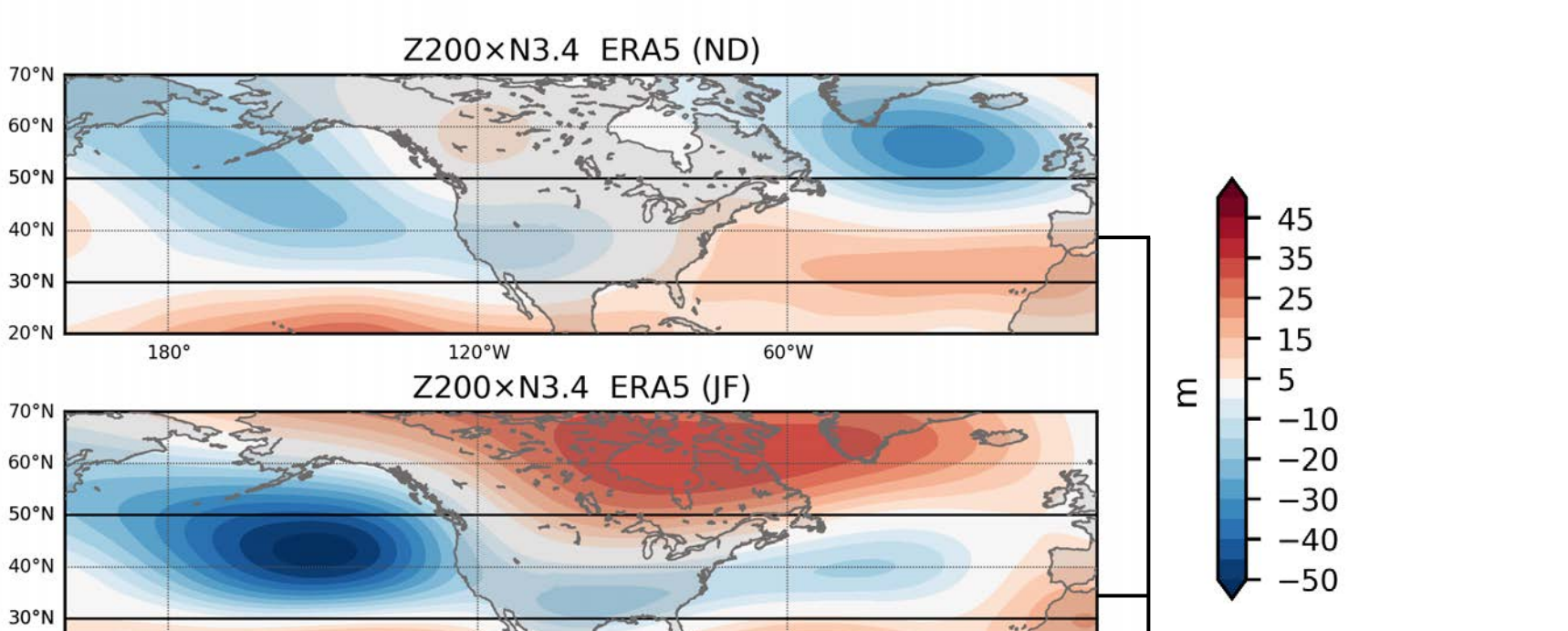
Compute average



- The RWS in the central tropical Pacific is underestimated in **IFS-FESOM** → consistent with velocity potential.
- It is overestimated in **IFS-NEMO** → role of mean flow ζ ?
- Still minor differences between **IFS-9km** and the ensemble mean from **IFS-28km**.

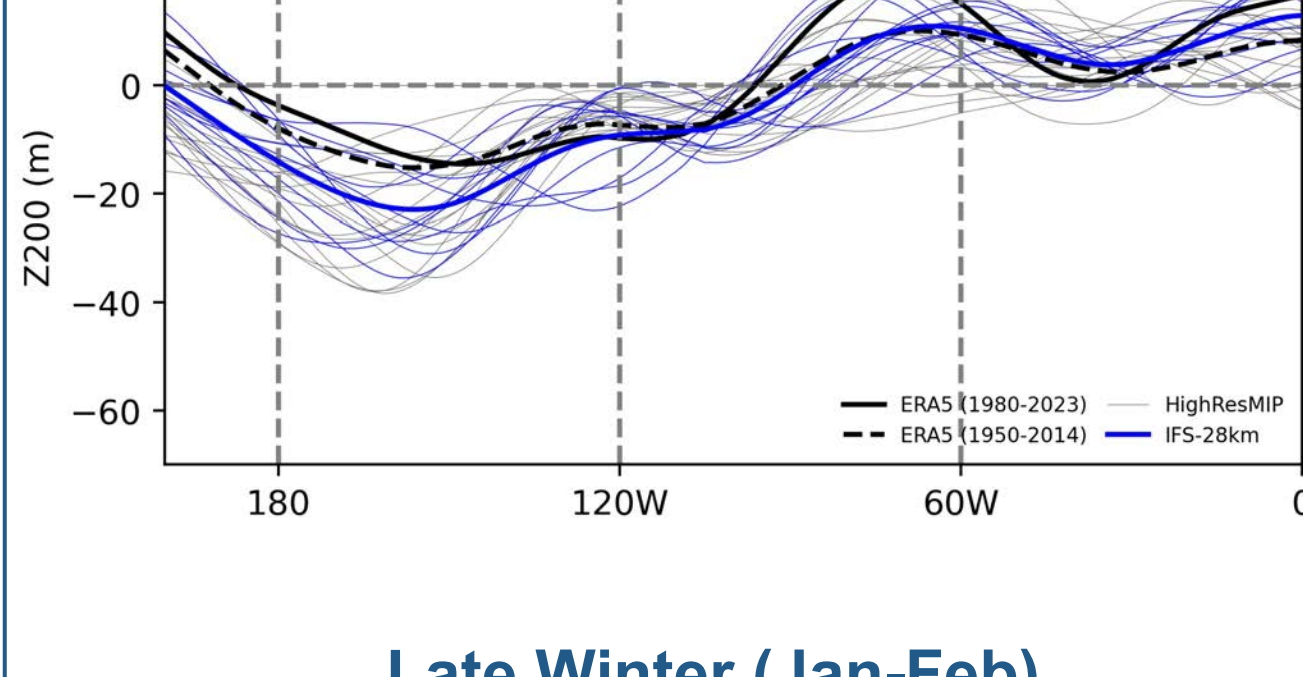
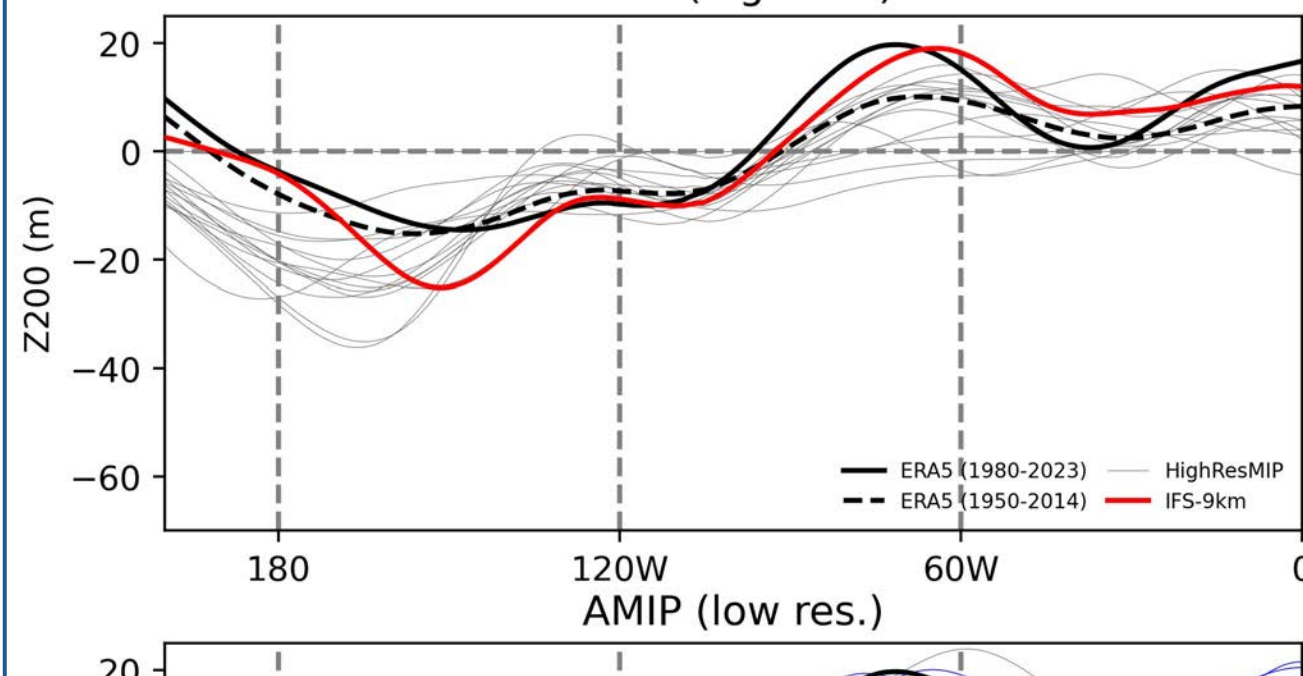
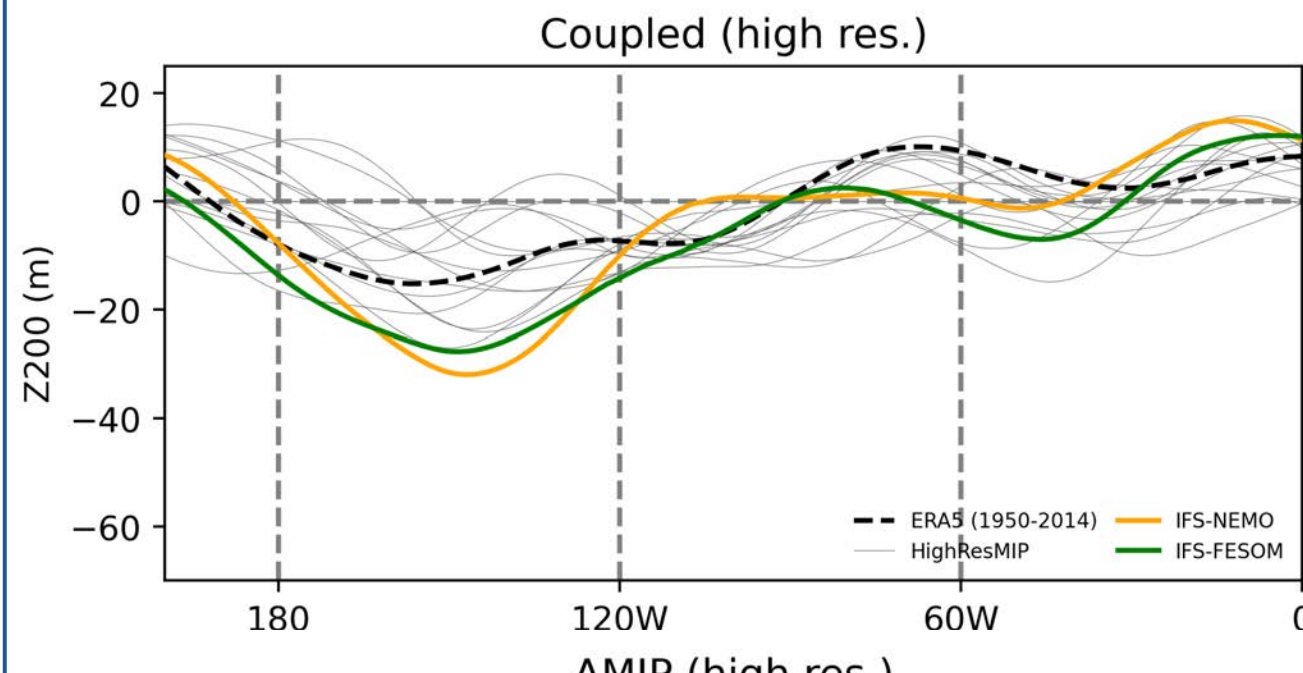
6. Rossby Wave Train (mid latitudes)

We can detect the wave train using the 200-hPa geopotential height, in early and late winter:



Mean over 30-50°N

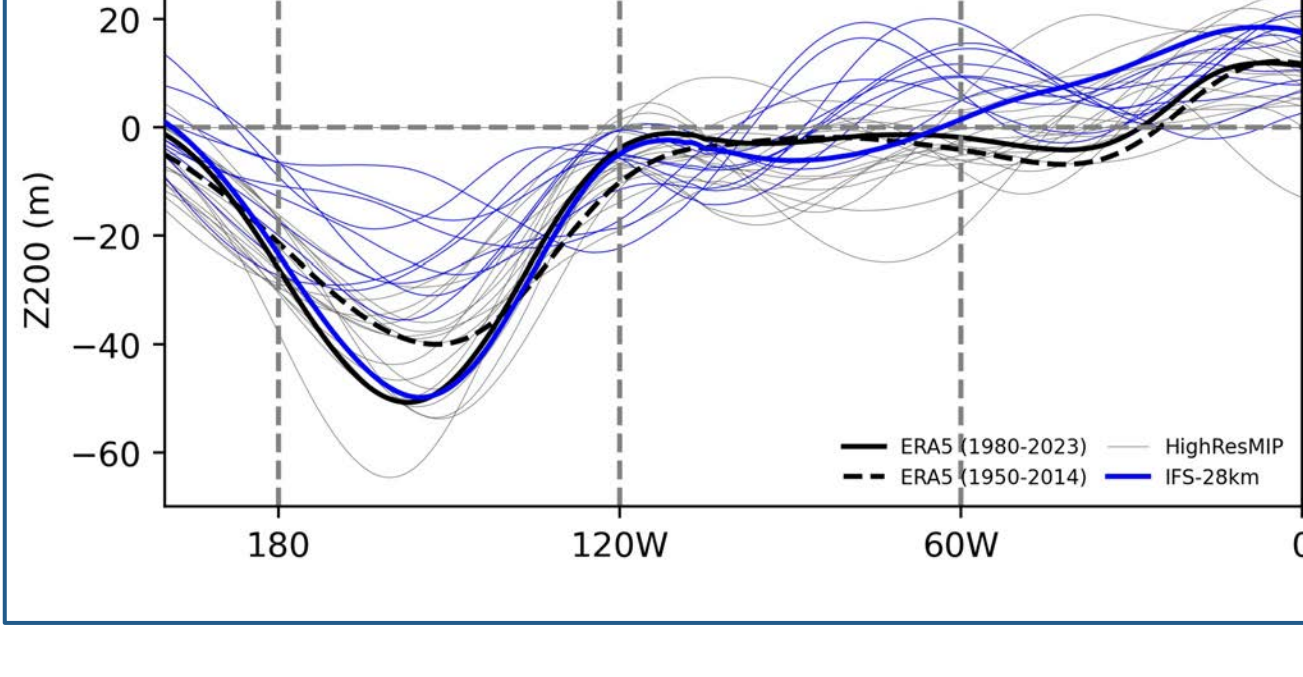
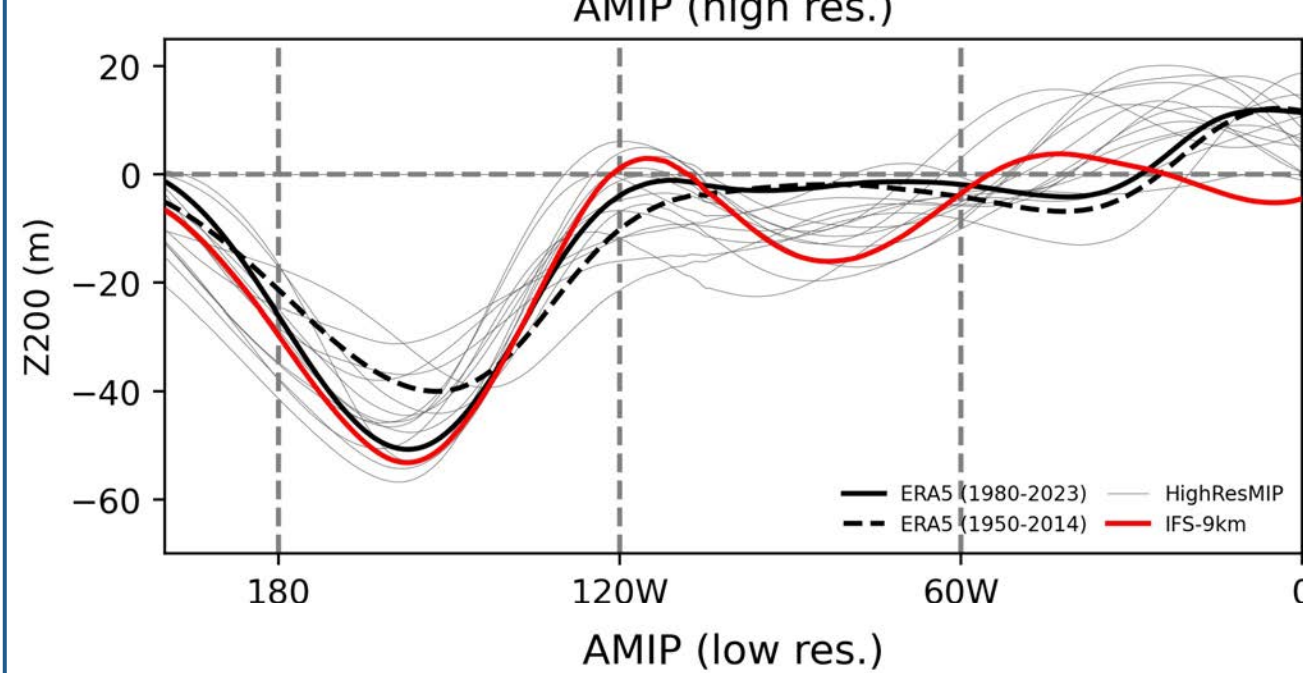
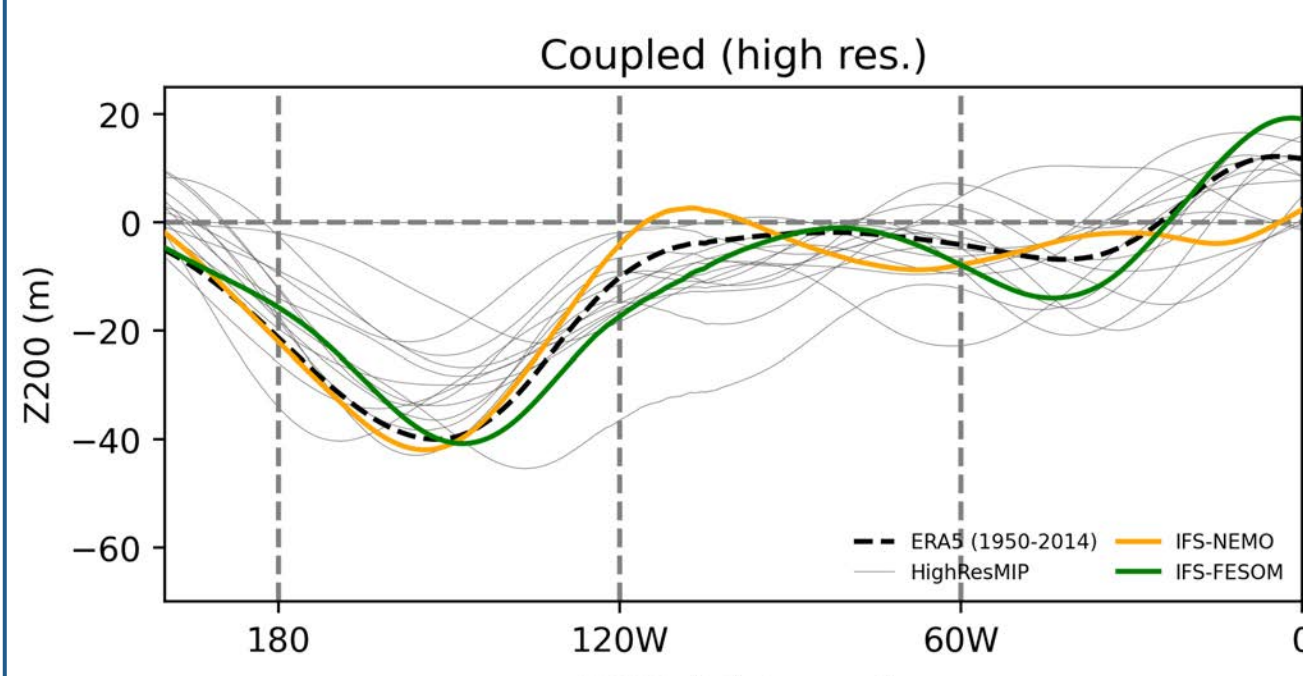
Early Winter (Nov-Dec)



IFS-NEMO and **IFS-FESOM** struggle in both the North Pacific and western North Atlantic → JF-like response ☹

Some differences between **IFS-9km** and the ensemble mean of **IFS-28km**.

Late Winter (Jan-Feb)



Despite some differences in the tropical response and RWS, **all models** capture properly the extra-tropical late-winter signal in the **North Pacific**.

The **North Atlantic** remains a challenge, and no model is a clear winner...

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