

# Predicting ocean oxygen: capabilities and potential

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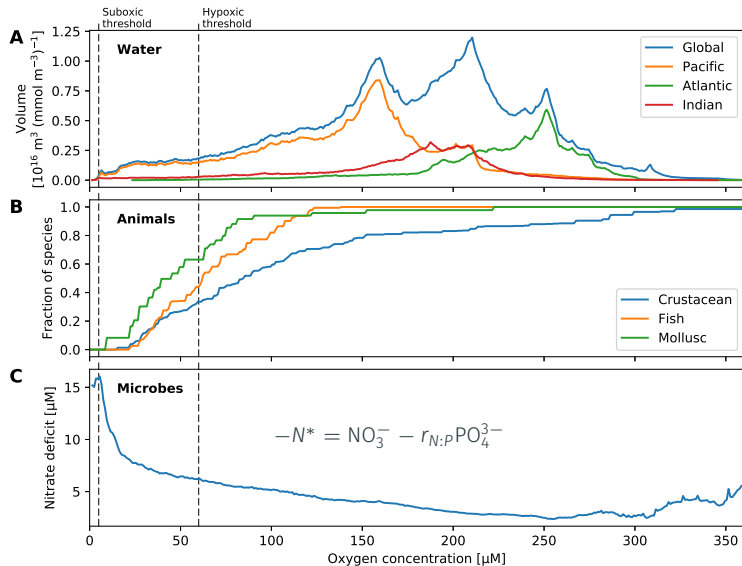


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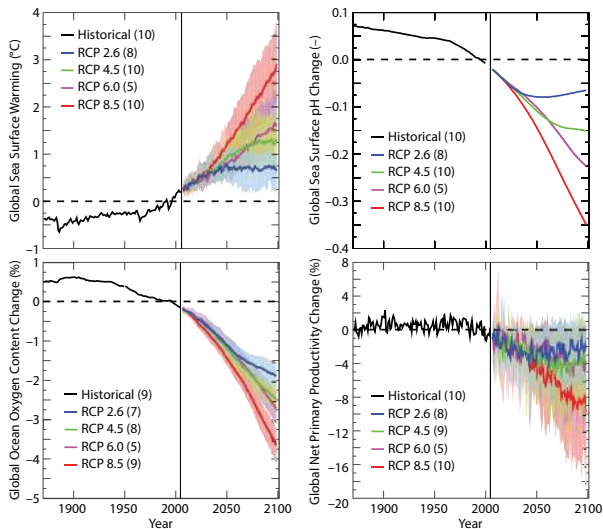


# Oxygen is a fundamental environmental constraint

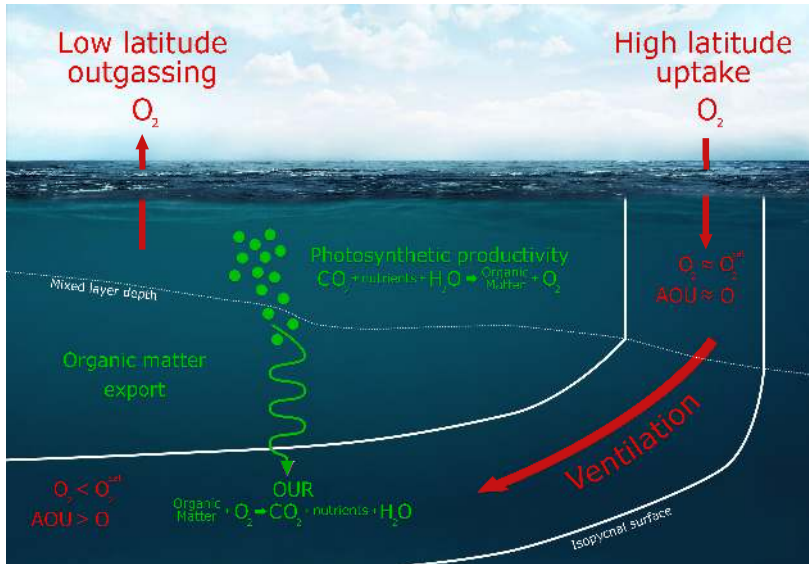


# Warming up, turning sour, losing breath\*

## CMIP5 multi-model global-mean projections



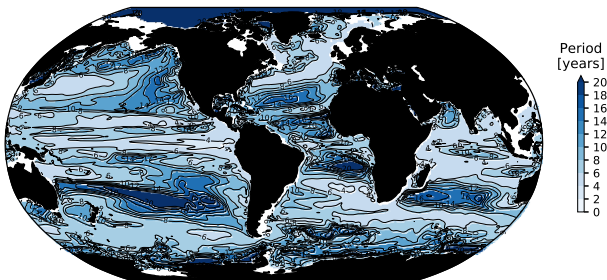
# Physical & biological controls on interior oxygen





# Timescales of natural variability in thermocline O<sub>2</sub>

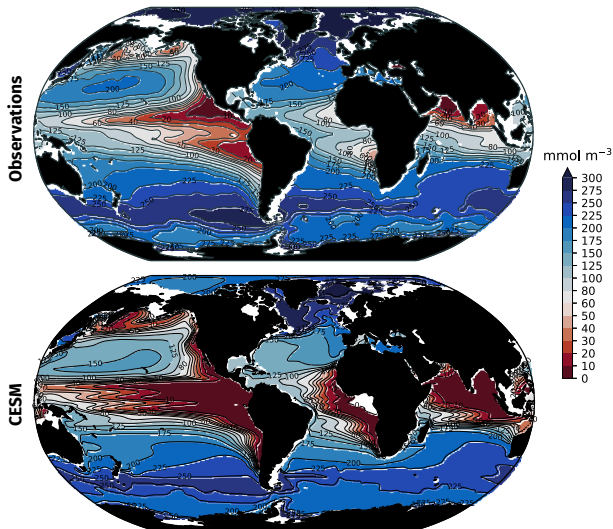
Variance-weighted mean period (CESM 1850-control)



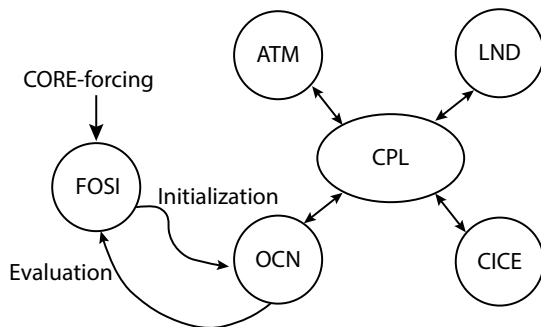
$$T_x = \sum_k V(f_k, x) / \sum_k f_k V(f_k, x)$$

# A persistent bias in Earth system models: Extensive OMZs

## Thermocline (400–600 m) $O_2$ distributions

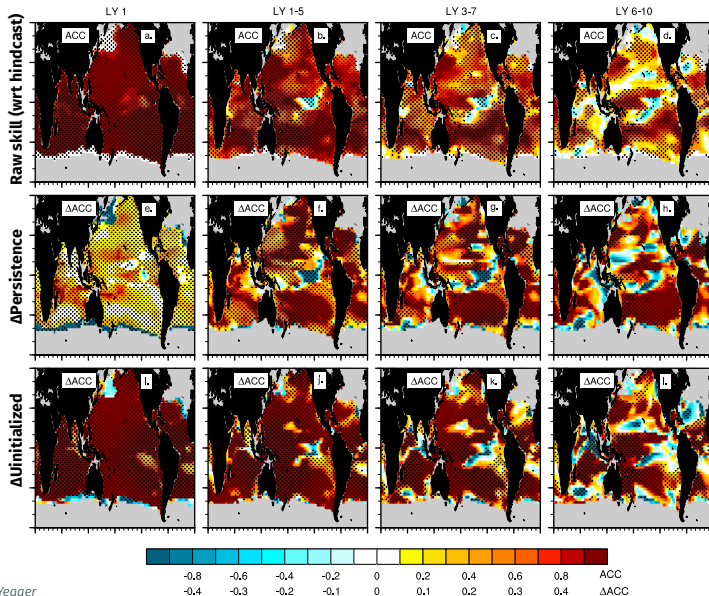


# CESM Decadal Prediction Large Ensemble



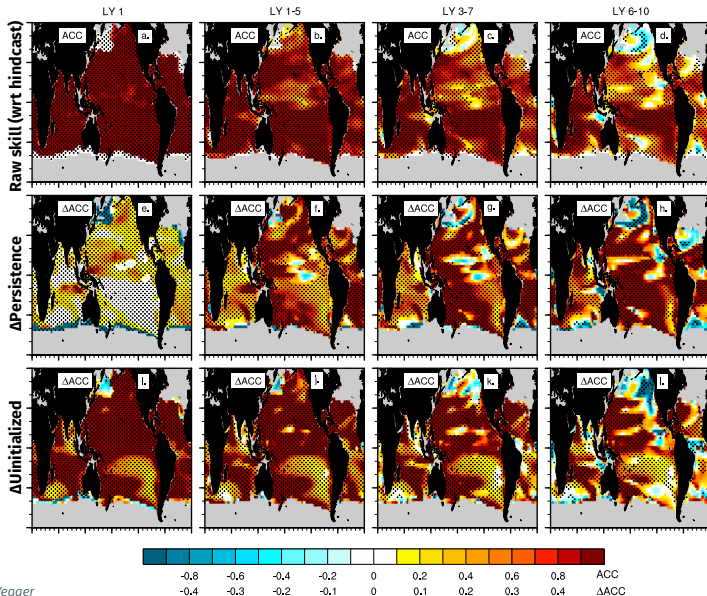
## Thermocline oxygen concentrations look to be highly predictable

Anomaly correlation coefficient:  $O_2$  on  $\sigma_\theta = 26.5$

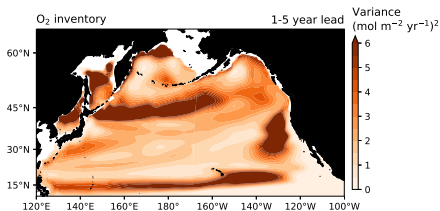
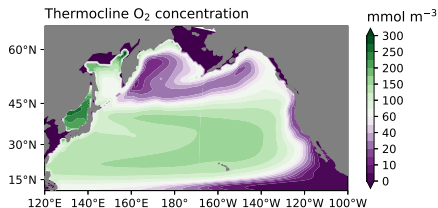


# Thermocline oxygen concentrations look to be highly predictable

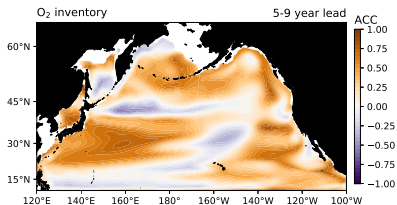
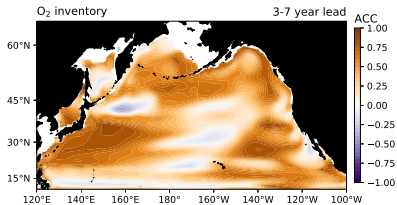
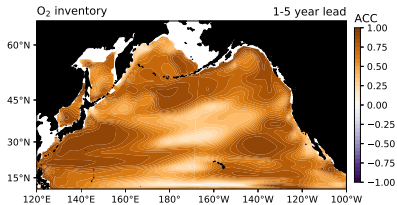
Anomaly correlation coefficient: Salinity on  $\sigma_\theta = 26.5$



# North Pacific dissolved oxygen is skillfully predicted

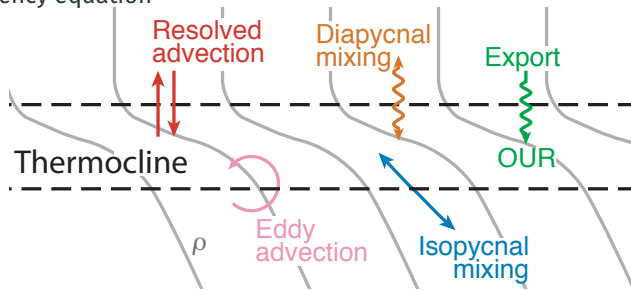


“Thermocline” := 200–600m mean



# What mechanisms provide predictability for O<sub>2</sub>?

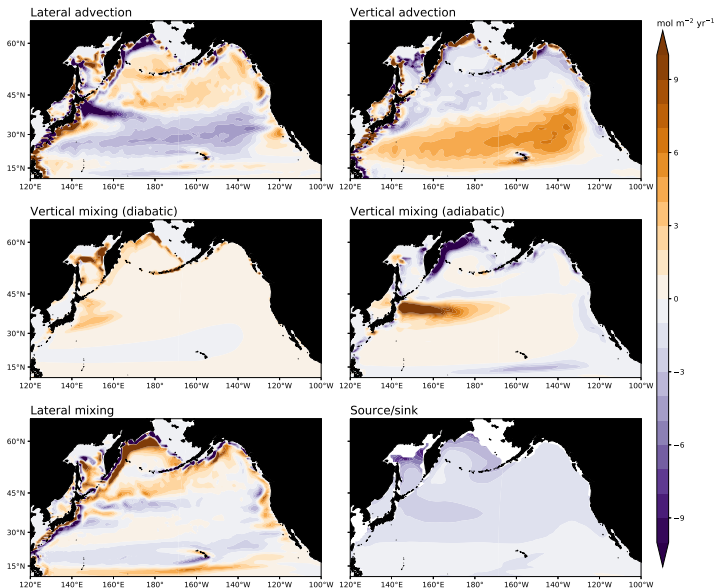
Tracer tendency equation



$$\frac{\partial O_2}{\partial t} + (\mathbf{u} + \mathbf{u}^*) \cdot \nabla O_2 = \mathcal{D}_{iso}(O_2) + \mathcal{D}_{dia}(O_2) + \mathcal{J}_{bio}(O_2)$$

# What mechanisms provide predictability for O<sub>2</sub>?

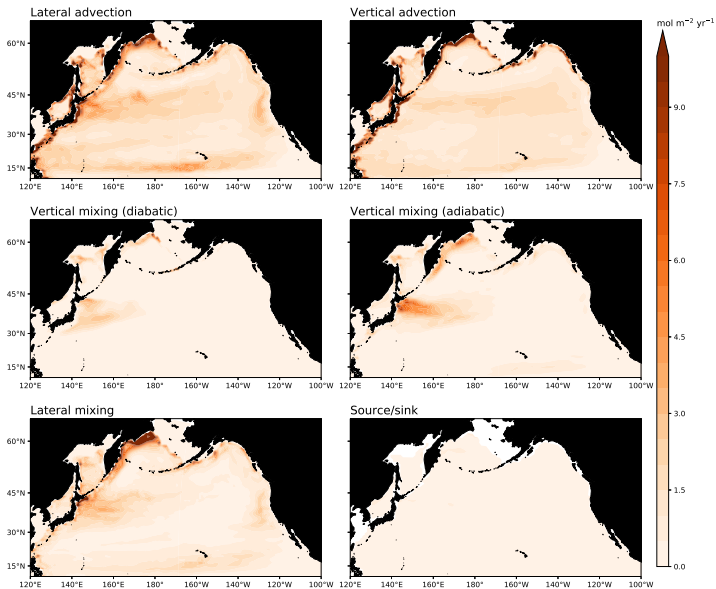
## O<sub>2</sub> term balance: annual mean





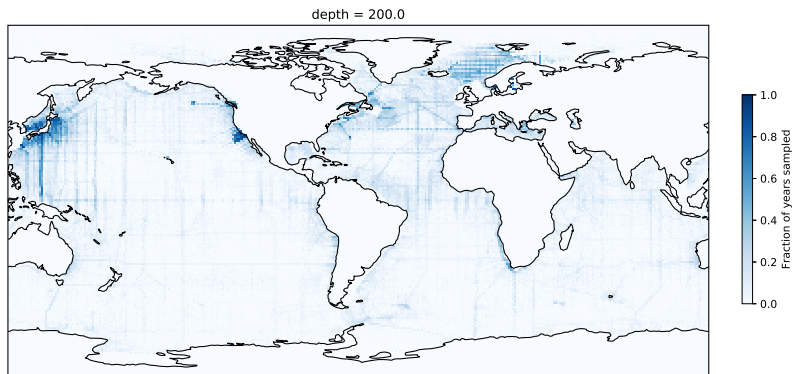
# What mechanisms provide predictability for $O_2$ ?

## $O_2$ term balance: interannual variability (std. dev.)



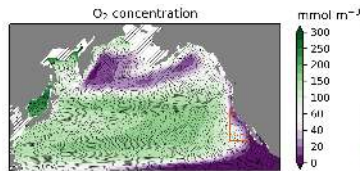
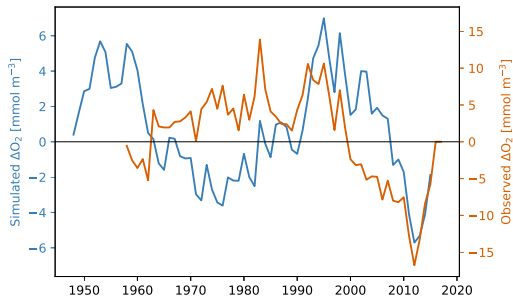
## Timeseries observations of oxygen are sparse

### World Ocean Database 2013: Fraction of years sampled (1958–2015)



# Model skill in CalCOFI\* region: questionable

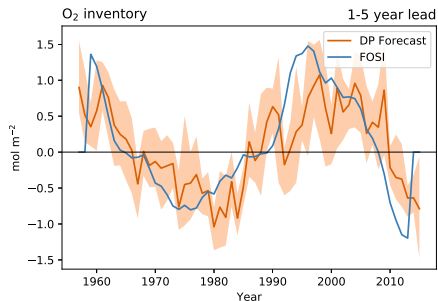
## Thermocline O<sub>2</sub>



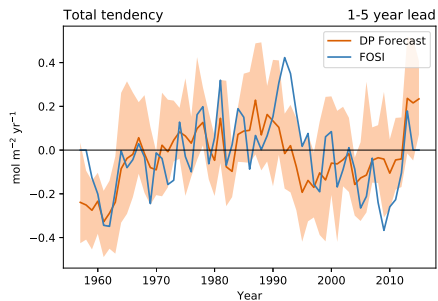
\* California Cooperative Oceanic Fisheries Investigations

# CalCOFI dissolved oxygen is skillfully predicted

## Thermocline O<sub>2</sub> in CalCOFI region

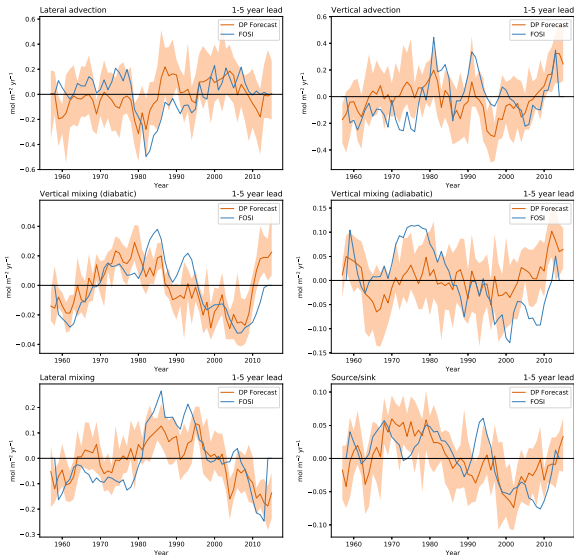


## Thermocline O<sub>2</sub> tendency



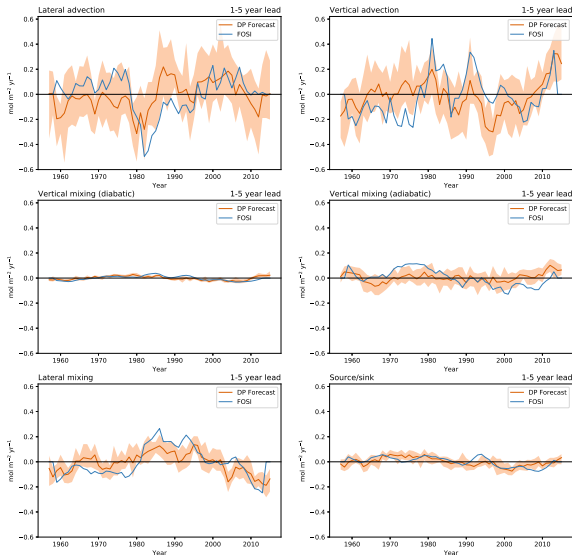
# What mechanisms provide predictability for $O_2$ ?

## $O_2$ term balance: annual mean



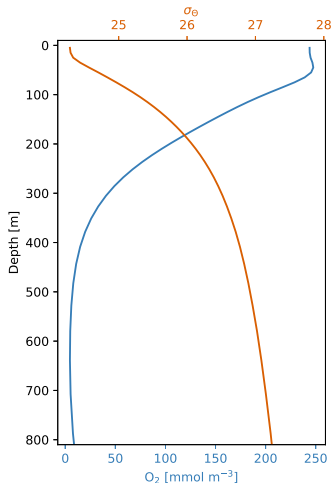
# What mechanisms provide predictability for $O_2$ ?

## $O_2$ term balance: annual mean



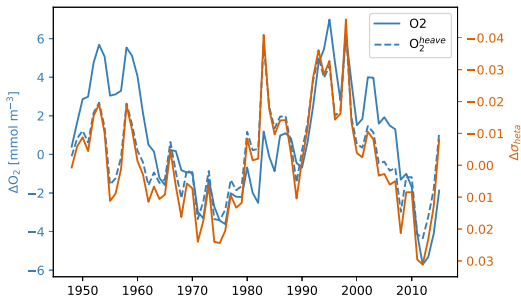
# What mechanisms provide predictability for O<sub>2</sub>?

## Mean vertical gradients



## “Heave”

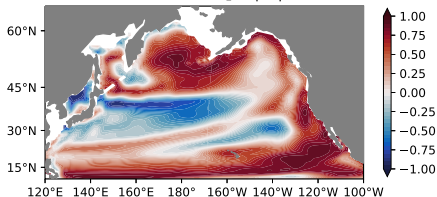
$$O_2^{heave} = \left( \frac{\partial \overline{O_2}}{\partial z} \right) \left( \frac{\partial \overline{\rho_\theta}}{\partial z} \right)^{-1} \rho'_\theta$$



# East-west difference in anomaly generation mechanism

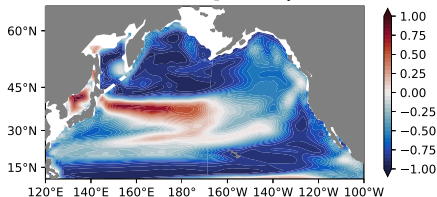
## “Ventilation regime”

Correlation: O<sub>2</sub> v. |PV|

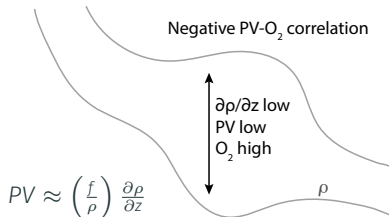


## “Heave regime”

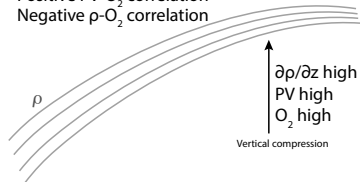
Correlation: O<sub>2</sub> v. density



Negative PV-O<sub>2</sub> correlation



Positive PV-O<sub>2</sub> correlation  
Negative  $\rho$ -O<sub>2</sub> correlation





- Thermocline dissolved oxygen concentrations are highly predictable on multi-annual timescales.
- Vertical displacement of isopycnals in response to basin-scale thermocline adjustment explains much of the variance in simulated CalCOFI O<sub>2</sub>.
- The “heave” regime of the eastern Pacific contrasts with a “ventilation” regime of the west, indicated by differing correlations with PV.
- Model skill remains a challenge.

Questions?

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