

Antarctic stratospheric ozone and seasonal predictability over southern Africa

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Drought in southern Africa

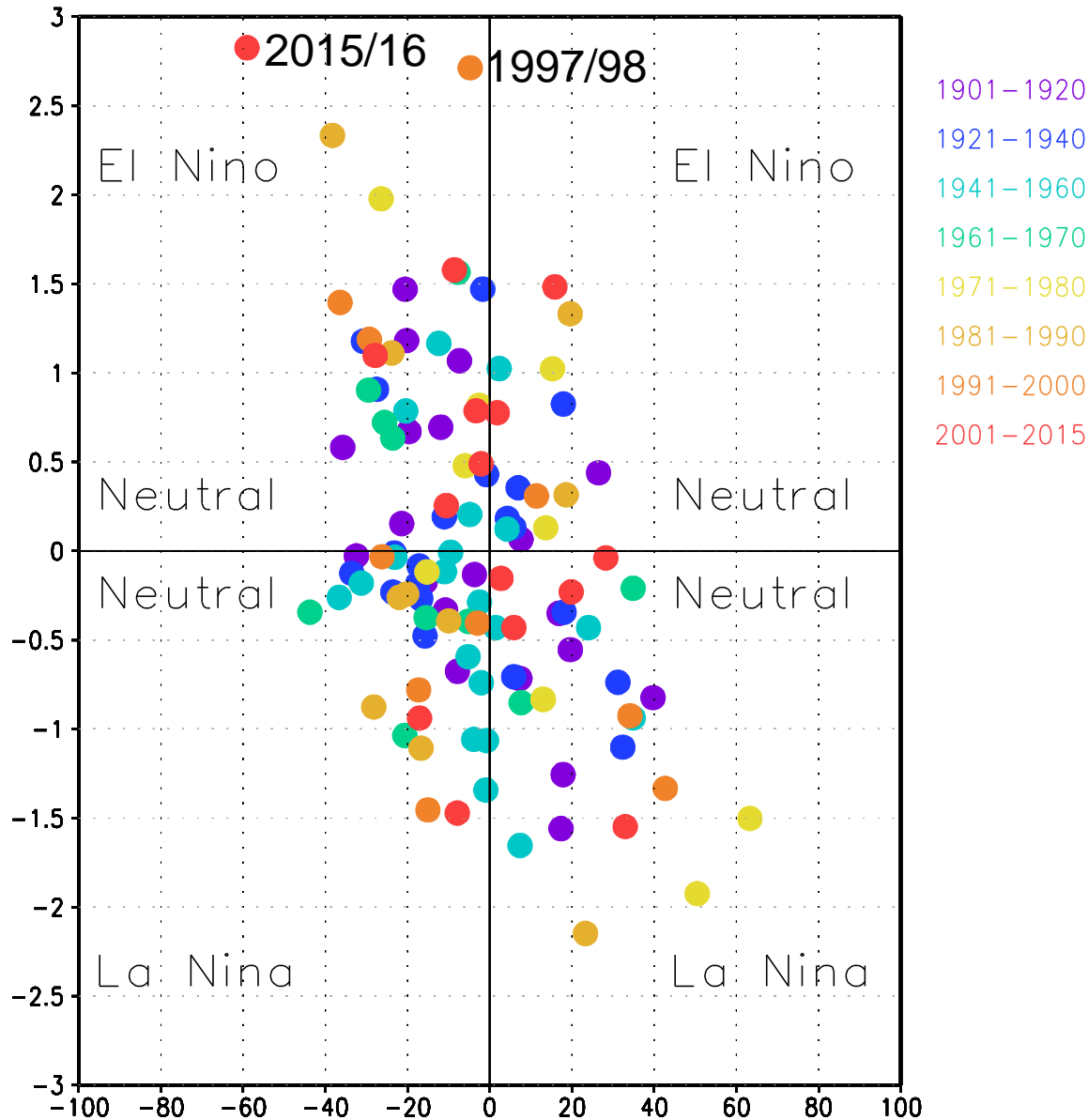


Depleted grazing in Kruger Park in September 2016.

By September 2016, The entire summer rainfall region was in a state of mild drought, or worse, after the 2015/16 super El Niño

The Free State, northern KwaZulu-Natal, eastern Mpumalanga and Kruger Park was in a state of severe drought .

NINO3.4 and FS+NW rainfall 1901–2015



Summer-season rainfall anomalies over the Free State and North West provinces (x-axis) and Niño 3.4 sea-surface temperature anomalies (y-axis) for 1901-2015.

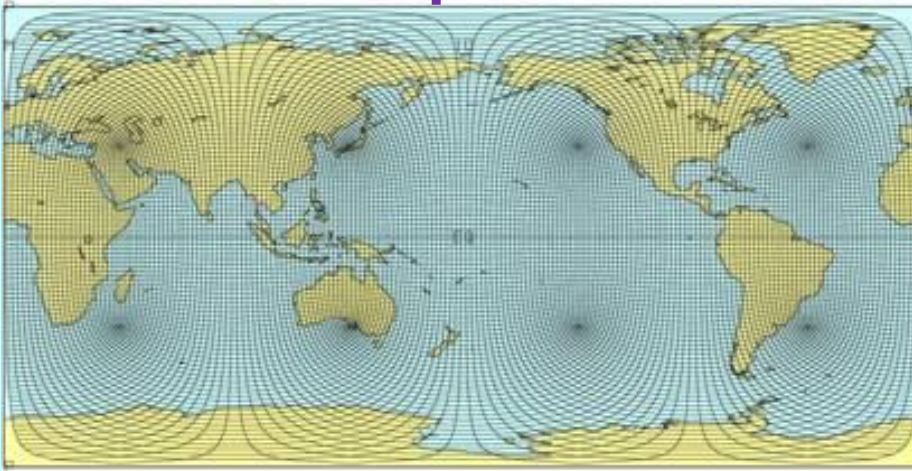
Rainfall anomalies from CRU and GCPC are for DJF. SST anomalies from AMIP II are for OND.

All anomalies were calculated with respect to the 1971-2000 baseline period.

CSIR

CCAM AMIP-style simulations

Control experiment



Climatological CO₂
Climatological ozone
Climatological aerosols
Observed SSTs and sea-ice

CCAM experimental design

C48 (200 km horizontal resolution)

27 sigma levels

Simulation period 1978-2005

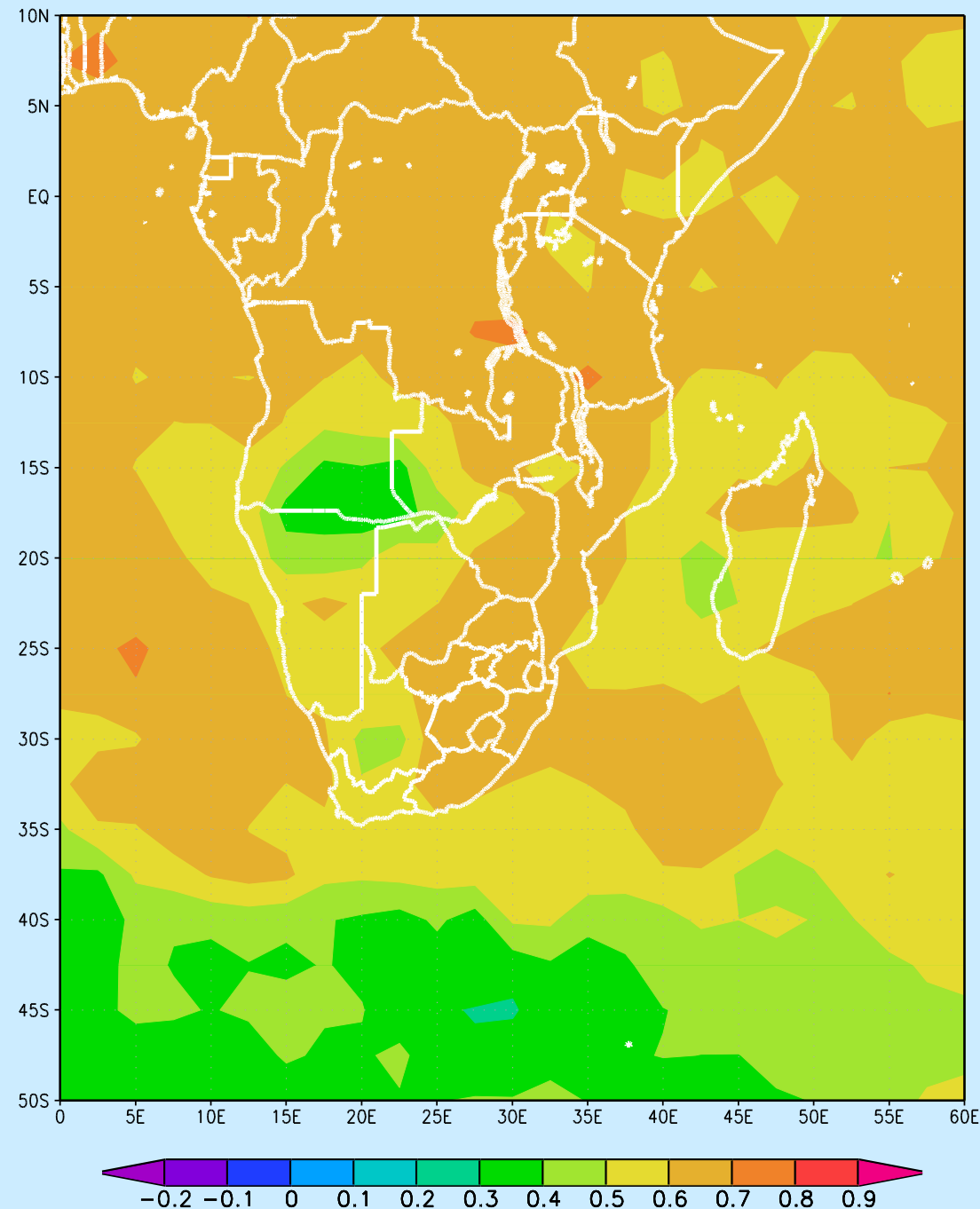
12 ensemble members

CHPC

ACCESS-IDEWS

WRC Radiative Forcing
K5/2163

DJF RPSS

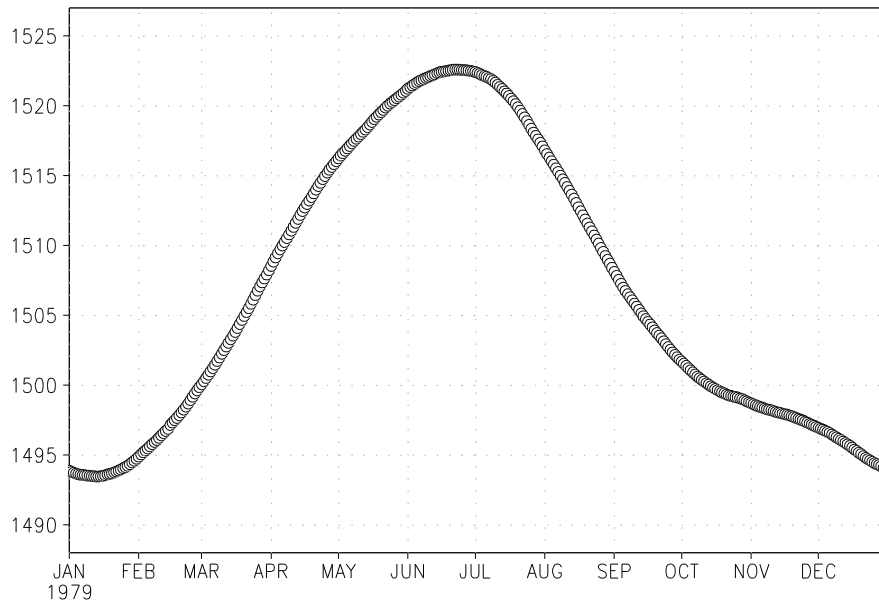


DJF CCAM AMIP-simulation skill in simulating the inter-annual variability in circulation (850 hPa geopotential heights) over southern and tropical Africa

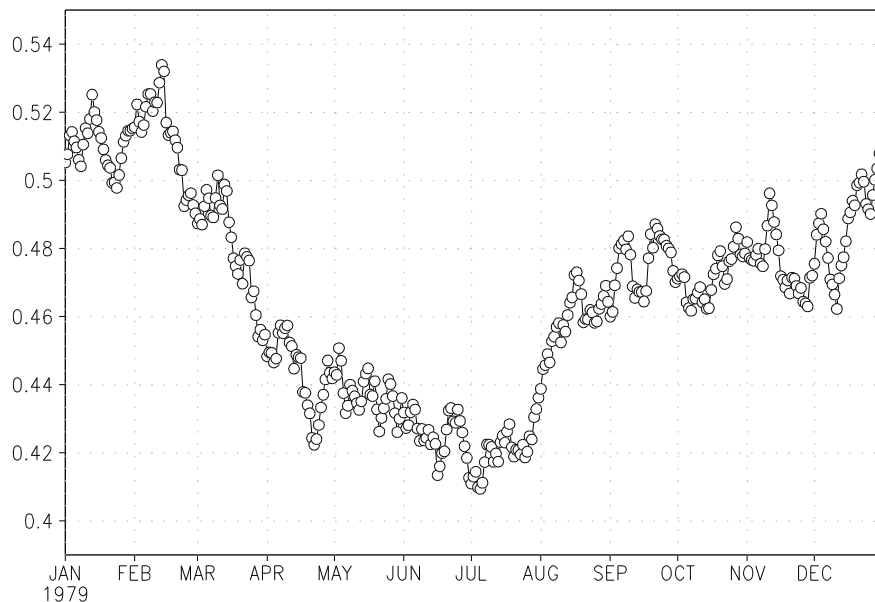
$$\text{RPS} = \sum_{k=1}^K (S_k - O_k)^2$$

$$\text{RPSS} = 1 - \frac{\text{RPS}}{\text{RPS}_c}$$

Seasonal cycle 850 hPa heights



Seasonal cycle RPS



Seasonal cycle in circulation and seasonal forecast skill as deduced from AMIP-style simulations

SA domain: 30 S to 10 S and 15 E to 35 E

Climatological values of radiative forcings (control experiment)

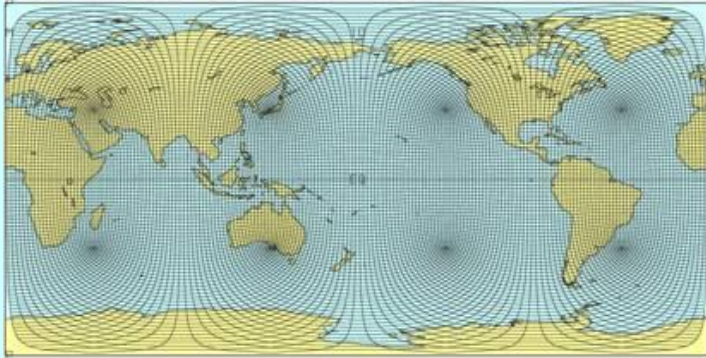
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AMIP-style simulations: Antarctic stratospheric forcing

Control experiment



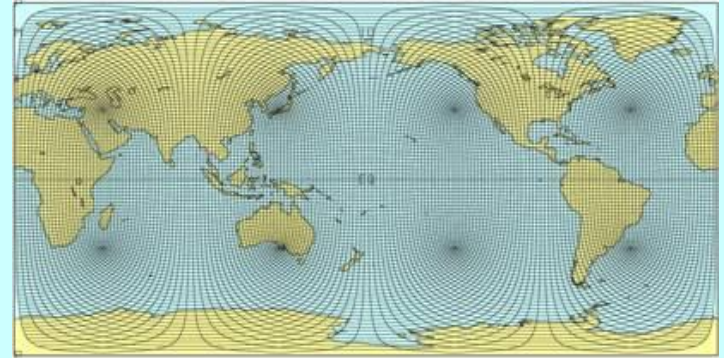
Climatological CO2

Climatological ozone

Climatological aerosols

Observed SSTs and sea-ice

Radiative forcing: ozone



Climatological CO2

Ozone time-varying

Climatological aerosols

Observed SSTs and sea-ice

CCAM experimental design

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Simulation period 1978-2005

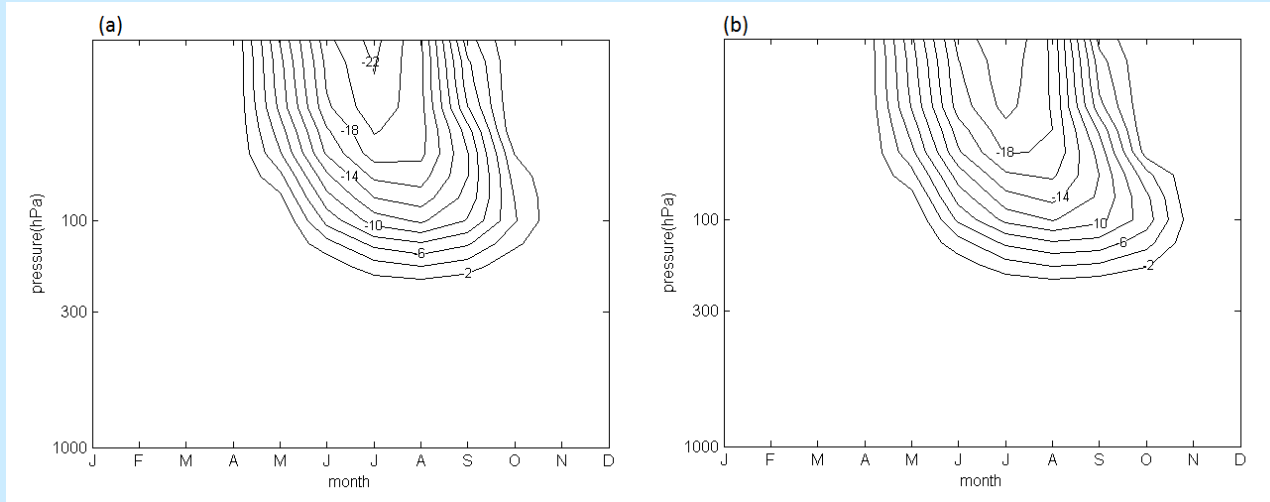
12 ensemble members

CHPC

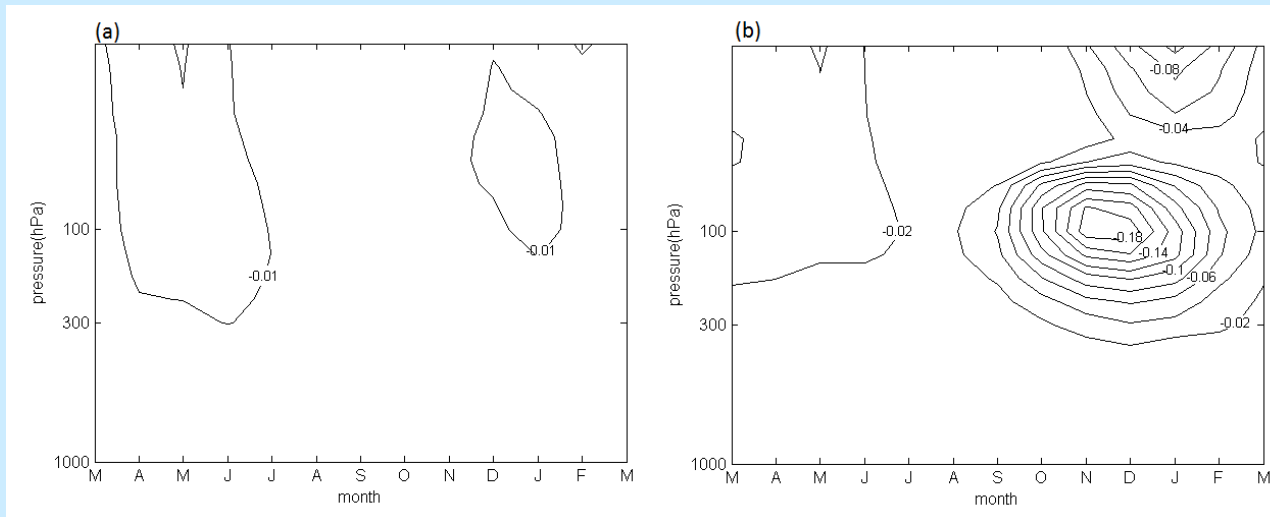
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CCAM simulations of polar cap temperatures (Temp -200 K)



Simulated seasonal evolution of polar-cap (70S-90S) temperatures

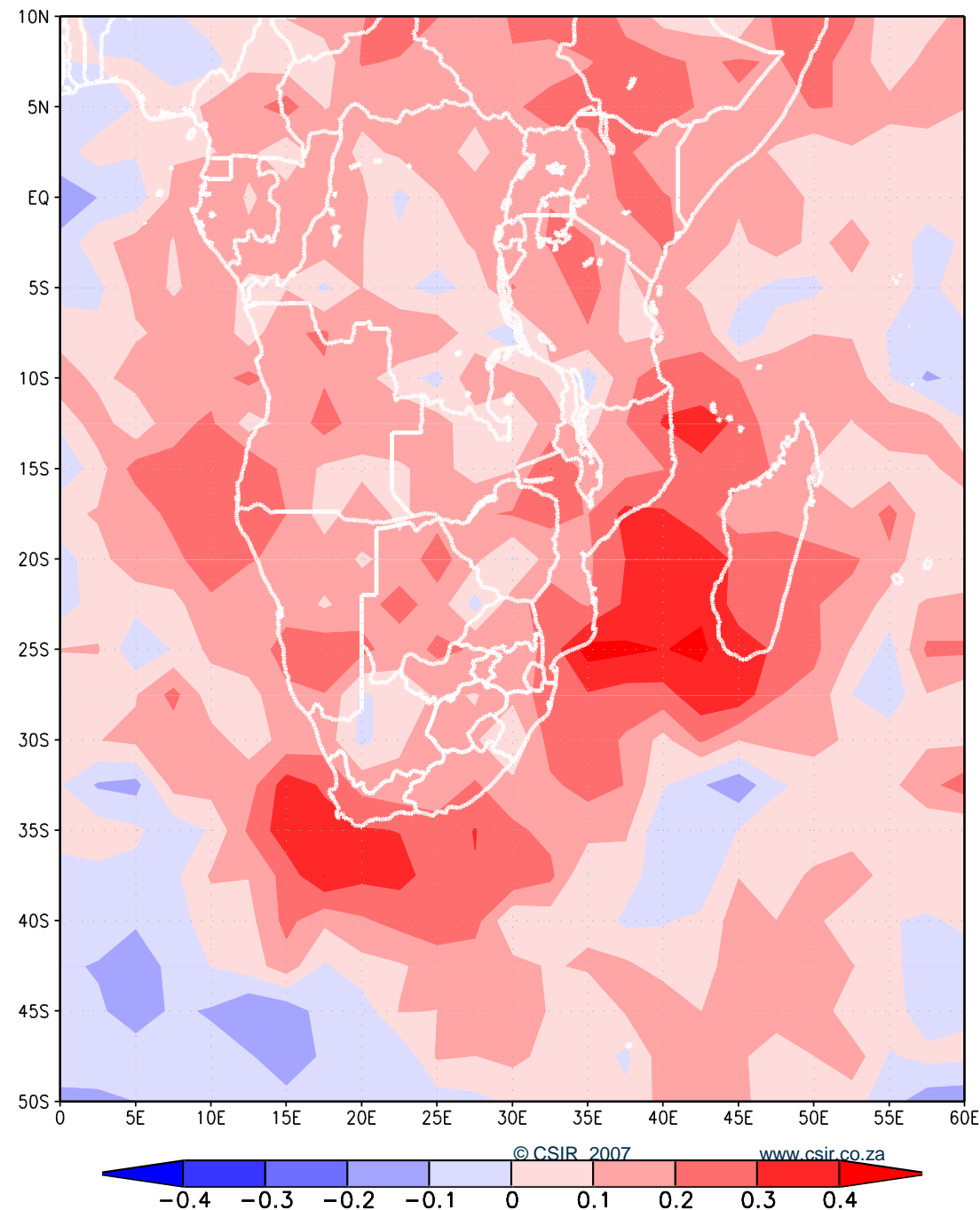


Simulated seasonal trends in polar-cap (70S-90S) temperatures

Control experiment

Ozone Radiative forcing experiment

DJF SS-RPS



**Change in skill for
DJF: Time-varying
ozone vs
climatological values
(control experiment)**

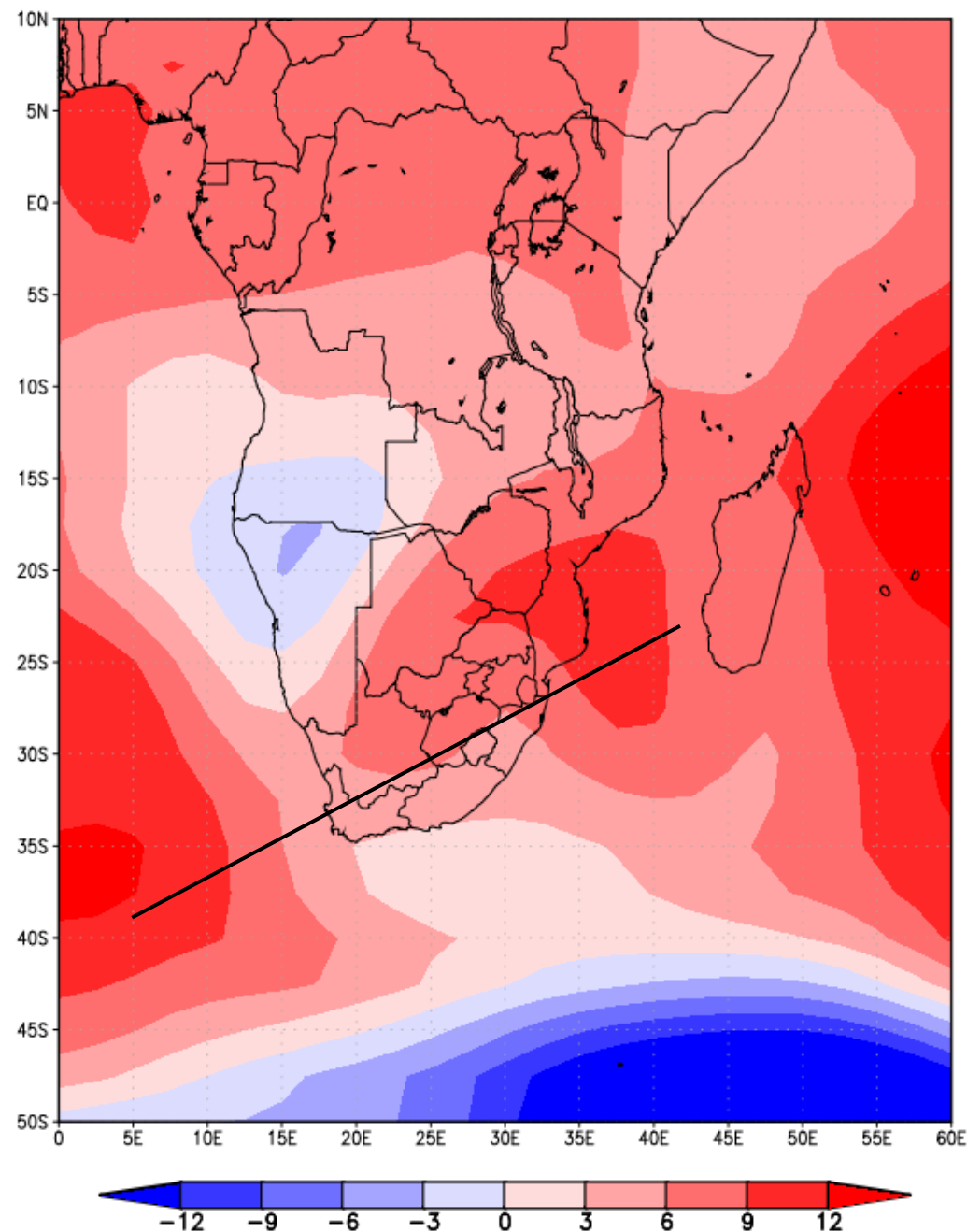
**Inclusion of time-varying
stratospheric ozone leads to a
step-up change in DJF
predictability – the ozone
signal is seemingly
sufficiently strong to
overcome the westerly-wave
chaos-barrier**

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**WRC Radiative
Forcing K5/2163**

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DJF 850 hPa anomalies



850 hPa circulation anomalies during December-February of 1997/98

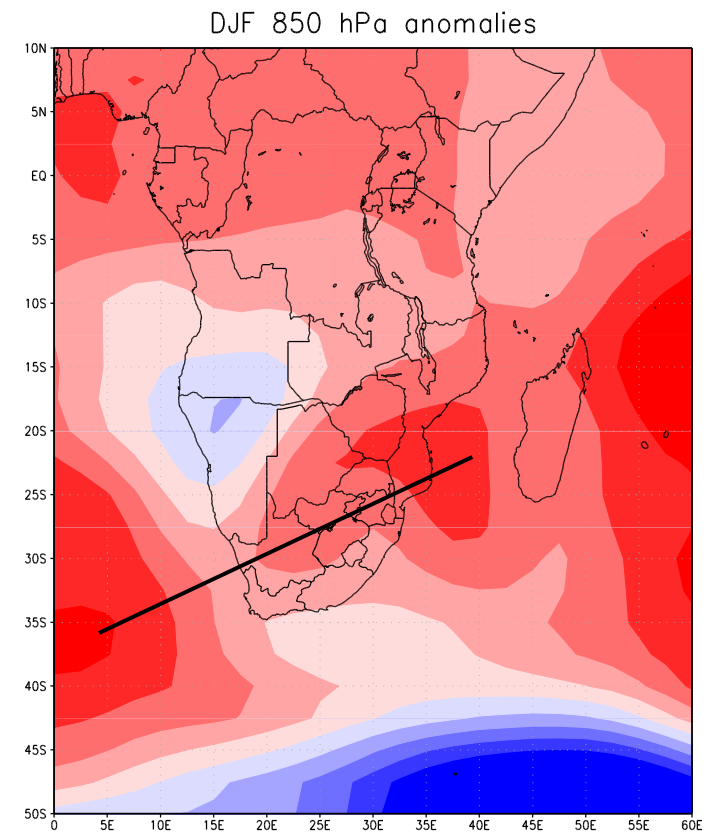
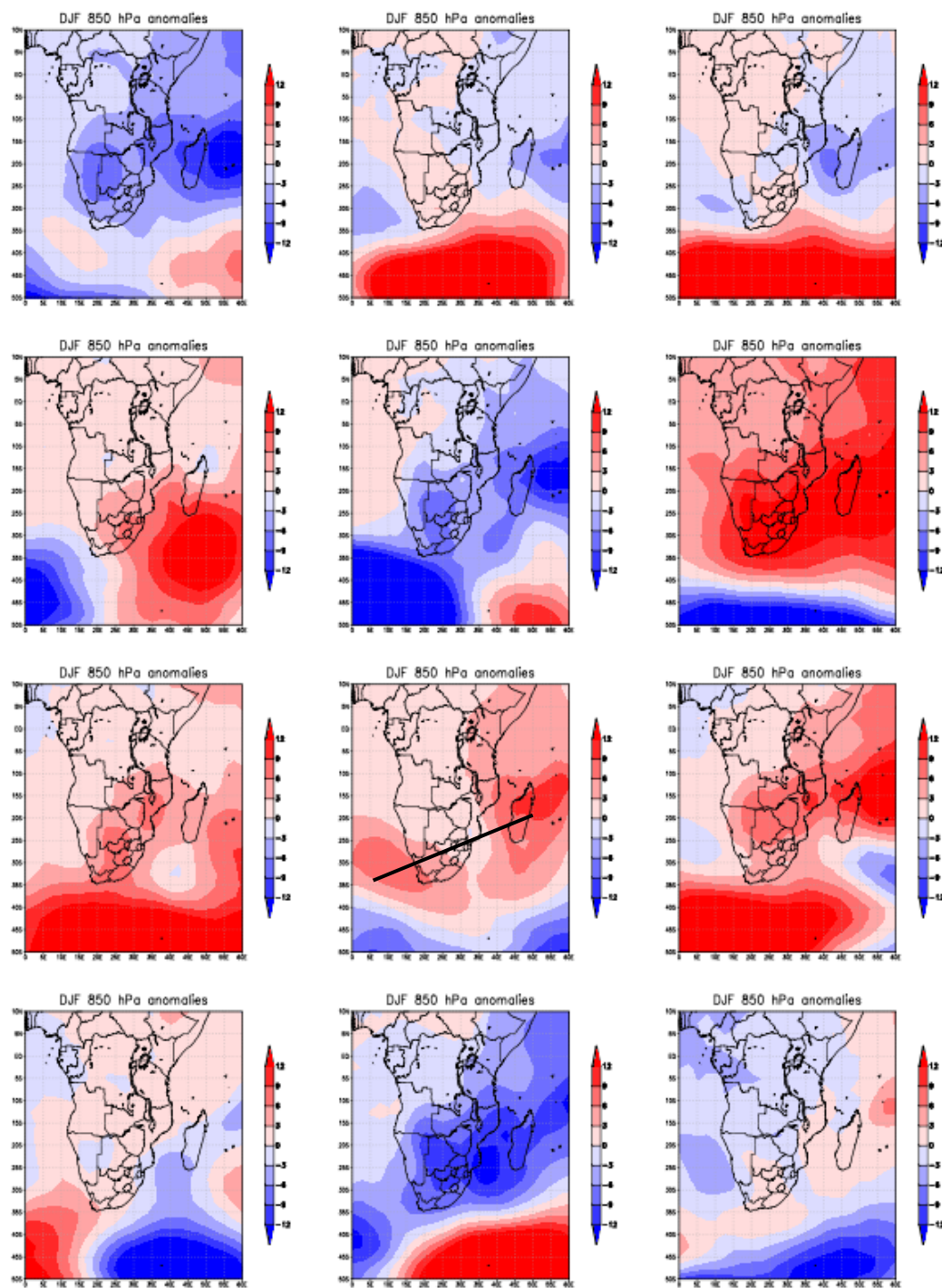
Pattern of trough formation and relatively deep Angola low in the west, with a ridge-axis from the southwest to the north-east over southern Africa occurred, which promoted moisture advection from the east

CHPC

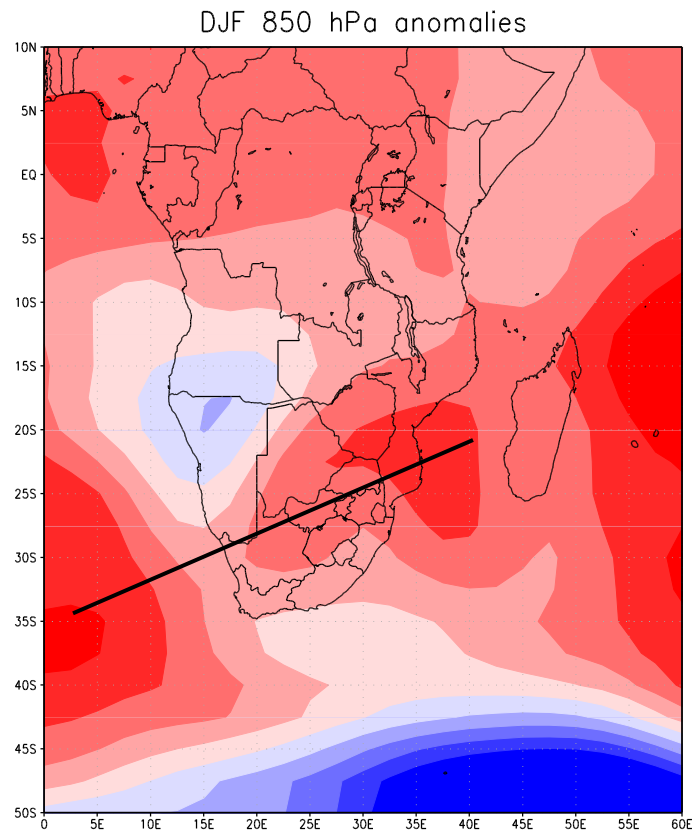
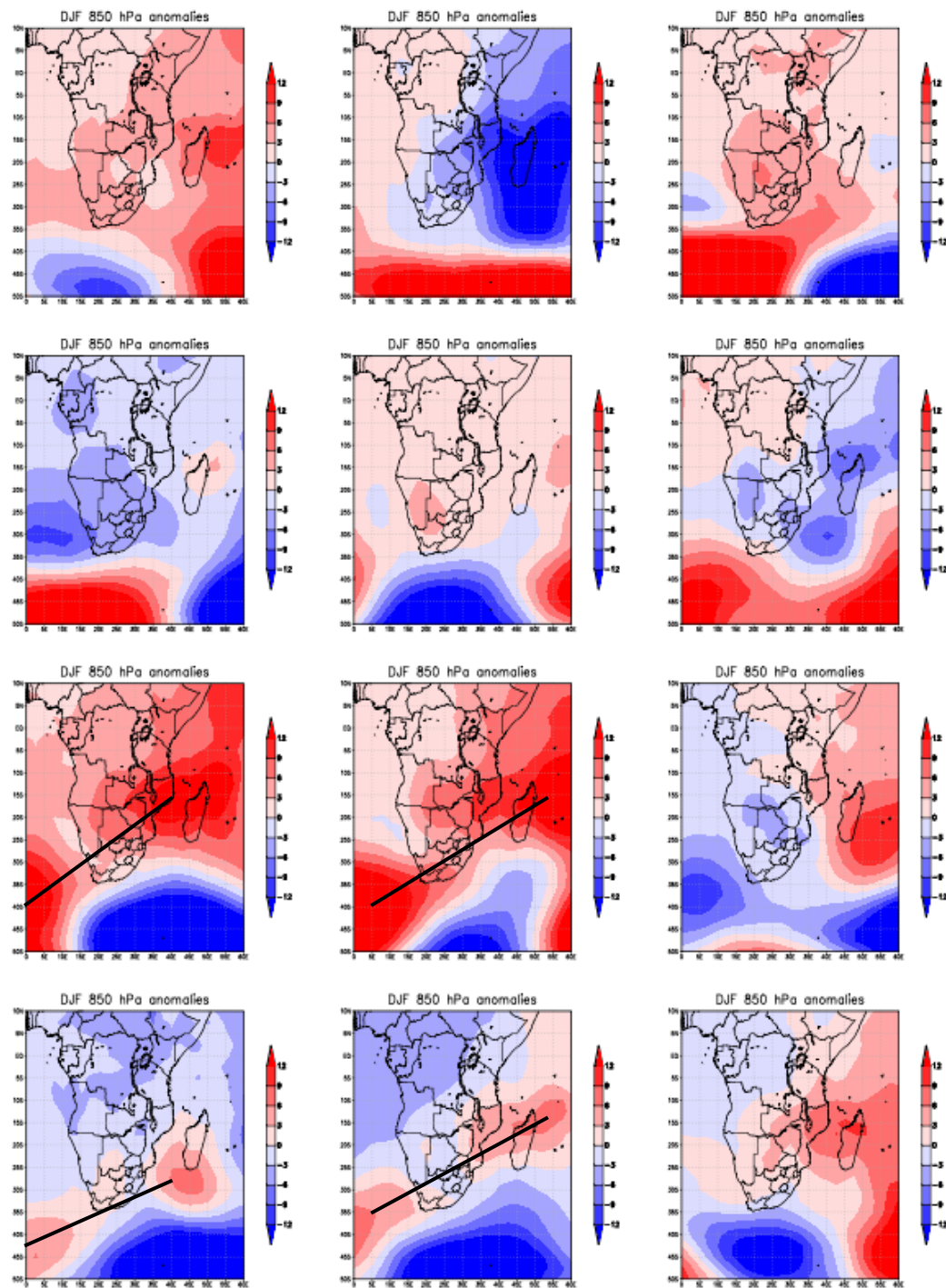
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The “control” simulation (climatological ozone) fails to simulate the circulation anomalies of the 1997/98 El Niño (12 ensemble members)



“Ozone experiment”:
Anomalous Antarctic
stratospheric ozone
concentrations may have
contributed to the “normal
rainfall” of the 1997/1998
El Niño over southern
Africa



Conclusions

- **CCAM forced with observed time-varying ozone concentrations is capable of simulating the observed cooling trends in SH polar-cap stratospheric temperatures that have occurred over the last few decades**
- **Simulations of inter-annual variability are most skilful for DJF over southern Africa**
- **High-latitude tentacles of low predictability reach southern Africa during MAM, JJA and SON**
- **Inclusion of time-varying stratospheric ozone leads to a step-up change in DJF predictability – the ozone signal is seemingly sufficiently strong to overcome the westerly-wave chaos-barrier**
- **Further investigations with 72 level versions of CCAM and additional climate models to confirm this finding**