

# Global overturning circulation: freshwater transports and down-gradient flow Lynne D. Talley and Joseph L. Reid (Italley@ucsd.edu)

Scripps Institution of Oceanography, UCSD La Jolla CA WCRP Open Science Conference Denver, 2011

#### **Abstract**

The relation between the observed global overturning circulation, observed steric height distributions at all depths, and freshwater transports is examined.

- (1) The global circulation is down-gradient in terms of the largest spatial scales of steric height distributions, at all depths from the surface to the bottom, based on observational syntheses. Diapycnal upwelling in both the Southern Ocean and the Indian/Pacific Oceans is integral to the global overturn, in addition to downwelling in the well-known deepwater source regions. The Indian/Pacific upwelling to mid-depth forms the deep waters that return to the Southern Ocean (22 Sv total), upwell to the sea surface, and feed the northward surface flows that eventually return to the northern North Atlantic (NADW overturning cell of 18-20 Sv). The remainder of the Indian/ Pacific Deep Waters join the upwelled North Atlantic Deep Water to form the large (34Sv) Antarctic Bottom Water overturning cell.
- (2) The distribution of diapycnal fluxes depends mainly on the small salinity differences between ocean basins that arise from the pattern of atmospheric water vapor transports. The equatorward freshwater transports from the high latitude southern and northern hemispheres are carried by the surface overturning cells in the south, and by North Atlantic Deep Water and North Pacific Intermediate Water formation in the north: the dynamical difference is due to the open Drake Passage in the south.

# **Global Overturning Circulation**





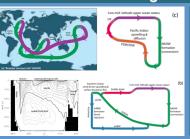


Upwelling of AABW in the Indian and Pacific, into Indian and Pacifi

Antarctic, feeding northward transport of SAMW/AAIW (with buoyancy gain) plus formation of AABW (with buoyancy loss)

Southern Ocean view of the overurning circulation, in the style of Schmitz (1995) and Lumpkin and Speer (2007), from Talley et al. (2011) DPO textbook, Chapter 14

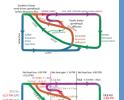
## **Global Overturning Circulation**



ncomplete overturning

IADW-Indian upwelling-Missing Southern Ocean and AABW)

NADW-AABW overturning cells in zonally-average circulation (missing Indian-Pacific)



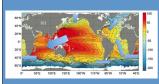
Essential elements of the global overturning circulation

North Atlantic and Antarctic dense water

Upwelling in the Southern Ocean Upwelling in the Indian and Pacific

Formation of AABW from upwelled NADW in the Southern Ocean Warm return path water from upwelled AABW in the Indian and Pacific, including S.O. and warmer return

# Down-gradient global flow



Surface height

paths for upwelled water

From surface drifters Maximenko et al., 2009)

Superimposed informal arrows are the "conveyor belt" plus "colo water path" through Drake Psg. direction at the largest scale



Deep Waters: 2000 dbar

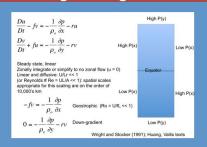
Adjusted steric height (absolute geostrophic flow) (Reid, 1994, 997, 2003). North Atlantic to South Atlantic and North Pacific



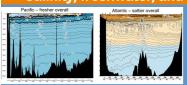
Bottom Waters: 4000 dbar

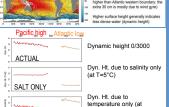
Adjusted steric height (absolute ostrophic flow) (Reid, 1994, Down-gradient direction from Intarctic to all other regions

### Down-gradient global flow



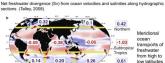
#### Salinity, freshwater, and overturn





S=34.9)

llinity difference counts for the gross



TEMPERATURE ONLY

0.61

Fun calculation: F =  $V\Delta S/S_o$ For F = 0.3 Sv and  $\Delta S$ = 35.2-34.6=0.6, then V = 17 Sv