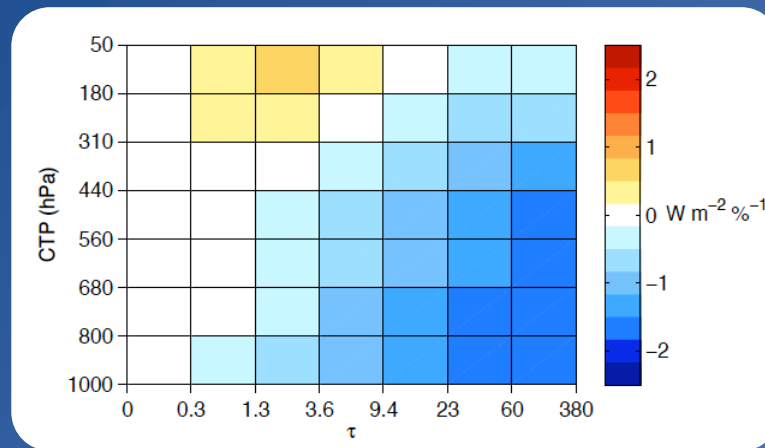


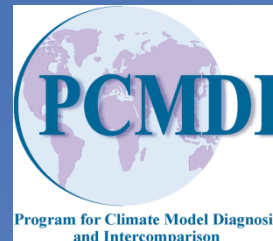
COMPUTING AND PARTITIONING CLOUD FEEDBACKS USING CLOUD PROPERTY HISTOGRAMS



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2. Department of Atmospheric Sciences, University of Washington



Goals

- To provide a **clean and simple** method of computing cloud feedbacks that is highly **informative**
- **Clean:**
 - compute cloud feedback from ISCCP simulator-interpreted cloud changes directly (not inferred)
 - standard definition of “cloud” and radiation code across models
- **Simple:**
 - no need to correct for non-cloud effects
 - no partial radiative perturbation calculations are needed
 - can use monthly mean model output
- **Informative:**
 - can **quantify** the contribution to cloud feedback from **changing amounts of individual cloud types** (high, middle, low) and from **individual processes** (Δ altitude, Δ optical depth, Δ total amount)

Data & Methodology

- Doubled CO₂ equilibrium slab ocean model simulations from 11 GCMs as part of CFMIP1
- ISCCP simulator (Klein & Jakob 1999) run inline during integration
 - Produce distribution of cloud fraction (as function of CTP and τ) that is consistent with how a satellite-borne passive sensor would “view” the model atmosphere
 - Simulated cloud fractions are defined consistently across models

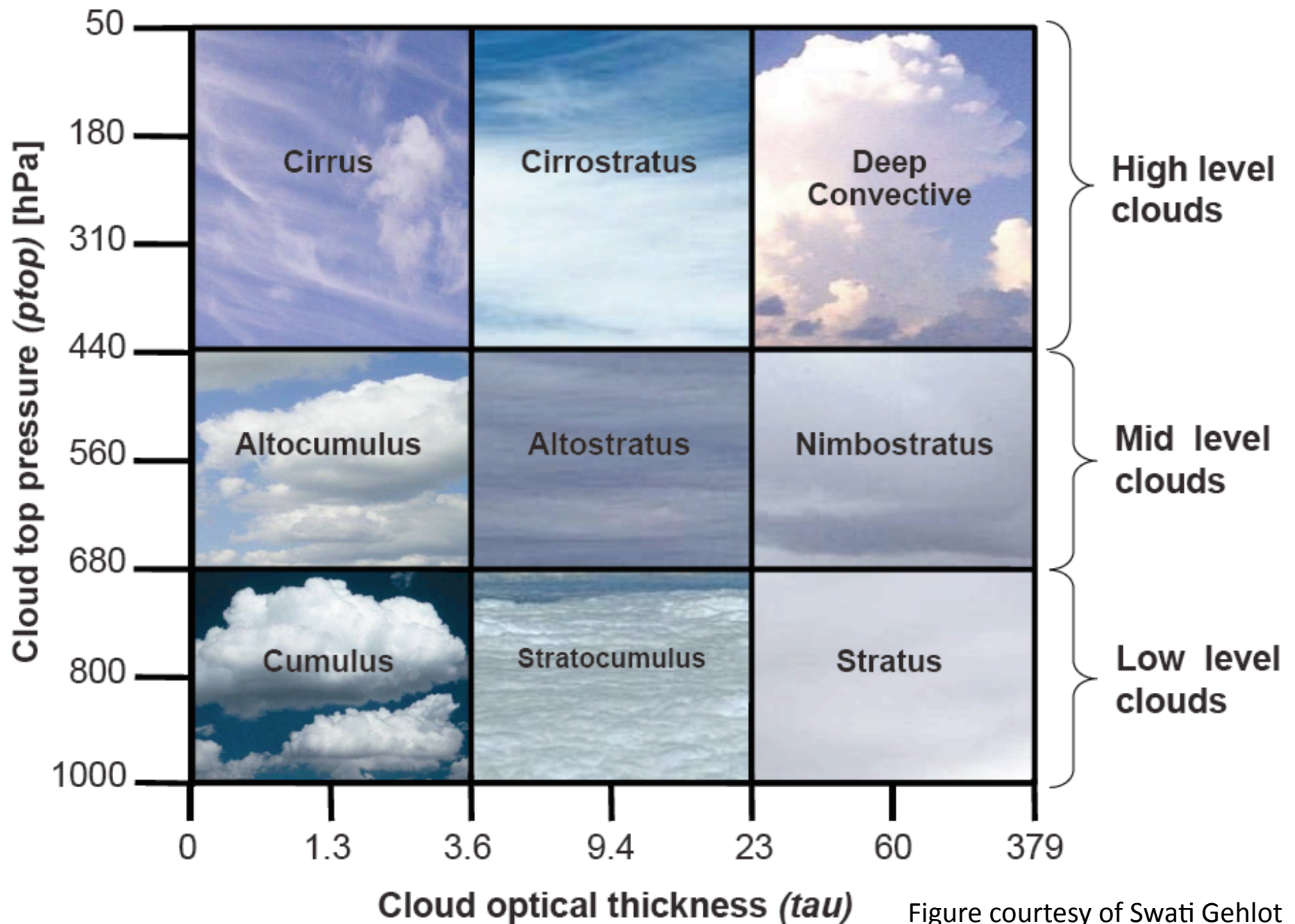


Figure courtesy of Swati Gehlot

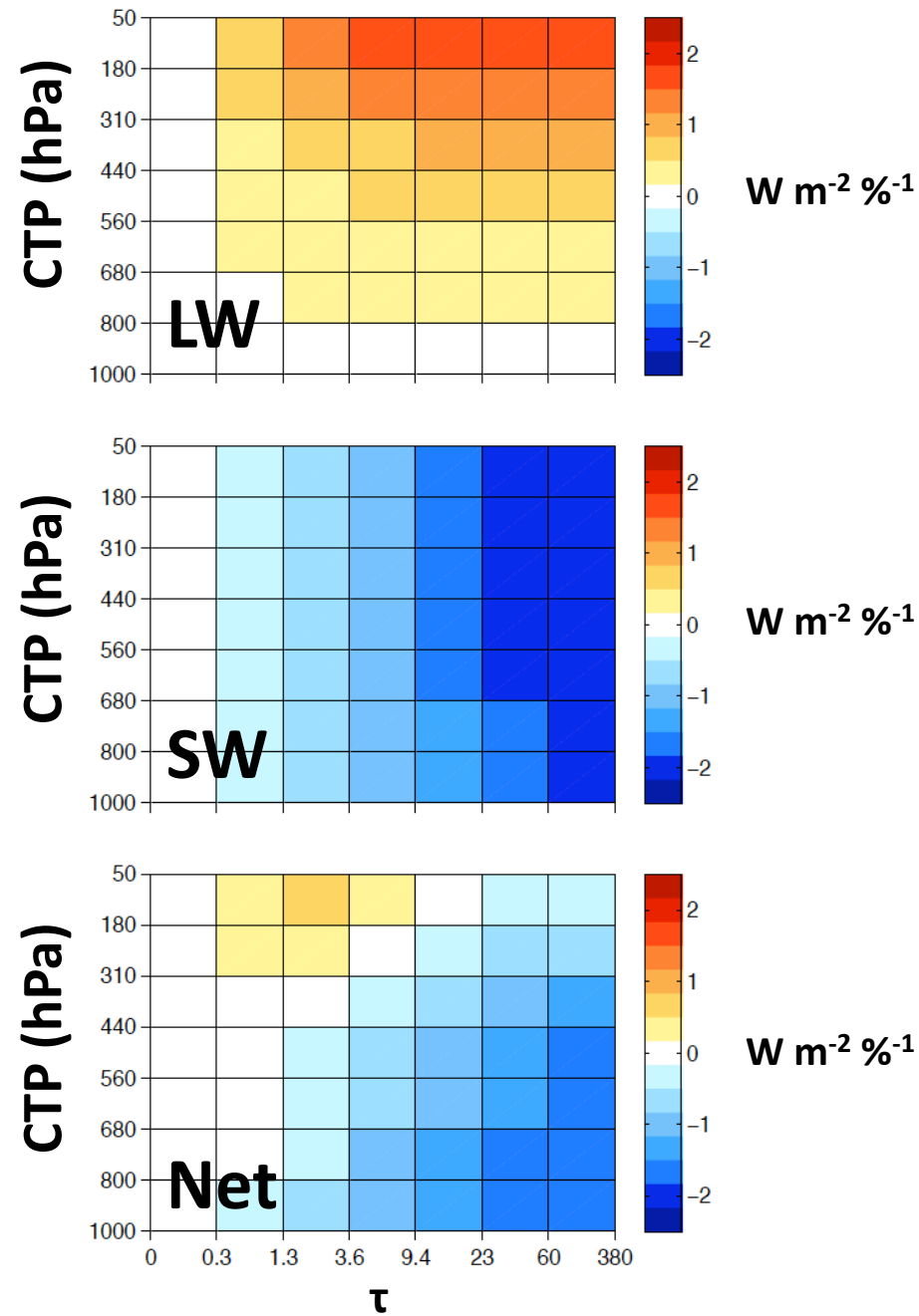
Data & Methodology

- Doubled CO₂ equilibrium slab ocean model simulations from 12 GCMs as part of CFMIP1
- ISCCP simulator run inline during integration
 - Produce distribution of cloud fraction (as function of CTP and τ) that is consistent with how a satellite-borne passive sensor would “view” the model atmosphere
 - Simulated cloud fractions are defined consistently across models
- We compute cloud radiative kernels $\rightarrow$ sensitivity of TOA radiation to cloud fraction changes in each CTP- τ bin
- Cloud feedback = Δ cloud fraction times cloud kernel normalized by ΔT_{sfc}

Recipe for Constructing Cloud Radiative Kernels

- ❑ Input model mean zonal mean T and q profiles to Fu-Liou code
- ❑ Compute clear-sky TOA fluxes
- ❑ Compute overcast-sky fluxes for each CTP and τ bin by setting the LWC / IWC profiles to values appropriate for each cloud type
- ❑ Subtract overcast TOA fluxes in each bin from the clear-sky flux to compute a matrix of overcast sky cloud forcing
- ❑ Divide by 100 to get $\text{W m}^{-2} \%^{-1}$
- ❑ Repeat every calculation for 24 solar zenith angles, all latitudes, 12 months, and 3 surface albedos (0.0, 0.5, 1.0)

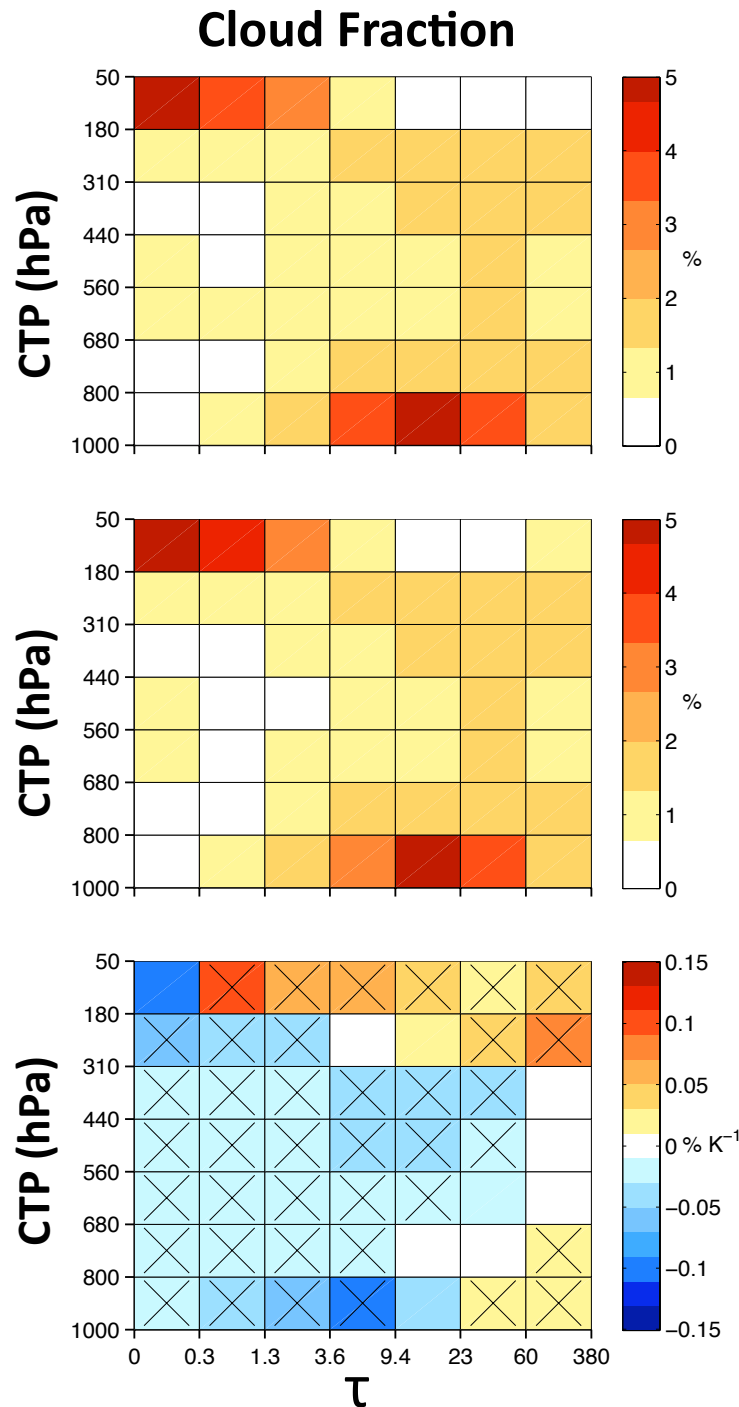
Global Annual Mean Cloud Kernels



1xCO₂
58.1 %

2xCO₂
56.8 %

Change
-0.5 % K⁻¹



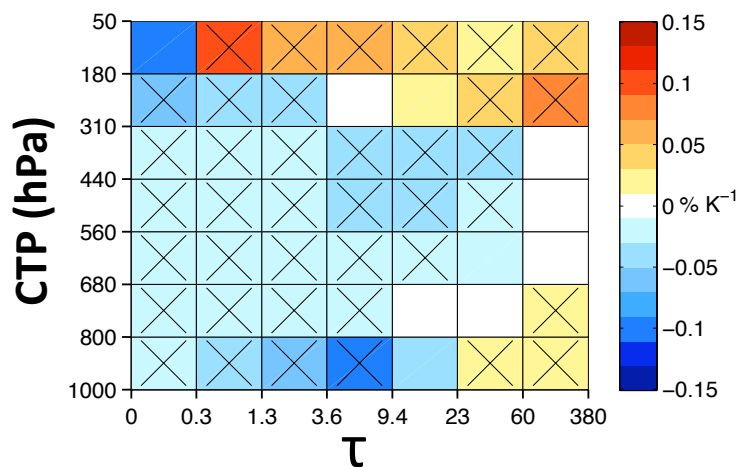
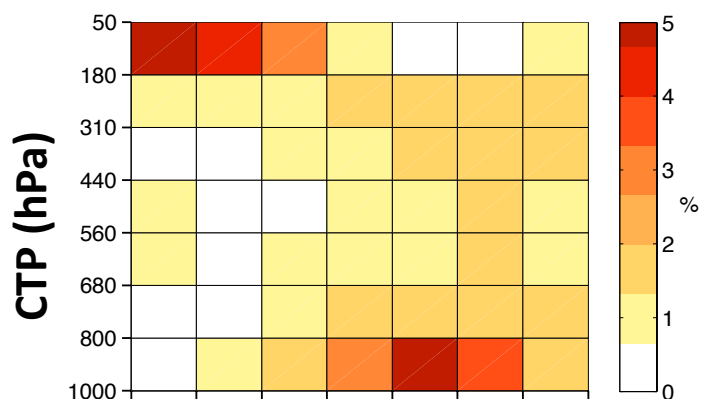
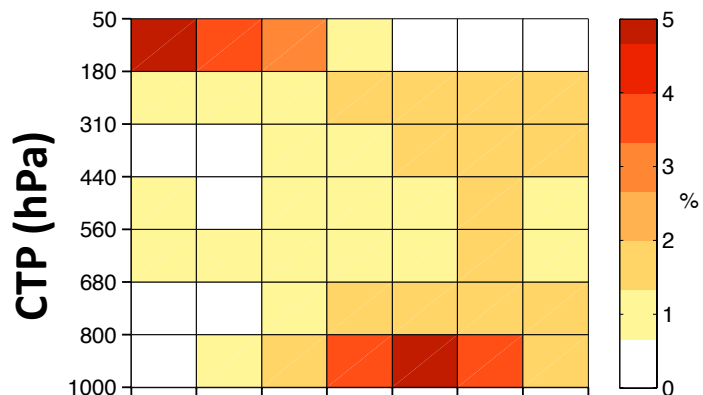
x Cloud Radiative Kernels
at each location and month,
then averaged annually,
globally, and across models...

1xCO₂
58.1 %

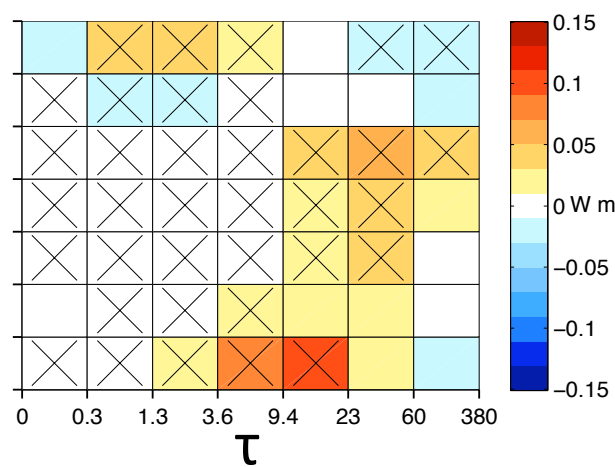
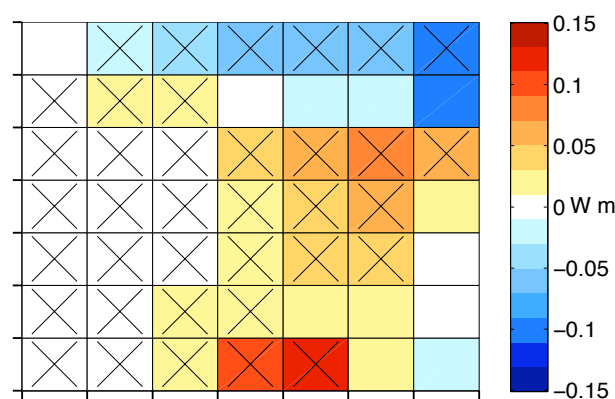
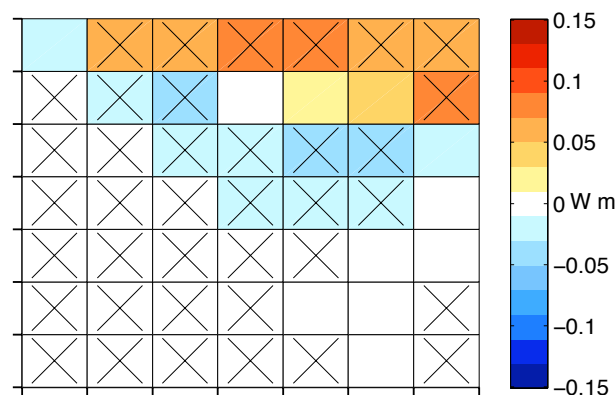
2xCO₂
56.8 %

Change
-0.5 % K⁻¹

Cloud Fraction



Cloud Feedback



LW
0.20 W m⁻² K⁻¹

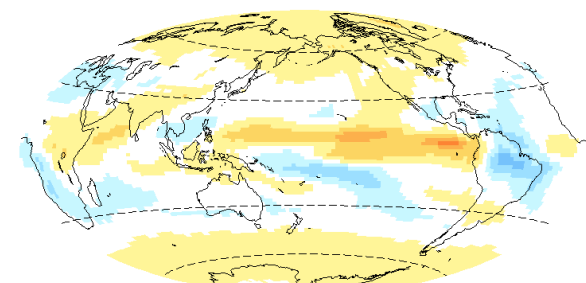
SW
0.37 W m⁻² K⁻¹

Net
0.58 W m⁻² K⁻¹

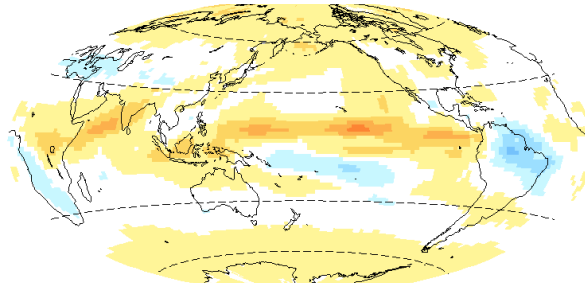
Cloud Kernel

Adjusted Δ CRF

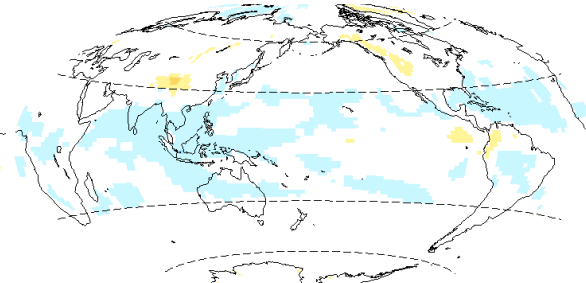
**Kernel minus
Adjusted Δ CRF**



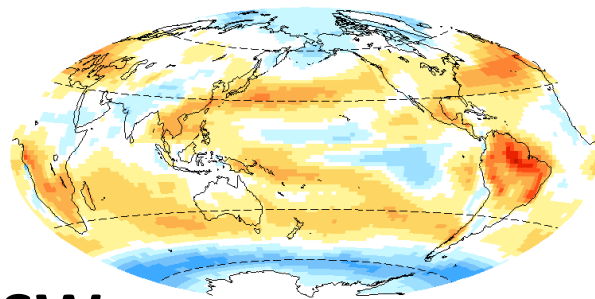
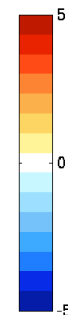
LW $0.21 \text{ W m}^{-2} \text{ K}^{-1}$



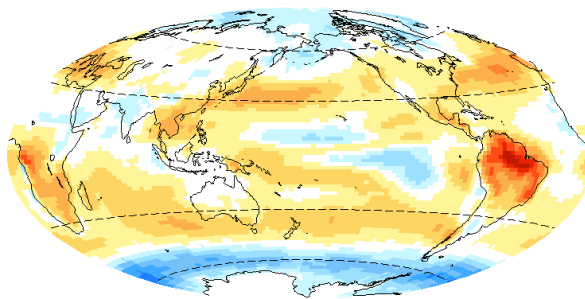
$0.36 \text{ W m}^{-2} \text{ K}^{-1}$



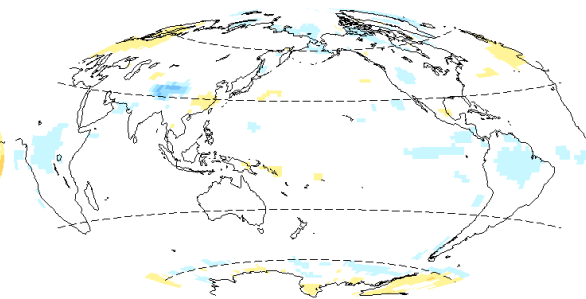
$-0.15 \text{ W m}^{-2} \text{ K}^{-1}$



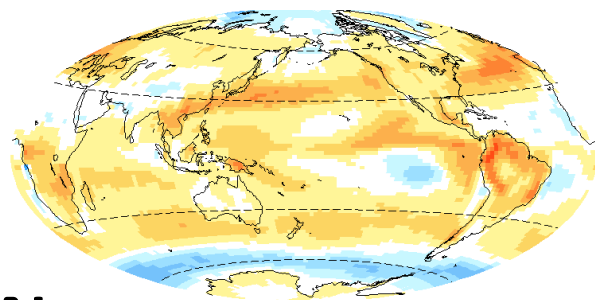
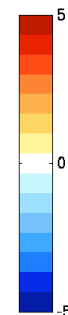
SW $0.35 \text{ W m}^{-2} \text{ K}^{-1}$



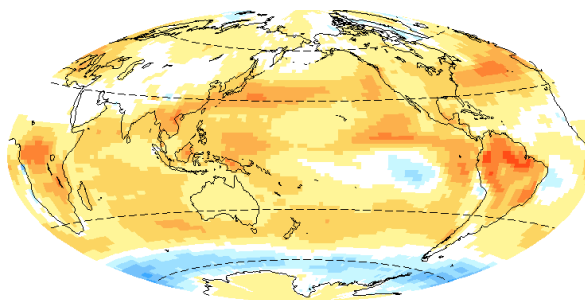
$0.39 \text{ W m}^{-2} \text{ K}^{-1}$



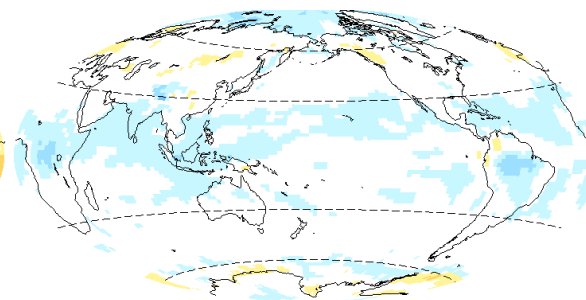
$-0.04 \text{ W m}^{-2} \text{ K}^{-1}$



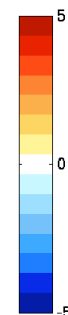
Net $0.56 \text{ W m}^{-2} \text{ K}^{-1}$

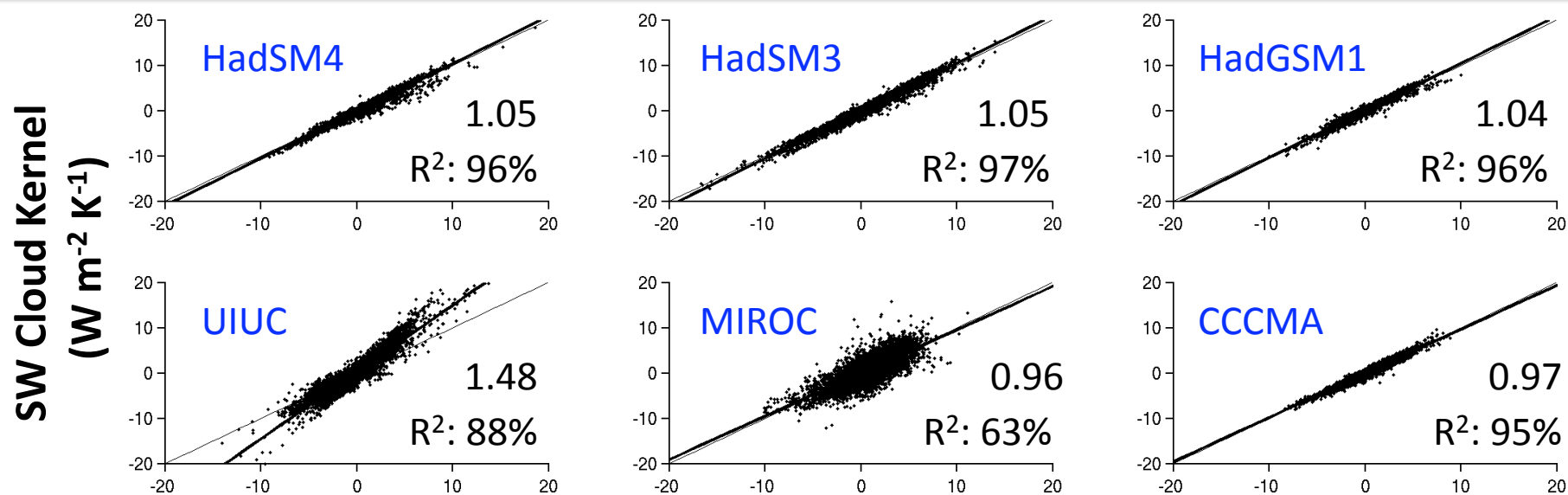
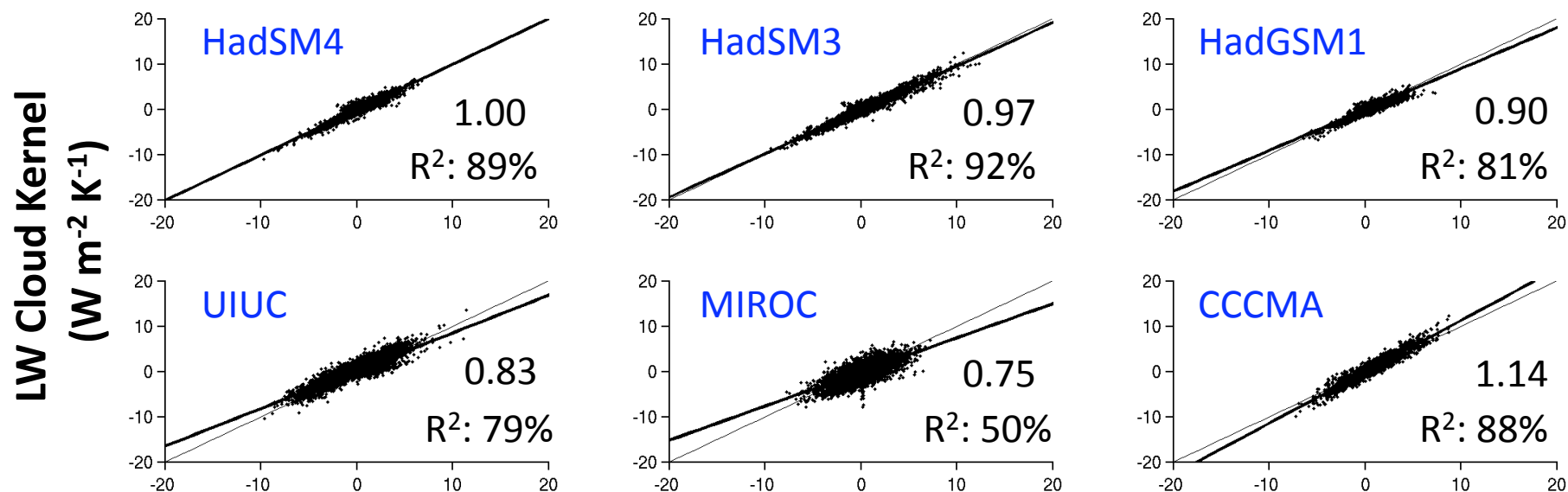


$0.75 \text{ W m}^{-2} \text{ K}^{-1}$

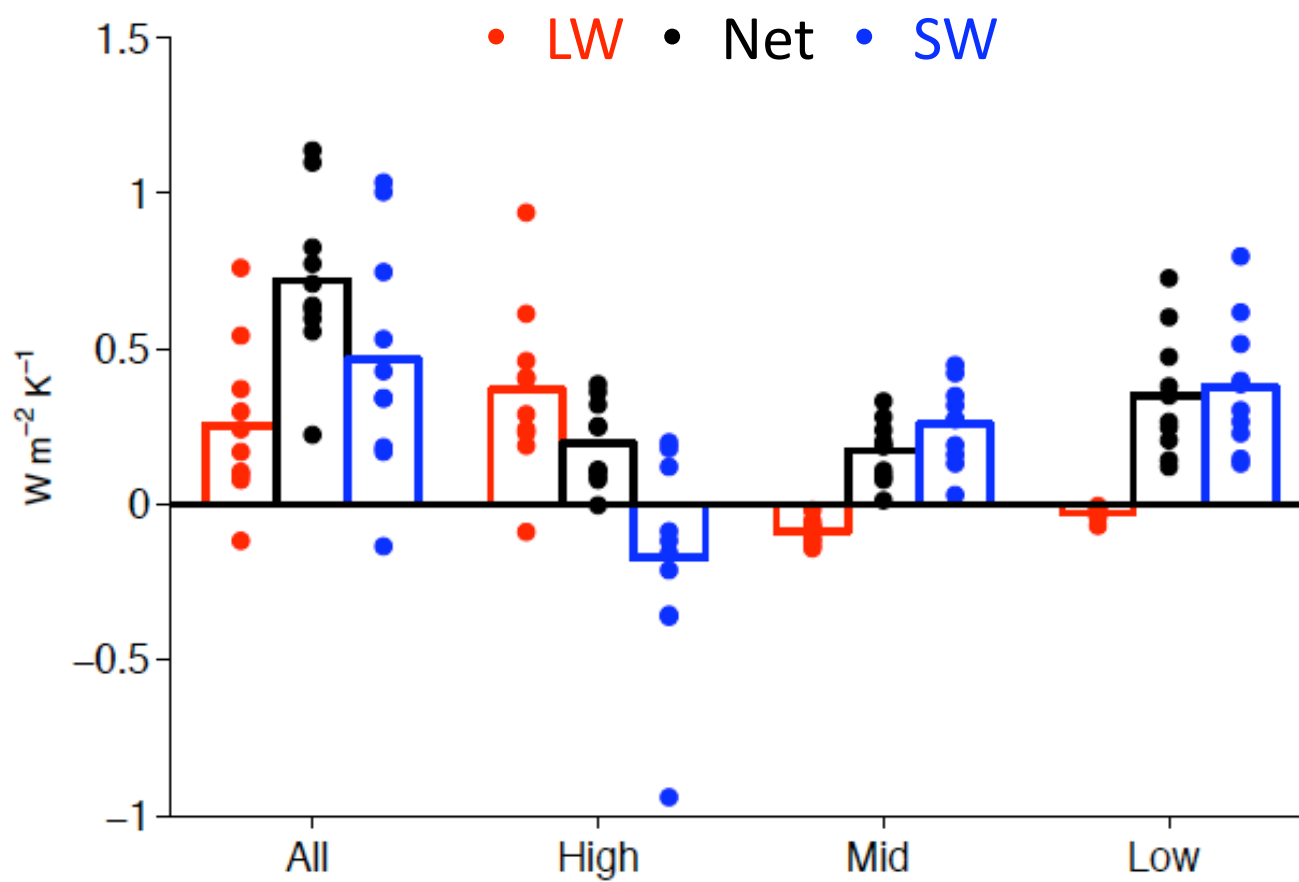


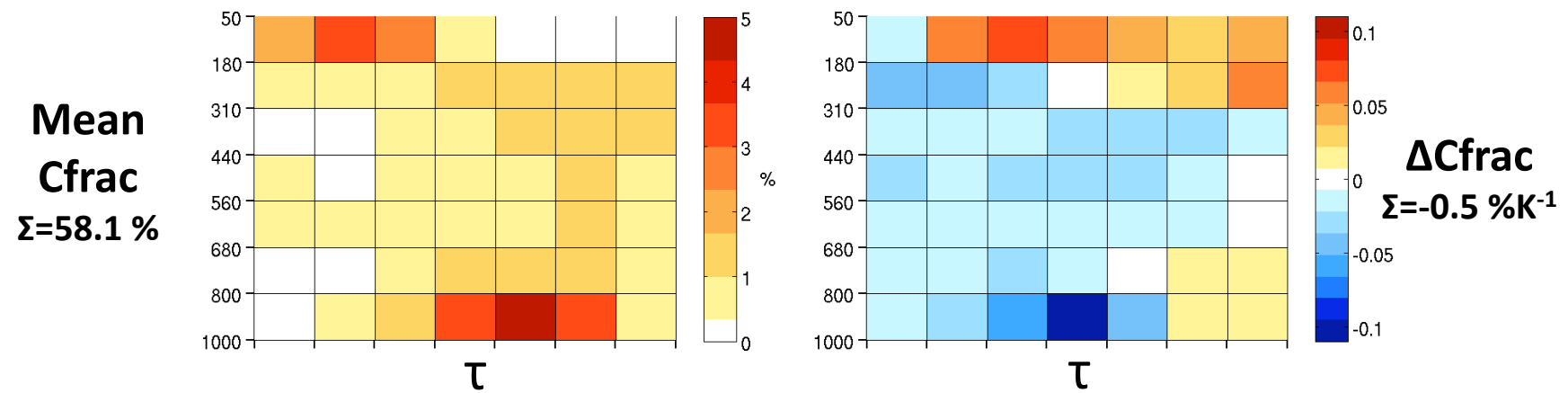
$-0.20 \text{ W m}^{-2} \text{ K}^{-1}$





Adjusted ΔCRF ($\text{W m}^{-2} \text{K}^{-1}$)





- Decompose the cloud changes into

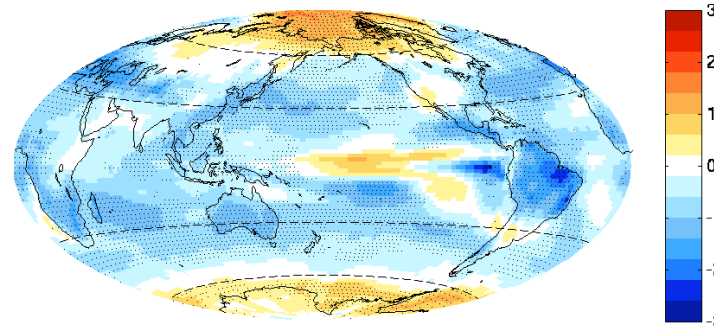
Δ AMOUNT

Δ ALTITUDE

Δ OPTICAL DEPTH

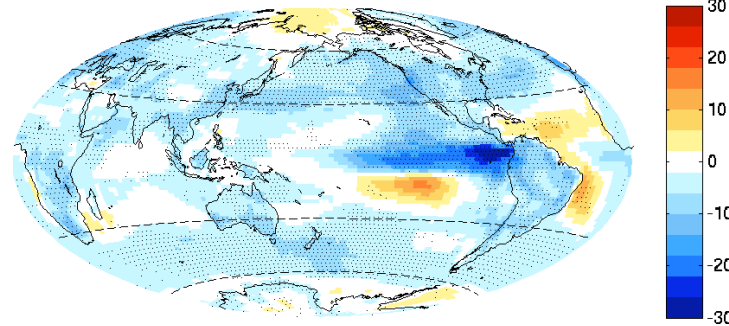
Clouds go away

Δ Total Cloud Fraction: $-0.44 \text{ \% K}^{-1}$



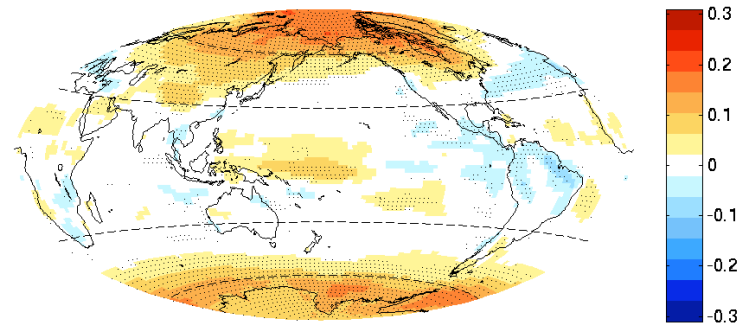
Clouds rise

Δ CTP: -3.68 hPa K^{-1}



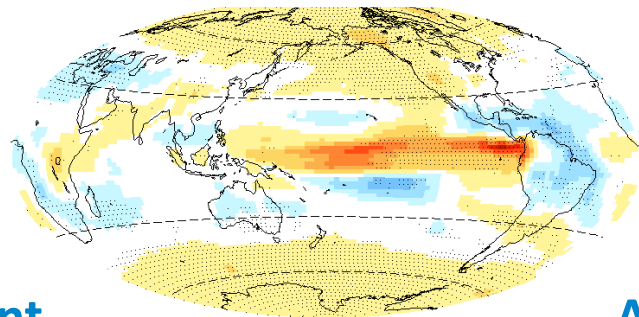
Cold clouds get thicker

$\Delta \ln(\tau)$: 0.03 K^{-1}



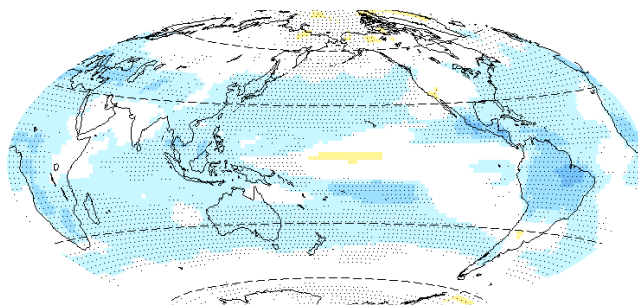
LW Cloud Feedback

$0.21 \text{ W m}^{-2} \text{ K}^{-1}$



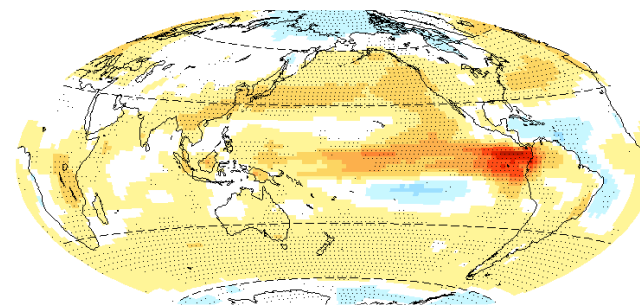
Amount

$-0.29 \text{ W m}^{-2} \text{ K}^{-1}$



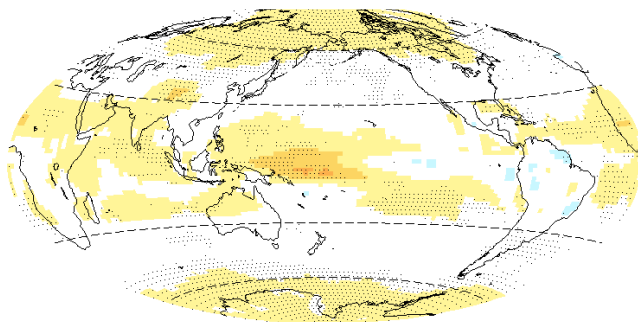
Altitude

$0.39 \text{ W m}^{-2} \text{ K}^{-1}$



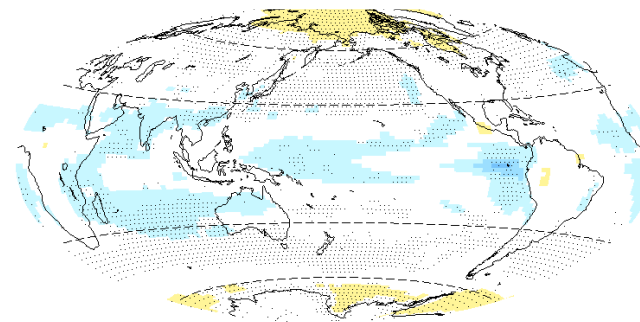
Optical Depth

$0.22 \text{ W m}^{-2} \text{ K}^{-1}$



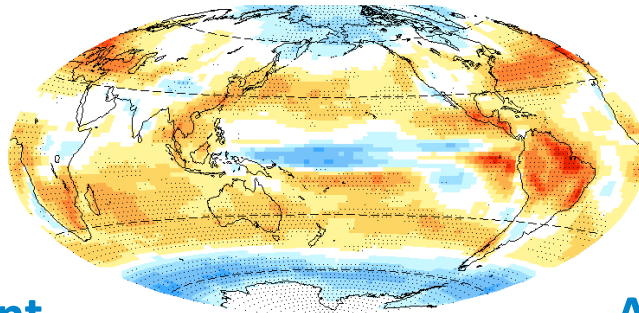
Residual

$-0.11 \text{ W m}^{-2} \text{ K}^{-1}$



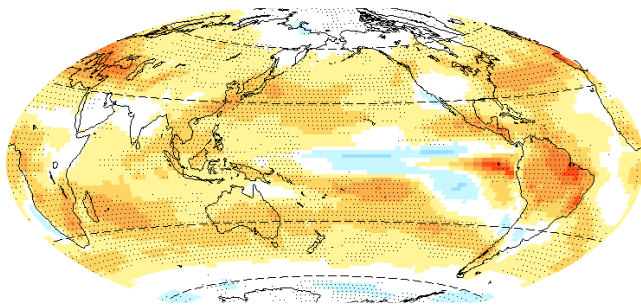
SW Cloud Feedback

$0.37 \text{ W m}^{-2} \text{ K}^{-1}$



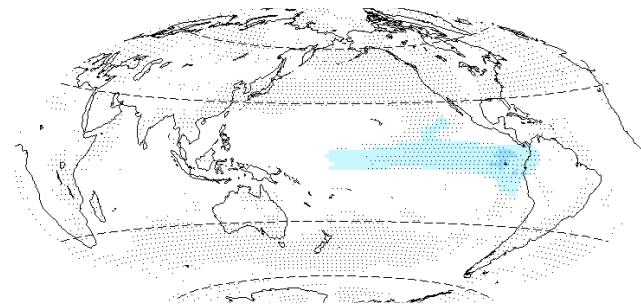
Amount

$0.57 \text{ W m}^{-2} \text{ K}^{-1}$



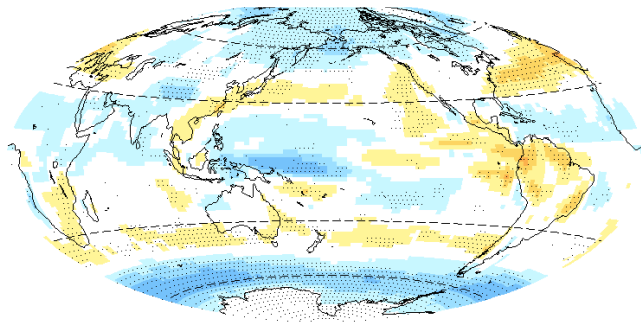
Altitude

$-0.07 \text{ W m}^{-2} \text{ K}^{-1}$



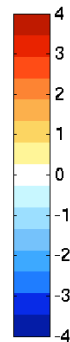
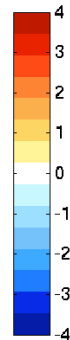
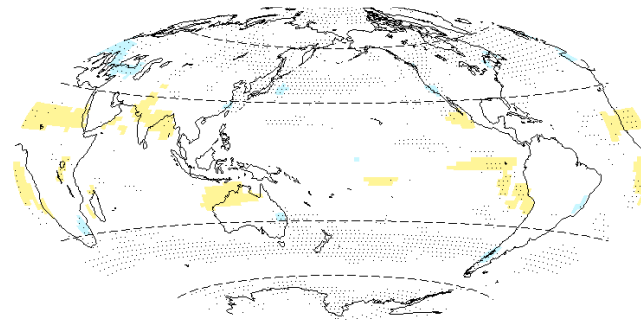
Optical Depth

$-0.14 \text{ W m}^{-2} \text{ K}^{-1}$



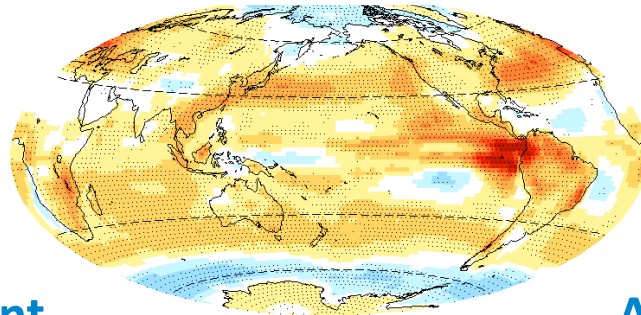
Residual

$0.01 \text{ W m}^{-2} \text{ K}^{-1}$



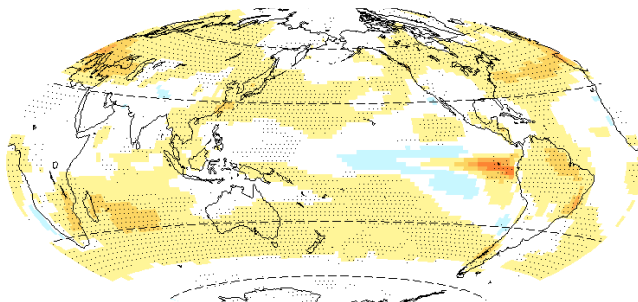
Net Cloud Feedback

$0.57 \text{ W m}^{-2} \text{ K}^{-1}$



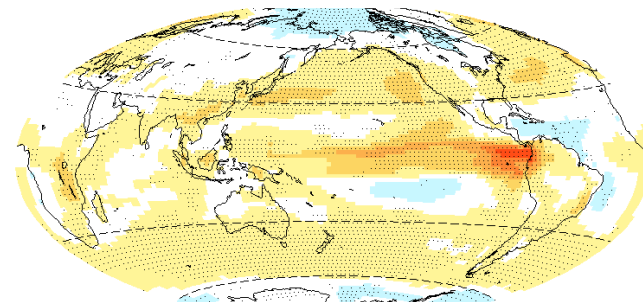
Amount

$0.27 \text{ W m}^{-2} \text{ K}^{-1}$



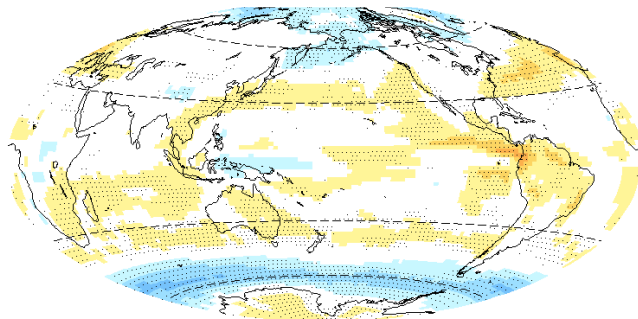
Altitude

$0.33 \text{ W m}^{-2} \text{ K}^{-1}$



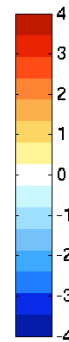
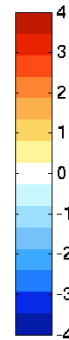
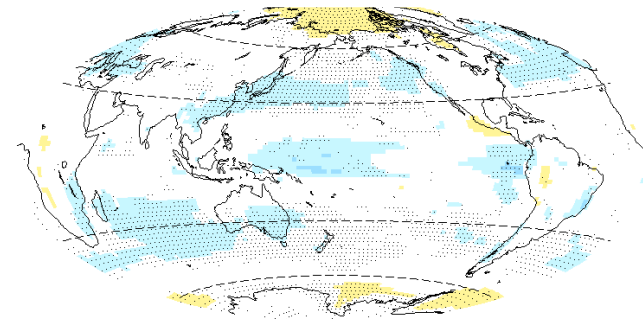
Optical Depth

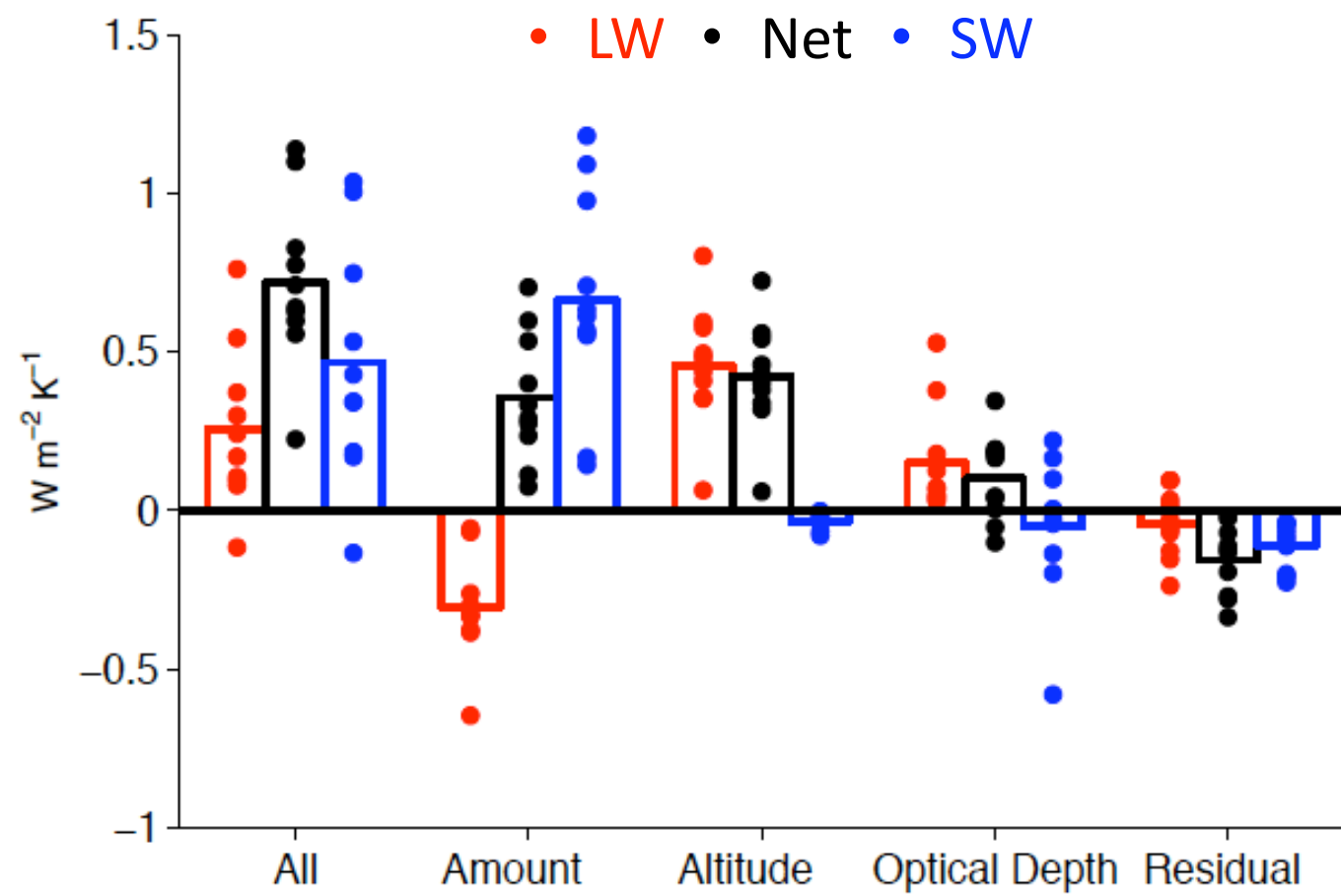
$0.07 \text{ W m}^{-2} \text{ K}^{-1}$



Residual

$-0.10 \text{ W m}^{-2} \text{ K}^{-1}$





Conclusions

- Cloud radiative kernels allow computation of cloud feedback directly from cloud fraction histograms produced by ISCCP simulator
- Feedbacks computed with cloud kernels compare very well with those computed by adjusting the change in cloud forcing
- More than half of the global mean net cloud feedback can be attributed to the combined response of middle- and high-level clouds
- High cloud changes induce a wider range of LW and SW cloud feedbacks across models than do low clouds
- Increasing cloud top altitude is dominant contributor to the positive global mean LW and net cloud feedbacks (positive in every model)
- Decreasing total cloud fraction is dominant contributor to global mean positive SW cloud feedback (positive in every model)
- Large negative net cloud feedback at high latitudes is caused by increased optical depth, not increased cloud amount

Thank you!

Drafts of papers: Google “Mark Zelinka”

- This work was performed under the auspices of the U.S. Department of Energy by Lawrence Livermore National Laboratory under Contract DE-AC52-07NA27344 and was supported by the Regional and Global Climate Program of the Office of Science at the U. S. Department of Energy and by NASA Grant NNX09AH73G at the University of Washington.

IM release # LLNL-ABS-474235

Conclusions (1 of 2)

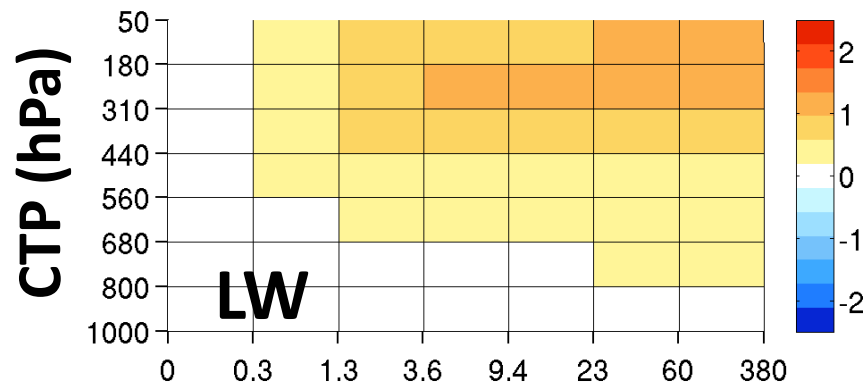
- Cloud radiative kernels allow computation of cloud feedback directly from cloud property histograms generated by ISCCP simulator
 - Standard radiative transfer and definition of “cloud” across models
 - No cloud masking adjustments necessary
 - Simple calculation (multiply two matrices) on monthly mean output
 - Can quantify contribution to feedback from individual cloud types / processes
- Feedbacks computed with cloud kernels compare very well with those computed by adjusting the change in cloud forcing
- Ensemble (10 model) mean results:
 - Low cloud changes are dominant contributor to SW cloud feedback
 - High cloud changes are dominant contributor to LW cloud feedback
 - More than half of the global mean net cloud feedback can be attributed to the combined response of middle- and high-level clouds
 - High cloud changes induce wider range of LW and SW cloud feedbacks across models than do low clouds

Conclusions (2 of 2)

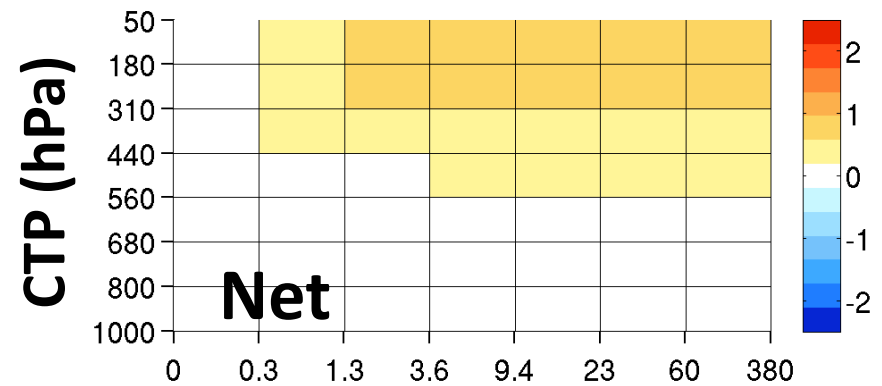
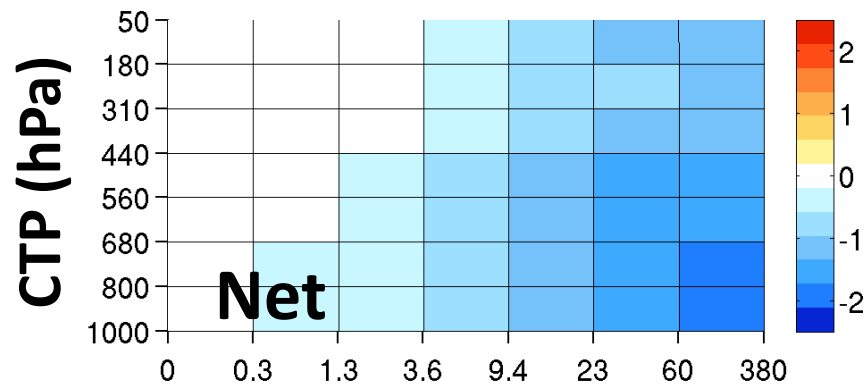
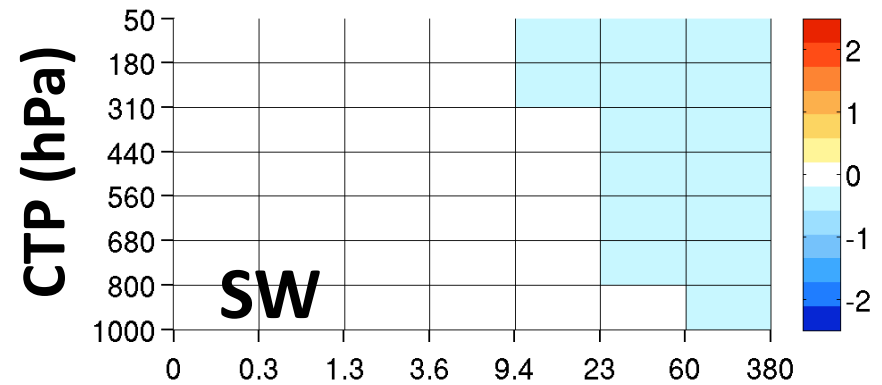
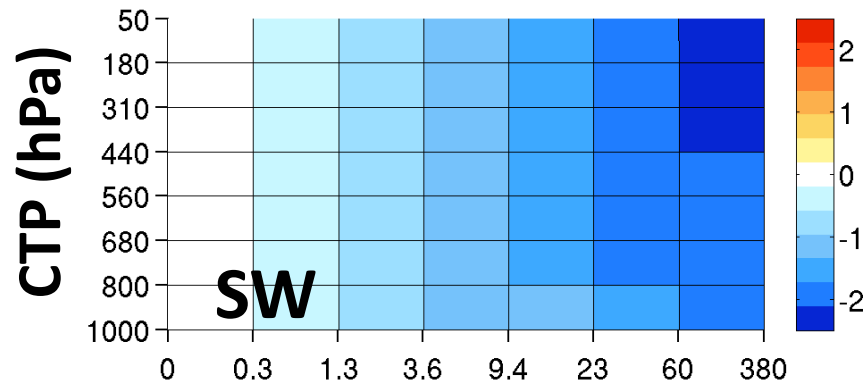
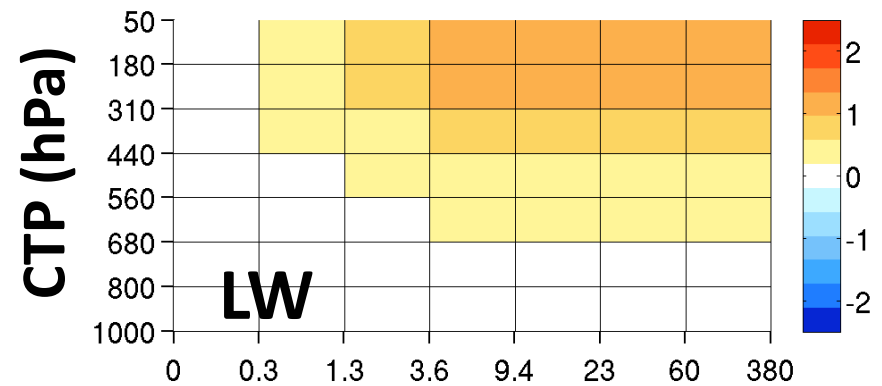
- **Increasing cloud top altitude** is dominant contributor to the **positive global mean LW and net cloud feedbacks (positive in every model)**
- **Decreasing total cloud fraction** is dominant contributor to global mean **positive SW cloud feedback (positive in every model)**
 - Inter-model **spread is greater** than for any other feedback component
 - Overall cloud amount reductions have **2x as large** an impact on SW as on LW fluxes
- **Large negative net cloud feedback at high latitudes** is caused by **increased optical depth**, not increased cloud amount
 - Results from **increased cloud water content and phase changes from ice to liquid**
- This work was performed under the auspices of the U.S. Department of Energy by Lawrence Livermore National Laboratory under Contract DE-AC52-07NA27344 and was supported by the Regional and Global Climate Program of the Office of Science at the U. S. Department of Energy and by NASA Grant NNX09AH73G at the University of Washington.

Extra Slides

50S-90S: DJF



50S-90S: JJA

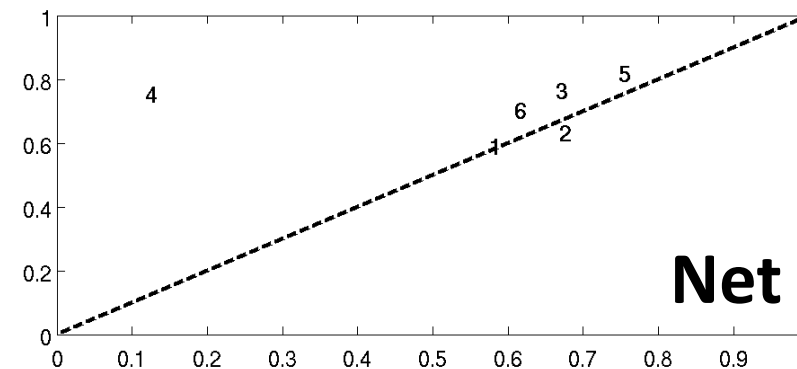
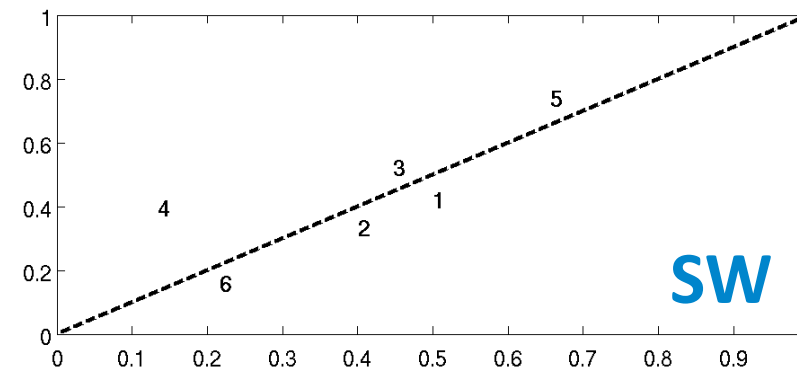
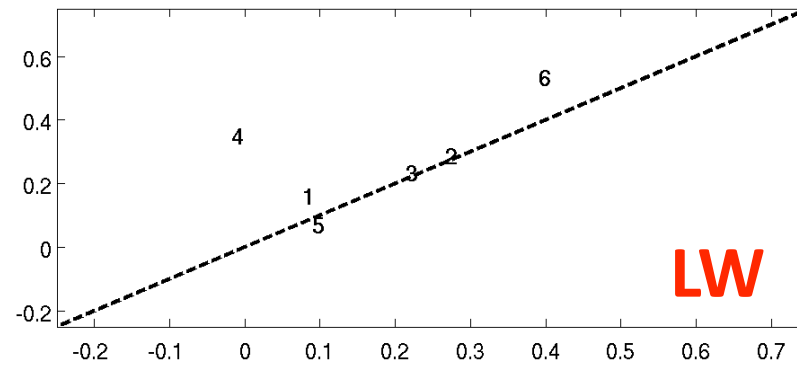


τ

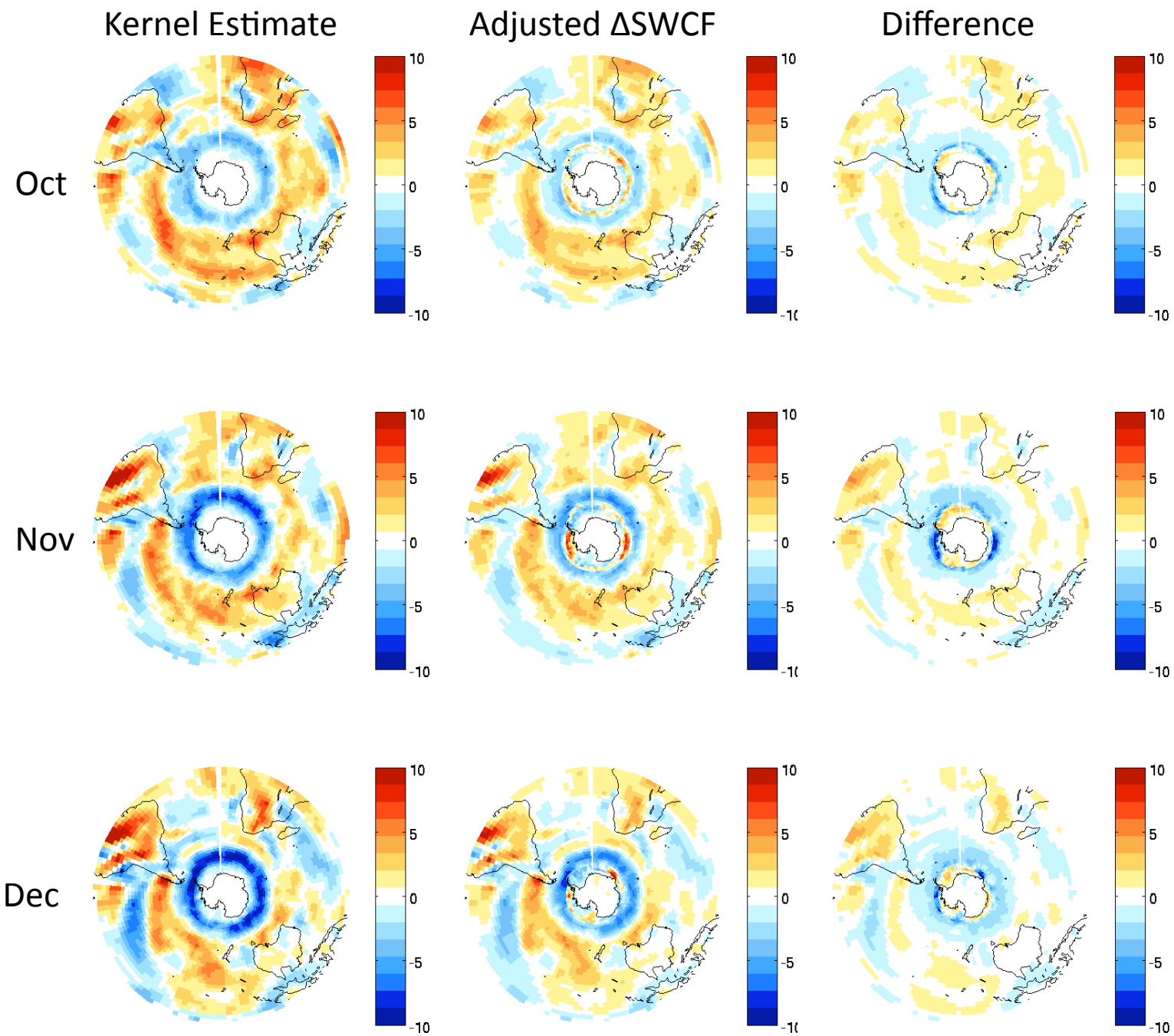
τ

1. ukmo_hadsm4
2. ukmo_hadsm3
3. ukmo_hadgsm1
4. uiuc
5. miroc_losens
6. cccma_agcm4.0

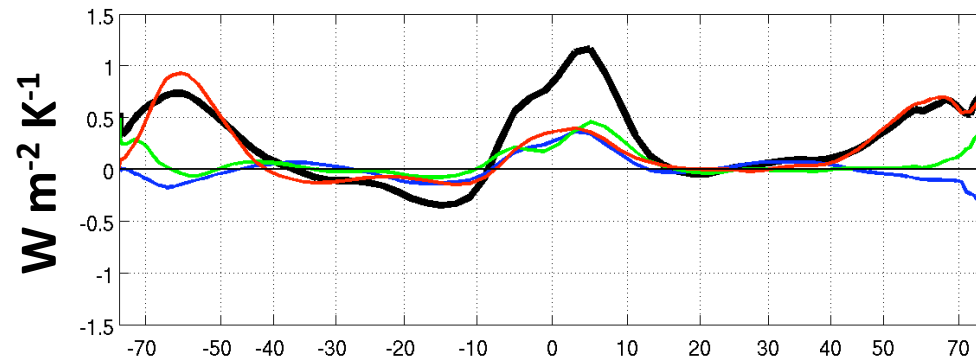
Cloud Kernel Estimate



"Truth"



LW



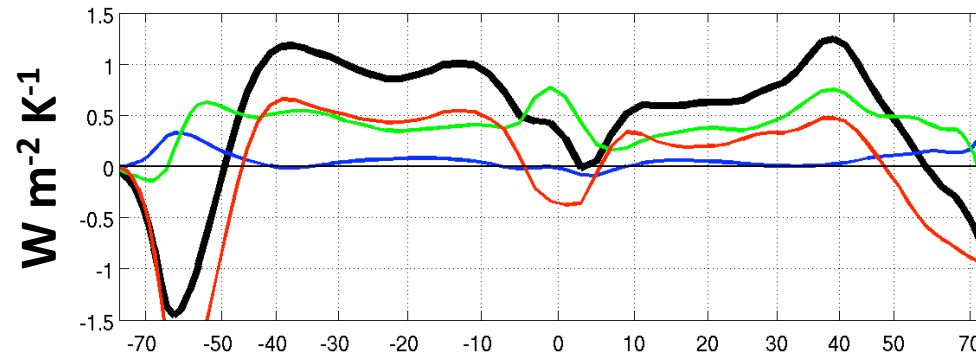
Total: 0.26

Thin: 0.01

Medium: 0.06

Thick: 0.19

SW



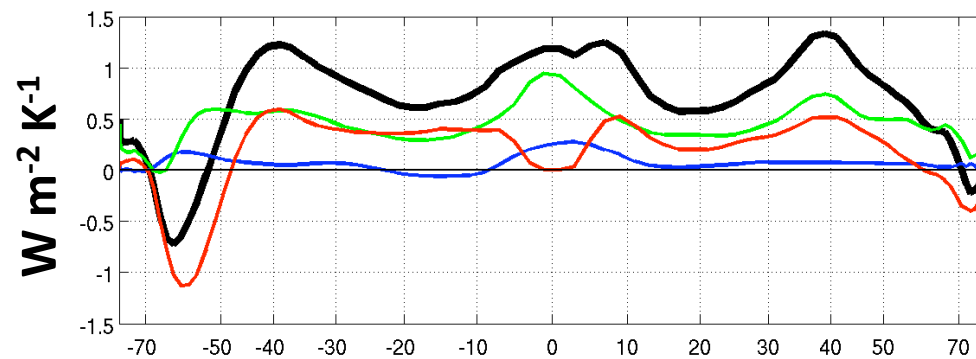
Total: 0.46

Thin: 0.06

Medium: 0.40

Thick: 0.00

Net



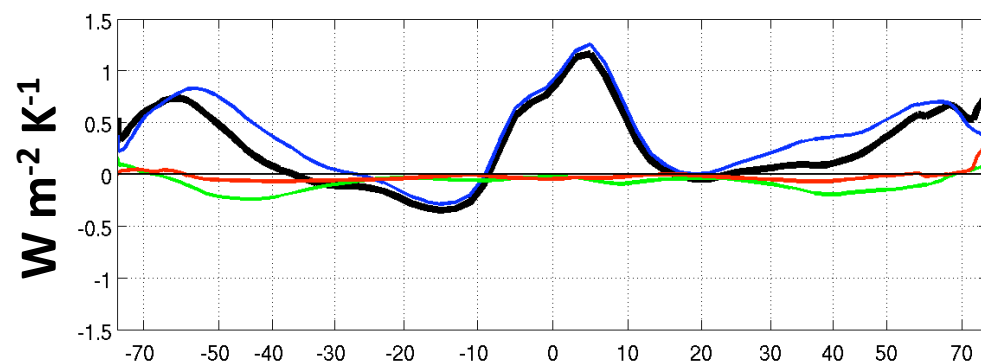
Total: 0.72

Thin: 0.07

Medium: 0.46

Thick: 0.19

LW



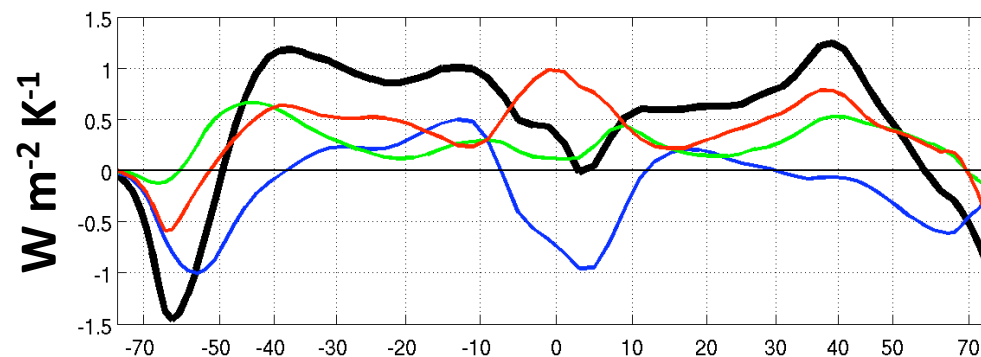
Total: 0.26

High: 0.37

Middle: -0.08

Low: -0.03

SW



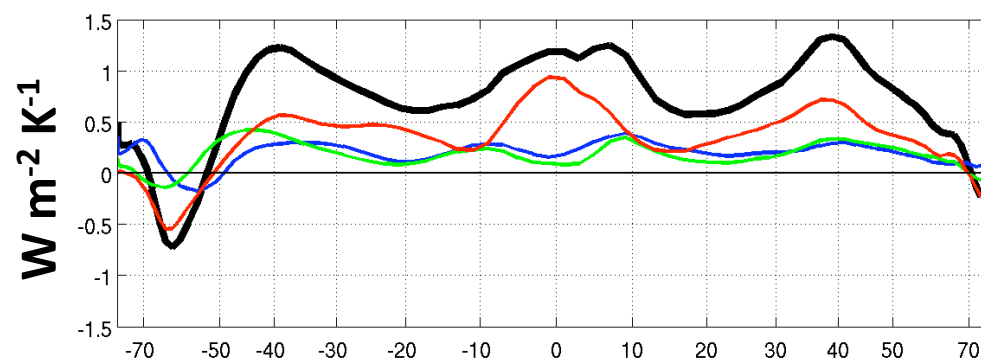
Total: 0.46

High: -0.17

Middle: 0.26

Low: 0.38

Net



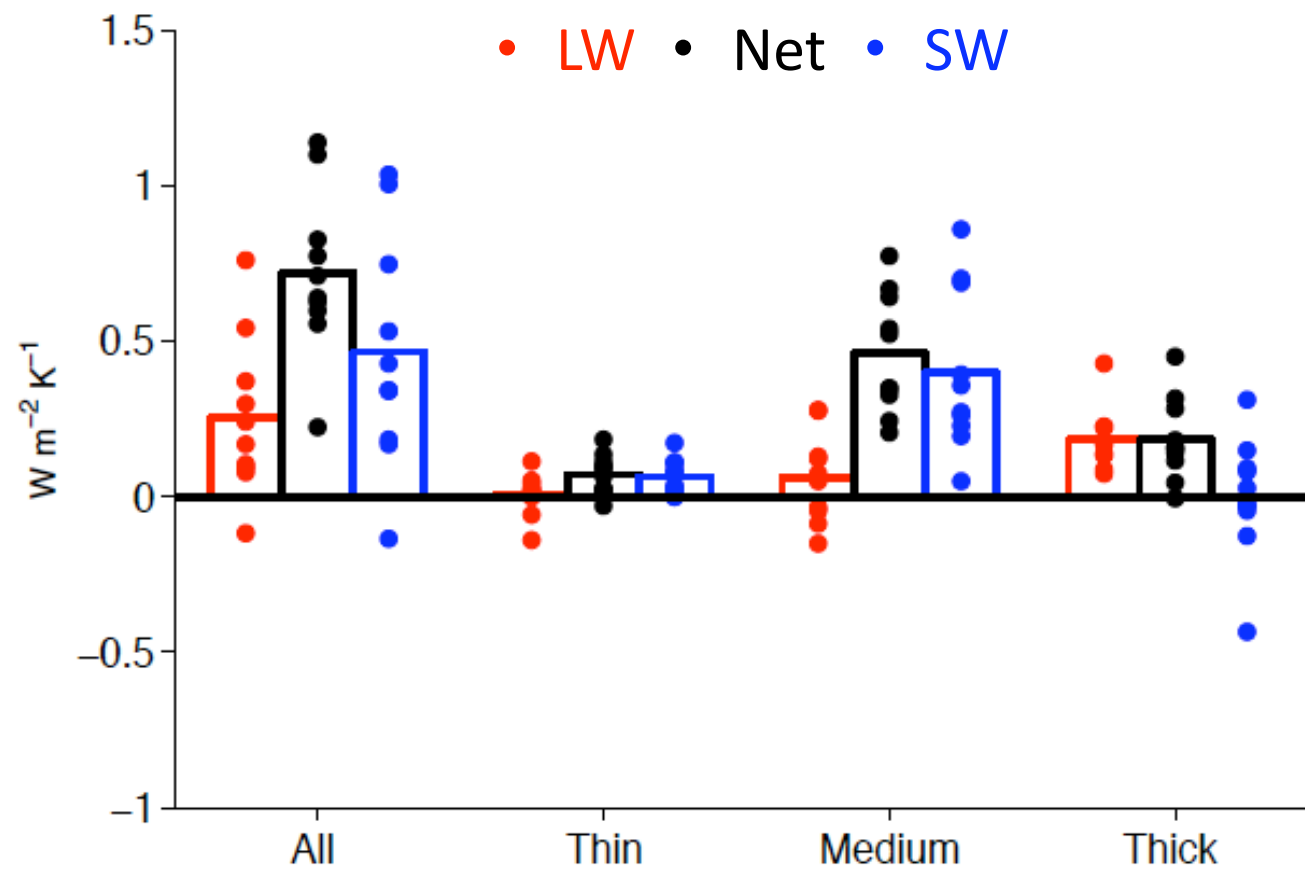
Total: 0.72

High: 0.20

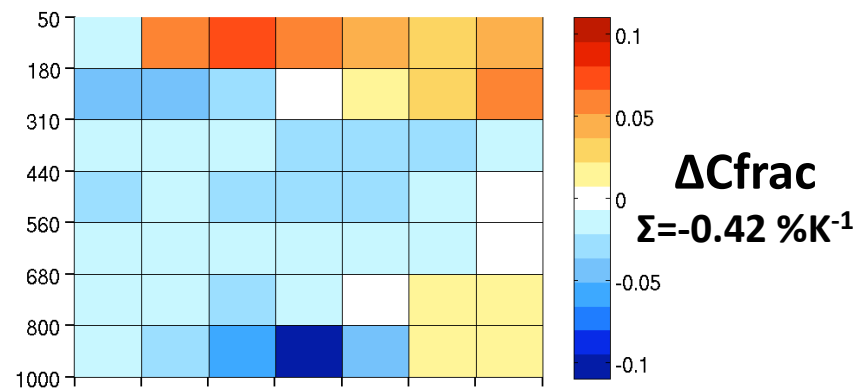
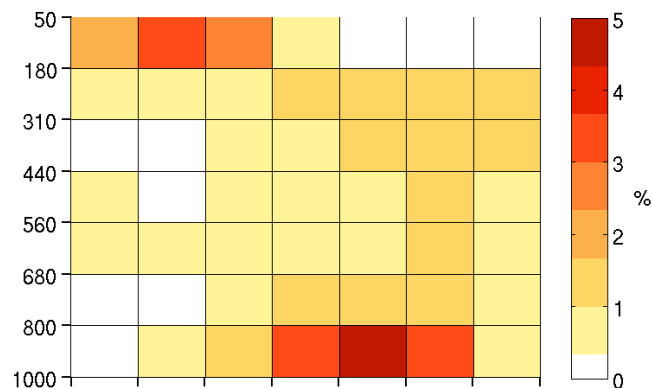
Middle: 0.17

Low: 0.35

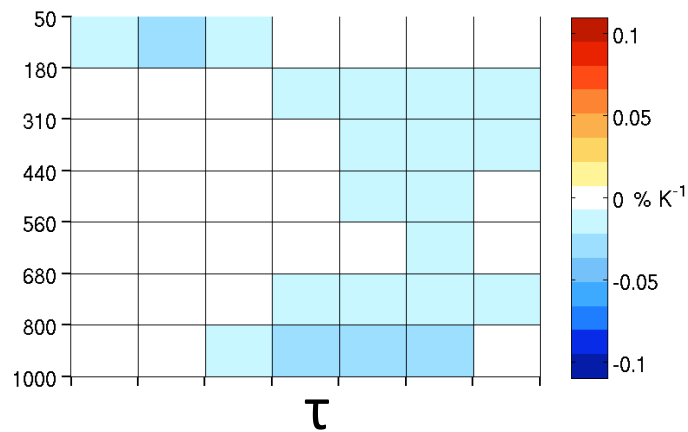
Latitude



**Mean
Cfrac**
 $\Sigma=52.5\%$

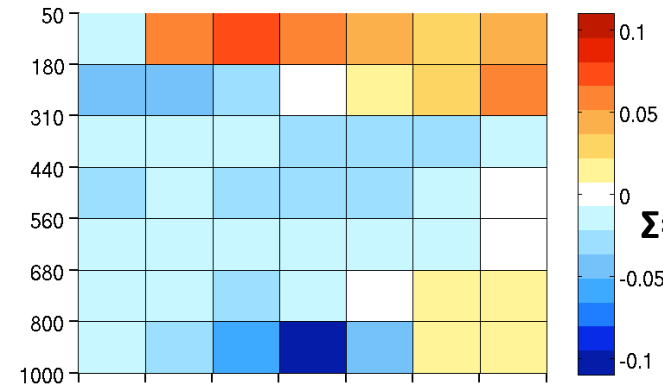
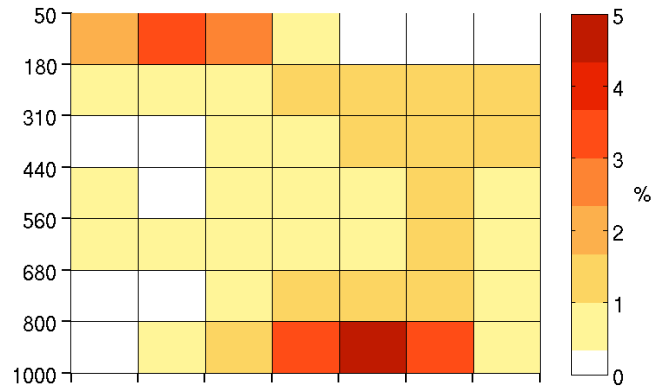


Δ Amount
 $\Sigma=-0.42\%K^{-1}$



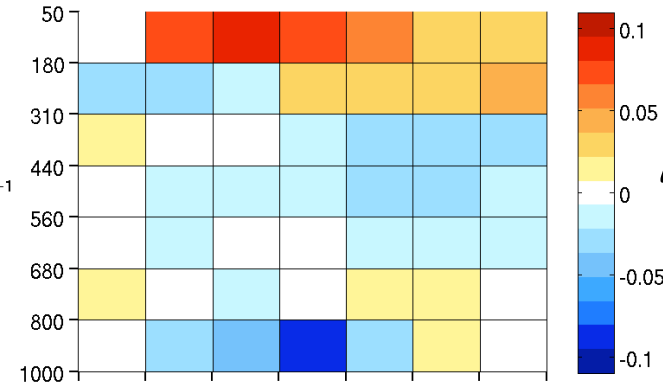
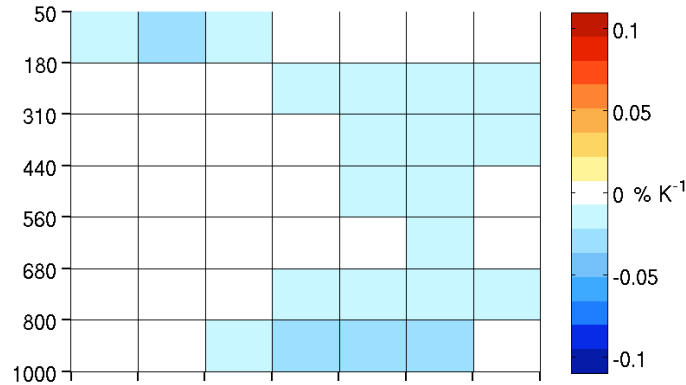
Δ AMOUNT = cloud fraction altered
in proportion to amount in 1xCO₂
histogram; no change in vertical or
optical depth distribution

**Mean
Cfrac
 $\Sigma=52.5\%$**



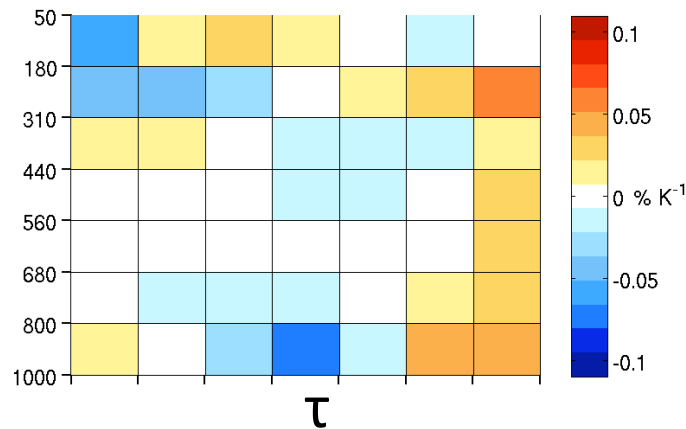
**$\Delta Cfrac$
 $\Sigma=-0.42\%K^{-1}$**

**$\Delta Amount$
 $\Sigma=-0.42\%K^{-1}$**



**$\Delta Altitude$
 $\Sigma=0\%K^{-1}$**

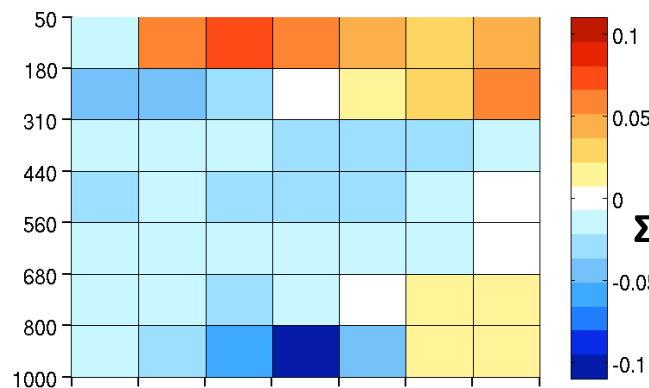
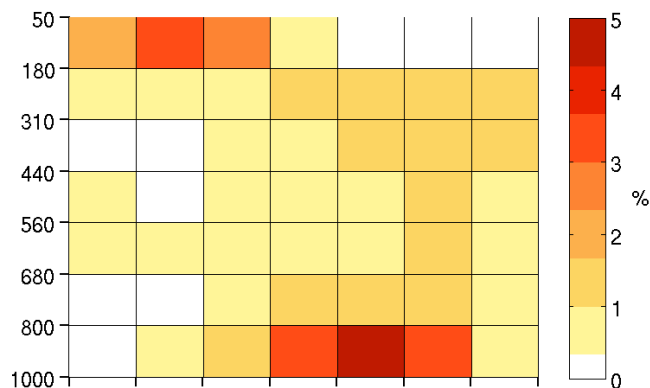
**$\Delta \tau$
 $\Sigma=0\%K^{-1}$**



$\Delta ALTITUDE$ = anomalous vertical
distribution within each τ bin

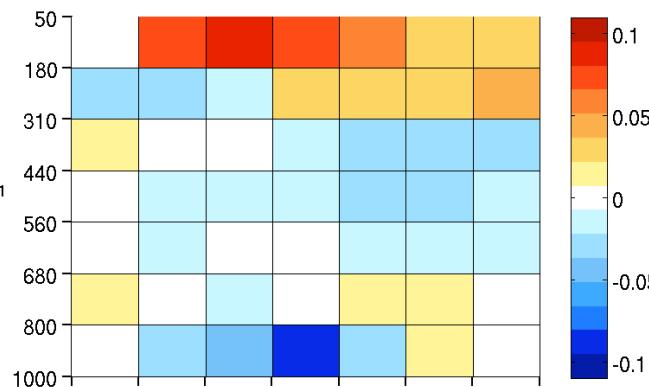
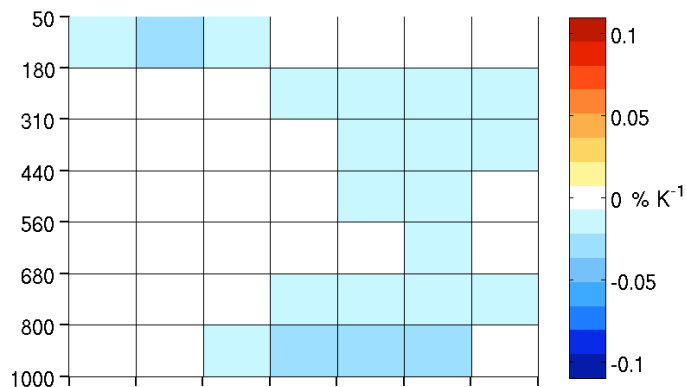
$\Delta OPTICAL DEPTH$ = anomalous optical
depth distribution within each CTP bin

**Mean
Cfrac
 $\Sigma=52.5\%$**



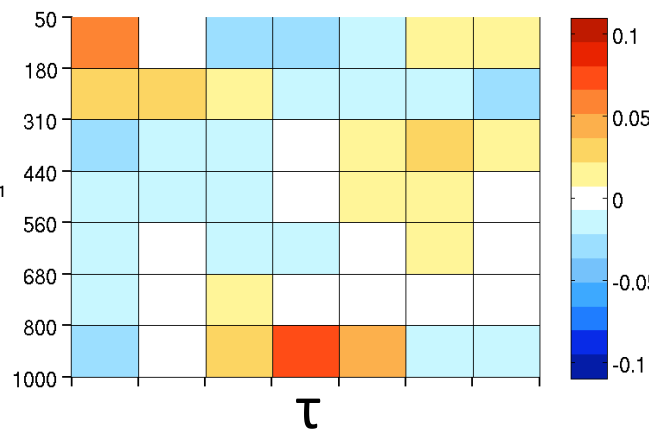
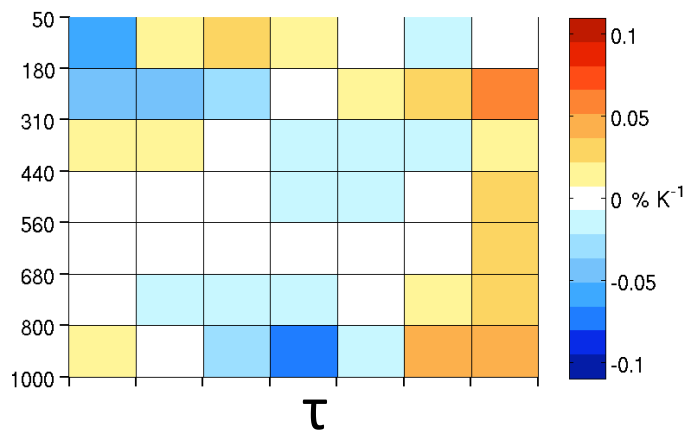
**$\Delta Cfrac$
 $\Sigma=-0.42\%K^{-1}$**

**$\Delta Amount$
 $\Sigma=-0.42\%K^{-1}$**



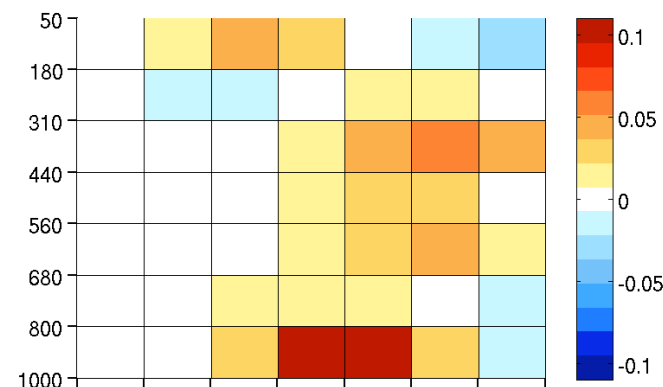
**$\Delta Altitude$
 $\Sigma=0\%K^{-1}$**

**$\Delta \tau$
 $\Sigma=0\%K^{-1}$**



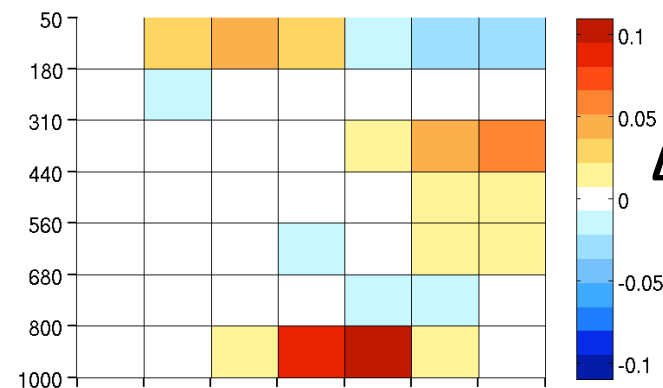
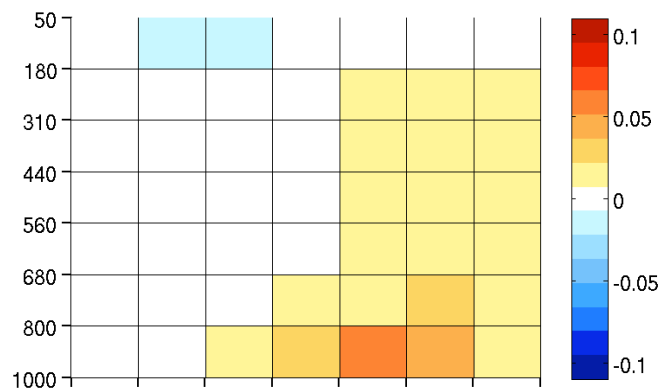
**Residual
 $\Sigma=0\%K^{-1}$**

Net Cloud Feedback Components (W m⁻² K⁻¹)



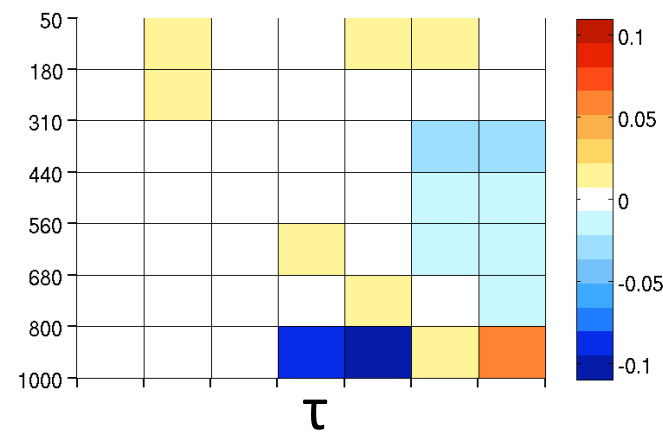
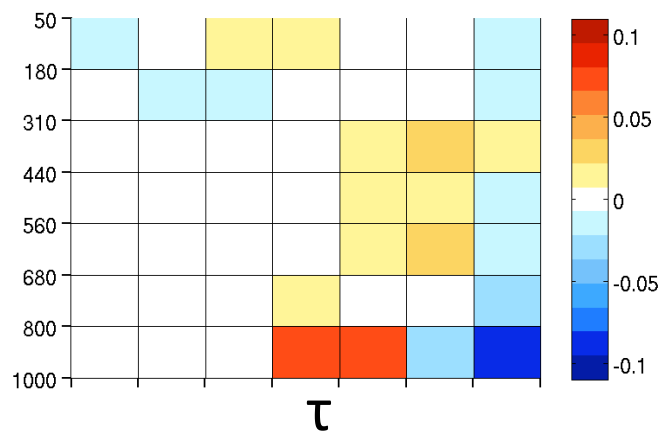
Total
 $\Sigma = -0.72$

ΔAmount
 $\Sigma = -0.36$

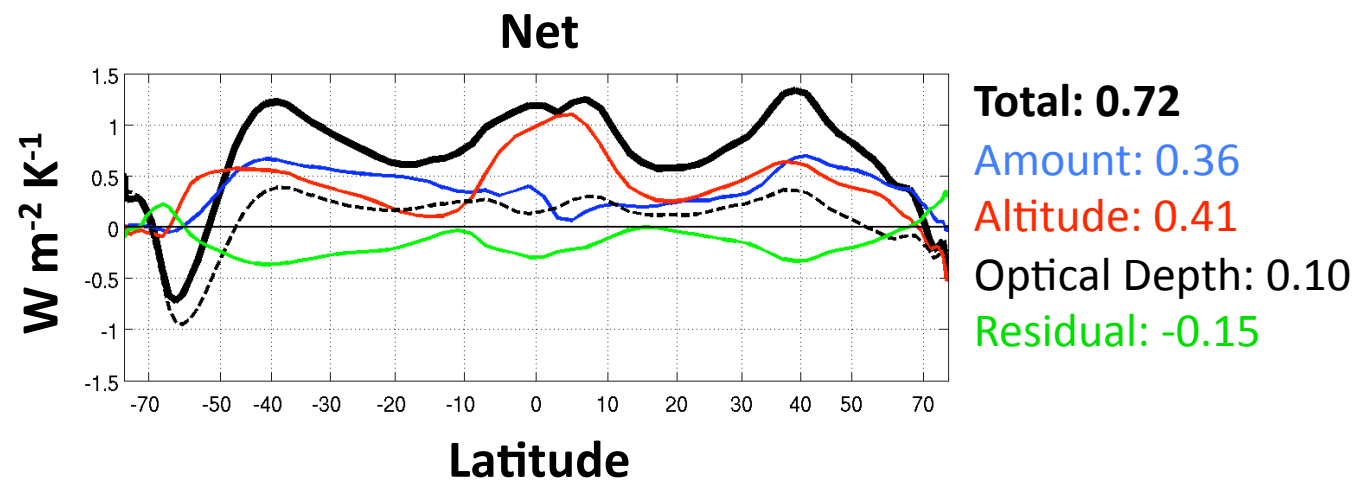
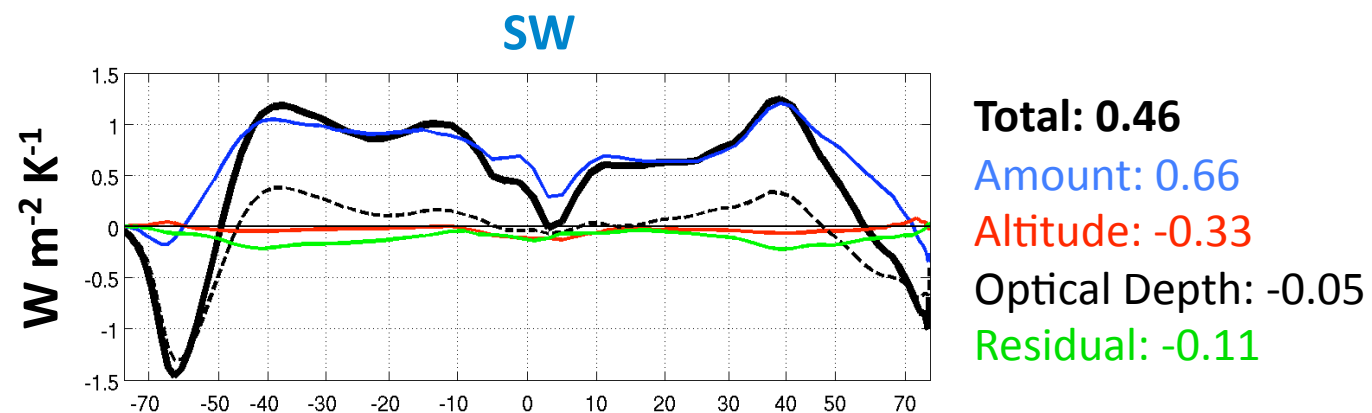
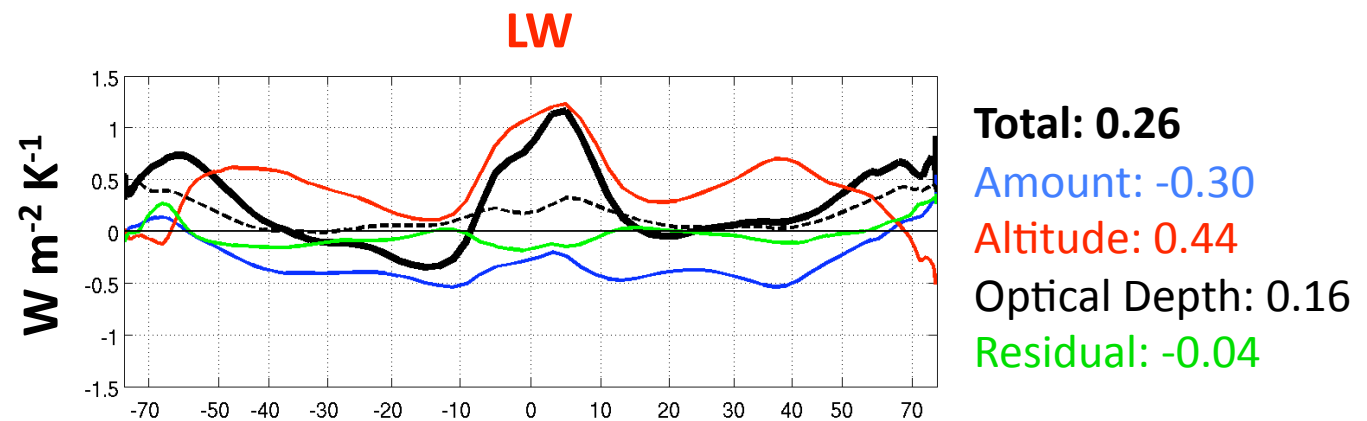


$\Delta\text{Altitude}$
 $\Sigma = 0.42$

$\Delta\tau$
 $\Sigma = 0.10$



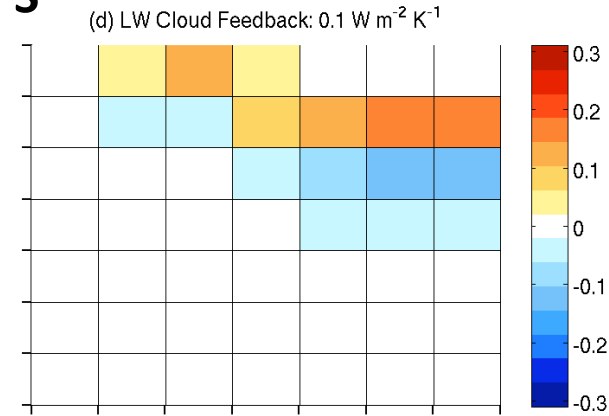
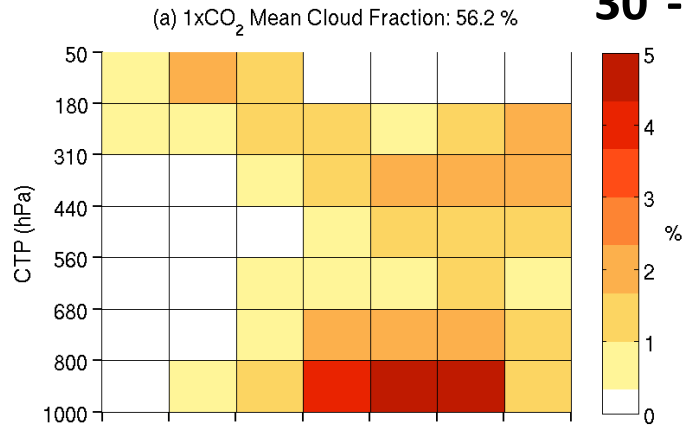
Residual
 $\Sigma = -0.16$



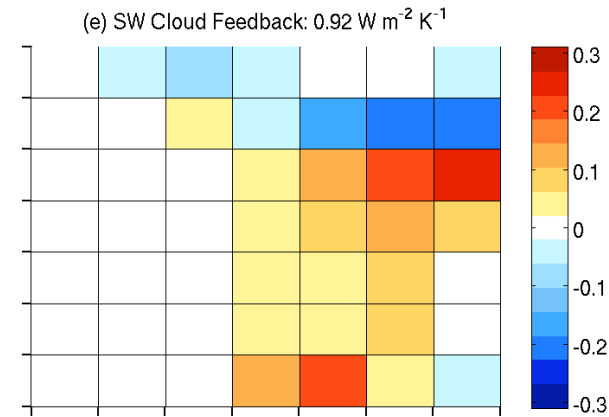
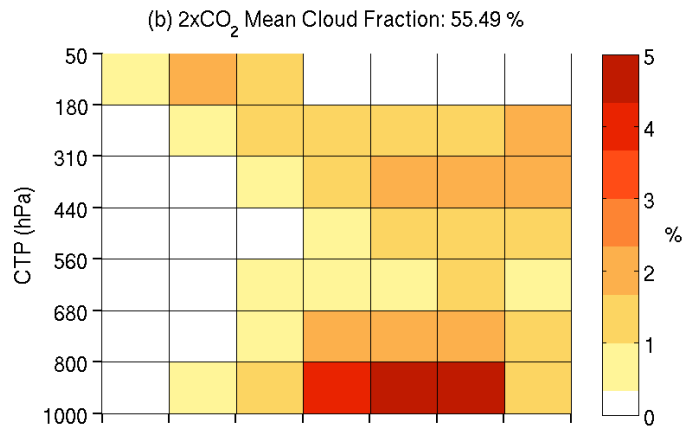
SW optical depth feedback: Negative at high latitudes, Positive at low latitudes

- Somerville & Remer (1984), Betts & Harshvardan (1987):
warmer → more liquid water in clouds → brighter clouds
(temperature sensitivity is twice as large at high latitudes)
- Tselioudis et al. (1992), Tselioudis & Rossow (1994), Chang and Coakley (2007): satellite observations indicate low cloud τ increases with T for cold clouds; opposite sensitivity for warm (low- and mid-latitude) clouds
- Senior & Mitchell (1993), Tsushima et al. (2006):
replacing ice with liquid → less efficient to precipitate (B-F effect)
→ more cloud water
- Timescale invariance of optical depth feedback?
→ Neil Gordon's talk yesterday

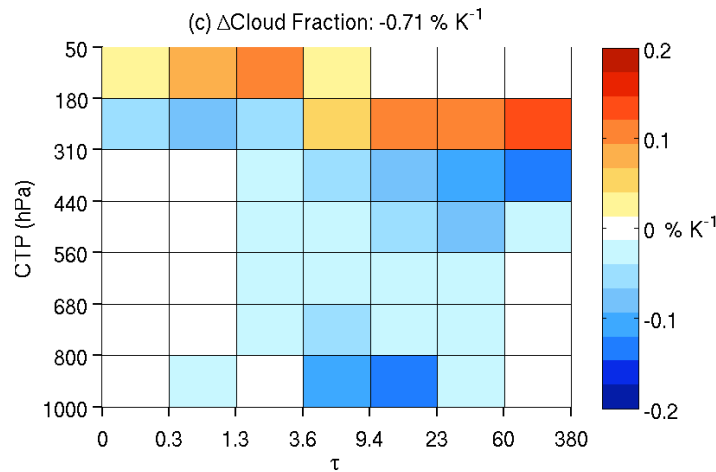
30°-50°S



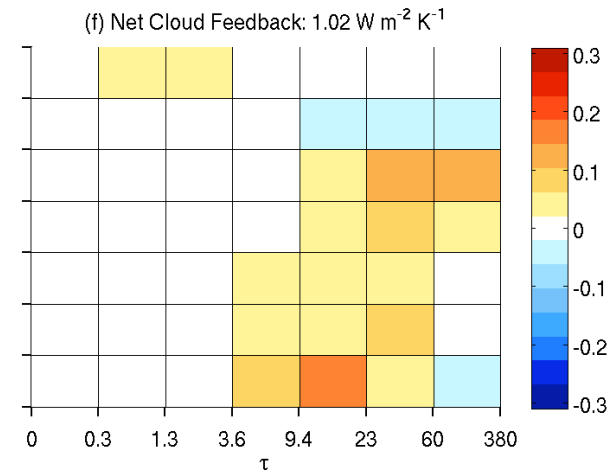
LW
0.10 W m⁻² K⁻¹



SW
0.92 W m⁻² K⁻¹



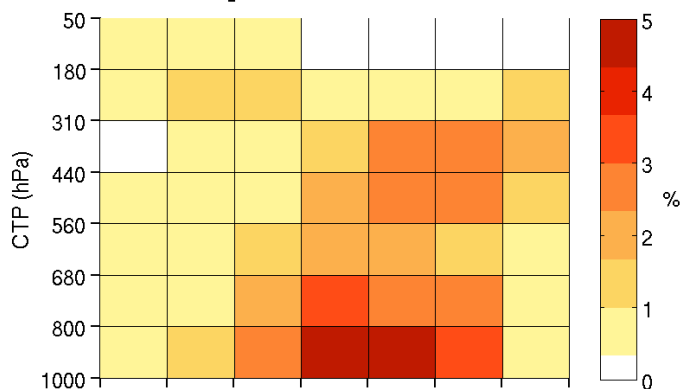
Δcfrac
-0.71 % K⁻¹



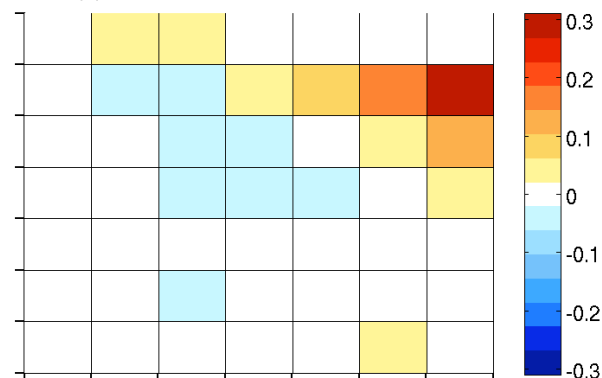
Net
1.02 W m⁻² K⁻¹

50°-70°S

(a) 1xCO₂ Mean Cloud Fraction: 74.63 %

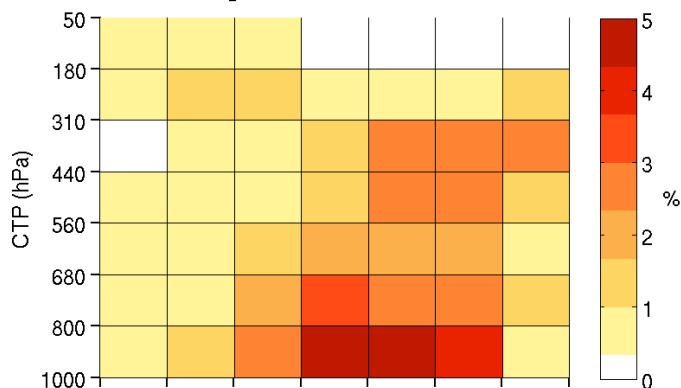


(d) LW Cloud Feedback: 0.66 W m⁻² K⁻¹

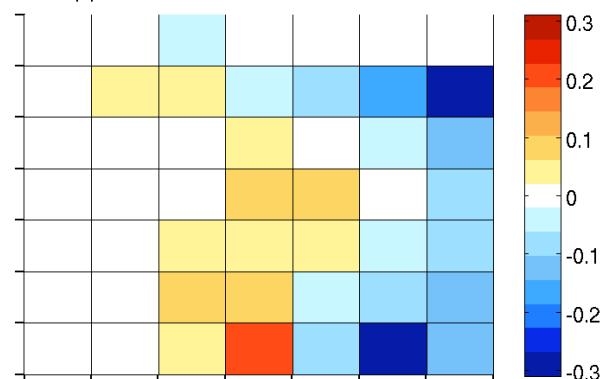


LW
0.66 W m⁻² K⁻¹

(b) 2xCO₂ Mean Cloud Fraction: 74.73 %

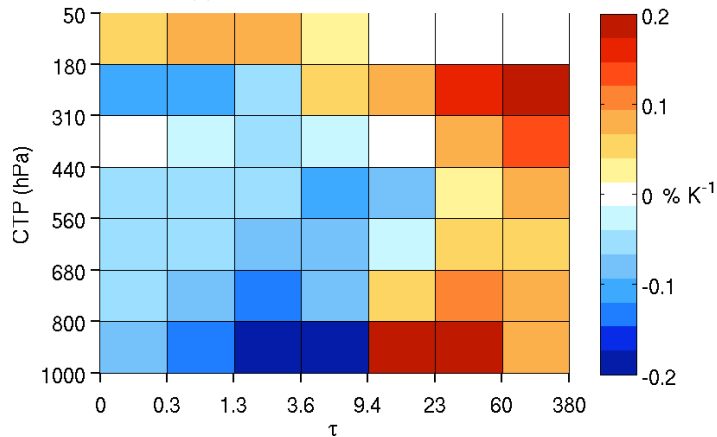


(e) SW Cloud Feedback: -0.95 W m⁻² K⁻¹

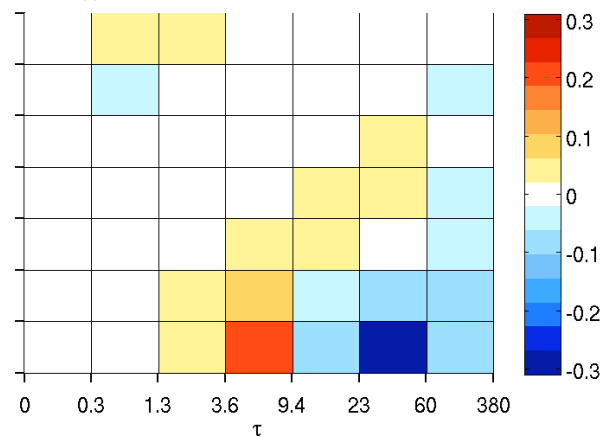


SW
-0.95 W m⁻² K⁻¹

(c) ΔCloud Fraction: 0.11 % K⁻¹



(f) Net Cloud Feedback: -0.3 W m⁻² K⁻¹



Net
-0.30 W m⁻² K⁻¹

Δcfrac
0.11 % K⁻¹

