



THE UNIVERSITY OF TOKYO

Modeling activity of the MIROC group for climate feedback and sensitivity studies

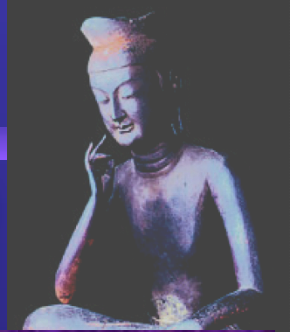
Masahiro Watanabe
Atmosphere & Ocean Research Institute
University of Tokyo
hiro@aori.u-tokyo.ac.jp

Collaborators: M. Yoshimori, Y. Kamae, M. Kimoto (AORI, Univ of Tokyo), H. Shiogama, T. Yokohata, T. Ogura, S. Emori (NIES), J. D. Annan, J. C. Haargreaves (JAMSTEC)



Background & Outline

Earth Simulator 2



Modeling activity by MIROC group

- ✓ Watanabe et al. 2010 JC (MIROC₅)
- ✓ Yokohata et al. 2010 JC (MIROC₃ PPE)
- ✓ Shiogama et al. 2011 CD (MIROC₅ PPE) *
- ✓ Watanabe et al. 2011 ASL (MIROC₅ 20C)
- ✓ Watanabe et al. 2011 CD (MIROC_{3/5}) #
- ✓ Watanabe et al. 2011 JC (MPE) *
- ✓ Yoshimori et al. 2011 JC (MIROC₃ PPE) #
- ✓ Yokohata et al 2011 CD (MIROC₃ PPE/MME) #

* submitted, # in press

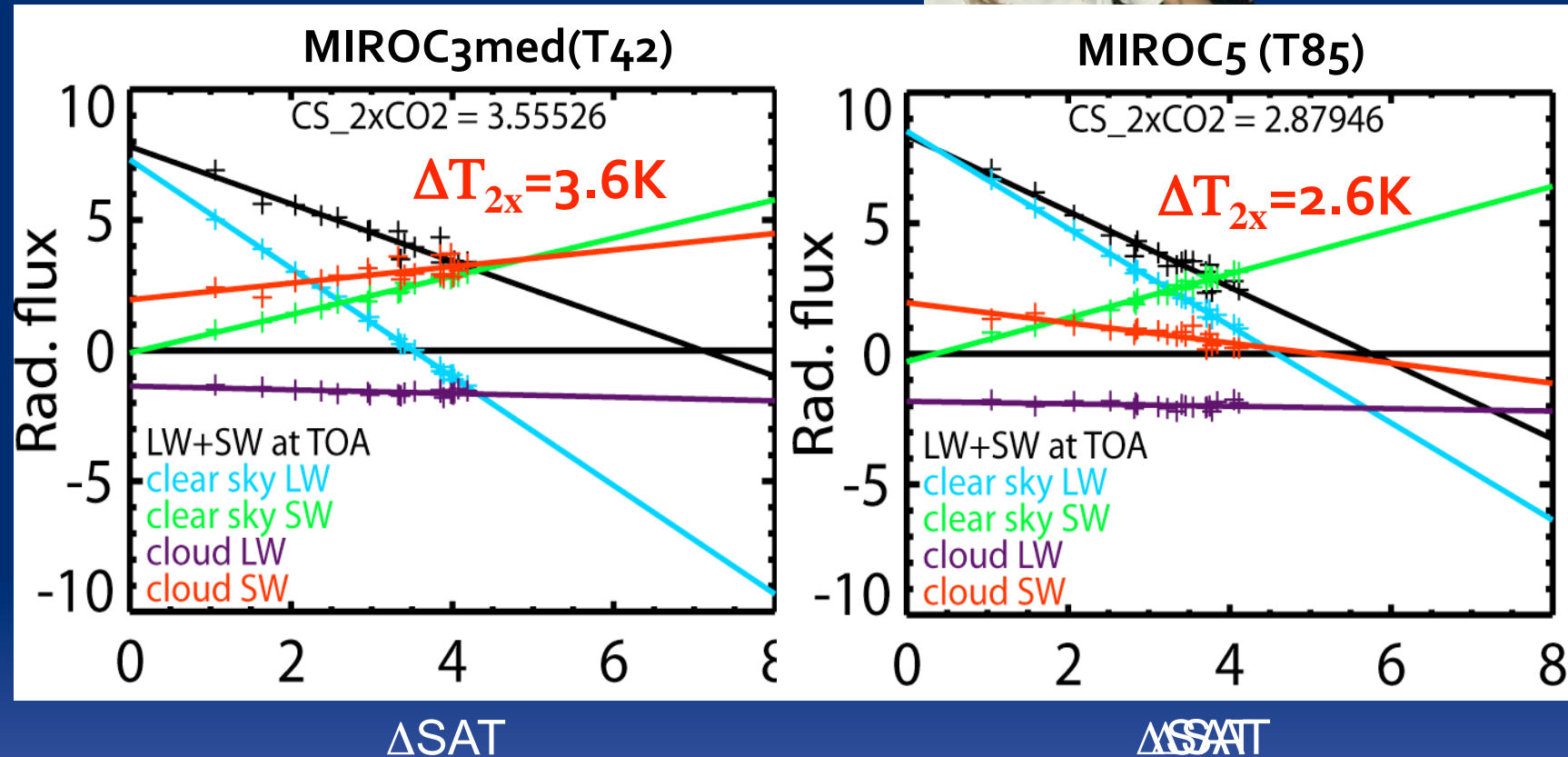
Outline of this talk

- Cloud feedbacks in MIROC_{3.2}/MIROC₅/CMIP₃ MME
- Various ensembles using MIROC_s

Equilibrium climate sensitivity

$$N = F -$$

Gregory plots for $4\times\text{CO}_2$

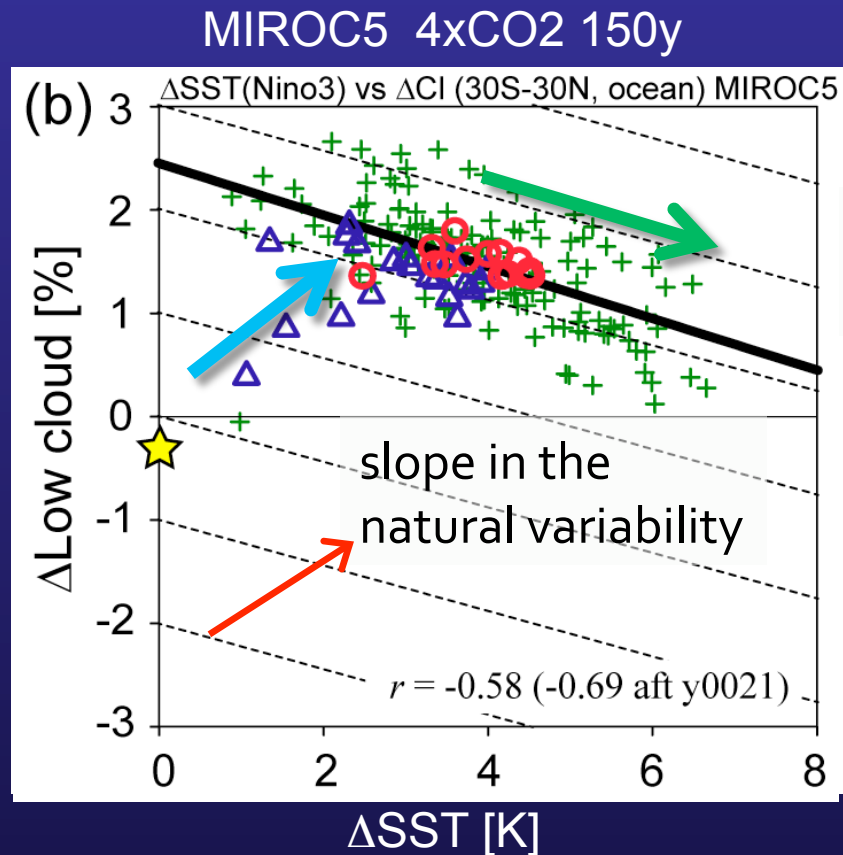


MIROC3 was a 'high-sensitivity model' in CMIP3

Watanabe et al. (2010, JC)

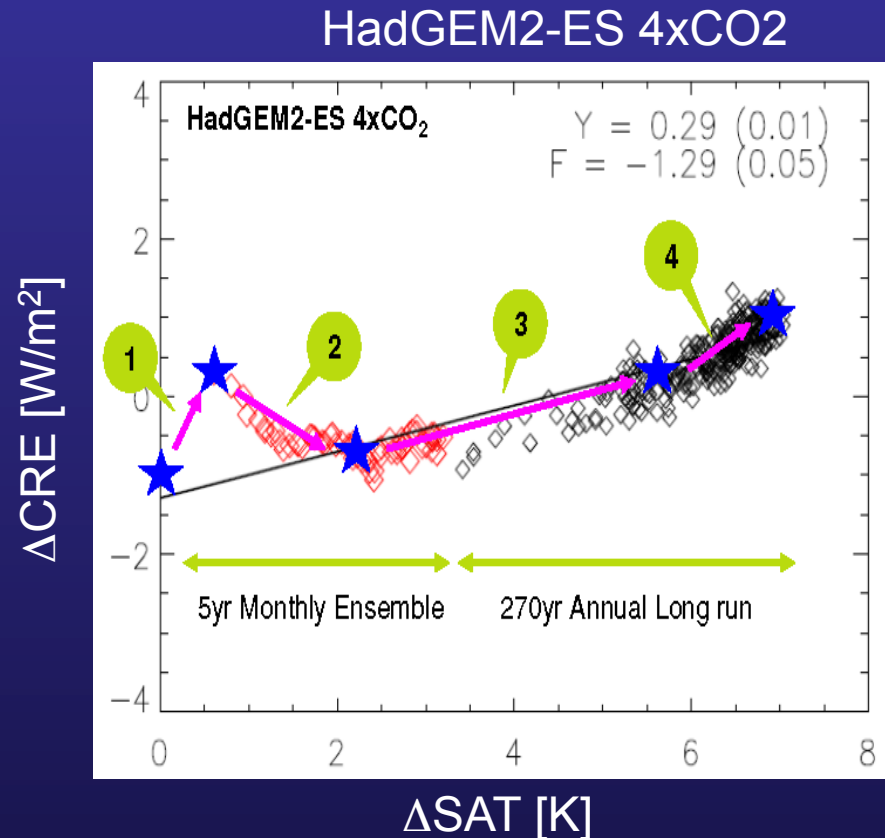
Cloud-shortwave (SWcld) feedback is the key component

Multiple timescales of the cloud feedback



Watanabe et al. (2011, Clim. Dyn.)

- + ann. mean, single run (150y)
- decadal mean, single run
- △ ann. mean, ensemble avg (20y)
- ★ initial month of △



Courtesy of Tim Andrews

Sign of the SWcld feedback is determined by the fast response of clouds during a few years in MIROC5

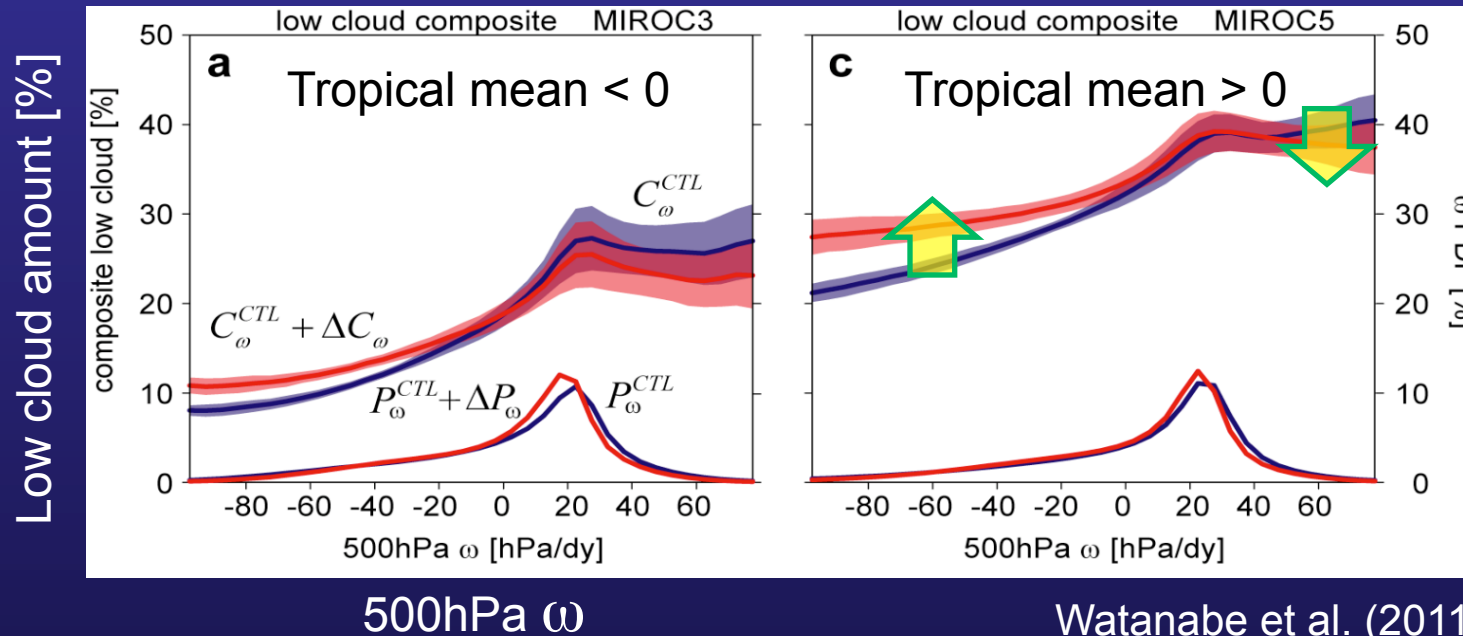
Fast response of low clouds

4xCO₂ runs
(20y ensemble)

Regime composite of low cloud wrt ω_{500}

MIROC3

MIROC5



$$\Delta \tilde{C}_l \approx \int \Delta P_{\omega} C_l^{CTL}(\omega) d\omega + \int P_{\omega}^{CTL} \Delta C_l(\omega) d\omega$$

Thermodynamic driving

Tropical-mean changes in C_l & SWcld are due to a subtle difference between C_l decrease / increase over the subsidence / ascent regime

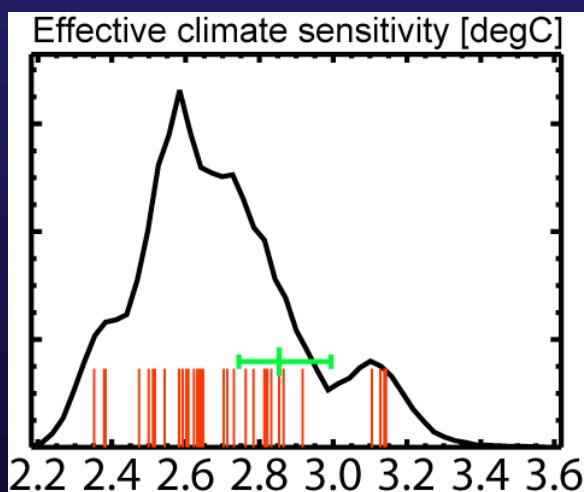
MIROC Perturbed Physics Ensembles (PPEs)

□ MIROC3.2 PPE (JUMP project, Annan et al. 2005)

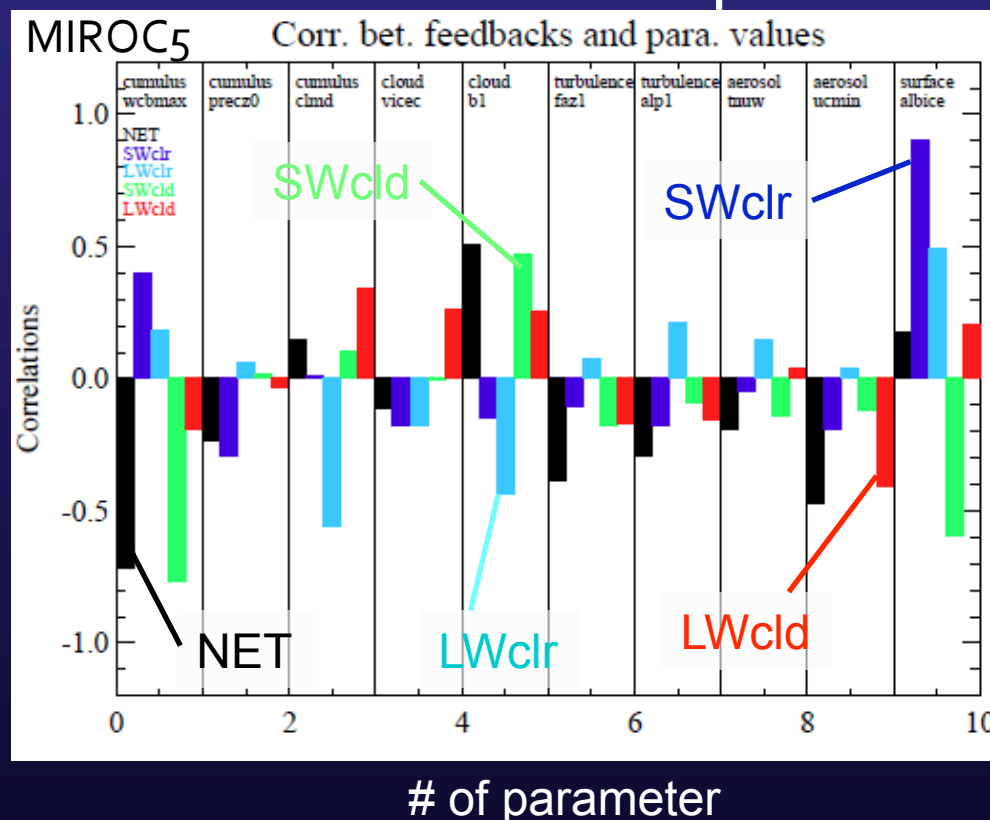
- ✓ T20L20 slab CGCM
- ✓ 13 parameters perturbed w/ EnKF

□ MIROC5 PPE (Shiogama et al. 2011, submitted)

- ✓ T42L40 full CGCM
- ✓ 10 parameters perturbed w/ Latin hypercube + emulator



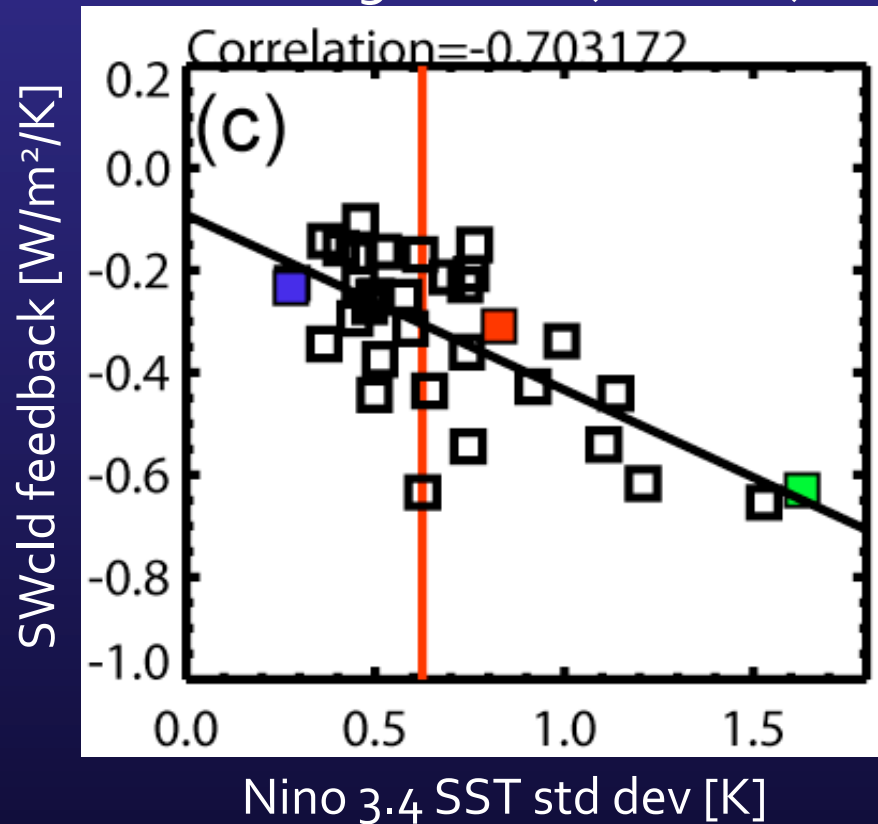
Corr. between feedbacks and parameters



ENSO and climate sensitivity?

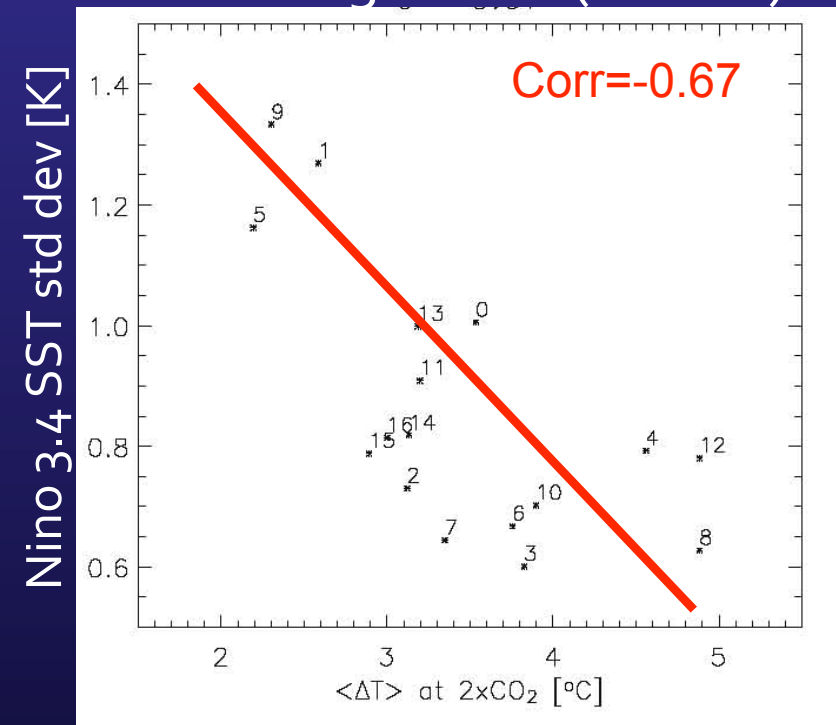
Reasons may be different in the two models, but a common feature of *strong ENSO in the control run <---> low CS*

MIROC5 PPE-C (JUMP2)



Shiogama et al. (2011, submitted)

HadCM3 PPE-C (QUMP)



Toniazzo et al. (2008)

ENSO and climate sensitivity?

Compare 10 models having

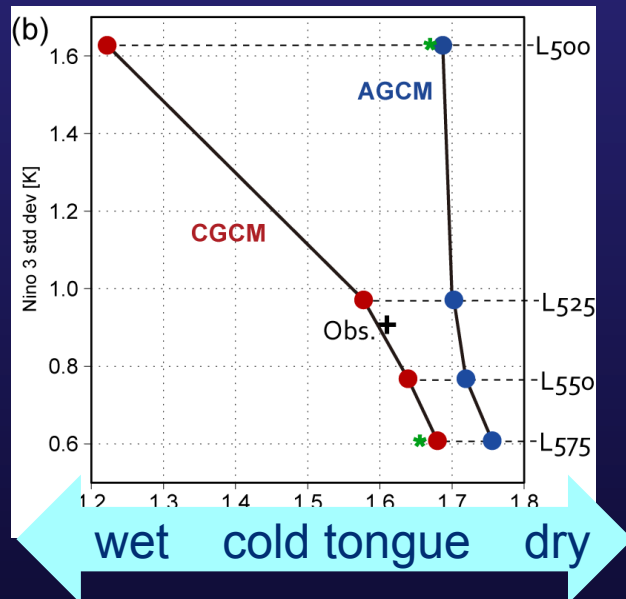
- * Largest negative SWcld feedback
- * Smallest negative SWcld feedback

Larger SWcld feedback

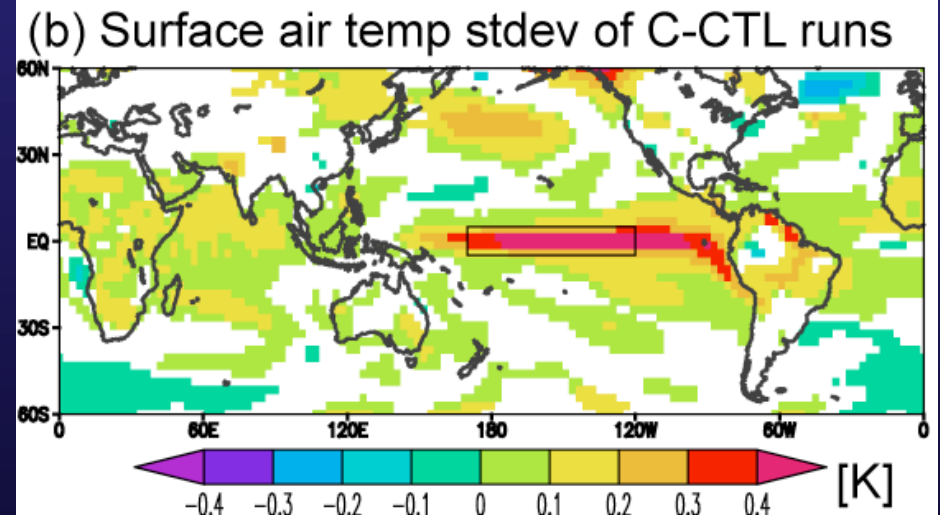
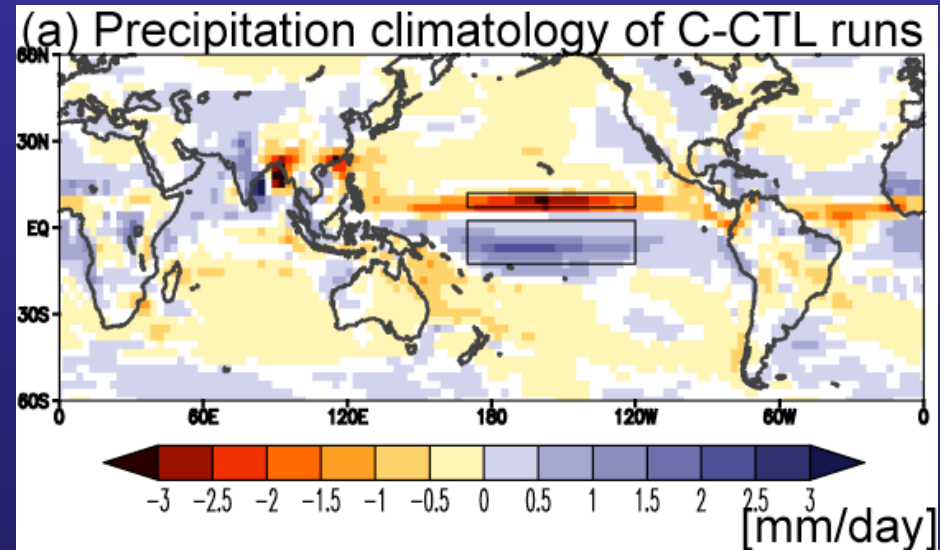


Wetter mean state over the central equatorial Pacific & stronger ENSO

ENSO amplitude [K]



Watanabe et al. (2011, JC)

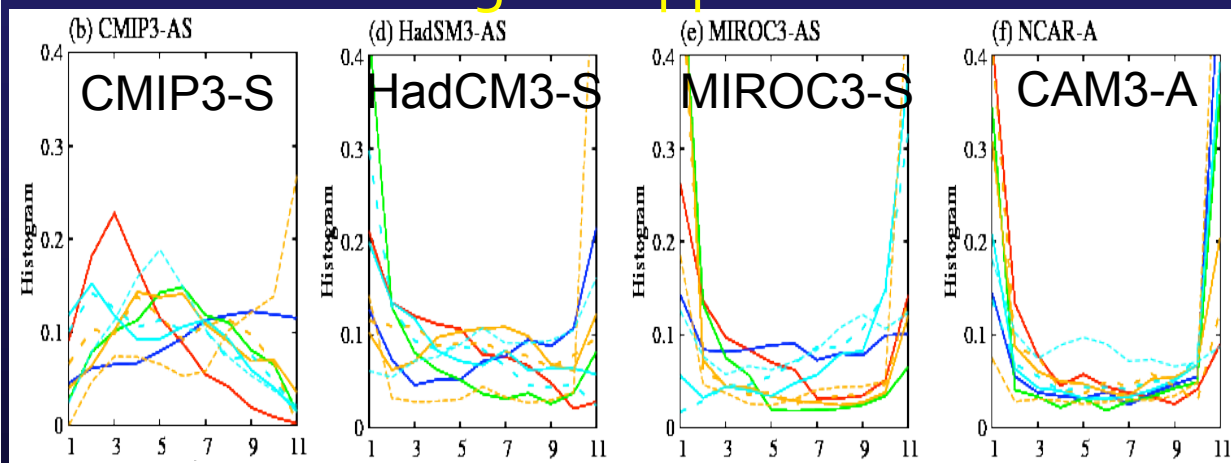


Shiogama et al. (2011, submitted)

Combined structural/parametric uncertainties

- ❑ PPE can be complementary to CMIP MME and has some advantages over MME
- ❑ PPE may also have disadvantages that
 - ✓ The ensemble depends on the base model (Yokohata et al. 2011 JC)
 - ✓ Model errors in the ensemble may not cancel each other
 - ✓ May not be suitable for exploring feedback *mechanisms*

Rank histogram approach

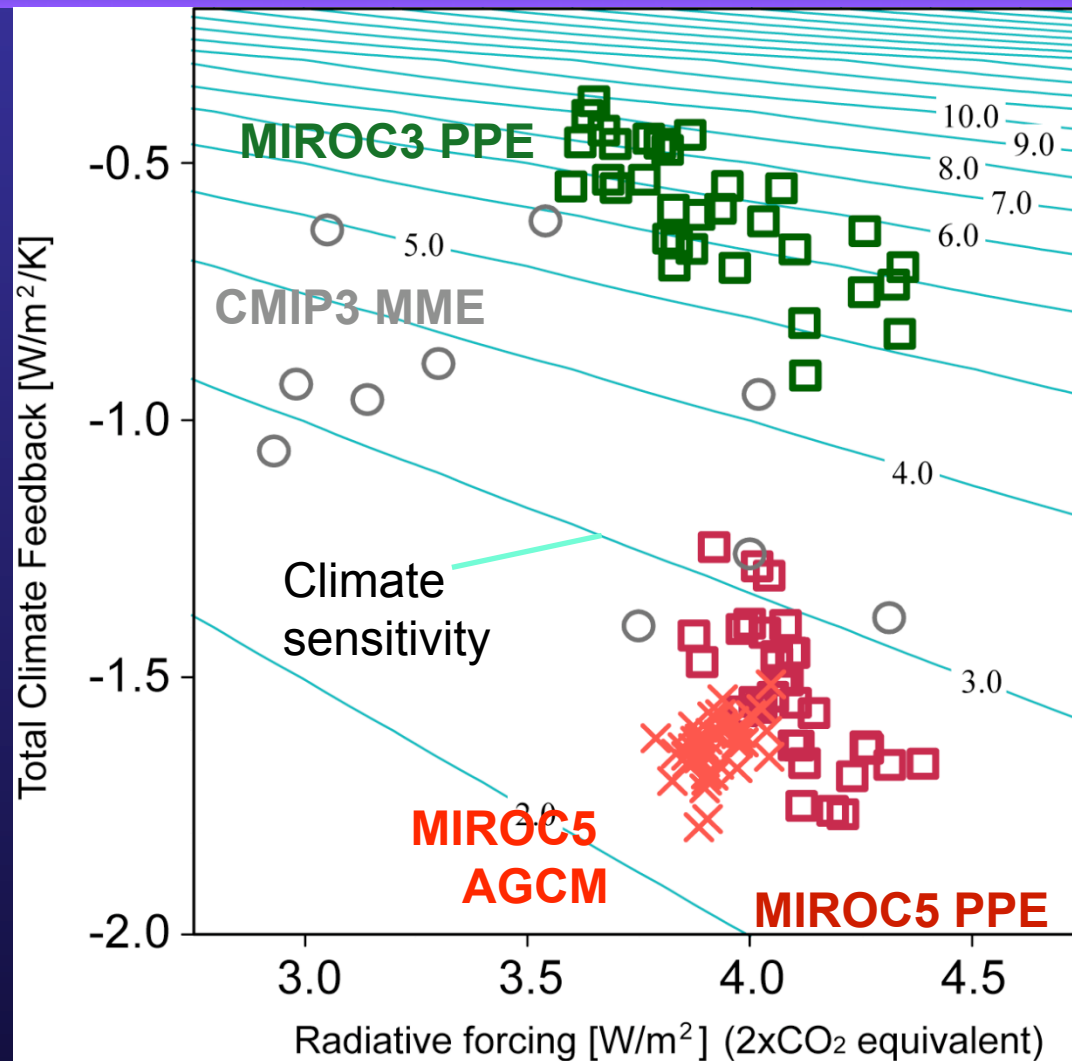


Rank

Yokohata et al. (2011, Clim Dyn)

- ✓ Ranks observations among ensembles
→ Histogram
- ✓ **Flat histogram**
= ensemble members are around the observation
- ✓ **Edge-peaked histogram**
= ensemble members over-/under-estimate observation

Forcing-Feedback in two PPEs



Multi-Physics Ensemble (MPE)

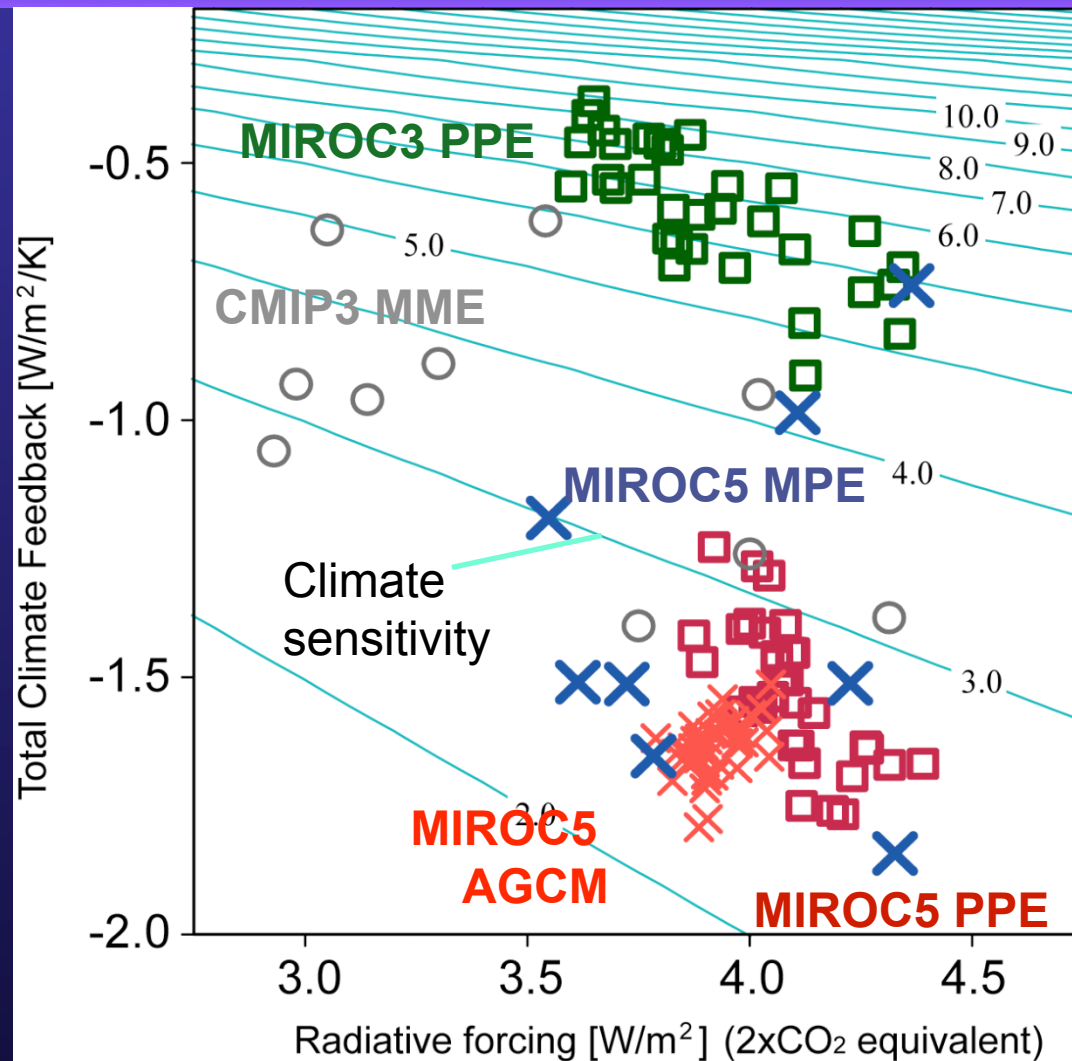
Replacing one or more schemes in MIROC5 w/ old ones:

- ✓ Std (=MIROC5)
- ✓ (old)CLD
- ✓ (old)CUM
- ✓ (old)VDF
- ✓ (old)CLD+CUM
- ✓ (old)CUM+VDF
- ✓ (old)CLD+VDF
- ✓ (old)CLD+CUM+VDF

- ✓ Structural difference > Parametric difference
- ✓ Any strategy to link them ?

Watanabe et al. (2011, submitted)

Forcing-Feedback in two PPEs



Multi-Physics Ensemble (MPE)

Replacing one or more schemes in MIROC5 w/ old ones:

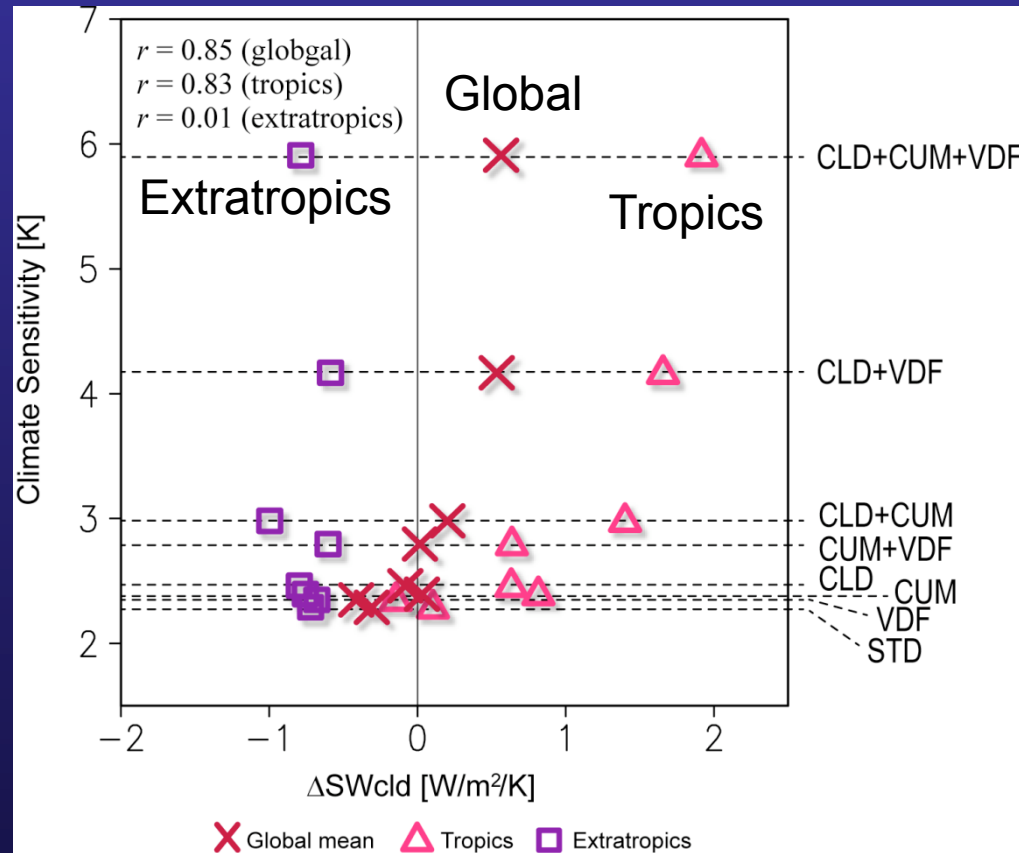
- ✓ Std (=MIROC5)
- ✓ (old)CLD
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- ✓ (old)CLD+CUM+VDF

- ✓ Structural difference > Parametric difference
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Watanabe et al. (2011, submitted)

MPE: Filling the gap between PPEs

Climate sensitivity vs SWcld feedback



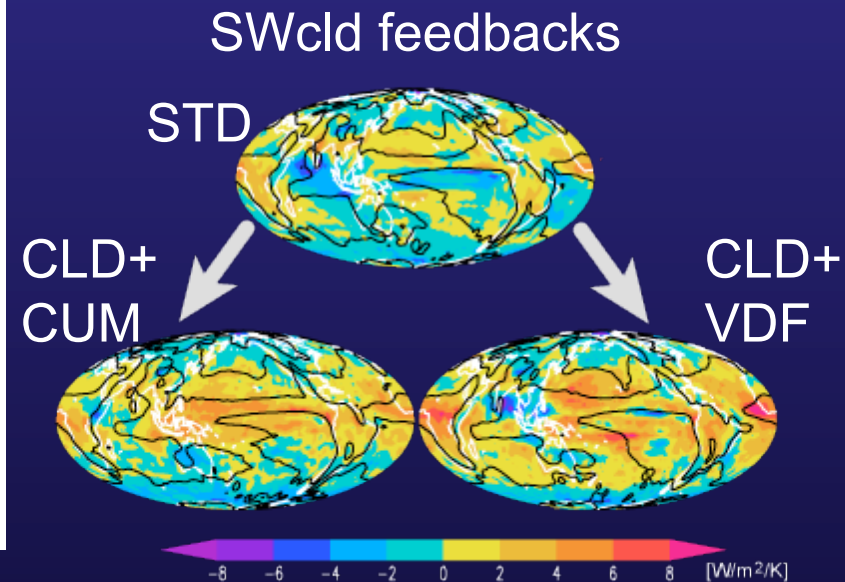
Watanabe et al. (2011, submitted to JC)

Processes are nonlinear, e.g.,

(old)CLD → small impact

(old)VDF → small impact

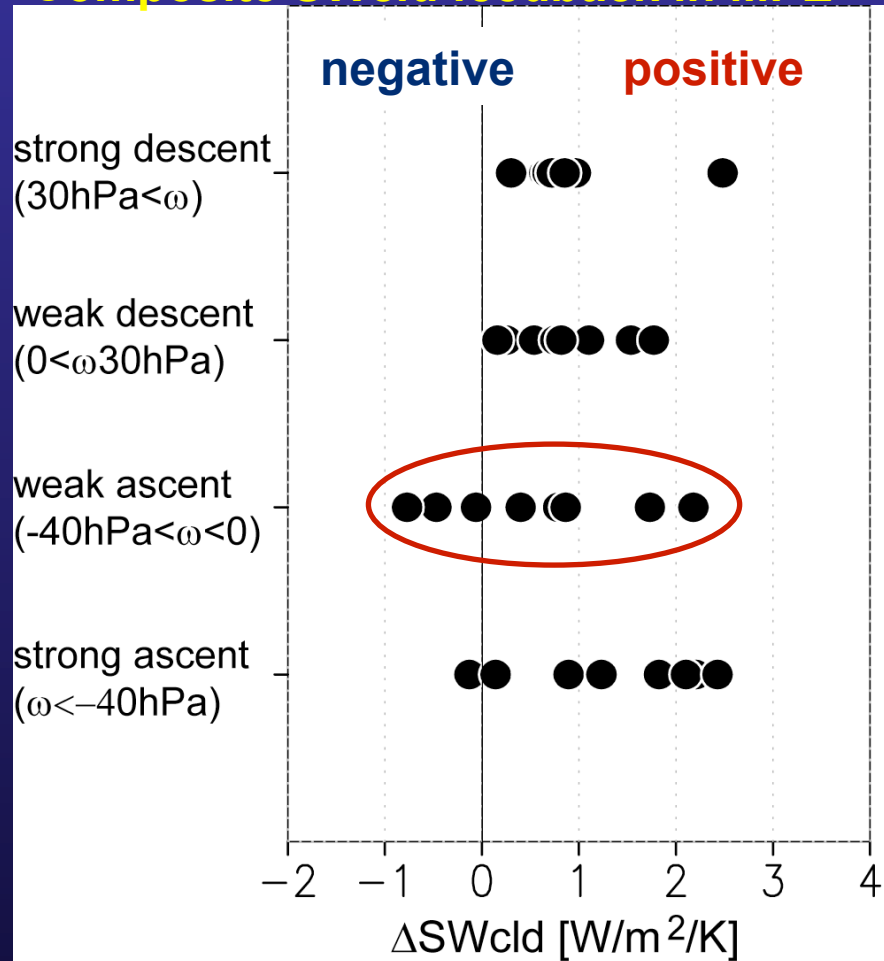
(old)CLD+VDF → large impact!!



‘Feedback occurs thru the interaction of a suite of parameterized processes rather than from any single process’ (Zhang & Bretherton 2008)

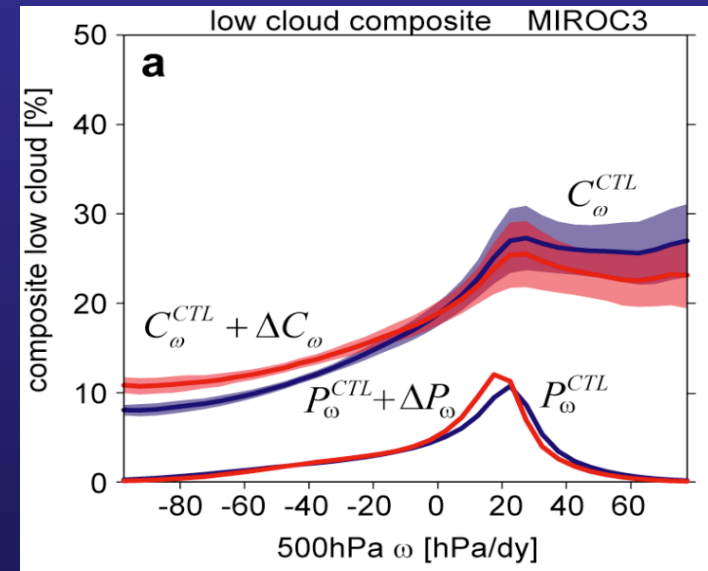
SWcld feedback in different circulation regimes

Composite SWcld feedback in MPE



Watanabe et al. (2011, submitted)

Composite low clouds in MiROC3 4xCO₂

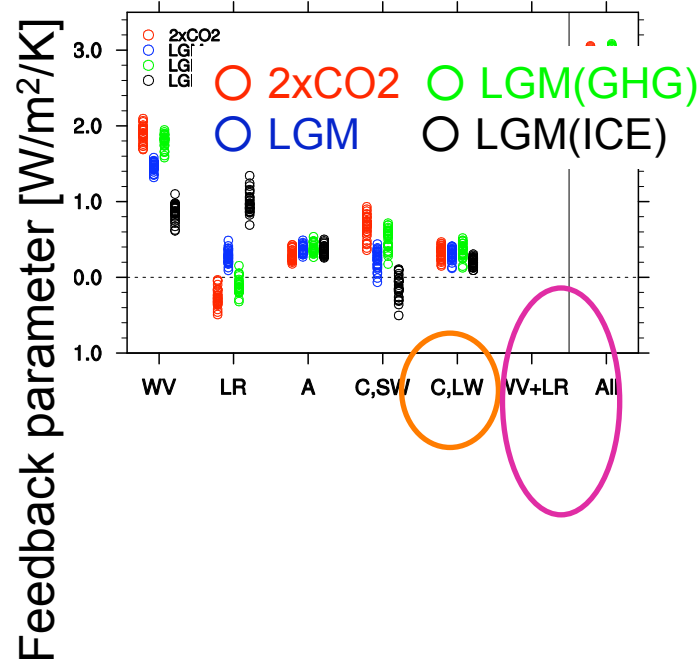


SWcld feedback appears to be robust (at least in MIROC3/5) for the subsidence regimes, but not for the convective regime!

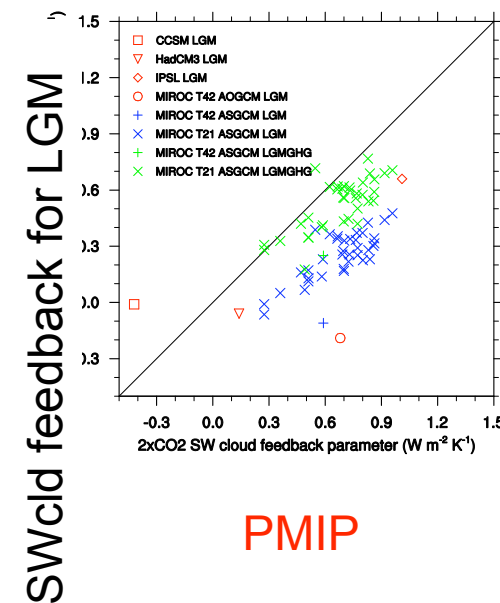
➔ Trade cumulus response?

Can climate feedback be constrained with palaeoclimate simulations?

Feedback analysis based on MIROC_{3.2} PPE



SWcld values correlated for LGM & 2xCO₂ in PPE (but not in PMIP2)



SWcld feedback for 2xCO₂

Yoshimori et al. (2011, JC)

- ✓ Different total feedback (hence ECS) for different forcing & mean climate
- ➔ caution to use palaeo data to constrain ECS

Conclusions

Toward understanding sources of model's "uncertainty" in terms of the climate sensitivity & climate feedbacks,

- We emphasize thru the analysis of MIROC_{3.2}/MIROC₅/CMIP₃ an importance of:
 - ✓ identifying timescales of cloud response/feedback
 - ✓ distinguishing robust/fragile parts associated w/ cloud response/feedback
- Extensive use of various ensembles based on MIROC reveals
 - ✓ a connection between apparently independent phenomena (ENSO & climate sensitivity)
 - ✓ attribution of inconsistency among GCMs to individual processes/interactions

Recent publications

- ✓ Watanabe et al. 2010 JC (MIROC₅)
- ✓ Yokohata et al. 2010 JC (MIROC₃ PPE)
- ✓ Shiogama et al. 2011 CD (MIROC₅ PPE) *
- ✓ Watanabe et al. 2011 ASL (MIROC₅ 20C)
- ✓ Watanabe et al. 2011 CD (MIROC_{3/5}) #
- ✓ Watanabe et al. 2011 JC (MPE) *
- ✓ Yoshimori et al. 2011 JC (MIROC₃ PPE) #
- ✓ Yokohata et al 2011 CD (MIROC₃ PPE/MME) #

* submitted, # in press

All available at

[http://amaterasu.ees.hokudai.ac.jp/
~fswiki/pub/wiki.cgi?page=CMIP5](http://amaterasu.ees.hokudai.ac.jp/~fswiki/pub/wiki.cgi?page=CMIP5)

ID guest

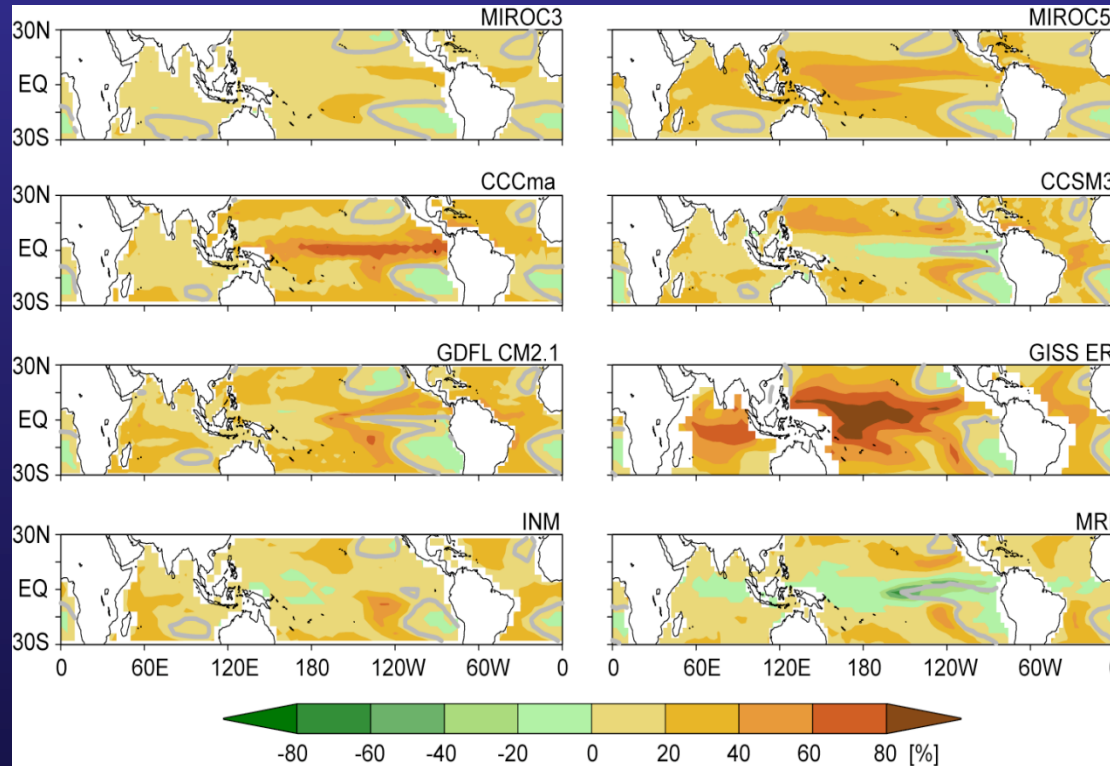
Password miroc



backup

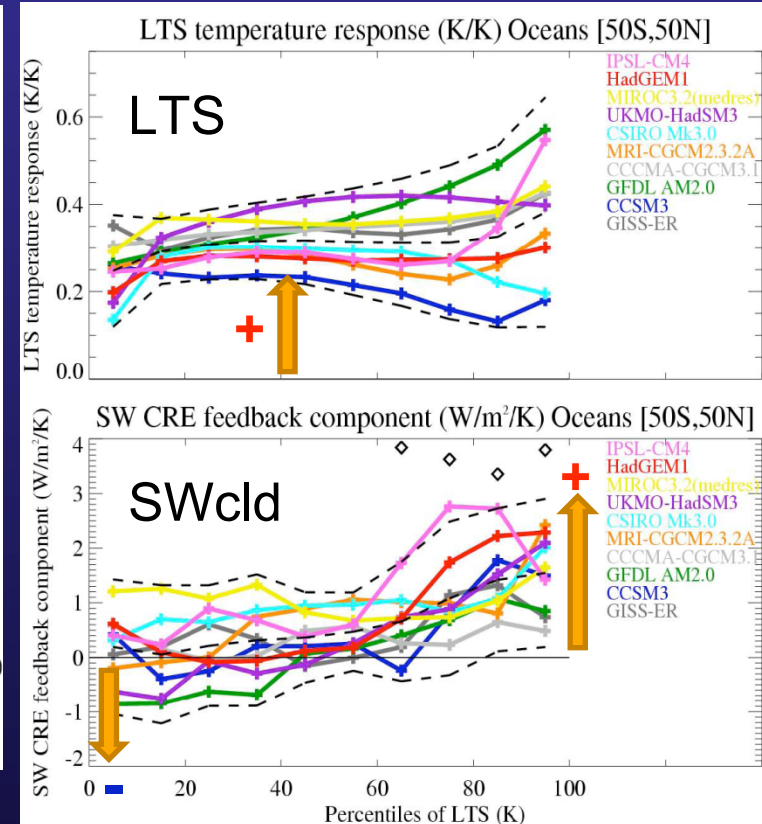
What is robust? What is not robust?

Change in stable regime frequency in CFMIP1



Watanabe et al. (2011, Clim Dyn)

Sorted responses in CFMIP1



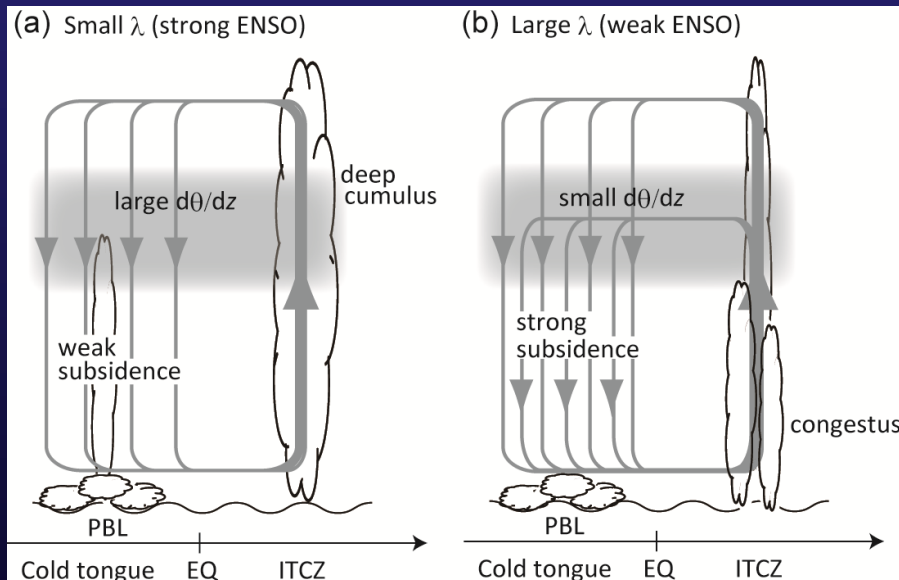
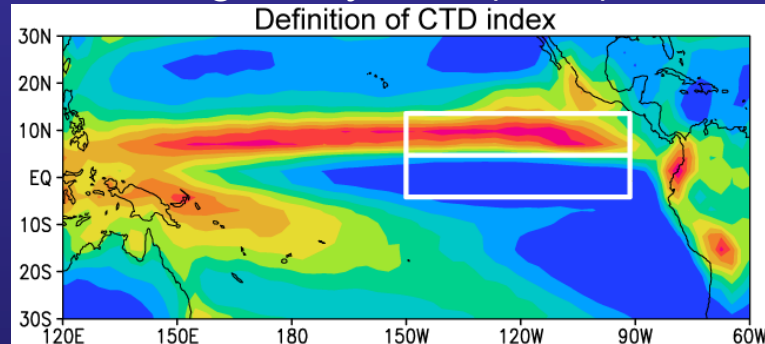
Unstable $\longleftrightarrow$ Stable

- ✓ Increasing LTS: sign robust (but not magnitude)
- ✓ Change in SWcld: fragile even sign

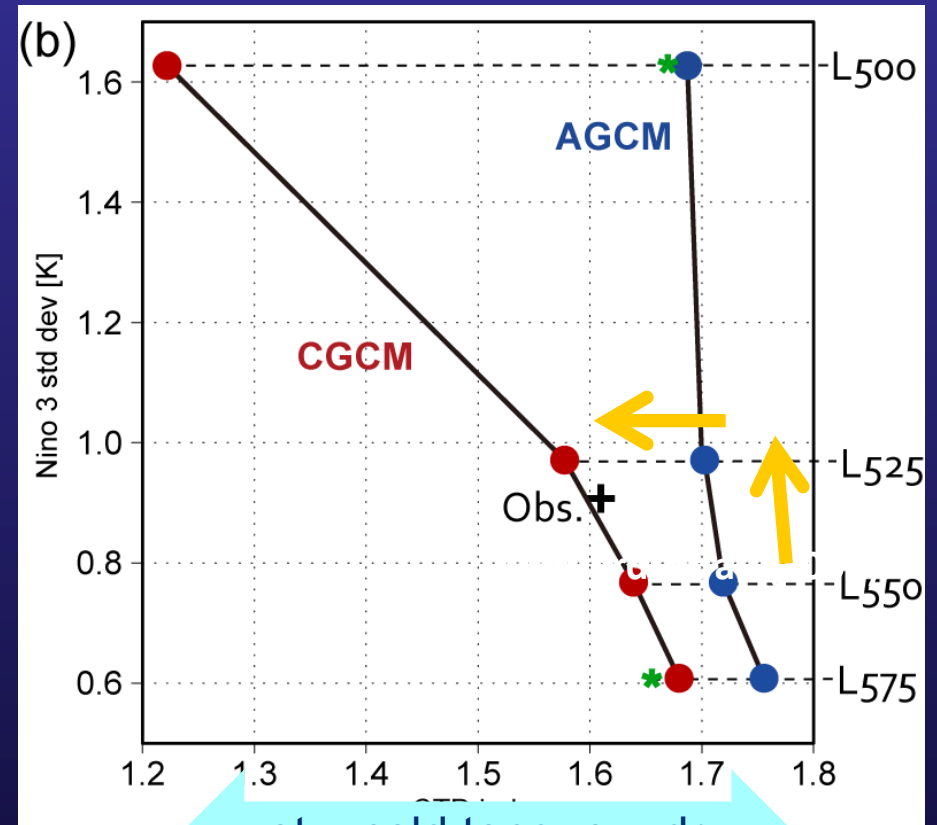
Courtesy of M Webb

ENSO control mechanism in MIROC5

Cold tongue dryness (CTD) index



ENSO & mean state error



wet cold tongue dry

Direct effect of convection

Coupled feedbacks

MIROC updates after AR4

MIROC3.2

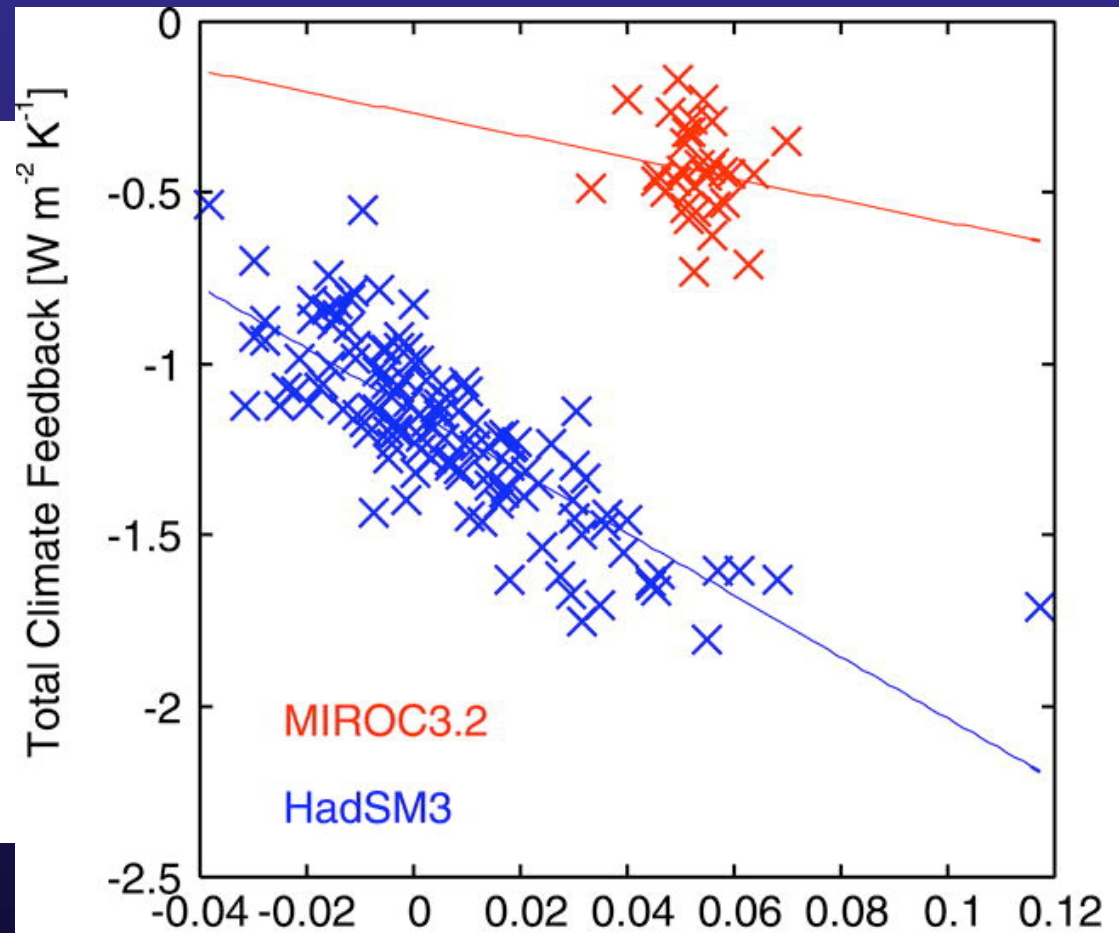
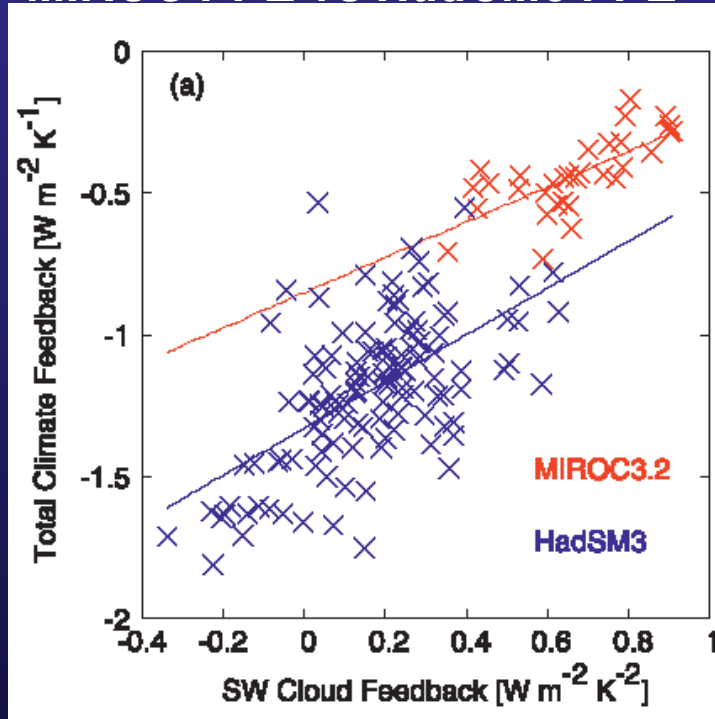
MIROC5

Atmos.	Dynamical core	Spectral+semi-Lagrangian (Lin & Rood 1996)	Spectral+semi-Lagrangian (Lin & Rood 1996)
	V. Coordinate	Sigma	Eta (hybrid sigma-p)
	Radiation	2-stream DOM 37ch (Nakajima et al. 1986)	2-stream DOM 111ch (Sekiguchi et al. 2008)
	Cloud	Diagnostic (LeTreut & Li 1991) + Simple water/ice partition	Prognostic PDF (Watanabe et al. 2009) + Ice microphysics (Wilson & Ballard 1999)
	Turbulence	M-Y Level 2.0 (Mellor & Yamada 1982)	MYNN Level 2.5 (Nakanishi & Niino 2004)
	Convection	Prognostic A-S + critical RH (Pan & Randall 1998, Emori et al. 2001)	Prognostic AS-type, but original scheme (Chikira & Sugiyama 2010)
	Aerosols	simplified SPRINTARS (Takemura et al. 2002)	SPRINTARS + prognostic CCN (Takemura et al. 2009)
Land/ River		MATSIRO+fixed riv flow	new MATSIRO+variable riv flow
Ocean		COCO3.4	COCO4.5
Sea-ice		Single-category EVP	Multi-category EVP

Combined structural/parametric uncertainty

Understanding how parameter-driven uncertainties in PPEs depend on structural properties of the model

MIROC PPE vs HadCM3 PPE



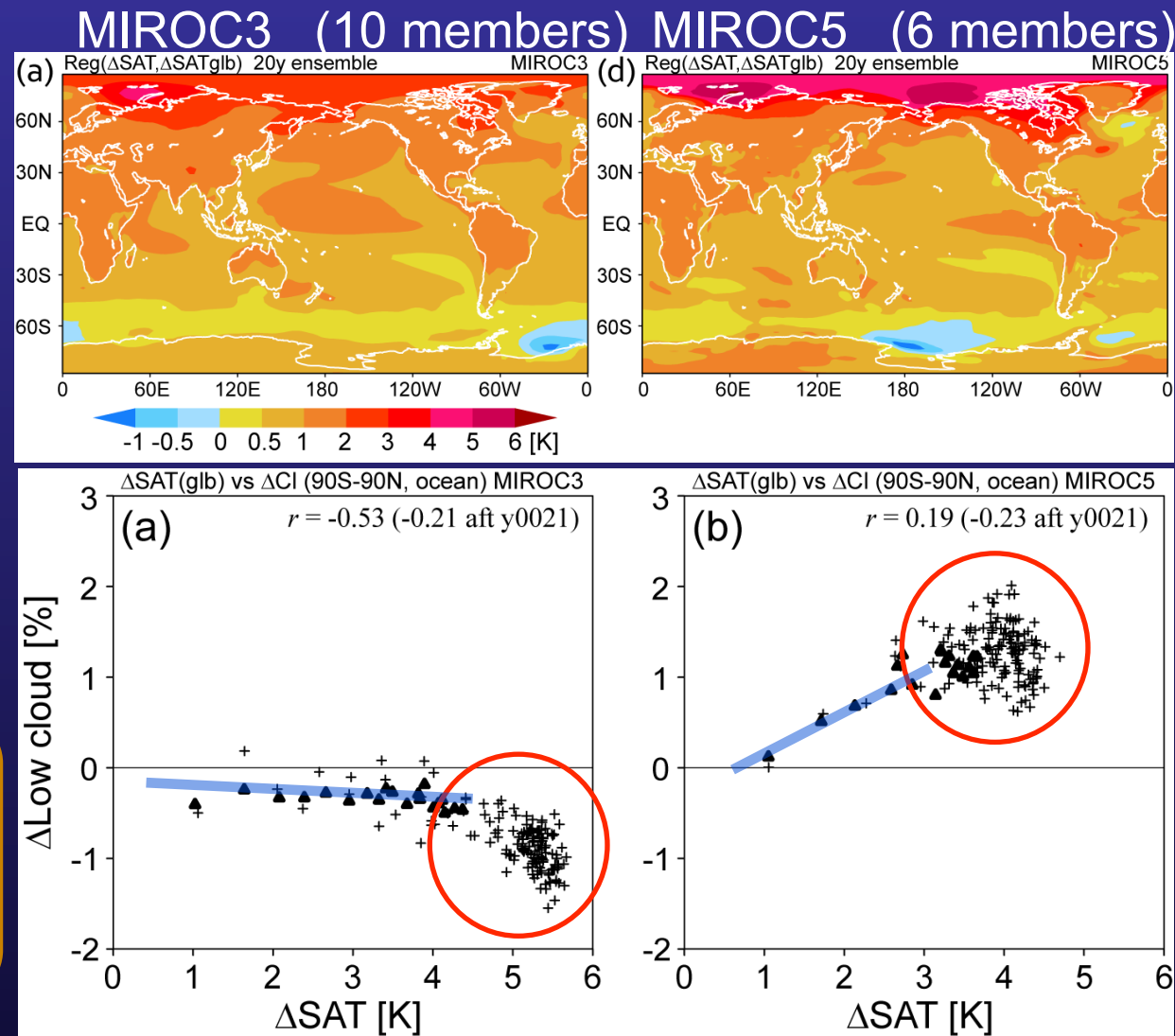
Cloud Albedo Bias at 680-1000 hPa in Control

Yokohata et al. (2011, JC)

Global warming patterns

4xCO₂ runs

Reg(Δ SAT,
 Δ SAT_g)



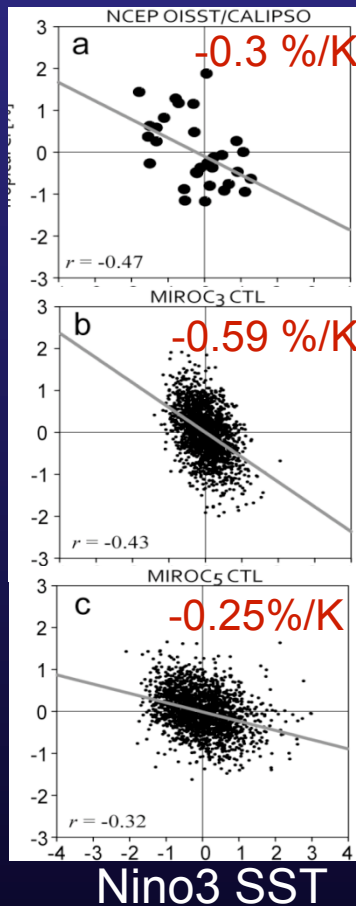
Δ SAT_g may not explain global-mean Δ C_l, except for initial years

Natural low-cloud (C_l) variability

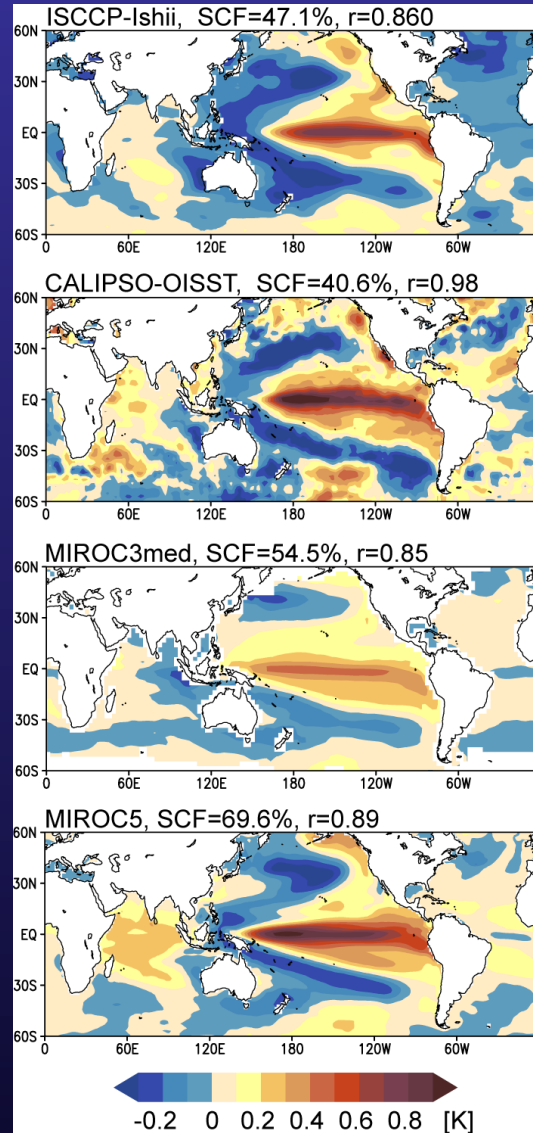
Obs/CTL runs

Leading SVD

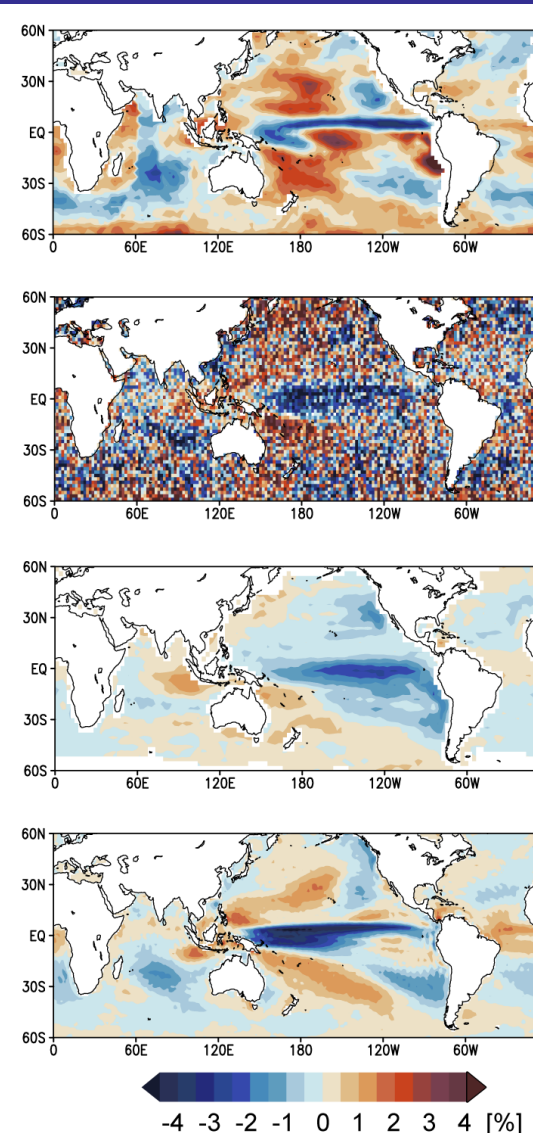
Tropical-mean C_l



SST anomaly



C_l anomaly



Ishii SST/
ISCCP

OISST/
CALIPSO

MIROC3 CTL

MIROC5 CTL

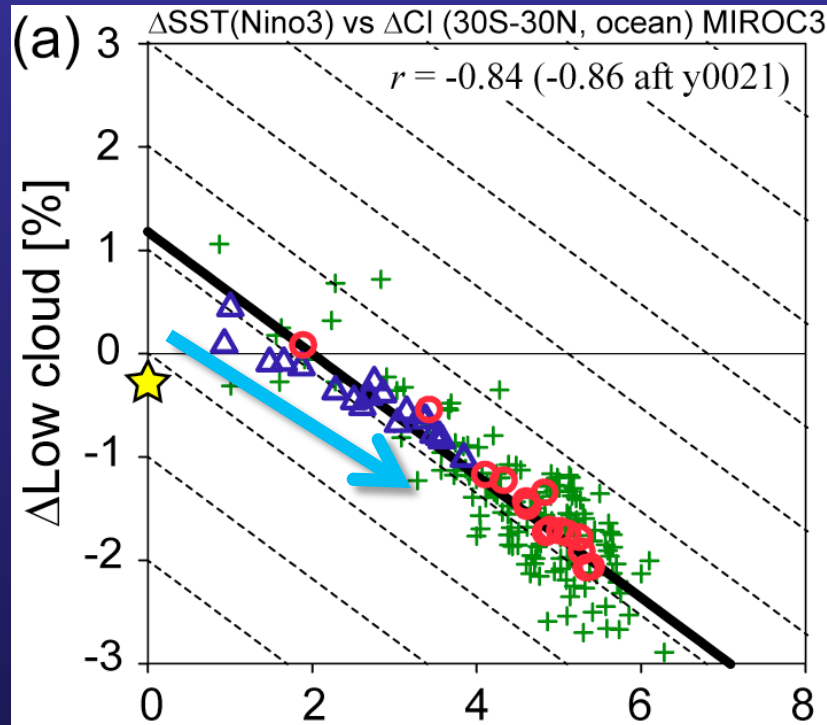
Watanabe et al. (2011, Clim Dyn)

Three timescales of low-cloud changes

4xCO₂ runs

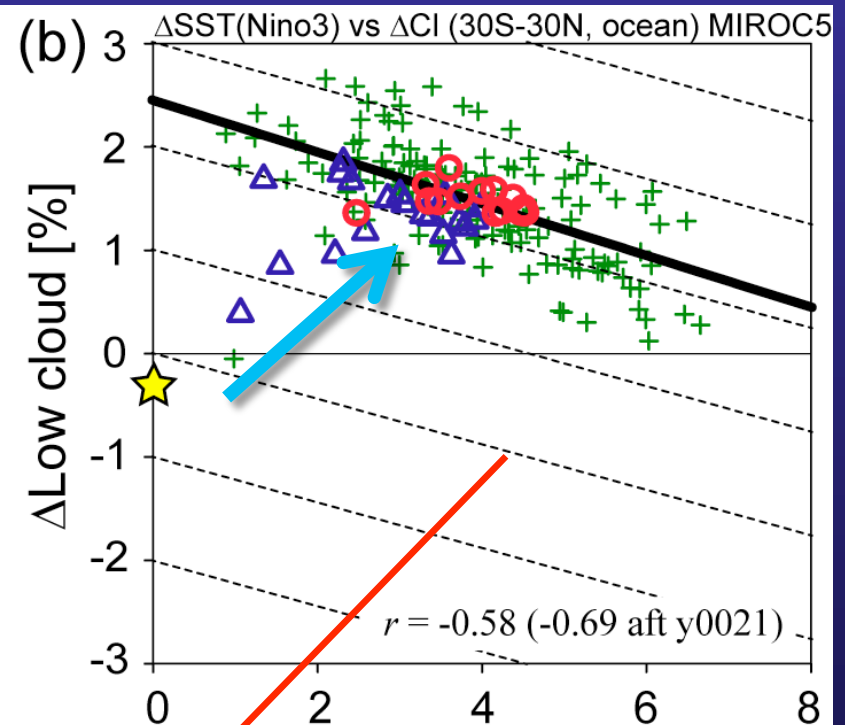
MIROC3

ΔC_l in the tropics (30S-30N)



ΔSST in Niño 3 [K]

MIROC5



- + ann. mean, single run (150y)
- decadal mean, single run
- △ ann. mean, ensemble avg (20y)
- ★ initial month of Δ

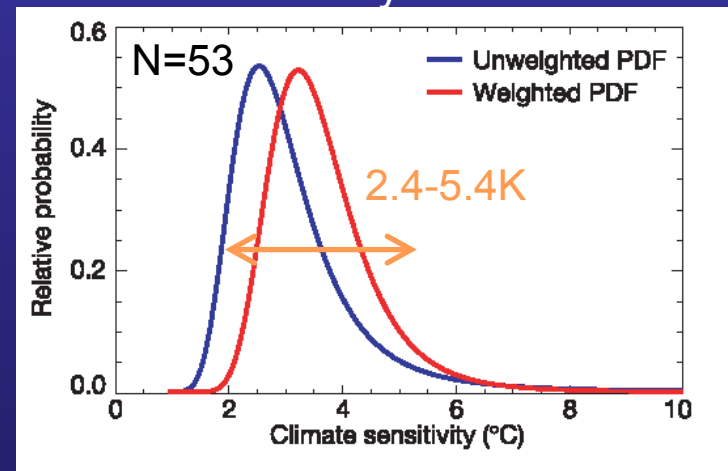
- ✓ Tropospheric adjustment
- ✓ Slow response ($> 20\text{y}$): $(\partial C'_l / \partial T'_s)_{CTL}$ matters!
- ✓ Fast response ($< 10\text{y}$): opposite between the models

Parametric uncertainty

Perturbed Physics Ensemble (PPE)

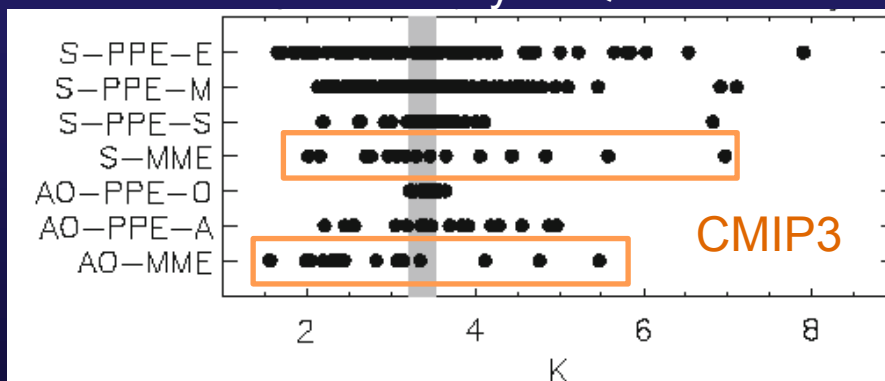
- ✓ Complementary ensemble to CMIP MME
- ✓ Designed for systematic sampling of uncertainty in a single model framework
- ✓ Some advantages over the “ensemble of opportunity”
- ✓ PPEs have been done with
 - HadCM3/HadSM
 - NCAR CAM3.x/4
 - MIROC3.2/5
 - ECHAM5
- ✓ Difference in parameter sets & perturbation strategy

Climate sensitivity in HadSM PPE



Murphy et al. (2004)

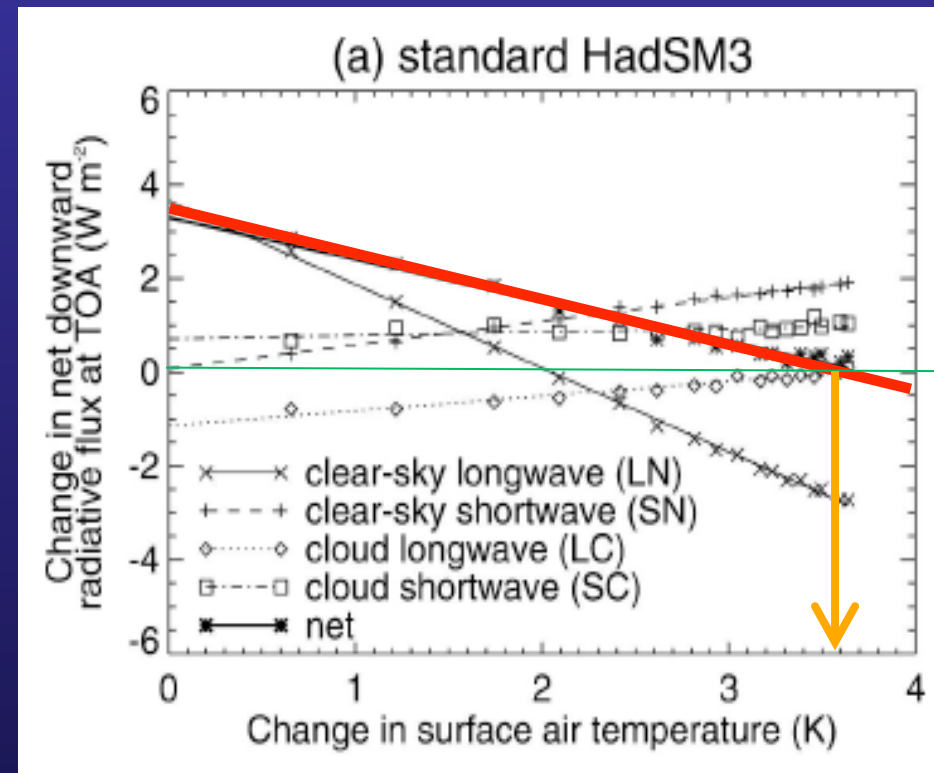
Climate sensitivity in QUMP & CMIP3



Collins et al. (2010)

How to evaluate climate sensitivity?

- In a CGCM $2\times\text{CO}_2$ experiment ($\Rightarrow 20\text{y}$),
 $N = F + \alpha \Delta T$
and plot N against ΔT
- Assuming constant F and α on short timescale
- N can be decomposed into SW, LW and their clear-sky, cloud components

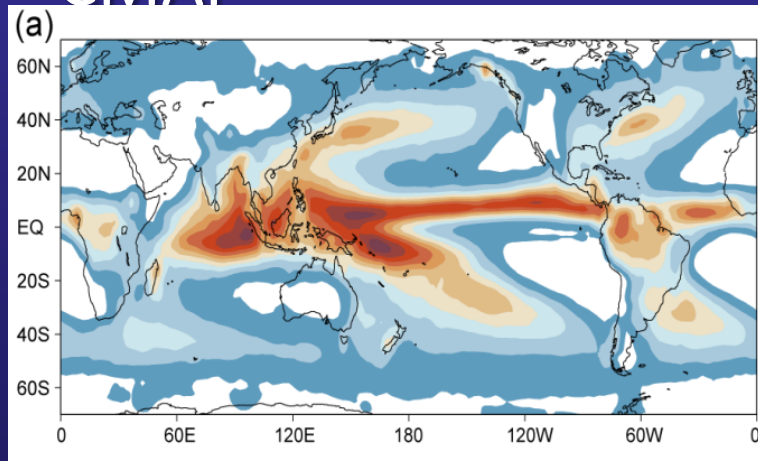


Gregory and Webb (2008)

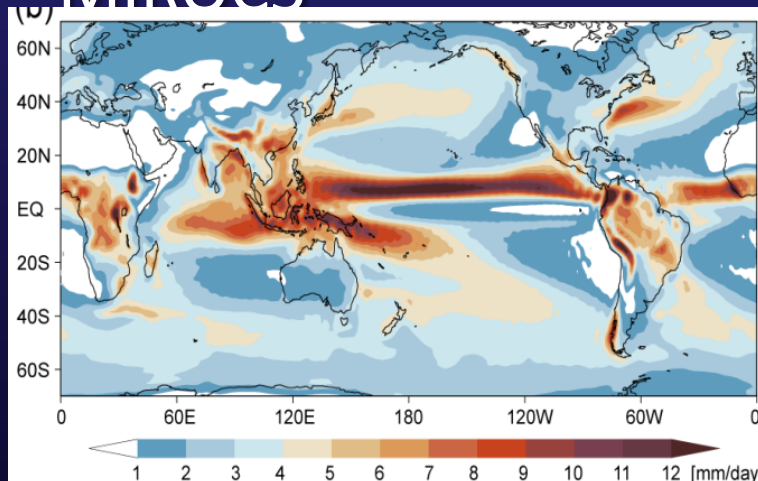
- **Cloud component is surprising: F_{LC} and F_{SC} are non-zero!**
- This caused by rapid local changes to the vertical temperature profile due to the altered radiative heating, with consequent changes to stability, vertical mixing, and the moisture profile.
- **Cloud contribution to F = “tropospheric adjustment” of CO_2 forcing**

Mean climatology

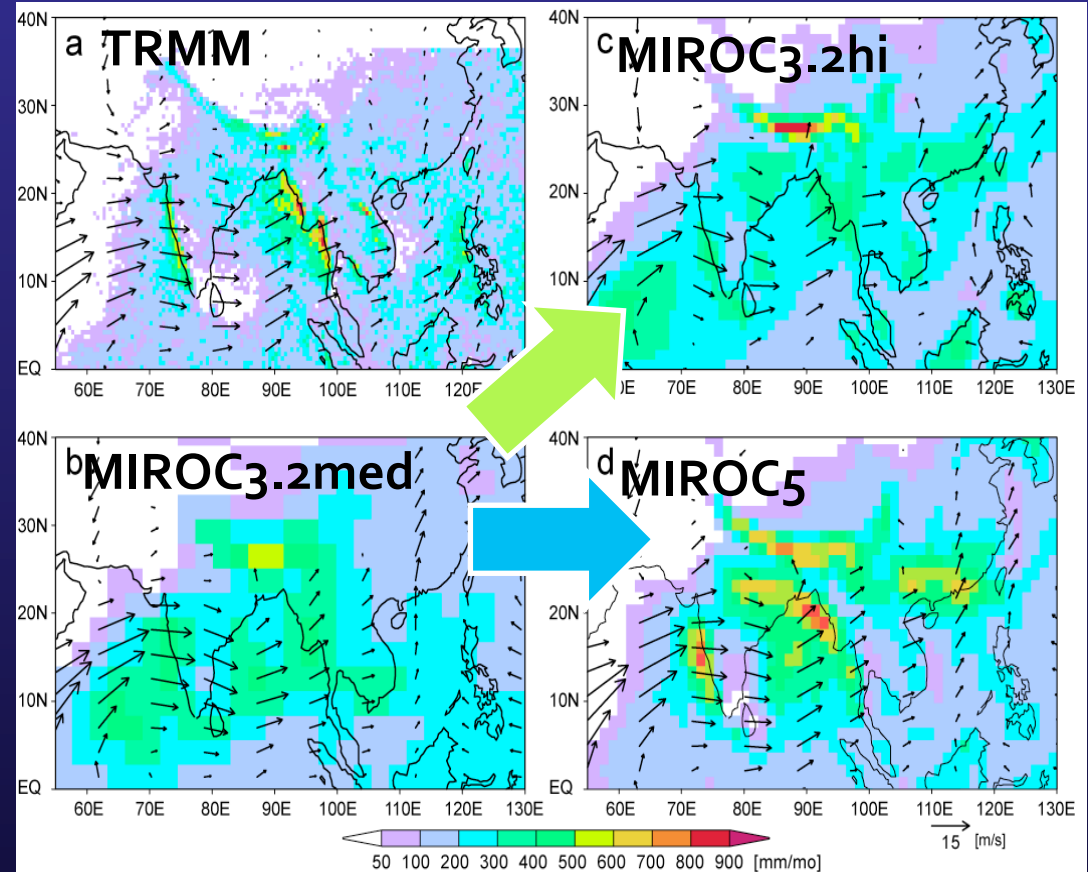
Annual mean precipitation
CMAP



MIROC5



JJA precipitation



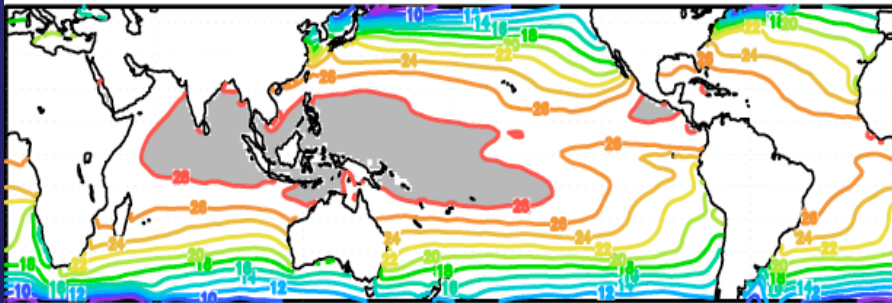
- ✓ Effect of increased resolution
- ✓ Effect of the new model physics

Watanabe et al. (2010, JC)

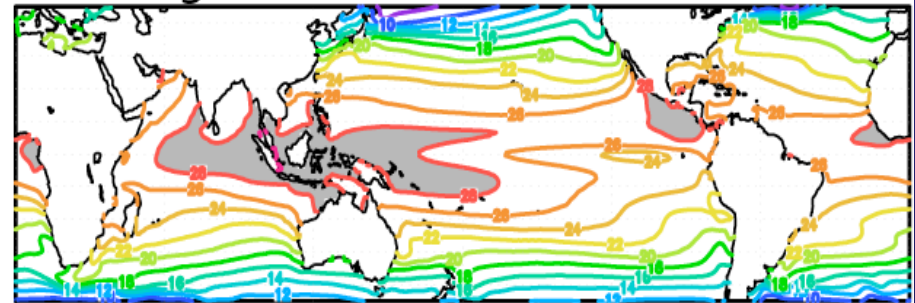
20C mean SST

1961-1990 annual avg. from 20C run

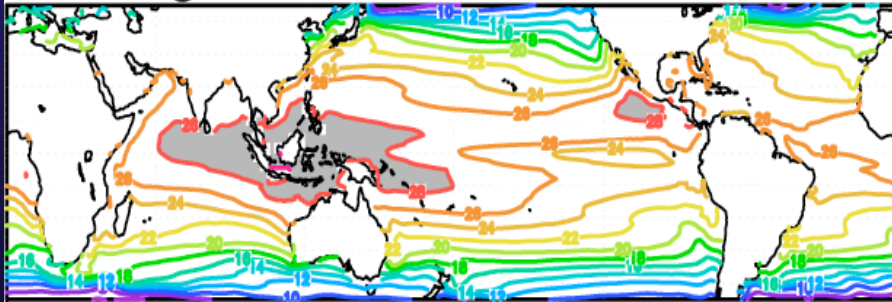
OBS



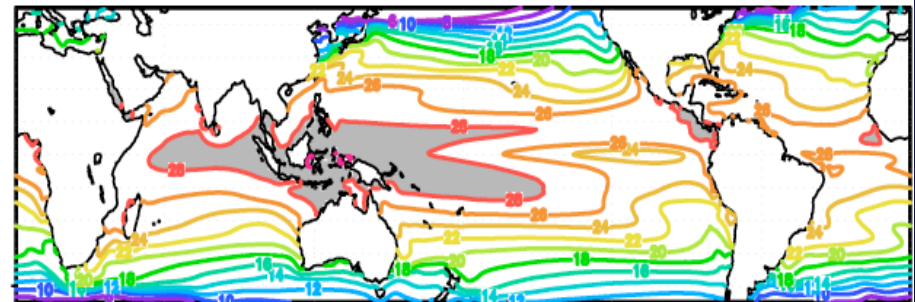
MIROC5



MIROC3med

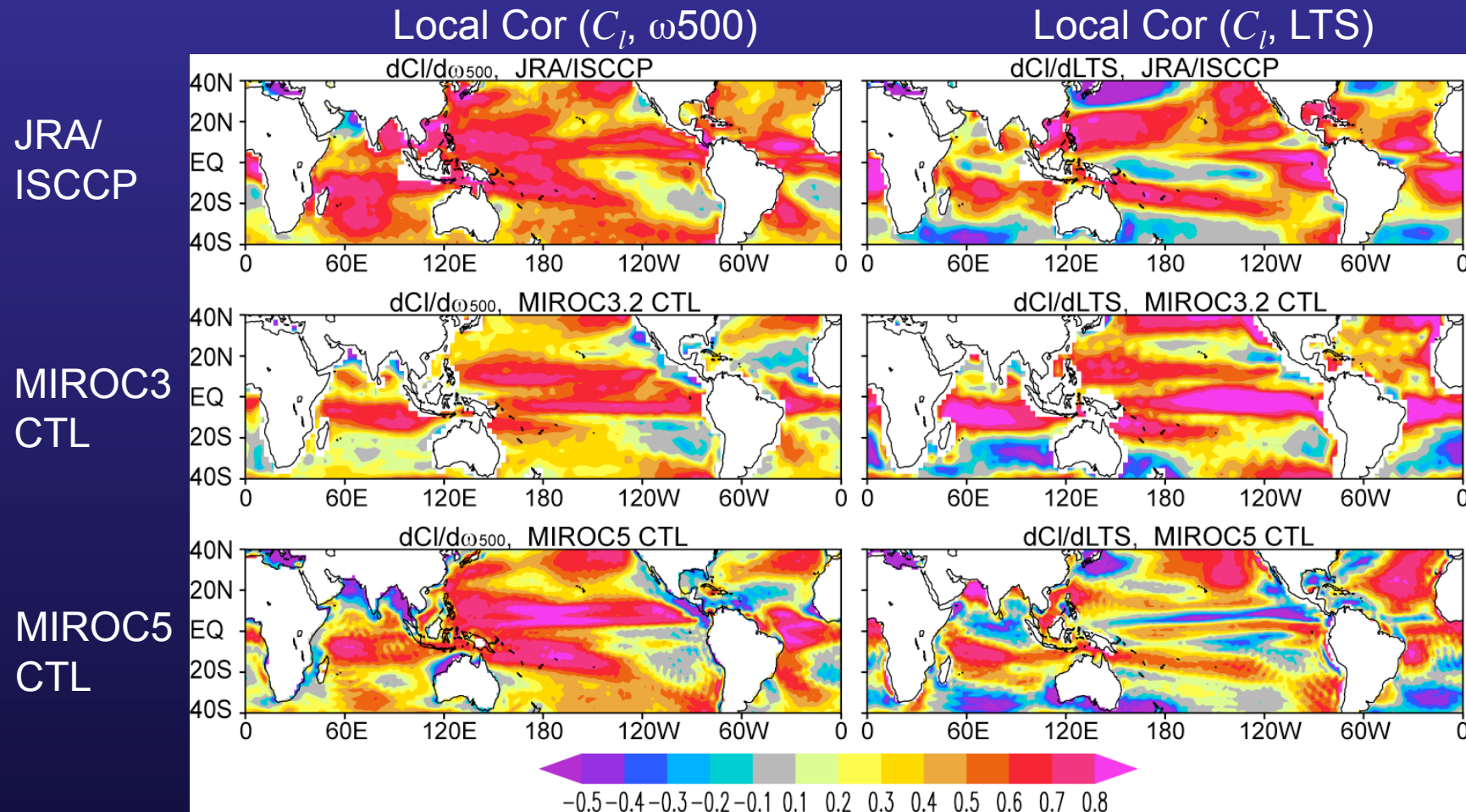


GFDL CM2.0



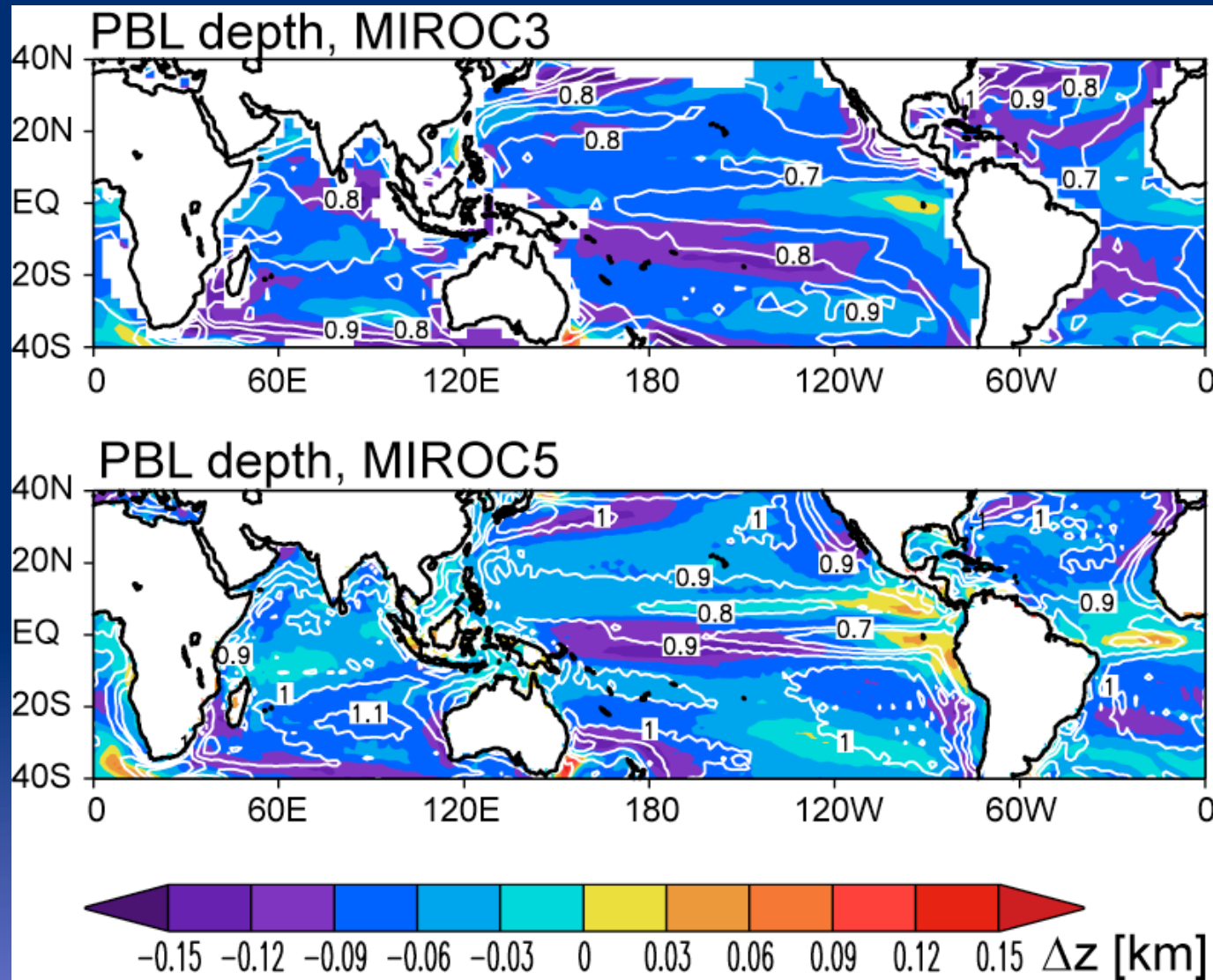
Natural low-cloud variability

Obs/CTL runs



Fast response of low clouds

Change in PBL depth



Fast response of low clouds

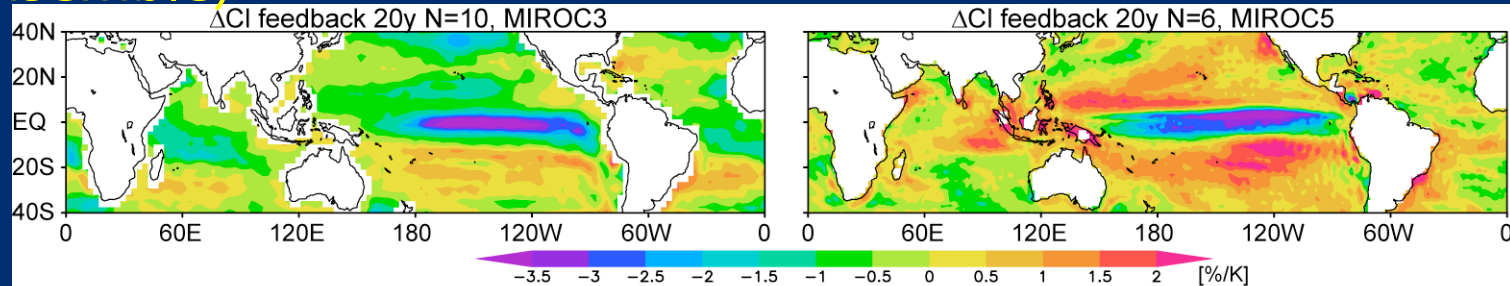
4xCO₂ runs

(20y ensemble)

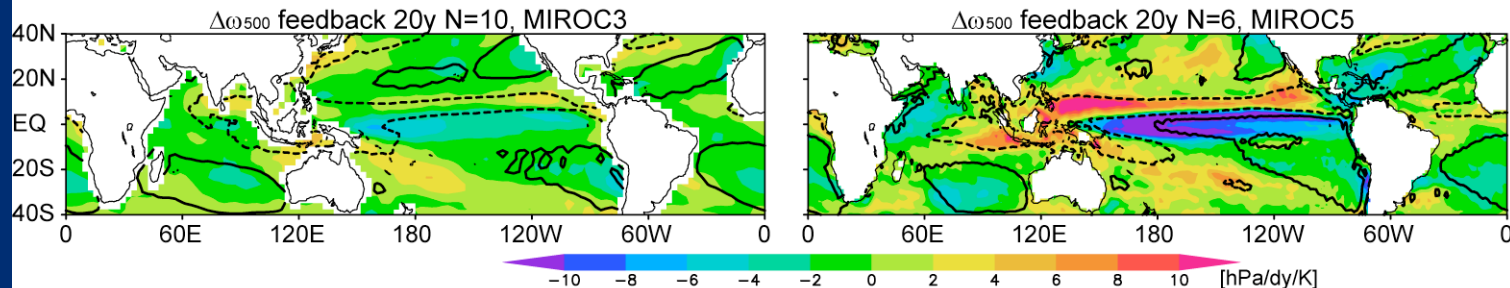
MIROC₃

MIROC₅

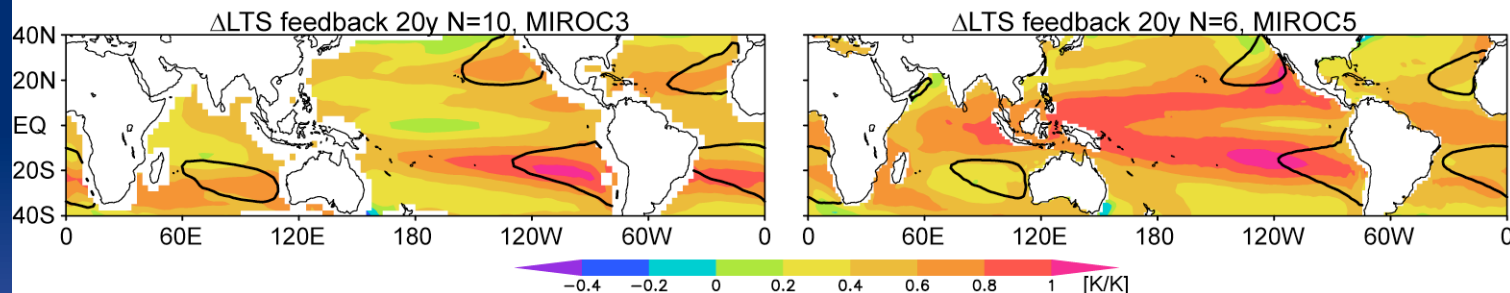
ΔCI



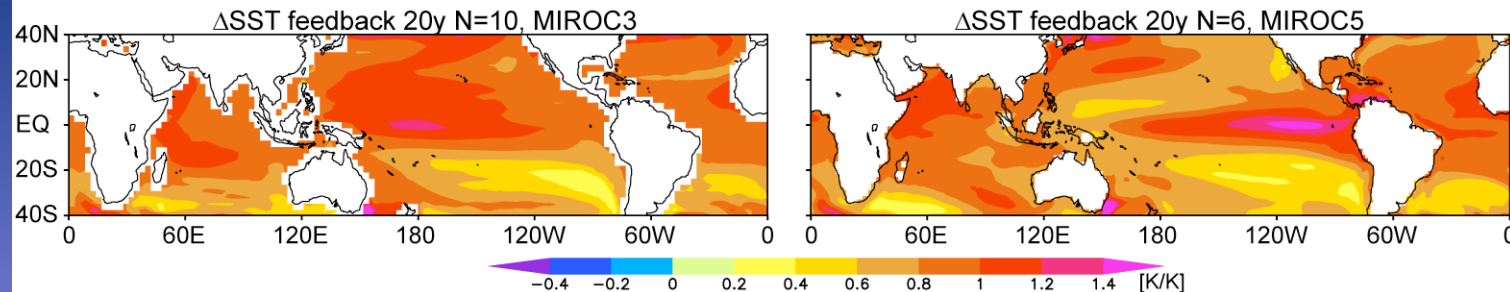
$\Delta \omega_{500}$



ΔLTS



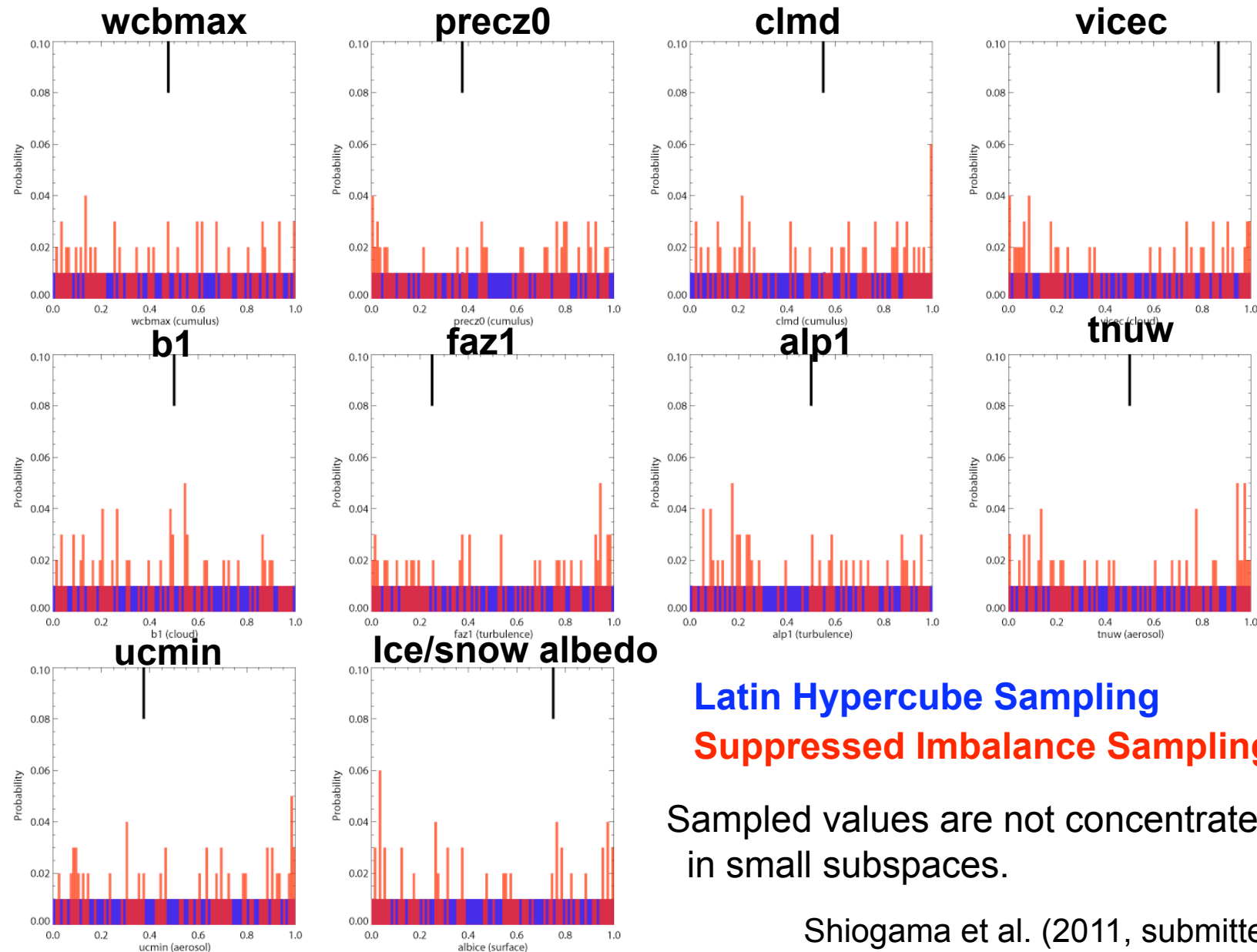
ΔSST



MIROC5 PPE: How to construct?

1. AGCM-CTL runs w/ min/max values of each parameter, estimating changes in the TOA net radiation imbalance
2. Generate the potential parameter sets by large Latin hypercube sampling (5000 samples)
3. Statistically emulate TOA imbalance at TOA for each sample by using ATL-CTL runs
4. Select samples w/ smallest TOA imbalance, and a post process eliminating "very close" parameter values.
5. Repeat (4) to select 100 subsets.
6. Perform AOGCM CTL and 4xCO₂ runs w/ selected 100 *low-imbalance* parameter sets without flux adjustment.

MIROC5 PPE: Sampled parameter values

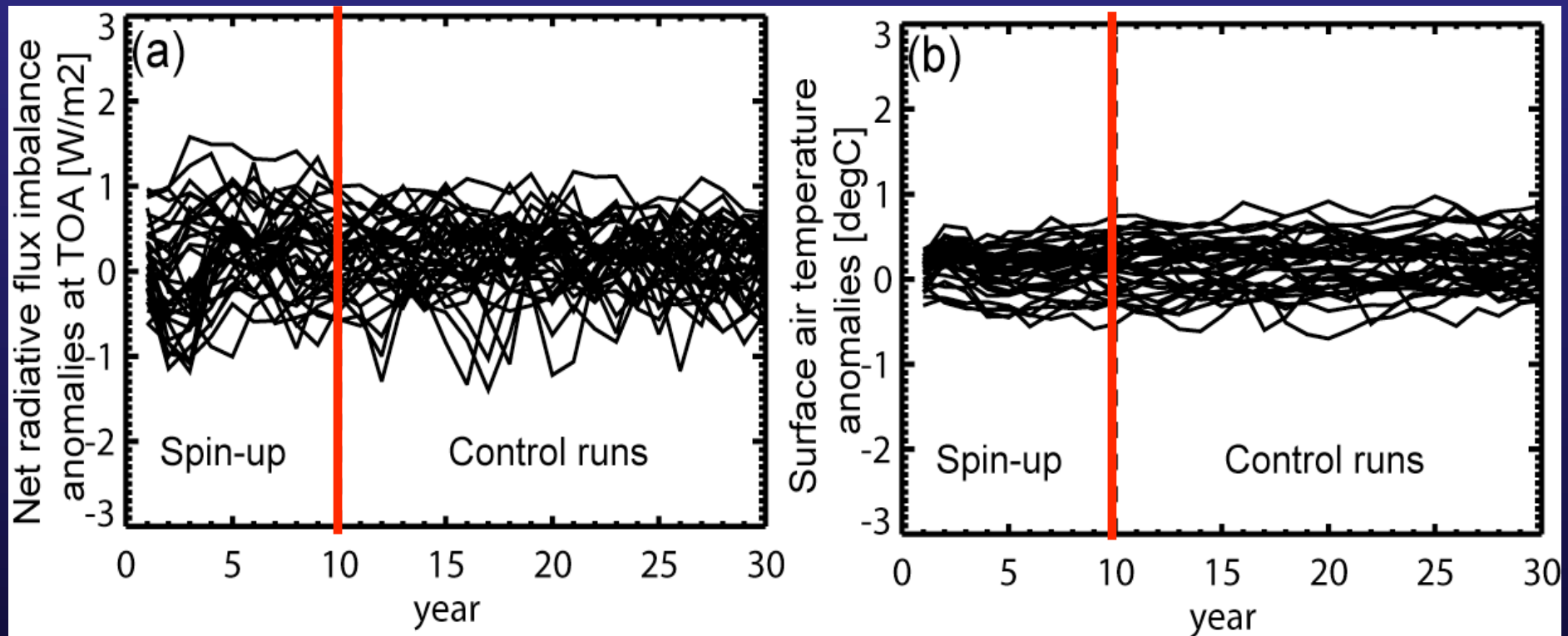


MIROC5 PPE

CGCM-CTL runs (anomalies from the run1=ref)

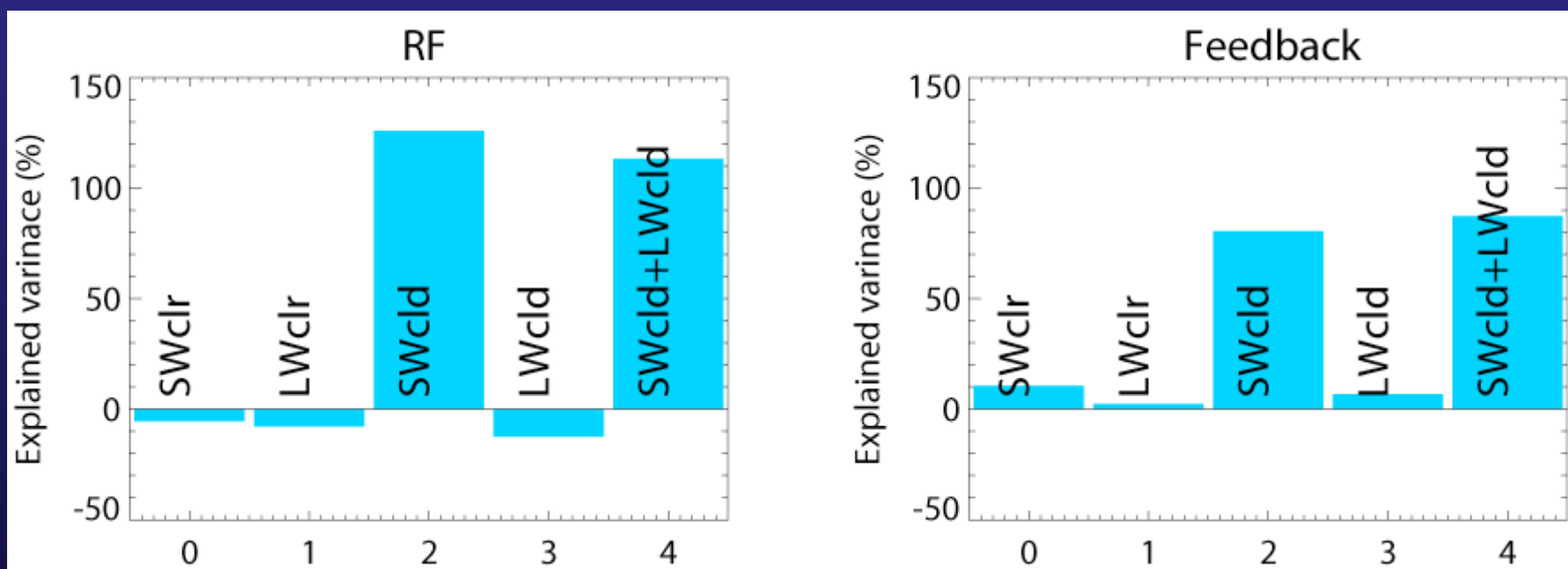
TOA radiative imbalance

Global-mean SAT



MIROC5 PPE

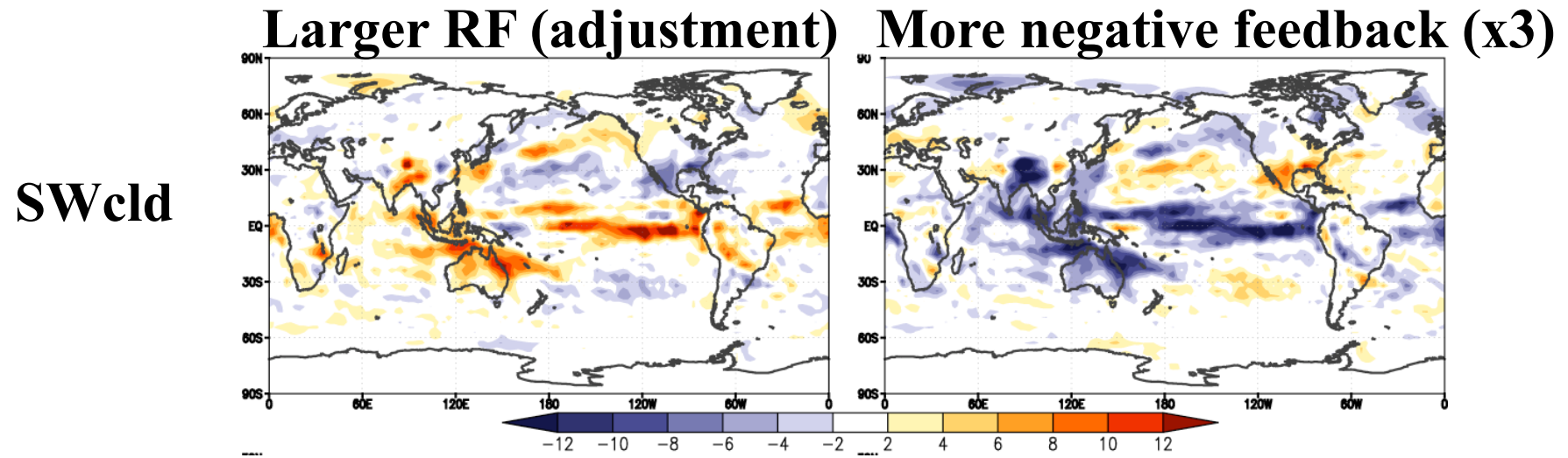
Inter-model variances of total forcing and feedback explained by each component



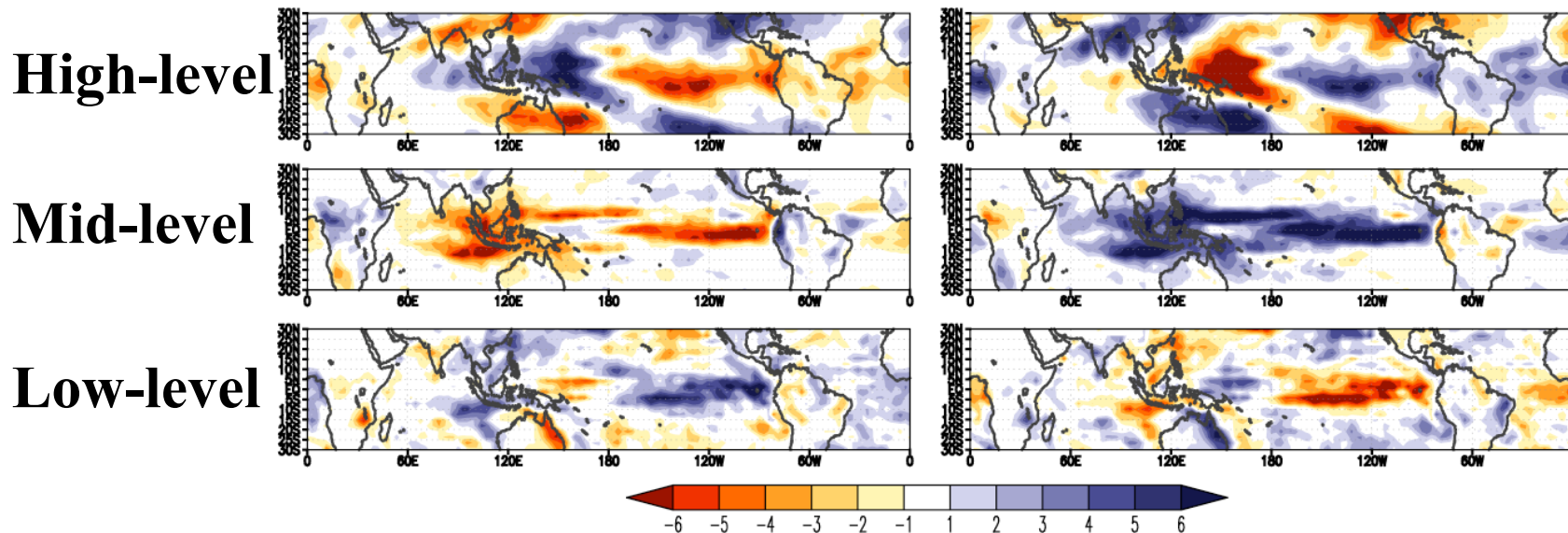
SWcld is the most important component for both the variances of forcing and feedback.

Differences of composite:

(large positive SWcld RF, more negative SWcld FDB) minus
(small positive SWcld RF, less negative SWcld FDB)



Cloud cover



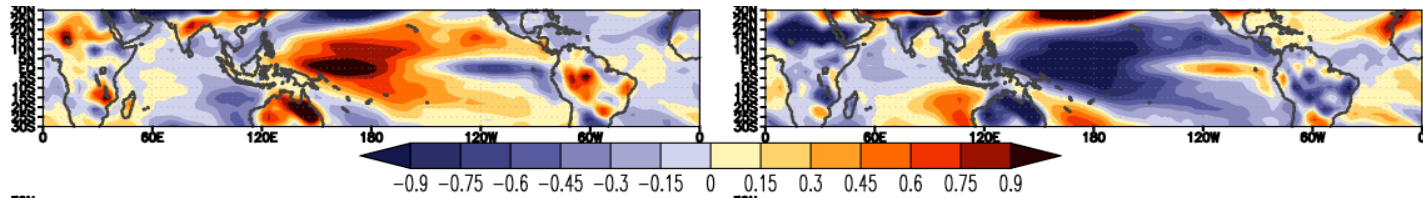
Differences of composite:

(large positive SWcld RF, more negative SWcld FDB) minus
(small positive SWcld RF, less negative SWcld FDB)

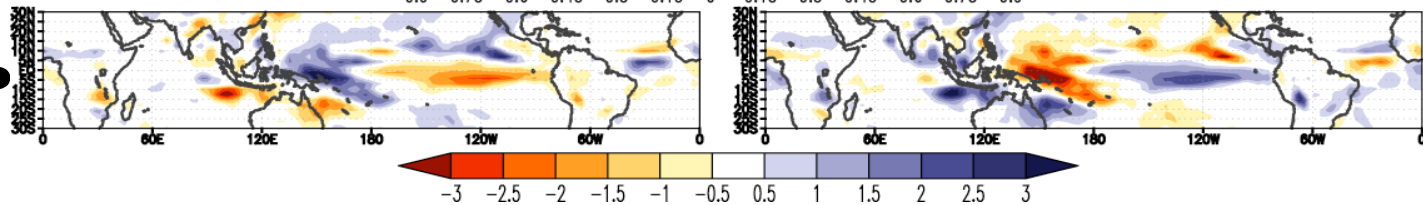
Larger RF (adjustment)

More negative feedback (x3)

Tsfc

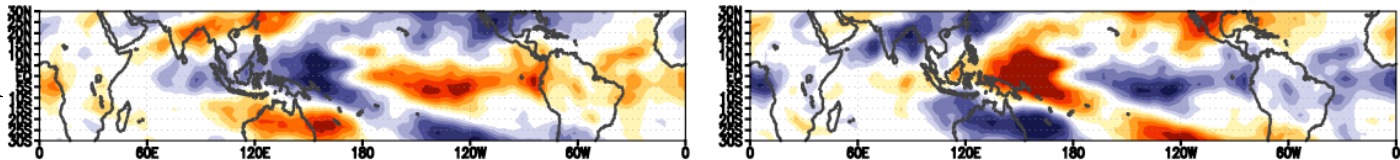


PRCP

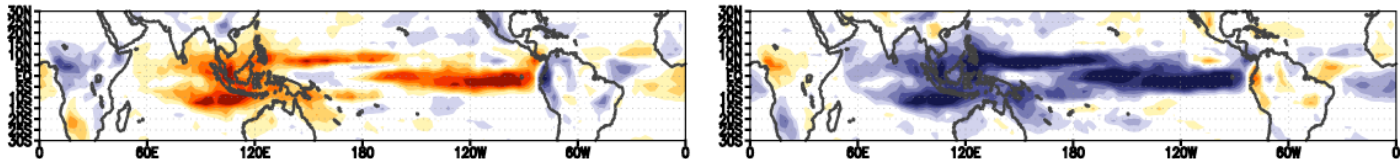


Cloud cover

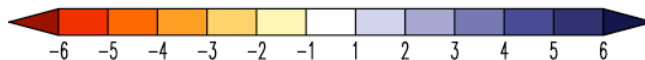
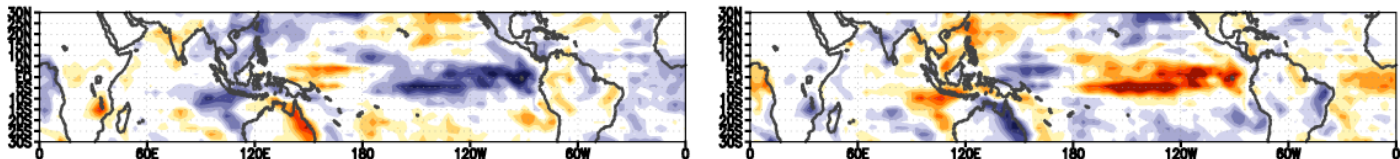
High-level



Mid-level



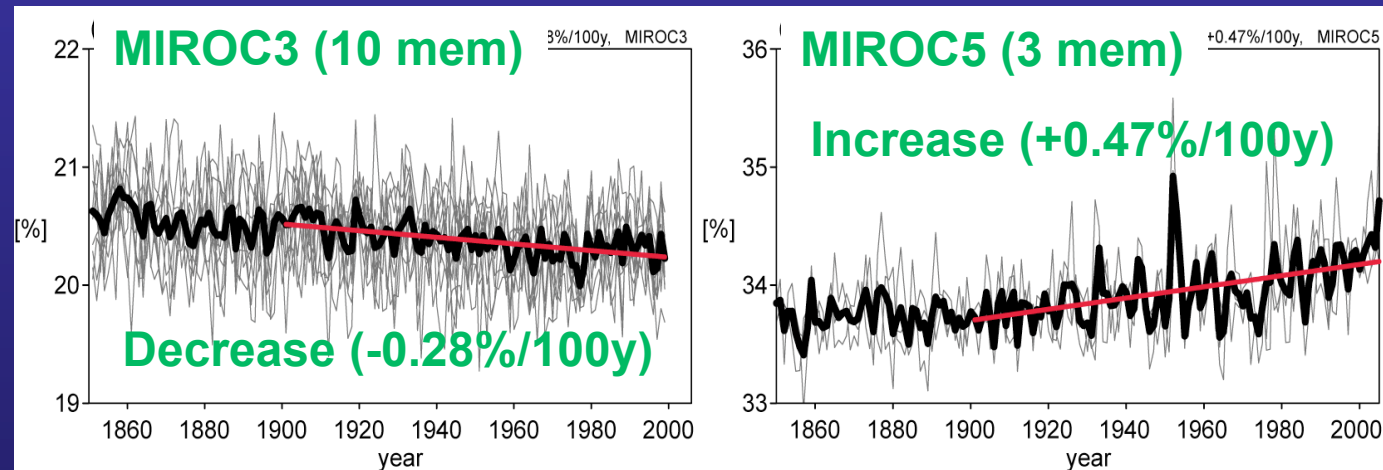
Low-level



Implication to 20th century trends

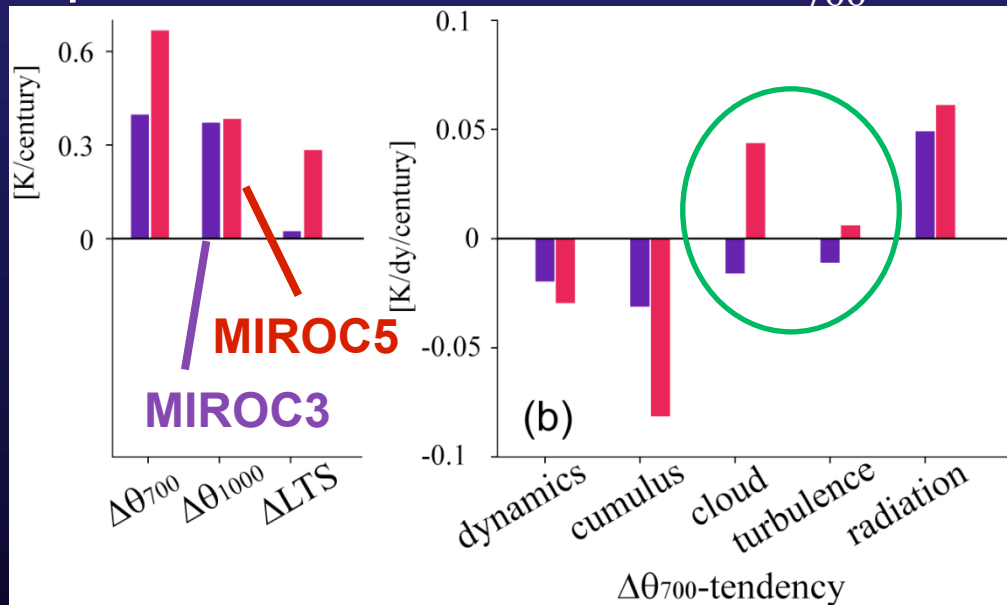
20C runs

Tropical
CI (30S-30N)



✓ Likely due to 'fast' response (but change is much slower)

Tropical-mean trends in θ & θ_{700} -tendency



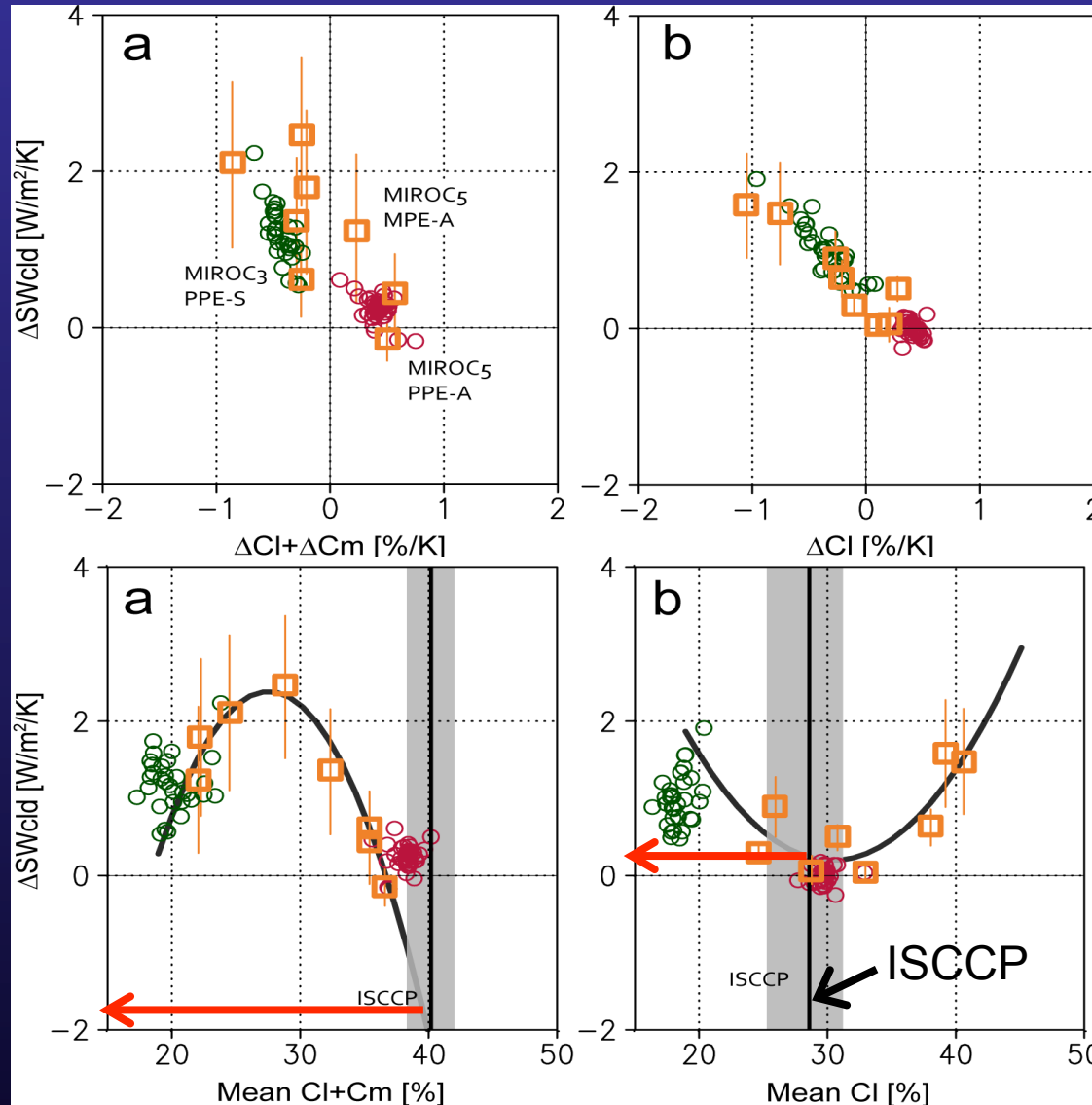
Diff. θ_{700} tendency
due to cloud physics &
turbulence is essential !

Watanabe et al. (2011, ASL)

Can we constrain the cloud feedback?

convective

subsidence



Approx linear relation
in SW_{cld} feedback vs
cloud fraction changes

Nonlinear relation
in SW_{cld} feedback vs
mean cloud fraction

→ can be used for
constraining ΔSW_{cld} ?