Geoengineering efficiency: Preliminary assessment for year 2100 with an energy-balance climate model

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One of the possible solution of global warming is a loading sulphur particles in the stratosphere to enhance the Earth's albedo (Budyko, 1974) recently entitled as geoengineering (Izrael, 2005; Crutzen, 2006; Wigley, 2006; Mokhov and Eliseev, 2008). In the present paper, we estimate geoengineering efficiency by using an energy-balance climate model.

The governing equation for globally averaged model reads

$$C\frac{dT}{dt} = S(1 - \alpha_A(T)) - (A + BT)\eta_C(q) - F_{strat}, \qquad (1)$$

where T is globally and annually averaged surface air temperature, C is heat capacity per unit area, t is time, S is one quarter of the insolation at the top of the atmosphere, α_A is planetary albedo, A and B are coefficients of the linear dependence of outgoing longwave radiation (OLR) on temperature (Budyko, 1974), $\eta_C = 1 - c_0 \ln \frac{q(t)}{a_0}$ is OLR greenhouse correction

factor, q is CO₂ atmospheric content, q_0 is its initial value, $c_0 = 1.4 \cdot 10^{-2}$ (Mokhov, Petukhov, 1978), F_{strat} is mitigation geoengineering forcing. In the linearised setting, this model has a solution which can be represented as a sum of two responses, one due to greenhouse forcing and another due to geoengineering mitigation:

$$T = T_C + T_{strat} . (2)$$

If the buildup of greenhouse gases in the atmosphere (being expressed via "effective CO₂" (IPCC, 2001)) has an exponential form: $q = q_0 \cdot \exp(t/t_p)$, (t_p is time scale of CO₂ atmospheric content change), then the solutions are:

$$T_C = -\frac{\Delta T_{2xCO_2}}{t_p \ln 2} \left(\frac{1}{p} (e^{pt} - 1) - t \right)$$
 (3)

$$T_{strat} = \frac{1}{p} \left(-\frac{a \cdot k_e M_{strat}}{4\pi R^2 C} \right) (e^{pt} - 1) . \tag{4}$$

In Eq.(3), ΔT_{2xCO_2} is equilibrium model's response to the doubling of CO₂ in the atmosphere, and $p = -\frac{S(1-\alpha_A)c_O\ln 2}{C\cdot \Delta T_{2xCO_2}}$. Eq.(4) is obtained assuming that stratospheric

aerosol mass is equilibrated, normaly, $M_{strat} = E_{strat} \cdot t_{strat}$. Here, E_{strat} stands for geoengineering emission and t_{strat} is lifetime of stratospheric aerosols, k_e is a coefficient of extinction of stratospheric sulphur aerosol (it's equal $7.6\text{m}^2/\text{gS}$).

With this model, an ensemble simulation is performed with E_{strat} is varied between 0.6 MtS/yr (Izrael, 2005) up to 5 MtS/yr (including values 1-2 TgS/yr as suggested by Crutzen (2006) and Wigley (2006)) but kept independent on time, t_{strat} is varied in the range 1-4 yr, t_p is varied from 50 yr to 250 yr, and ΔT_{2xCO_2} is varied in the range 1.5-4.5 K which is slightly wider than the range figured in (IPCC, 2007).

Without a geoengineering mitigation, at the end of 21st century temperature changes by 0.5-14.0°C depending on the model parameters (Fig.1).

Fig.1 Temperatures changes in year 2100 without a geoengineering mitigation. The time scales corresponding to the SRES scenarios (taking into account CO₂, CH₄, and N₂O under the "effective CO₂" approximation) (IPCC, 2001) are depicted by horizontal lines.

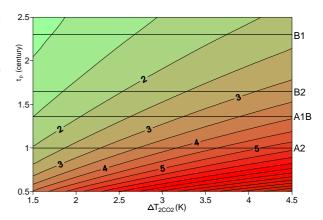
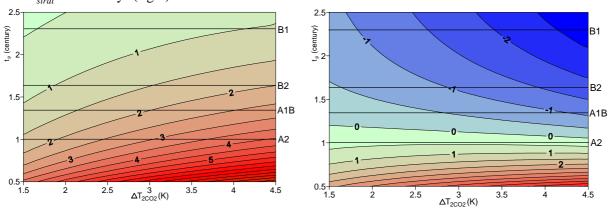


Fig.2 Temperatures changes in year 2100 with a geoengineering mitigation for E_{strat} =1.0 MtS/yr (left) and E_{strat} =4.0 MtS/yr (right).



According to obtained results, it is possible to slow down current anthropogenic warming by applying a geoengineering approach (Fig.2). This mitigation is very efficient (and even excessive) if geoengineering emissions and/or life time of sulphates in the stratosphere are large enough and, additionally, the ${\rm CO_2}$ atmospheric buildup is not too rapid. However, for $E_{\it strat}$ from the lower part of the studied range, the residual warming is still substantional, especially for the scenarios with small t_p (e.g., for the SRES A2 and A1B scenarios).

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