Ocean-Atmosphere Characteristics of Tropical Instability Waves(TIW) Simulated in the NCEP Climate Forecast System Reanalysis

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Impact of TIW on Climate & Marine Ecosystem

- Momentum and heat budget (Jochum and Murtugudde 2006)
- Nutrient distribution (e.g. Strutton et al. 2001)
- Interannual variations, i.e ENSO (e.g. An 2008; Zhang and Busalacchi 2008)
- Mean ocean and atmosphere states

Challenges:

In situ observations

 ---- too sparse

 Satellite Observations

 ---- limited surface variables and record
 NWP reanalysis products (e.g. NCEP R1,R2)

 ----- coarse resolution, SST boundary conditions



<u>CFSR:</u>

- Much higher resolution : 1979-2009 Atmosphere:T382(~38km) Ocean:0.25° ~0.5°
- The guess forecasts was generated from a 6-h coupled system
- Assimilation of ¼ degree daily OI SST Analysis

Purpose: Assess capability of the CFSR in capturing ocean-atmosphere interactions associated with TIWs

Data:

- CFSR : SST, HFLX,τ,P,T2m, WS,q, currents Daily data 2001-2008
- Satellite observations: TMI SST and QuikSCAT winds
- In situ: TAO



0.01 0.02 0.04 0.06 0.08 0.11 0.14

Coupling of SST and Surface Wind



Coupling of SST and Surface Wind



•150° W – 110° W, 2°N, June-January, 2001-2008

•Both SST and wind perturbations are about 70% of the satellite observations.

Coupling Strength (coefficient ms⁻¹ per degree)



•The CFSR well reproduced the observed linear SST-Wind relationship

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Coupling Strength (coefficient ~0.6 ms⁻¹ per degree)



•The CFSR well reproduced the observed linear SST-Wind relationship

Wind Stress Response and Feedback



Does the perturbations of wind stress curl feed back to TIW through Ekman dynamics?

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-0.2 •W_e is much weaker than W, indicating
-0.4 the small feedback effect on TIW growth
-0.6 •Caveat: (1) Difficult to estimate Ekman pumping near the equator; (2) Impact of surface current on wind stress is neglect in the CFSR



3-D Atmosphere Response



Uncertainties in Atmosphere Response



CFSR: ~ 10 Pa °C⁻¹ TAO: ~ 10 Pa°C⁻¹ (Cronin et al. 2003) Radiosonde : little TIW signal (Hashizume et al. 2002)



CFSR: Mainly confined in the PBL Radiosonde: Strong temperature inversion (Hashizume et al. 2002)

Response and Feedback of Surface Heat Flux

Solar heat flux and SST

Satellite Visible Cloud and SST



Deser et al.(1993) :

- <u>4.2%</u> per degree
- Over and upstream of positive SST anomalies
- Change in solar radiation of
- ~ <u>10 W/m² per degree</u>
- •Damping of local SST



CFSR:

- •<u>5%</u> per degree
- Down stream of positive SST anomalies
- Change in solar radiation of
- ~<u>9W/m²</u> per degree
- •Damping of local SST

Response and Feedback of Surface Heat Flux

Latent heat flux and SST





CFSR:

- <u>25 W/m² per ^oC</u> of latent heat flux
 40 W/m² per ^oC of net heat flux, resulting in -0.5 ^o C per month (MLD=50m)
- Zhang and McPhaden(1995): ~ <u>50 W/</u> <u>m² per °C</u> of latent heat flux
 Thum et al. (2002): ~ <u>40 W/m² per °C</u> of latent heat flux, resulting
 -0.5° C per month (MLD=50m)

Summary

An in-phase relationship between SST and surface wind was well represented in the CFSR;

- The CFSR exhibits coherent patterns associated with TIWs both in the ocean and the atmosphere ;
- The feedback of wind stress curl perturbation to TIWs is negligible;
- Uncertainties exist in surface pressure and vertical structure;
- Surface net heat flux are dominated by latent heat fluxes and have a large negative feedback to TIW SSTs variability.

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Annual mean (2001-2008) of (a) EKE (cm2 s-2) at 5m; (b) EKE (cm2 s-2) at 1100 W; (c) EKE (cm2 s-2) along 0oN. In (b) and (c), blue line presents the annual mean 20oC isotherm depth.



Yan et al. 2011