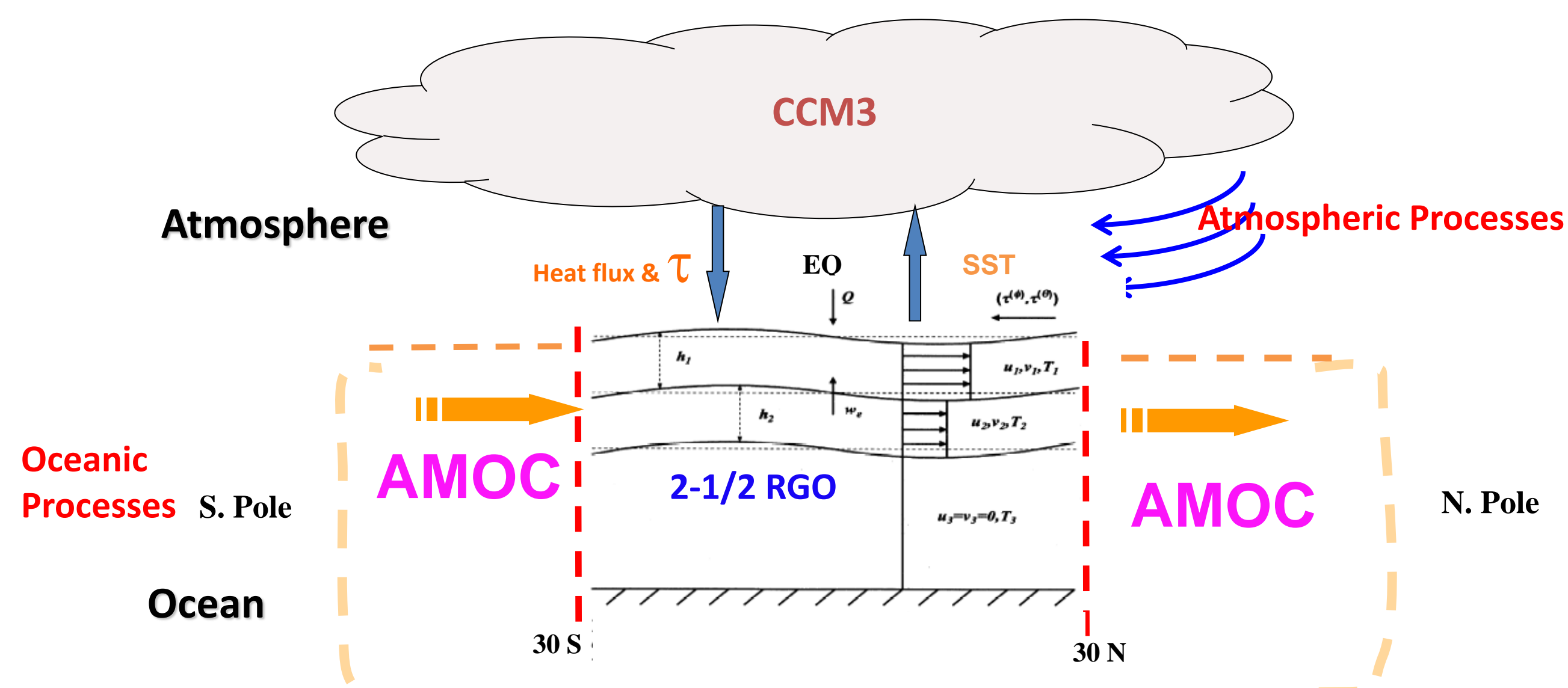


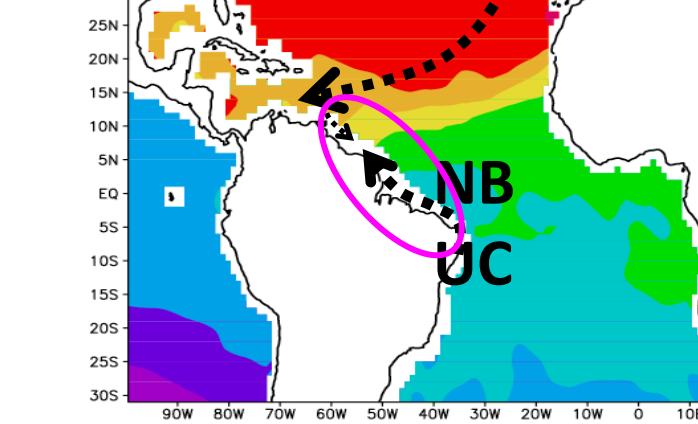
Introduction:

The pronounced dipole like SST pattern and the southward shift of ITCZ are robust climate response to a slow-down of AMOC, coming from both coupled climate model simulations and paleoclimate studies. However, the detailed mechanisms by which the AMOC exerts its influence on TAV remain to be elucidated. The goal of this study is to advance our understanding of the role of AMOC on abrupt climate change, and explore the impact of AMOC on TAV and predictability. Specifically, objectives of this study are (1) identify the relative contribution of the oceanic and atmospheric processes to the response of tropical Atlantic to AMOC changes; (2) explore the impact of AMOC changes on the annual cycle of the Atlantic cold tongue/ITCZ complex; (3) assess the sensitivity of the tropical Atlantic coupled system to changes in AMOC strength.

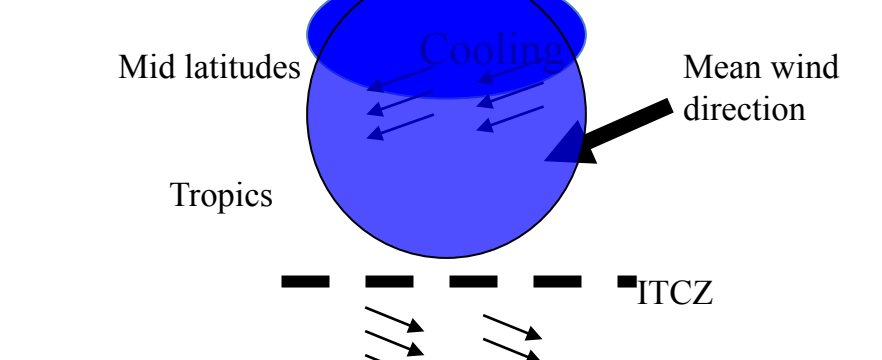
Regional Coupled Model Structure and Experiments



Oceanic gateway (Chang 2008)



Atmospheric Processes (Chiang 2005)



(Mechanisms to be examined)

Experiments:

- To elucidate the relative role of oceanic and atmospheric processes

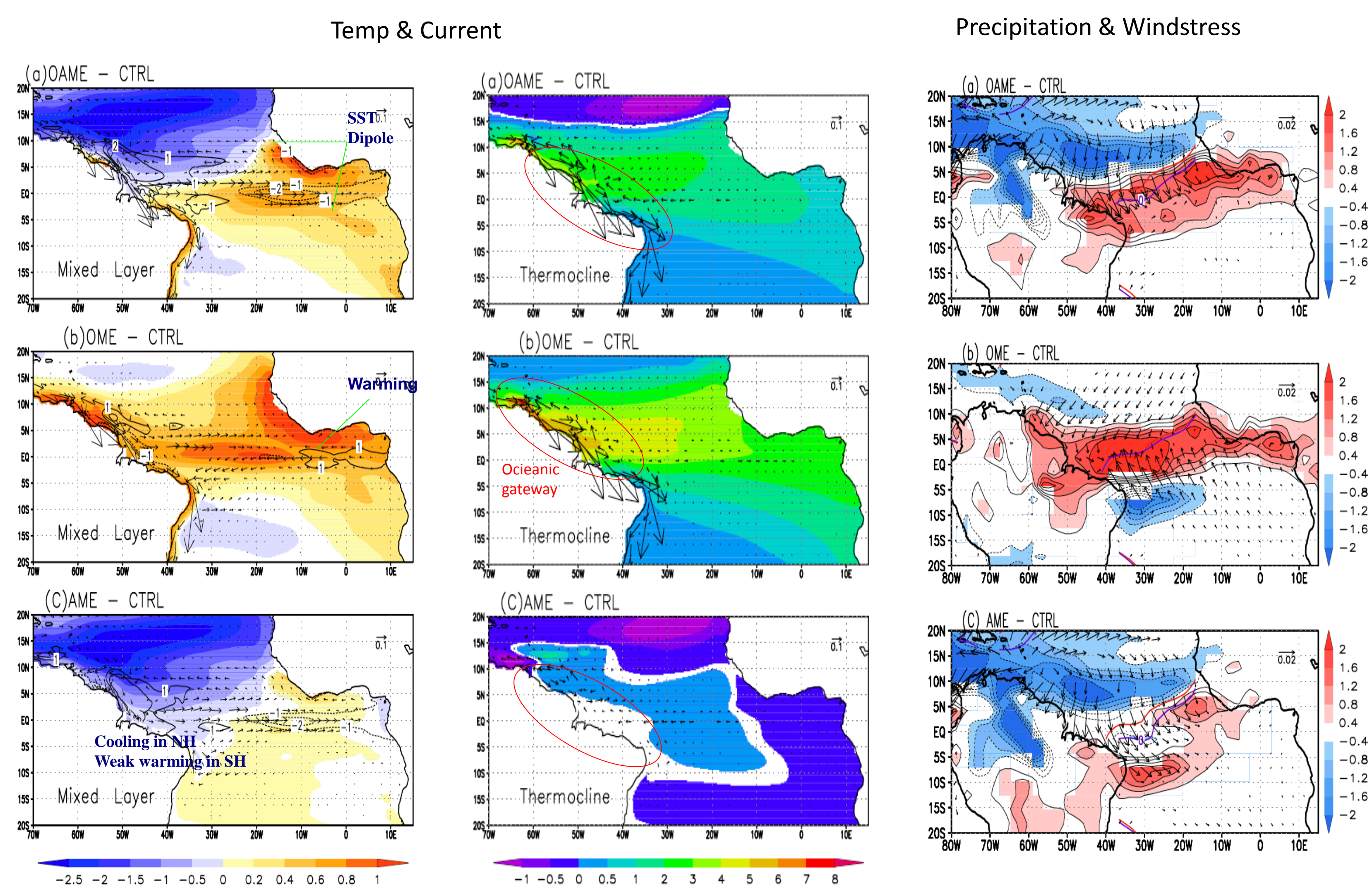
	OAME	OME	AME
Oceanic Processes: (Mass transport set to 0Sv)	X	X	
Atmospheric Processes (heat flux anomalies applied between 10N-30N)	X		X

- To assess the sensitivity of TAV to changes in AMOC strength

--imposed northward mass transport is decreased systematically from 14Sv to 0Sv

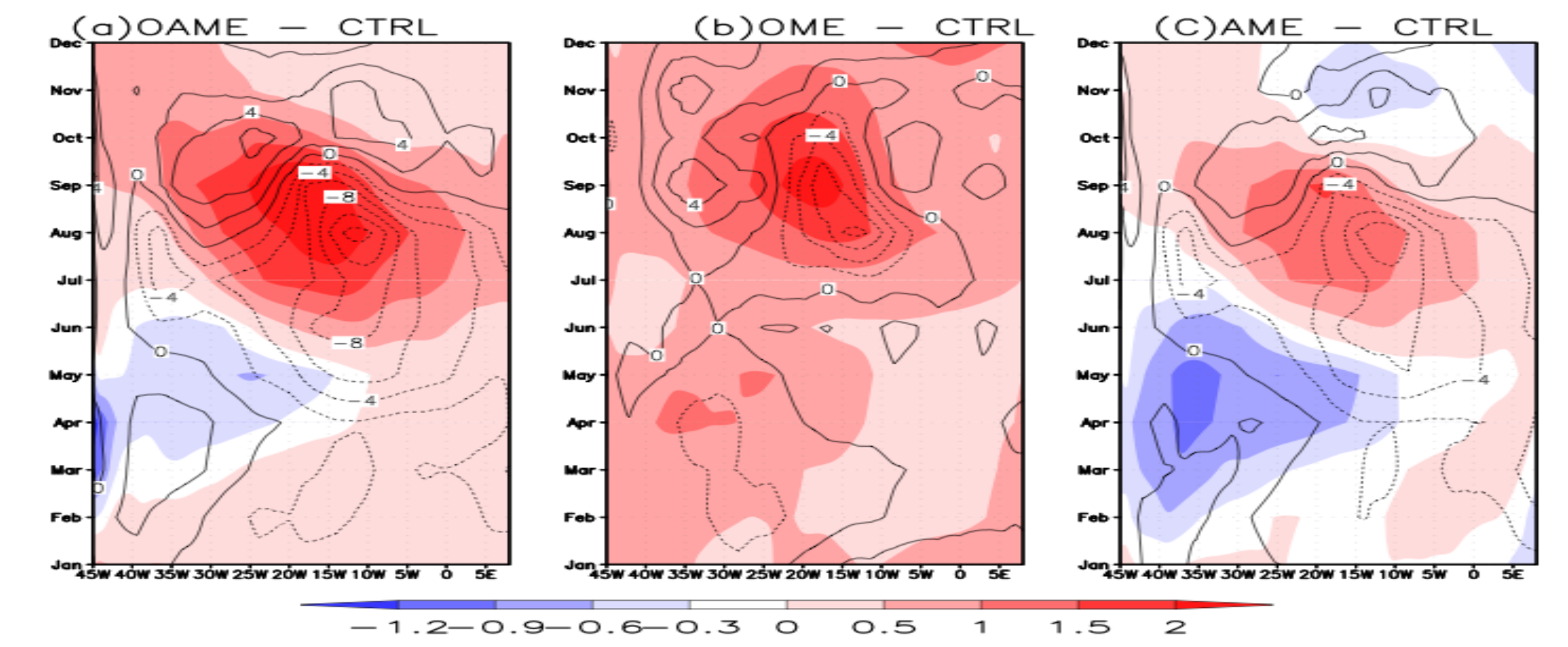
Impact of AMOC changes on TAV

Annual mean response

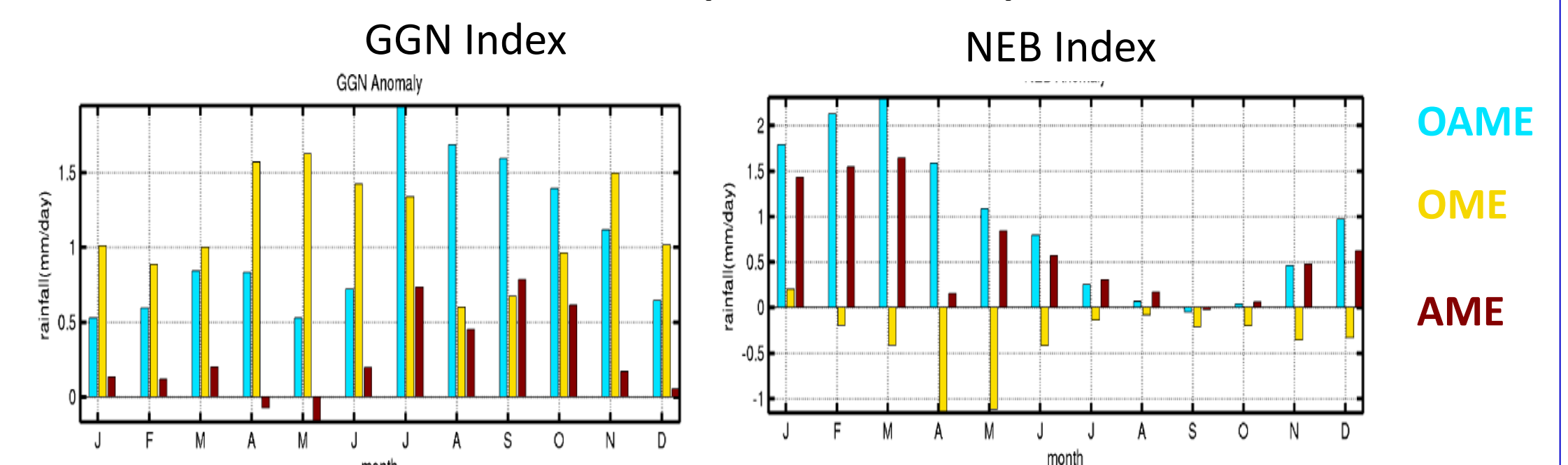


Seasonality of response

SST(shaded) and Entrainment(contour) response along EQ

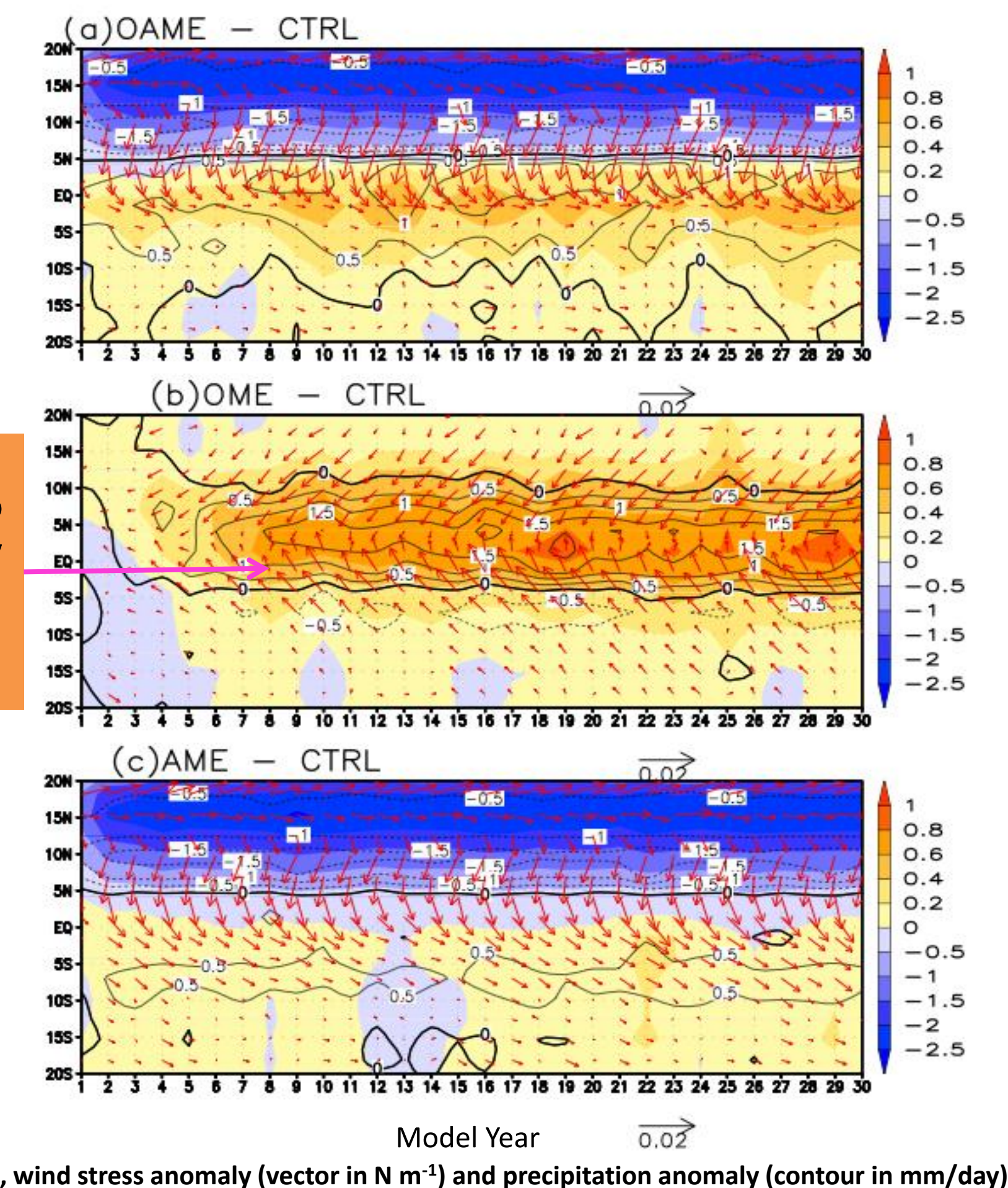


Precipitation response



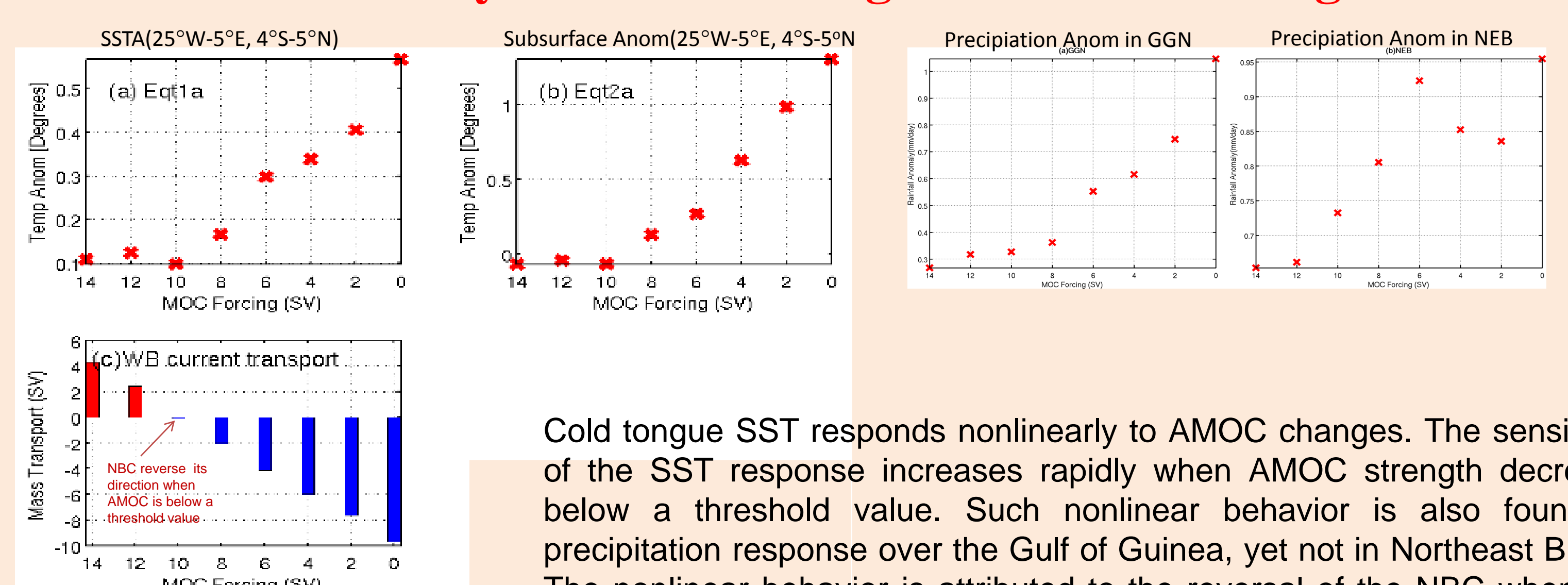
GGN: Precipitation anomaly averaged over Gulf of Guinea (10°W-10°E, 0°-7°N)
NEB: Precipitation anomaly averaged over Northeast Brazil (42°W-35°W, 12°S-4°S)

Evolutions of Zonally Averaged Response



Oceanic processes causes the equatorial warming to lag the AMOC changes by 8yr. This may give rise to certain predictability at decadal time scales.

Sensitivity of TAV to changes in AMOC strength



Cold tongue SST responds nonlinearly to AMOC changes. The sensitivity of the SST response increases rapidly when AMOC strength decrease below a threshold value. Such nonlinear behavior is also found in precipitation response over the Gulf of Guinea, yet not in Northeast Brazil. The nonlinear behavior is attributed to the reversal of the NBC when the western boundary current carrying AMOC return flow becomes weaker than the northern STC return flow.

Conclusions:

- Oceanic processes are a primary factor contributing to the warming at and south of the equator and the precipitation increase over the Gulf of Guinea, while atmospheric processes are responsible for the surface cooling of the tropical north Atlantic and southward displacement of ITCZ.
- The response of tropical Atlantic to AMOC changes exhibits a pronounced seasonality. The atmospheric processes and the oceanic processes interact in a complex manner.
- Equatorial warming lags the AMOC change by about 8 years. This delayed response may give rise to certain predictability of equatorial SST in response to AMOC changes at decadal time scales.
- SST over the cold tongue region responds nonlinearly to AMOC changes. Similar behavior is also found in precipitation response over the Gulf of Guinea. These nonlinear behaviors are attributed to the oceanic gateway mechanism.

Reference:

- Wen, Caihong, Ping Chang, Ramalingam Saravanan. (2011) Effect of Atlantic Meridional Overturning Circulation on Tropical Atlantic Variability: A Regional Coupled Model Study. J. Climate, 24, 3323-3343
 Wen, Caihong, Ping Chang, Ramalingam Saravanan. (2010) Effect of Atlantic Meridional Overturning Circulation Changes on Tropical Atlantic Sea Surface Temperature Variability: A 2½-Layer Reduced-Gravity Ocean Model Study. J. Climate, 23, 312-332.
 Chang et al., 2008: Oceanic link between abrupt changes in the North Atlantic Ocean and the African monsoon. Nature Geoscience, 1, 444.
 Chiang, J.C.H. and C.M. Bitz, 2005: Influence of high latitude ice cover on the marine Intertropical Convergence Zone. Climate Dynamics, 25, 477-496.