Atmospheric Blocking in the CMCC CMIP5 simulations



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Atmospheric blocking describes a mid-latitude weather pattern where a quasi-stationary high-pressure system modifies the westerly flow, "blocking" (or at least diverting) the eastward movement of the migratory cyclones (Rex, 1950a).



Blocking occurs throughout the year, even if it is more frequent during winter and spring, and it typically develops of the underlying at the end of the Pacific and Atlantic eddy jet streams, and it affects significantly the weather regions, sometimes leading to cold spells in winter and heat waves in summer (Trigo et al. 2004, Buehler et al., 2010).

Correct detection and forecast of blocking is still an open issue for both climate and NWP models: the underestimation of the blocking frequency, especially the one occurring over the Euro-Atlantic sector, is generally accepted as a common: these biases usually lead to significant errors in the representation and prediction of temperature and precipitation patterns over Europe. Moreover, blocking appears to be connected to the many different components of the climate system, as Sudden Stratospheric Warmings (Martius et al. 2009) and the North Atlantic Oscillation (Woollings et al, 2008)

This work goes in the direction of exploiting the CMCC model simulations in order to get insight of the interconnection between blocking and other components of the climate system, and, in the same time, trying to establish what is the cause behind the biases in blocking representation.

2. Data and Model simulations 3. Blocking Detection Historical simulation of CMCC-CM CMCC-CMS CMCC-CESM For the blocking detection is adopted the bidimensional extension of the three different OPA 8.2/LIM 2°x2° - 31 lev OPA 8.2/LIM OPA 8.2/LIM Tibaldi and Molteni (1990) index. coupled versions of 2°x2° - 31 lev ECHAM5 2°x2° - 31 lev ECHAM5 2°x2° - 31 levels ECHAM5 Blocking events are recognized when the inversion of the meridional the CMCC model has Atmospheri gradient of geopotential height at 500 hPa occurs. An additional constrain **î** ĵ been evaluated in order T159 (0.8 deg) Atm. Hor. Re T63 (1.8 deg) T31 (3.75 deg) applied northward to exclude fake blocking events. to compare the results COUPLER OASIS3 Blocking index is calculated from 35° N to 75° N: minimum spatial Atm. Ver. Resolution 31 levels 95 levels with NCEP / NCAR Model Ton 10 hPa 0.01 hPa 0.01 hPa dimension (15° lon), quasi-stationarity and temporal persistence (5 11 00-1 bPa 44 les reanalysis. Two of them have a days) are applied. full resolved Cathon NO NO VES Terres Cycle $GHGS = \left[\frac{Z(\phi_o) - Z(\phi_o)}{\phi_o - \phi_o}\right]$ stratosphere. N HI $GHGN = \left[\frac{Z(\phi_n) - Z(\phi_o)}{\phi_n - \phi_n}\right]$ 4. NCEP/NCAR Reanalysis Blocking Climatology (1) GHGS > 0(2) GHGN < -10 m/deg latitude</p> Blocking frequency was calculated for North Pacific the winter season (DJF) from 1951 5. CMCC Models Blocking Climatology Blocking to 2005. CMCC-CMS смсс-см Greenland CMCC-CEMS The two classical area of blocking over T63L95 T159L31 Blocking T31L39 Pacific and Atlantic basin emerges as totally different. Figure 1: Blocking frequencies in DJF NCEP/NCAR Reanalysis 1950-2005 as shown by bidimensional extension of the TM90 index Black contours shows the position of low tropospheric jet streams (U850 more than 8 m/s). European Blocking 6. New Blocking Diagnostics ency for cyclonic (left) and anti ng for DJF Blocking Intensity Index: a measure of how the meridional circulation is perturbed, a 2D extension of the index developed by Wiedenmann et al. (2002). Meridional Gradient Index: a measure of the intensity of the easterlies associated with the blocking. Average duration of the blocking events Wave Breaking Index: provide information of whether the blocking is associated to cyclonic or anticyclonic Rossby Wave Breaking. CMCC models underestimate the bulk of events on **Europe** (Figure 2), while there is agreement representation of the Greenland blocking. anomalies with respe to NCEP/NCAR The diagnostics suggest dynamical differences eanalysis (lower inels) in the 3 CMCC Biases in the climatological jet streams are tightly (in terms of duration, wave breaking, intensity and linked to the frequency of mid-latitude blocking een High Latitude Blocking and Mid-, more) bet detected I atitude Blocking **Conclusions** Models replicate the main properties of the 3: From left to right, Blocking Instensity index, Meridional Gradient sal and Duration as computed for the DJF 1951-2005 of NCEP/NCAR blocking events but duration, suggesting that the bias is related mainly to persistence and frequency. The adoption of new bidimensional index allows a more detailed study of different physical featurse of blocking events: • Pacific and Greenland blocking shows similar diagnostics (Fig. 3 and 4) and can be considered as high-latitude blocking (HLB) in agreement to the definition by Berrisford et al. (2007). On the other side, European Blocking 7. Blocking and Atlantic Jet Displacements behaves differently than HLB, showing that just the European one can be considered as a "real" blocking events and suggesting the possibility that Greenland Iberian Wave Breaking (IWB) Euro-Atlantic Blocking appears to be strictly linked to the Atlantic Europe jet displacements, studied via the Jet Latitude Index (JLI) introduced 50W-20W 30N-40N 15W-20E 70W-20W 47.5N-62.5N 62.5N-72.5N 30W-10W 37.5N-47.5N by Woollings et al. (2010). they are governed by different dynamics. The Euro-Atlantic winter variability. expressed via the JLI (Fig. 5 and 6). CEP.let Latitude Index (.II.I) Southern jet displacements are linked with Greenland blocking, can be associated with the occurrence/absence of blocking: this is confirmed by CMCC model bias, even if the impact on the jet is more clear for Greenland blocking than for European blocking. confirming the NAO/Blocking hypothesis of Woollings et al. (2008) 8 Greetland MR No Slock Northward jet CMCC models (Fig. 2) show general underestimation of European blocking, even if they possess good representation of the North Atlantic lex (JLI displacements are linked with blocking Oscillation. 80 LI Europe Orientard WB No Block occurring over the As European Blocking appears to be not correlated to the phase of the IAO, studying blocking events is a key element to capture the variability over 8 southern side of the jet. NAO Days 100 50 Latitude ė. the Euro-Atlantic region in climate simulations. Models bias Figure 5: PDF for Jet Latitude Index (JLI, black line) as introduced by Woollings et al. (2010) for NCEP/NCAR Reanalysis in DJF 1951-2005. Blue, green and yellow line represent JLI when blocking is occurring over Greenland, Europe and IWB sector respectively. Dotted line represents JLI distribution when no blocking is occurring in the 3 sectors. Main References strengthen this idea, main Références erratord, P., B. Hoskins, and E. Tyrlis, 2007: Biocking and Rossby Wave Breaking on the Dynamical Tropopa Hemisphera. Journal of the Atmospheric Science, 64, 248-2486. Scherrer et al. Two-dimensional indices of atmospheric blocking and their statistical relationship with winter clin Atlantic region. In J. Cirradal. (2004) vol. 26 (2) pp. 233-2409 Tabal 5. and F. Moteni. On the operational predictability of blocking and Atlanter Region. In Aver Moting et al. A New Rossby Wave-Streaking Interpretation of the North Atlantic Coscillation. Journal of the Atl Woollings, T. A. Ipareadul, and B. Hoskins. 2010: Venation suggesting that blocking acts 40 60 onship with winter climate patterns in the Euro perturbing jet from As Figure 5, but for CMCC-CMS Dashed line is the .II I for NCEP/NCAR the its average "non

perturbed" state