stratosphere (Cohen et al., 2007; Saito et al., 2001; Orsolini and Kvamstø, 2011)

high northern latitudes

- season, to "snow"



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Surface pressure differences at one-month lead: Ensemble-mean difference in SLP (Series1-Series2) through the second forecast month (color shading) for OCT 15 (left) and DEC 1 (right). Contours denote the ensemble mean SLP for Series 1 (hPa). Indicative of circulation changes : high snow \rightarrow deeper Aleutian Low, stonger Siberian High, lower SLP over Arctic

• Forecast skill increment for T_{2m}. Forecast skill

increment (Series1-Series2) r² defined as the square of the anomaly correlation coefficient (ACC) between observation and ensemble-mean forecast. Each r² is multiplied by the sign of the ACC. Model anomalies are deviations from model climatologies valid for each start date and lead time. Observational anomalies are taken from ERAINT climatology over the years 2004-2009.

- Initial (0 lead) weak positive increment over snowcovered land
- ■Very large (~0.7) increment over Arctic at 30-day lead

Note: GLACE2 \rightarrow soil moisture skill increment ~0.3 Teleconnection influence 30-day lag consistent with remote forcing through planetary wave propagation (Saito et al., 2001; Cohen et al., 2007) Note: a downward stratospheric influence probably cannot be seen in our 2-month forecasts

Snow pack initialisation has great potential to improve forecast skill in surface temperature over the Arctic, at monthly lead times. The patchiness of skill increment remains to be explained

Heavy snow pack has cooling effect on lower atmosphere, decoupling the lower atmosphere from the soil layer below. Snow albedo feedback has limited role due to low insolation in autumn

Our high-resolution coupled forecasts partly confirm results from earlier studies about:

Eurasian autumn snowpack and Aleutian Low/Siberian High co-variability

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DEC 1



