Constructing an Ozone Long-Term Climate Data Set (1979 – 2010) from V8.6 SBUV/2 Profiles
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Introduction: Much of our understanding of ozone depletion issues is based on measurements and analyses of total ozone. However, to fully understand the processes of ozone depletion we must consider the effect throughout the ozone unit. Several high quality long-term ozone profile datasets exist and the ozone profile. The full record is provided by a series of satellites which must yield consistent results in order to provide a trend quality dataset. The recent release of Version 8.6 provides improved calibrations to ensure inter-satellite agreement.

Version 8.6 Changes:
- Climatology changed from SAGE based to Ozonezone in the Inoposphere and AURA/MLS in the stratosphere
- Cross sections changed from Cisl and Paur to Briton-Daumont-Malvet – to provide a more spectral resolution, extended wavelength range and better characterization of temperature dependence
- Includes OMI-based cloud-height climatology
- Updated calibrations including “no local time” calibration technique to cross calibrate between satellites instead of the previous reliance on SBUV

Time-of-day: Complicating the issue is the longevity of the NOAA satellites, which have a drifting equatorial crossing time (Figure 1). Profile ozone values change during the daylight hours at some levels producing a false trend in the long-term satellite record.

Figure 1: Equatorial crossing time of NOAA satellites. Note ascending and descending nodes have significantly different measurements times

Table 1: Satellite Comparison Dates

<table>
<thead>
<tr>
<th>Satellite</th>
<th>Date Range</th>
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<tbody>
<tr>
<td>NOAA 7</td>
<td>1/1/79 – 12/31/85</td>
</tr>
<tr>
<td>NOAA 11</td>
<td>1/1/84 – 12/31/93</td>
</tr>
<tr>
<td>NOAA 9</td>
<td>1/1/84 – 12/31/93</td>
</tr>
<tr>
<td>NOAA 11</td>
<td>1/1/84 – 9/30/10</td>
</tr>
<tr>
<td>NOAA 14</td>
<td>1/1/94 – 12/31/00</td>
</tr>
<tr>
<td>NOAA 16</td>
<td>1/1/94 – 12/31/00</td>
</tr>
</tbody>
</table>

Table 2: Satellite Periods

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<thead>
<tr>
<th>Period</th>
<th>Time</th>
<th>Node</th>
</tr>
</thead>
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<tr>
<td>1/1/79 – 12/31/79</td>
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<td>NOAA 9</td>
</tr>
<tr>
<td>1/1/84 – 12/31/93</td>
<td>Ascending</td>
<td>NOAA 11</td>
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<tr>
<td>1/1/84 – 9/30/10</td>
<td>Ascending</td>
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Figure 2: Correlation graph of Nimbus 7 and NOAA 11 at 10 hPa, 15S. Note the need beyond a simple bias correction.

Note that the overlap period for N11 to N16 is very short in time, therefore a bias only calculation is used. Also we adjust both ascending and descending N11 to Nimbus 7, and then adjust N14 to the N11 leg. This avoids propagating the non-physical trends in the N9 data. The final combined dataset uses a single satellite in each time period as shown in Table 2.

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Figure 3: N14-N11 period mean and std. dev. of the intercept (left), slope (center), and correlation coefficient (right) of the zones between 80S and 80N for each SBUV/2 pressure level between 50 hPa up to 0.5 hPa.

An examination of the means and standard deviation of each period intercepts, slopes, and correlation coefficients indicate that there is large variability in the vertical of how well the SBUV/2(2) data sets agree with each other during the overlap periods. Figure 3 shows these N14-N11 period statistics. Note that at 10 and 5 hPa there are large intercepts as well as deviations of slopes from unity. The correlations indicate that the best agreement occurs in the upper and lower parts of the profile and less agreement between 20 and 3 hPa. The standard deviation indicates that at some levels there is large variability with altitude. Figure 4 shows the intercept, slope and correlation coefficient for each zone and pressure level. Note the large variability in the intercept and slope in the Southern Hemisphere between 10 and 5 hPa. Also note the decline in correlation in the tropics in the low pressure levels.

Figure 4: N14-N11 period intercepts (left), slopes (center), and correlation coefficient (right) for each latitude zone between 80S and 80N for each SBUV/2 pressure level between 50 hPa up to 0.5 hPa.

Conclusions: We have shown that the combination technique can improve comparisons to SAGE II especially in the Northern Hemisphere. More work needs to be done for the N16 tie on, perhaps by using N14 as an intermediary. Comparisons with ground data are plentiful with mixed results. More statistics need to be considered to get a complete picture. Time-of-day corrections need to be added to resolve the drifting satellite local time effects.

References:

Sage II Comparisons: Comparisons are made to zonally averaged SAGE II data. Figure 5 shows the average differences for the periods defined in Table 2 for the V8.6 dataset before and after adjustments. There is moderate improvement in the intra-period consistency as compared to SAGE II in the equatorial and Northern Hemisphere regions. But at high-Southern latitudes the adjustment from N11 to N16 shows a large error in N16 and subsequently N17 as compared to SAGE II at 5 to 10 hPa.

Looking into the above successes and failures in more detail, we show in Figure 6 two sets of time series of the combined SBUV(2) product and SAGE II observations for the Eq at 2 hPa and 45S at 10 hPa. The adjustment at the former zone/pressure shows an improvement in the consistency of the SAGE II – SBUV/2 differences notably for the N16 and N17 periods. However, at the later zone/pressure the combination technique appears to fail for N14, N16 and N17. Also note that the differences are offset for the ascending and descending nodes of N11. This is likely a manifestation of the differing time of measurement of these two branches.

Figure 5: Average differences SAGE II – SBUV(2) in relevant periods before (left) and after (right) adjustment.

Ground instrument comparisons: A wealth of ozone profile data is available from the Network for the Detection of Atmospheric Composition Change (NDACC). Figures 7 shows comparisons to lidar and microwave instruments at Lauder, NZ, 49S, in the middle of the zone of concern. These comparisons are complicated in that we are now comparing zonal data with point data, and the station may show variability not exhibited in the zonal average. A better method would be to compare SBUV(2) profiles near the site. We can apply the above derived adjustments for the zone to these profiles, but this has not yet been done. In lieu of that we compare the point measurement to the zonal average. For the microwave data at 10 hPa, N16 is slightly too high before adjustment, but perhaps slightly too low after. N17, however, remains about right after adjustment. As compared to lidar at Lauder, N16 at 10 hPa seems appropriately adjusted.

Also shown are comparisons before and after adjustment to lidar at Hohenpeissenberg. In this case N14 is clearly better at most levels after adjustment, but again N16 is problematic. N17 is reasonable. In general areas can be found where the adjustment is an improvement, and where it fails. In particular, the N16 tie on must be more carefully examined.

Figure 6: Presentation of cases where adjustments to the SBUV(2) improved (i.e. made the bias more similar) the inter-satellite comparisons with SAGE II (Eq. 2 hPa) (right) and where the adjustments made comparisons with SAGE II worse (45S, 10 hPa) (left). Unadjusted SBUV(2) is on top and adjusted are on the bottom.

Figure 7: Comparisons of combined SBUV(2) with NDACC ground instruments: Lauder Lidar (left), Lauder Microwave (center), and Hohenpeissenberg (right). Unadjusted are on the top and adjusted are on the bottom.